Floods and Droughts in the Tulare Lake Basin
Second Edition

By John T. Austin

Figure 1. A dry pasture near Alpaugh, July 7, 2014.
Photograph by Matt Black

COVER PHOTO
Flood on the Kaweah River, January 2, 1997. Photograph by Tony Caprio

Sequoia Parks Conservancy
Three Rivers, CA 93271
Floods and Droughts in the Tulare Lake Basin

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Floods and Droughts in the Tulare Lake Basin

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The following report began as an effort to understand the hydrologic cycles of Sequoia and Kings Canyon National Parks, but it has turned into something even more important. The pages that follow provide priceless insights into an entire region: the Tulare Lake Basin of Central California. This distinct geographic zone contains not only the southern Sierra Nevada with its twin national parks but also major cities, a significant part of the richest agricultural area in the United States, and the bed of what was less than two centuries ago the largest freshwater lake in the western half of the conterminous United States.

Water, through both its presence and its absence, affects this region profoundly and in distinctive ways. The Tulare Lake Basin has a number of characteristics that make this particularly true. First, the region is close to the Pacific Ocean and thus well within range of intense oceanic storms. It has a Mediterranean climate which sees the great majority of its precipitation fall during the winter months of November through April. A further factor is that, because the region falls within the mid-thirties latitude range, it occupies the highly variable frontier between the wet winter climate of the Pacific Northwest and the often very dry winter climate of Southern California and northwestern Mexico. Adding more interest is the presence of the high-elevation terrain of the Sierra Nevada (including Mt. Whitney), which means that when the right kind of disturbances do arrive, extremely heavy precipitation can be extracted as the storms move eastward. And finally, the region’s major rivers — the Kings, Kaweah, Tule and Kern — flow into an interior basin rather than into the ocean. This last fact is true of no other significant area along the immediate Pacific Coast of the United States.

Assemble these characteristics and the complexity and significance of the region comes into focus. The Tulare Lake Basin has a highly variable climate, irregularly endures oceanic storms of ferocious magnitude, and (naturally at least) collects and holds all the runoff that occurs within its watershed. And, we must not forget, it is the home of large numbers of human beings and their institutions, everything from national parks to cities and corporate farms.

Author John Austin has approached the problems inherent in this report from the perspective of history: that is, he has sought out historical evidence from the many sources that document the Tulare Lake Basin’s highly variable patterns of flood and drought. As the pages that follow will document, he has thrown his metaphorical net wide, taking in everything from newspaper accounts to the literature regarding geological sediment cores. In doing so, the author has brought together a large and rich body of knowledge that has simply never been looked at before as a part of a single, unified pattern.

The value of such a unified perspective is immense. In many ways, we modern humans are just coming to know the Tulare Lake Basin. In less than two centuries we have settled the region and harnessed it to our needs. Yet, as this report so clearly demonstrates, we have not been here nearly long enough to know how the basin actually works. We have yet to experience either floods or droughts of the intensity found within the 2,000-year-long period documented. The report warns us how much we have yet to learn if we are to build a sustainable civilization within the basin.

Adding importance to this study is the accelerating presence of global climate change. The management policies of everything from cities and farms to national parks assume a “normal” world, one where averages can be defined and counted on. Discerning a core of climatic normality in a place as variable as the Tulare Lake Basin is no easy feat, and now we face the challenge that even such normality as we have known is inevitably evolving into something else. It is in this final context that this report adds yet more value by giving us a longer-term context in which to consider those things that will yet occur.

\textit{Floods and Droughts in the Tulare Lake Basin} is a report of significant long-term value to all who live in or care about this important region. In a way not seen before, it provides a historically powerful climatic overview of the region and how it works. It should be studied carefully by all who intend to manage lands or make their homes living within this dynamic region.

Wm. Tweed
Chief Park Naturalist, Sequoia and Kings Canyon National Parks
1996–2006
Purpose and Scope

Purpose

This document has a variety of purposes:

- To tell the story of water in the Tulare Lake Basin, to make it meaningful to the public. Why should residents of the Tulare Lake Basin care about the nearly 2,000-year history of the hydrology of this basin when considering modern day agriculture, dams, public health and safety, etc.?
- To provide a human dimension to the long-term climate record in an easily readable format. This document is meant to be read by the general public, not just by scientists and public land managers.
- To provide a context for understanding the predictions of the various climate models. Those models predict that the future will be different relative to the recent past. This document tells us what that past really looked like.
- To provide a single source for what is known about the history of floods and droughts within the Tulare Lake Basin. However, it is not intended to be a scientific treatise on that subject.
- To provide information so that the reader can better understand the risks that we face in preparing for future floods and droughts. To raise awareness of the seriousness of those risks.
- To provide context for understanding the link between storm precipitation and flooding.
- To provide a resource for interpreters and education specialists — to serve as a basic sourcebook for answering visitor questions as well as for building programs, exhibits, and other interpretive media.
- To provide a context for understanding and interpreting Sequoia and Kings Canyon National Parks’ collection of flood photographs. In this document, Sequoia and Kings Canyon National Parks are generally referred to simply as the “national parks.”

Scope

The intention of this document is to present the historical record of floods and droughts that have occurred within the Tulare Lake Basin over approximately the last 2,000 years. To the extent possible, this history is based on records specific to this basin. However, it has often been useful to include records from outside the basin for one or more of the following reasons:

- Records from within the basin are sometimes inadequate to describe a particular flood or drought, particularly in the early years of Euro-American settlement.
- Including records from outside the region is useful for major flood and drought events because those are larger than regional events. Examples of floods that affected an area much larger than the Tulare Lake Basin were the floods of 1861–62, 1916, 1938, 1964, and 1969.
- It’s useful for us to have an understanding of low-frequency events; that is, events that occur infrequently. For example, what does a 1,000-year storm look like? Or what happens when an 8-inch-per-hour storm hits a recently burned slope? By their nature, it takes a long time to observe such events in any given area, especially in a basin with as few gages and monitoring sites as the Tulare Lake Basin has. By looking beyond the boundaries of our basin, we can get a sense of what risks we might face in this time of climate change. Therefore, there are a number of 1,000-year events described in this document, and even some as rare as 300,000-year events.
Summary

Overview of the Document’s Contents
This document consists of two parts:

1. Overview and background material useful in understanding our history of floods and droughts, such as:
   - Maps of the Tulare Lake Basin and the adjacent basins.
   - Description and history of Tulare Lake and the neighboring lakes.
   - The types and causes of floods, and the terminology used to describe them.
   - Description of the federal reservoirs and how the conveyance structures work below the dams.
   - Summary graphs and tables showing runoff, floods, droughts, and temperatures.
   - Description of the different types of droughts.
   - Description of the consequences of using more water than we have.

2. A history of each of the floods and droughts, over approximately the last 2,000 years, for which we were able to find records.

Quick Start Guide
This document was intended to be read in order, starting at the beginning. That works for some readers, but others like to jump right into the history section, skipping the background material. If you like to jump right in, here is a suggested path for doing just that:

1) A good place to start is to look at the maps, especially Figure 2, Figure 4, and Figure 5. Then check out Figure 18 on page 111 to get a sense of how widely runoff varies from year to year in this highly variable climate.

2) Read a flood story from the modern era. A good one to start with involves Bobbie McDowall and her dad on the North Fork Kaweah. This drama took place late one night during the big November 1950 flood.

3) Then read about the huge 1861–62 flood. This is usually viewed as a Central Valley flood, but it was even bigger than that. The weather conditions during 1861–62 resulted in above-average precipitation between the Columbia River and the Mexican border. Major flooding was widespread throughout this area. The atmospheric mechanisms behind the storms of 1861–62 are unknown; however, the storms were likely the result of an intense atmospheric river, or a series of atmospheric rivers. Atmospheric rivers are relatively narrow regions in the atmosphere that are responsible for most of the horizontal transport of water vapor outside of the tropics. A strong atmospheric river can create major flooding when it makes landfall. The Tulare Lake Basin has had at least five large atmospheric river floods that we know of. Be sure to check out the link to Figure 30, where a satellite caught the remains of Super Typhoon Melor sitting over Japan, while simultaneously pummeling the Southern Sierra and Sequoia and Kings Canyon National Parks with an atmospheric river on October 14, 2009.

4) A representative drought to read about from the early pioneer days is the 1863–65 drought. It was severe, especially by the third year. There was no state or federal water system, so every rancher and farmer was on his own. The saving grace in many ways was Tulare Lake. Thanks to the 1861–62 flood, the lake was brimful when the drought set in. The lake served rather like a water hole on the Serengeti, albeit a 40-mile-long water hole. Vast herds of cattle would spread over the country for miles, traveling as far back from the lake as they could go without water in search of the scant grasses. Then they would rush back to the shore each day to quench their thirst.

5) Another good section to read is the one on California megafloods. These floods are even bigger than the 1861–62 flood and recur on a regular cycle of approximately 200 years. If they hold to their past schedule, the next one is expected to return within the next few decades.

6) Then go back to the beginning of the document and read any background material that interests you.

Using the Document Electronically
This document is all hot-linked from the Table of Contents, so there is no need to print it out, let alone read it in order. The body of the document also contains many hot-links. All figure and table references are hot-linked. There are also hot links to resources such as stream gage databases.
Key Findings about Runoff and Floods

The 121-year average runoff (1894–2014) of the four rivers in the Tulare Lake Basin (Kings, Kaweah, Tule, and Kern) is 2,941,237 acre-feet. It can be hard to grasp just how much water that is. The four federal reservoirs aren’t designed to capture the total runoff; that isn’t how they operate. Water continuously enters and is more or less continuously released from reservoirs. Ideally, reservoirs are drawn down most of the way before a flood, freeing up the flood-control pool.

The size of the reservoirs is useful for visualizing the volume of the runoff. The four reservoirs have a combined current capacity of 1,627,900 acre-feet. It’s useful to compare that capacity against the size of the historic runoff:

- The 1,627,900 acre-feet in combined current capacity can hold 55% of the 121-year average runoff (1894–2014) of the four rivers.
- The combined runoff of the four major rivers in the Tulare Lake Basin in 1983 was 8,746,222 acre-feet (see Table 83 and Figure 18). That is 5.4 times the combined current capacity of the federal reservoirs on those four rivers.

A look at Figure 18 on page 111 will show how widely runoff varies from year to year. Wet years and dry years commonly alternate, at least to some extent. The Tulare Lake Basin doesn’t have normal conditions in the sense of a statistical average. What is reliable about our climate is its extreme and relentless variability. That is our real normal; that is the lesson of Figure 18. Likewise, floods are amazingly commonplace in our area. A look at the Table of Contents or Figure 25 on page 163 will show just how commonplace; this document describes what we know about approximately 188 floods that have occurred during the last 2,000 years.

Floods occur at all manner of times. They occur in wet years, and they occur during multi-year droughts. They occur during the winter wet season, and they occur during the summer dry season. When they occur varies so widely because there are such a variety of causes for floods.

There is a surprising variety in what constitutes a flood (see the section of this document that describes the Causes of Flooding). This document contains a definition of what constitutes a drought, but it does not have an all-encompassing definition of a flood. That has proven too messy a concept to define.

Some of our floods are obvious: a river overflows its banks or a downpour overwhelms a city’s drainage system. At the other extreme, some of our floods have two components: hydrologic and socioeconomic. Society decides what its tolerance for natural processes is; that is, where to allow a river to flow. Thus some of our floods are the result of water appearing at the wrong place at the wrong time. They’re an inconvenience. For example, farmers wanted to drain Tulare Lake so that the lakebed could be used for agricultural purposes. They viewed Tulare Lake as an inconvenience, a nuisance to be prevented. Their viewpoint has prevailed. As a result, society has defined the presence of excess water in the lakebed as a flood. Water managers go to great efforts to minimize that type of flood.

Preparing for the Next Big Flood

Recalling our long-term flood history can be highly instructive. Historical information on floods can be used to prepare for taking future actions. That was the original impetus for preparing this document. One of the big lessons of this document is that our rivers have been relatively quiet of late. Table 30 on page 161 shows that the Tulare Lake Basin hasn’t experienced any 50-year floods or 100-year floods in over 40 years. The Kaweah and Tule Rivers haven’t even seen any 20-year or larger floods during that time period. This finding is based on the unimpaired flow of the rivers without factoring in the flood control provided by the reservoirs.

The fact that our rivers have been relatively quiet during the last 40 years probably doesn’t mean anything; it’s just a statistical coincidence. The problem is more psychological. We have become complacent. When we don’t experience a big flood for a while, we tend to forget just how big our floods can be. We have come to think of the federal reservoirs and our levees as protecting us from the effects of big floods, and that isn’t necessarily realistic when we consider our flood history.
The last really big flood in the Tulare Lake Basin was the December 1966 flood. It’s sobering to reflect back on the experience of that flood. Fifteen-foot waves were reported to have been common on the mainstem of the Kaweah in Three Rivers. Today we think of Dry Creek below Terminus Dam as not much more than a quiet foothills stream. However, in the 1966 flood, Dry Creek carried 44% more water than the Merced River in Yosemite Valley in the much more famous January 1997 flood.

The take-away message is that it would be prudent to prepare for big floods, floods much larger than we have been experiencing during the last 40 years. This is particularly important for those of us who live and work in areas that aren’t protected by a federal reservoir. We cannot assume that emergency plans have already been prepared by the county emergency agencies charged with planning for floods. Check just to be sure, ask to see the plans. If the county has an emergency plan, ask whether it includes a flood warning system, evacuation routes, etc.

People who live below the reservoirs tend to think that they’re safe, that the reservoirs are so big that they can catch and hold the floodwaters of the biggest events. Those reservoirs have been very effective at protecting downstream communities since their construction, but they do have their limits.

Authorized flood-control reservoirs are designed to provide a particular flood-control pool. That flood-control pool is used to store high inflows from a flood event so that flows downstream of a dam do not exceed the stated channel capacity. Hydrologists manage this flood-control pool to temporarily store the rain-flood runoff which would otherwise pass by a dam. The goal is to keep flows downstream of the dam within their stated channel capacity so that flooding conditions are avoided. The maximum size flood that a dam can control is termed a dam’s “level of flood protection.”

A reservoir’s level of flood protection can change over time. An example is Terminus Dam which forms Lake Kaweah. When originally constructed in 1962, Lake Kaweah’s storage capacity was estimated to be sufficient to provide a 60-year level of flood protection downstream. However, as sediment accumulated in the reservoir, the level of protection had decreased by 1978 to only a 46-year level of flood protection. When fuse gates increased the flood-control pool size of the reservoir in 2004, the level of protection increased to a 70-year level of flood protection (see Table 8). To learn more about the ins and outs of the level of flood protection, see the section of this document that describes Flood Rate and Flood-risk terminology.

In a large flood like the 1966 flood, reservoirs such as Lake Kaweah and Lake Success were pushed to their limit or beyond. In a huge flood like the 1867–68 flood, even greater flows would be passed downstream to the communities that sit below the reservoirs. The reservoirs were never designed to fully control floods of this magnitude. Sedimentation since construction has further reduced their flood control abilities.

Communities that sit below the reservoirs have to rely on levees for their fallback flood protection. And many of those levees have a history that dates back to the 19th century. Visalia is a prime example. At the time that Visalia was founded in 1852, the flow of the Kaweah River was distributed largely along the south side of the Kaweah Delta. That all changed thanks to the huge 1861–62 flood and the even bigger 1867–68 flood. One of the legacies of those floods was the creation of the St. Johns River which rerouted the majority of the Kaweah River floodwaters along the north side of the delta, to the north of Visalia.

It didn’t take Visalia long to erect a levee along the south bank of the newly formed river. That levee was built using material pulled up out of the river channel. That levee has failed numerous times since its initial construction, but none of the repairs corrected the levee’s significant structural shortcomings. The Visalia Chamber of Commerce hosted a thorough after-action review after a particularly bad levee failure that occurred during the 1945 flood. That review disclosed that the ability of the levee to protect Visalia was now significantly degraded, and it wasn’t obvious how the levee could be upgraded sufficiently to protect Visalia in the event of another flood of similar magnitude. The St. Johns River is subject to floods that are much larger than the 1945 flood.

Over six decades have passed since that after-action review, but there are still major concerns about the ability of the St. Johns channel to safely pass floodwaters. The 2005–06 Tulare County Civil Grand Jury investigated the St. Johns levee and found that it was not constructed to U.S. Army Corps of Engineers (USACE) certification standards, it was not being adequately maintained, and there was no adequate source of funds for its maintenance.
The Tulare County Resource Management Agency surveyed property owners in the levee district in 2002, but those owners were generally uninterested in levee maintenance and did not want to put more of their tax dollars into maintenance. Because the south-bank levee was in such bad shape, $17 million was then needed to bring it up to USACE certification standards. However, no source for those funds has yet been found, and the levee is in approximately the same shape now that it was in when the grand jury assessed the situation.

In June 2009, the Federal Emergency Management Agency (FEMA) found that the levee was in such bad condition that it provided essentially no reliable flood protection for Visalia. The USACE has also noted that many levees in our area were originally built to protect agricultural lands, but now protect urban development. As a result, they are under-designed for the purpose that they now serve.

The takeaway message is that, like the dams, we shouldn’t assume that all of our levees are designed, constructed, and maintained to provide protection from reasonably foreseeable floods. They are not. We have become complacent. We have not planned for flood events that are relatively common from the long-term perspective.

To help make historical knowledge applicable to future catastrophic events, the USGS Multi Hazards Demonstration Project (MHDP) applies science to improve the resiliency of communities in Southern California to a variety of major natural hazards. The MHDP assembled experts from a number of agencies to design a large, but scientifically plausible, hypothetical storm scenario that would provide emergency responders, resource managers, and the public a realistic assessment of what is historically possible. One of the MHDP’s full scenarios, called ARkStorm, addresses massive West Coast storms analogous to those that devastated California in 1861–62. This is a particularly reasonable assumption because storms of this magnitude are projected to become more frequent and intense as a result of climate change.

The ARkStorm scenario is patterned after the 1861–62 historical events. The ARkStorm scenario draws heat and moisture from the tropical Pacific, forming a series of atmospheric rivers that approach the ferocity of hurricanes and then slam into the West Coast over several weeks, resulting in large scale flooding. With the right alignment of conditions, a single intense atmospheric river hitting the Sierra east of Sacramento could bring devastation to the Central Valley. The U.S. Geological Survey (USGS) strongly urges risk management agencies to plan for the return of a flood as big as the 1861–62 event. That was the purpose of creating the ARkStorm scenario. The website for the USGS’s Multi Hazards Demonstration Project warns that an ARkStorm is plausible, perhaps inevitable. The 1861–62 storm was not a freak event and was not the last time that California will experience such a severe storm.

Building dams and levees is not the only way to prepare for big floods. Rivers need room to spill, and the first rule for surviving an irresistible force is to get out of the way. Preventing development in floodplains engineers out the risk; it gives rivers room to roam. The National Flood Insurance Program administered by FEMA tries to achieve that goal.

The California Legislature has given additional attention to this type of flood management in the Sacramento and San Joaquin River Basins. The Central Valley Flood Protection Board has real power. DWR is currently developing the 2017 Central Valley Flood Protection Plan as required under the Central Valley Flood Protection Act of 2005. That plan gives a lot of attention to lands subject to flooding. The plan also incorporates regional planning into the long-term flood-management effort. However, that planning area ends immediately north of the Tulare Lake Basin.

Preparing for the Next Big Landslide

Very large storm events occasionally result in not just floods, but also very large landslides. The ARkStorm scenario was based in large part on the 1861–62 flood. That was a huge flood in the Tulare Lake Basin. The force of the flood was so great that all four of our major rivers (Kings, Kaweah, Tule, and Kern) relocated and cut new channels. USGS has urged local risk management agencies to prepare for a return of a flood as big as the 1861–62 flood. We definitely do not at this time have plans in place for dealing with such an event.

However, the Tulare Lake Basin experienced a flood even bigger than the 1861–62 flood just six years later: the 1867–68 flood. It was a bigger flood on all four of our major rivers. In addition to being a major flood, the storm that brought on that flood was a deep soaking rain that lasted for upwards of six weeks. Hillslopes in the middle elevations of the Sierra are typically quite steep. Some are weathered in place, but many consist of unconsolidated colluvial debris slopes. When these hillslopes are soaked to depth, huge landslides can be triggered.
Floods and Droughts in the Tulare Lake Basin
Summary

In a short intense storm like the December 2010 storm, we get many relatively small landslides and debris flows. By contrast, in an extended event like the 1867–68 storm, the mid-elevation zone can experience cataclysmic landslides. In the past, some of those have formed landslide dams across our major rivers that were up to 400 feet high. When dams such as those fail, the results downstream can be catastrophic. For example, the residents of Bakersfield woke on New Year’s Day, 1868, to a 200-foot-high flood coming out of the Kern Canyon.

In the 1867–68 storm, the landslide dams on the Kaweah and Kern held the flooding rivers back long enough for the residents downstream to react and get out of the floodplain. In contrast, the landslide dams on the San Joaquin River and Mill Flat Creek presented a less clear signal downstream, partly because those events happened at night.

The residents of Old Kernville and Weldon had about 24 hours’ notice because the river stopped running. They were able to evacuate their towns before the Kern River submerged them under about 50 feet of water. The residents of Millerton, the county seat of Fresno at the time, weren’t aware of what was happening. The disintegrating remnants of one or more landslide dams hit their town just before midnight on Christmas Eve, 1867, destroying it. That is why Fresno is now the county seat of Fresno County.

Just as USGS is urging us to prepare for a return of a storm similar to the 1861–62 flood, it might be prudent to prepare for a return of a storm similar to the 1867–68 flood, complete with large landslides and landslide dams. This is especially true in high landslide hazard zones.

Preparing for the Next Megaflood

The website for the USGS’s Multi Hazards Demonstration Project warns that the 1861–62 storm was not a freak event, was not the last time that California will experience such a severe storm, and was not the worst case. The geologic record shows that six mega-storms more severe than the 1861–62 flood have struck California in the last 1,800 years. There is absolutely no reason to believe that similar events won’t occur again.

As detailed in Table 1, a huge flood strikes Southern California approximately every 200 years (for more information on these floods, see the section of this document that describes the California megafloods). The flood presumably strikes Northern and Central California as well, but the research has largely been done in Southern California. That storm cycle appears to be associated with the roughly 208-year cycle of solar activity (the Suess Cycle). The result seems to be that as solar activity decreases, the climate cools; and this shifts the prevailing wind patterns and associated storm tracks toward the Equator. Whatever the mechanism for these megafloods, the most complete record that we have of their occurrence is from the Santa Barbara Basin off Southern California.

<table>
<thead>
<tr>
<th>Approximate Date of Flood</th>
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<tbody>
<tr>
<td>212</td>
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<tr>
<td>440</td>
</tr>
<tr>
<td>603</td>
</tr>
<tr>
<td>1029</td>
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<td>1418</td>
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<td>1605</td>
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The bicentennial flooding in Southern California was skipped only three times since 212 and never twice in a row. The last skip was in the early 1800s, leading researchers in the Santa Barbara Basin to conclude that we foresee the possibility for a historically unprecedented flooding in Southern California during the first half of the 21st century.

The skip in the early 1800s may have only been a skip from the local perspective of the Santa Barbara Basin. The Central Valley experienced a huge flood in 1805, one that was even bigger than the huge 1861–62 flood. Perhaps the storm track that year just didn’t extend far enough south to be recorded in the Santa Barbara Basin.

Based on past experience, we can expect floods resulting from huge storms like those of 1861–62 and 1867–68 to last for up to three months. That is very serious, and we certainly aren’t prepared at this time for such events. However, they’re similar to the type of floods that we have experienced in more recent times, just much bigger.
Judging from what we know of the 1605 flood, megafloods appear to be a much longer-term type of event. Once such an event begins, it can last for up to 10 years. During that period, multiple episodes of flooding and extreme runoff can occur as well as other unusual climatic events. The 1605 flood was associated with a large-scale change in climate that affected the Northern Hemisphere from roughly 1600–1610.

The decade 1600–1609 stands out as the coldest in a 570-year (A.D. 1400–1970) comprehensive record of summer temperatures across the Northern Hemisphere, based on tree-ring and ice-core data. The years 1601 and 1605 produced unusually narrow tree-rings in the Sierra, suggesting very cold growing seasons.

The 16-year period from 1597–1613 in the Sacramento River Basin had the maximum reconstructed riverflow for the 420-year (1560–1980) time period. There was a major flood on several Northern California rivers, including the Salmon and Klamath, in about 1600. Those rivers wouldn’t see another flood that big until the December 1964 flood.

The year 1602 was the end of a 37-year drought (1566–1602) megadrought in the Southern Sierra. Mono Lake rose to elevation 6456, the highest level of the past millennium, around 1650. The Mojave River terminates at the Silver Lake playa in the Mojave Desert. (That playa is located along Interstate 15, just north of the town of Baker.) At very infrequent times, the Mojave River delivers so much water to the playa that it forms a perennial lake. The last time that this happened was likely during approximately the 1600–1610 time period.

Along the coast of California in the Santa Barbara area, 1604 was the fourth wettest year, and 1601–1611 was the third-wettest 11-year period in a 620-year (A.D. 1366–1985) reconstruction of precipitation. There was a major flood on the Santa Ana River in Orange County in about 1600. That river wouldn’t see another flood that big until the January 22, 1862 flood. Severe flooding occurred around Mexico City in 1604 and 1607.

It’s tempting to think that there would also have been floods in the Tulare Lake Basin during the 1600–1610 time period since the regions to our north, east, and south were experiencing immense precipitation events at that time. The Tulare Lake Basin would have been under the same general storm tracks. However, we haven’t found records of such floods.

If the mega-storms hold to their past schedule, the next one is expected to return within the next few decades. It may or may not be prudent to prepare for a return of a mega-storm similar to the 1605 or 1805 events. That is a question for risk managers to decide. The good news is that by preparing for events such as a return of the 1861–62 or 1867–68 storms, we’d be in a much better position to deal with a mega-storm should it materialize.

Key Findings about Droughts
This document describes what we know about 36 multi-year droughts. One of the findings is that it’s surprisingly common for floods to occur during droughts. That seems counter-intuitive, but it happens repeatedly.

The term “drought” is commonly used in two different ways in the Tulare Lake Basin: meteorological and socioeconomic. Traditionally, the term “meteorological drought” has been used to reflect precipitation that is significantly less than average. That is generally how the term drought is used in this document — a period of significantly less than average precipitation. Typically this condition has to last for at least two years before it is recognized as a drought. This type of drought ends when precipitation returns to approximately average or above-average conditions. The 1976–77 drought is an example of this type of drought.

Socioeconomic drought is when the available water supply fails to meet our needs (the amount of water we choose to apply). One could also say that our society’s water use has imposed a socioeconomic drought on the environment. For example, diversions from the Delta have imposed critical drought conditions (measured by alteration of unimpaired runoff out of the Delta) in more than 50% of the years or 80% of the years just on the San Joaquin River inflow to the Delta. The result of this environmental drought, which in turn is partially responsible for the decline in abundance of critical fish species, then leads to cutbacks in Delta exports or diversions from the San Joaquin River which in turn exacerbates the socioeconomic drought described herein. This category of drought often has a political component to it. It is used to describe conditions when, for various reasons, we don’t have all the water that we feel we are entitled to. In this kind of drought, 1) available water supply fails to meet our needs (the amount of water we choose to apply), and 2) we perceive that water that is rightfully ours is being used somewhere else. If “they” would just let us have our water, we would be better able to meet our needs. The Tulare Lake Basin has had a lot of the socioeconomic type of drought in recent years.
Floods and Droughts in the Tulare Lake Basin

Summary

That results in large part because our basin relies on a great deal of supplemental water imported from the San Joaquin River (via the Friant-Kern Canal) and from the Delta via the state and federal canals. Reduced imported supplies can stimulate a socioeconomic drought even when precipitation and runoff in the Tulare Lake Basin is not in a meteorological drought. It is more challenging to mark the end of this type of drought. Precipitation can return to average or even above-average conditions, but there still isn’t enough water to meet our needs (the amount of water we choose to apply). The latter part of the 2007–09 drought is an example of this type of drought.

In the 2007–09 drought, water years 2007 and 2008 were meteorological drought years. The runoff was so low in those years that the state’s water year index rated those years as critically dry. The 2007–09 drought was California’s first drought for which a statewide proclamation of drought emergency was issued.

That turned out to be critical. When precipitation returned to near-average or above-average, it was hard politically for the governor to declare an end to the drought. There clearly wasn’t enough water to go around. It wasn’t until March 30, 2011, after an incredibly wet winter, that the state of emergency was finally rescinded. That was long after the end of the hydrologic drought. The 2007–09 drought had morphed from the traditional type of drought (below-average precipitation) into the socioeconomic type (we’re entitled to more water than we’re getting).

When surface supplies are inadequate to meet our needs (the amount of water we choose to apply), we pump out of the groundwater aquifer. When surface supplies allow, water districts work to recharge the groundwater aquifer. In recent years, it has become increasingly apparent that we are withdrawing more than we are returning, we have a groundwater overdraft. Our demand for water is dramatically higher than the surface supply, so we are mining the underground supply. By that definition, we are effectively in drought conditions most of the time, even when precipitation is above average.

The groundwater supplies of the San Joaquin Valley are being depleted by an average of over 2.8 million acre-feet per year. For perspective, that 2.8 million acre-feet overdraft is nearly as great (80%) as the combined annual flow of the two largest rivers in the southern San Joaquin Valley. The San Joaquin River produces a long-term average annual flow (measured at Friant Dam) of 1.8 million acre-feet per year. The Kings River has an average annual flow (measured at Pine Flat Dam) of about 1.7 million acre-feet. Together, these two rivers produce an average of 3.5 million acre-feet of water.

In the Tulare Lake Basin, water for agriculture, cities, rivers, wildlife refuges, etc. comes from three sources:

1. The most sustainable and local of these three is the Sierra, mainly in the form of water from the Tuolumne, Merced, San Joaquin, Kings, Kaweah, and Kern Rivers.
2. The second source of our water is exports from the Sacramento–San Joaquin Delta, which is the immediate source of water that we import from Northern California via the Delta-Mendota Canal and California Aqueduct. That water, moved south at considerable expense, is increasingly fought over and hard to get. See the section of this document on the Role of the Endangered Species Act in Reducing Delta Exports for a discussion of some of the issues being fought over that limit our ability to increase Delta water exports.
3. The third source is what we pump from the groundwater aquifer, much of which is never replaced. A century ago, much of the valley had groundwater almost to the surface; artesian wells were common. Now, many areas have been mined for water to a depth of several hundred feet. The groundwater situation in Tulare County is particularly well studied and understood.

We consistently, decade after decade, use more water than our available surface supplies. The Tulare Lake Basin has by far the largest groundwater overdraft of any region in the state. The implications of this are hard to escape. Even if our temperatures and snowpack were to remain stable, the groundwater table in the Tulare Lake Basin will continue to drop due to the overdraft. One of the consequences of this is that the wells will continue to go ever deeper and the cost of pumping will continue to rise. This race to the bottom is not sustainable. Eventually the pumping will be limited by supply and demand. Agriculture (the valley’s single biggest water user) will be forced to reduce its reliance on groundwater.

The consequences of overdrawing the groundwater aquifer aren’t just financial. Subsidence in the San Joaquin Valley is one of the great changes that human activity has imposed on the environment. The San Joaquin Valley has the largest subsidence in the world due to groundwater withdrawal. (Although this was the case in the 1970s, it is conceivable that some other area elsewhere in the world may have surpassed us since.) By 1977, the subsidence had reached a maximum of 29.6 feet vertically and more than 5,200 square miles of irrigable land; one-half the entire valley floor (10,000 square miles) had been affected. Further significant subsidence
has occurred since 1977, and some areas are currently subsiding at upwards of one foot per year. Subsidence rates have accelerated in recent years because of droughts and increased reliance on groundwater.

With such a huge shortage of water in the Tulare Lake Basin, it would seem that we would be anxious to hold onto all the water that we could. However, it turns out that in wet years, we go to considerable efforts to move water out of the basin. This is done in order to keep the rivers from following their natural course back to the Buena Vista and Tulare Lakebeds.

As discussed above, we as a society have decided that we would rather use those lakebeds for growing crops than for storing water as the natural reservoirs that they were until the late nineteenth century. At the time of Euro-American arrival in the region (1840s), the Kings was flowing down the south side of its delta and into Tulare Lake. Now we have constructed waterworks so that it is possible to divert part of the Kings River flow north through the San Joaquin River into San Francisco Bay.

As described in the section of this document on Pine Flat Dam, the first 4,750 cfs of flood release water from that reservoir is directed through the North Fork / Fresno Slough / James Bypass channel to the San Joaquin River and, ultimately, San Francisco Bay. It keeps that water out of Tulare Lake, but it is essentially a loss from the point of view of Tulare Lake Basin water users. With the construction of Pine Flat Dam in 1954, the need to divert water through this system was greatly reduced. Even so, diversions through this system have occurred in 38% of the years since the dam was completed. Likewise, the Kern naturally flowed into Buena Vista Lake and then overflowed into Tulare Lake. Now it is possible to divert part of the Kaweah, Tule, and Kern River floodwaters and send them to the Los Angeles area.

Runoff is much larger in some years than in others, resulting in a greater need to export water out of the basin. For example, because of a huge snowpack, the runoff in 1983 was the largest since record-keeping began in 1894. As a result, a record 3.1 million acre-feet was exported from the Tulare Lake Basin that year. That was almost as much as the 3.5 million acre-feet combined average annual flow of the San Joaquin and Kings Rivers. That isn't to say that exports (or inter-basin transfers) are necessarily a bad thing; just that we need to recognize that diversions have consequences to the groundwater aquifer. Every time that water is transferred out of the Tulare Lake Basin, there is that much less water available for use or for groundwater recharge.

One of the big lessons learned from preparing this document is that our historic droughts haven't been nearly as bad as some that occurred in California prior to settlement. The 17-year-long 1918–34 drought was arguably the longest and most severe drought to strike the Tulare Lake Basin during historic times. It was one of only three megadroughts to occur in our basin since the Little Ice Age began in about 1450 (see Table 19).

However, California has experienced two prolonged, epic droughts (see the section of this document that describes Megadroughts before the Little Ice Age on page 164). The first of these droughts is thought to have lasted 243 years from AD 832–1074; the second drought lasted 178 years from AD 1122–1299. Evidence of those megadroughts is still surprisingly visible in places like Yosemite. The takeaway message is that what we think of as a long-term drought today is mild compared to these earlier megadroughts.

Studies show that although our precipitation is very variable, there is no long-term pattern to our mixture of wet years and dry years. In addition, California's average precipitation has been relatively stable for more than a century; it has not been declining. What has been changing is an increase in temperature, especially in recent decades.

Severe droughts have occurred approximately twice as often in California in the past two decades as in the preceding century. Why? Most severe droughts have occurred when conditions were both dry (precipitation less than the long-term average) and warm (temperature above the long-term average). Similarly, dry years were much more likely to produce a severe drought if they occurred in warm years. Years that were both warm and dry were about twice as likely to produce a severe drought as years that were cooler than average and dry.

The probability of dry years occurring during a warm year has been greater in the past two decades than in the preceding century. There has been more than a doubling of the frequency of warm-dry years in California. Because of increasing temperatures, the probability of dry years producing severe droughts has been approximately twice as great in the past two decades as in the preceding century.

More area in the Western U.S. has persistently been in drought during the 15-year period from 2000–14 than in any other 15-year period in more than 850 years, since the 1150s and 1160s. The year 2015 will almost certainly be a drought year, extending this to a 16-year period. The Western U.S., taken as a whole, has been
Floods and Droughts in the Tulare Lake Basin

Summary

in drought for the 16-year period 2000–15. However, the specific area affected by the drought has moved around each year. We have been on the edge of the drought and have only been affected by it for 12 of those 16 years. From our perspective, this means we haven’t been in a megadrought. Instead, we have experienced three droughts of relatively average duration: the 2000–04, 2007–09, and the 2012–15+ droughts. However, if we were to step back and look at the bigger picture, we are really on the edge of a record-setting megadrought.

What makes the 2012–15+ drought so severe is that it has consisted of four years (so far) of well-below-average precipitation coupled with well-above-average temperatures. The combination of very low precipitation and high temperatures has resulted in record low drought stress (PDSI), creating unprecedented stress on vegetation. It is possible, perhaps even probable, that the effects of the current drought on native vegetation may be unprecedented in at least the last century.

The Effect of Floods on Tulare Lake

For various reasons, there is no longer a lake in the Tulare Lakebed, at least in most years. For an explanation of those reasons, see the section of this document: Why is there no lake in the Tulare Lakebed today?

The story of Wildlife in and around Tulare Lake is summarized in that section of this document. The lake and its associated wetlands used to provide very valuable habitat for a wide variety of wildlife. In addition, the lake and wetlands provided biological connections among the various river and stream courses in the Tulare Lake Basin. It is hard for those of us living in the 21st century to wrap our minds around what that 19th-century ecosystem was like; it was truly extraordinary.

One of the lessons learned in preparing this document was that the flood cycle of the Tulare Lake Basin was critical in maintaining that ecosystem; the floods provided sufficient water storage to keep the lake going through the drought or non-flood years (see Figure 15 and the section of this document on the Role of Floods in Maintaining Tulare Lake).

Once Tulare Lake (and the other four valley lakes) had been dried up, disintegration of this remarkably complex system was sealed with the damming of the four main rivers. The functioning infrastructure of this formerly biodiverse ecoregion was so badly broken that it resulted in the loss of most of the wetland habitat and nearly all of the biological connectivity between the watersheds in the high country and the lowland floodplains. Water-dependent habitats on the adjacent land, particularly on the Kaweah Delta and other riparian corridors, were also significantly degraded during the ensuing decades.

The loss of the Tulare Lake ecosystem affected even protected areas like the national parks. For example, we speculate that the relict populations of beavers, mink, and river otters that hung on in the national parks became isolated from populations elsewhere in the parks and the basin. Their numbers gradually declined, and some of those species may now be extinct in the parks. We speculate that a similar problem occurred with many populations of birds, fish, and other animals.

Despite the best efforts of water managers, floodwaters still make it to the Tulare Lakebed on occasion, especially in heavy runoff years. Migratory waterfowl and shorebirds still return in large numbers when Tulare Lake has a major flood. Very large numbers returned when the lake had its last great reappearance in 1937–46. There was a significant but smaller appearance of waterfowl and shorebirds during the floods of 1982–83 and 1997. While exciting to see, these bird congregations are ephemeral and move on once the floodwaters recede.

Floods — with the water that they brought — created a marvelous ecosystem in the Tulare Lake Basin. Reminders of that ecosystem survive in disjointed preserves in the valley, in the foothills, and in the Sierra. The framework of the hydrologic system that powered that ecosystem still exists today. On occasion, flooding can recreate a portion of Tulare Lake. However, just adding water to that lakebed is not enough to recreate the complex ecosystem that once existed. The associated habitat is highly degraded, and the ability of the ecosystem to provide connections among the various river and stream courses in the Tulare Lake Basin has largely been lost.

Conclusion

The Tulare Lake Basin has a fascinating hydrologic history. We can and should learn from that history. Society increasingly relies on a stability of climate, but that is not supported by the last 2,000 years of history in the area covered by this study. What is reliable about our climate is its extreme and relentless variability. We should not be complacent about floods and droughts. We need to learn from our past as we plan for the future. The author hopes you enjoy and learn from this document.
Acknowledgements

We owe a debt of gratitude to all the National Park Service (NPS) employees who have recorded their observations in the superintendent’s monthly and annual reports over the years. Ward Eldredge was very helpful in going through the parks’ archives to provide access to those old reports.

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Tony Caprio, George Durkee, Annie Esperanza, Dave Graber, Linda Mutch, and other National Park employees provided insights on Sierra climate change and fire history. Karen Folger, Anne Birkholz, and Tony Caprio researched fire histories in the national parks and throughout the Tulare Lake Basin. Jon Keeley informed and reviewed the section on fires and drought.

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About 97% of Sequoia and Kings Canyon National Parks is designated wilderness. We know about the storms and floods in that part of the parks in large part because of the wilderness rangers who live and work there during the summer. Wilderness rangers Dave Alexander, George Durkee, Erika Jostad, Bob Meadows, and Cindy Wood all contributed to this document. Many others contributed by recording their observations each year. The meadow monitoring staff, especially Erik Frenzel, have become an important go-to source for trail conditions in the wilderness. Trail crews and their supervisors Tyler Johnson and David Karplus have been another valuable source of information on storms and floods in the wilderness.

Several retired national park employees have contributed from their years of knowledge. Among those were Manuel Andrade, Leroy Maloy, and Harold Werner. Steve Moffit, Sequoia National Park’s former trails foreman, was particularly helpful.

Thank goodness for newspapers. Those have proved an invaluable source for information on past floods and droughts in the Tulare Lake Basin. The Visalia Times-Delta and its predecessors were particularly helpful. Several newspapers put together specials on particular floods. For example, John and Sarah Elliot collected stories of the 1955 flood and published those in The Kaweah Commonwealth. Dody and Tom Marshall at the Three Rivers Historical Museum were helpful in providing access to two old issues of the Three Rivers Current in the museum’s archives. The Exeter Sun produced a special edition on the 1955 flood.

Los Tulares, the quarterly bulletin of the Tulare County Historical Society, was a treasure trove of information.

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Joyce Fernandez in the Sacramento Office of the U.S. Bureau of Reclamation (USBR) found what appears to be the last surviving copy of a key Tulare Lake document. She spent many hours scanning that document so that others could benefit from the information in it.

Josh Courter, the Western District Divide hydrologist on the Sequoia National Forest, assisted with analysis of the debris flows and landslides on the Kern River. Fletcher Linton, the forest botanist for the Sequoia National Forest also contributed to that analysis. In addition to being a talented field botanist, Fletcher has skills as a geologist. Jim Roche, the hydrologist at Yosemite National Park, provided insights on flooding and erosion on the Merced River.
Floods and Droughts in the Tulare Lake Basin

Acknowledgements

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Michael Dettinger (USGS/Scripps) is one of the country’s leading atmospheric river researchers. He gave his time to consult on atmospheric rivers and the 1905–06 snowpack. Arndt Schimmelmann (Indiana University) is one of the country’s leading paleo-climatologists. He gave his time to review the California megaflood portion of this document. Toby Ault (Cornell University) is one of the country’s leading climate scientists. He reviewed the section on the potential for future megadroughts.

Sarge Green has many years of experience working with agencies such as the California Regional Water Quality Control Board, resource conservation districts, and irrigation districts. He is currently program director of the California Water Institute at CSU Fresno. Sarge was the perfect person to review the Groundwater Overdraft section of this document.

Peter Vorster has nearly 40 years of experience as a hydrogeographer, much of it focused on water resource management in the Central Valley, Eastern Sierra, and Southern California. He has been at The Bay Institute since 1996, and is a senior advisor to the California Water Foundation on their Sustainable Water Management Profile project. Peter has a particular expertise in the Tulare Lake Basin hydrology. He was the lead author of the 2007 report, Tulare Lake Basin Hydrology and Hydrography: A Summary of the Movement of Water and Aquatic Species. Peter was an excellent reviewer; he really understands the hydrology of our basin, including its history.

John Shelton used to work with the Department of Water Resources (DWR) and he has wide knowledge of water resource issues. He is currently a senior environmental scientist with the California Department of Fish and Wildlife. John shared information and provided advice on several issues related to water use. John seems to live in two worlds; he is an environmental scientist who understands the system from a water user’s point of view. He is always worth listening to; he has the most perceptive insights. They help to connect the dots, explain how the system works, explain why water users do what they do, and explain why the ecosystem behaves the way that it does.

Eric Osterling works with the Kings River Conservation District. He is very knowledgeable about water issues and was so helpful in preparing the sections on water demand and water rights. Jim Sullins, the UC director for Cooperative Extension for Kings and Tulare Counties, helped straighten out the section on water supply.

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Mark Tilchen and Valerie McKay put together the very relevant book: *Floods of the Kaweah*. Denis Kearns contributed an out-of-print book that the State Disaster Office prepared on the 1955 flood. Bill Templin (a former USGS hydrologist) was instrumental in documenting the flood and debris flow that occurred on Lewis Creek in 2008, citizen science at work. Bill also knew the history of stream gages on the South Fork Kings and was very helpful in providing access to that data.

Two histories of Three Rivers were recently published. Sophie Britten wrote *Pioneers in Paradise: A Historical and Biographical Record of Early Days in Three Rivers, California 1850s to 1950s.*¹ Earl McKee, Jr. wrote *Echoes of Blossom Peak: Cowboys, Horsemen, and History of Three Rivers.*² Those books provided lots of good information about how floods have affected the community of Three Rivers.

Rob Hansen was generous in sharing his treasure trove of information about Tulare Lake and the San Joaquin Valley. Terry Ommen was incredibly helpful in providing historical accounts of floods in Tulare County. I really couldn’t have pulled this document together without the support of people like Rob and Terry who have spent years collecting all that primary source material and then entrusting that to my care.

Tulare Basin Wildlife Partners was supportive of this document in many ways. Several of the people active in that organization (particularly Rob Hansen, Carole Combs, Sarah Campe, and Niki Woodard) contributed material for the document, reviewed it, and generally provided encouragement. In addition, they were the first organization to post the electronic version of this document, doing so through their Tulare Basin Watershed Initiative (TBWI). Niki was the webmaster for the TBWI. She also volunteered her considerable design skills to laying out the front and back covers of both the first and second editions of this document.

I am honored that Matt Black donated his beautiful photographs for use in the second edition of this document.

More than 65 people contributed their time to reviewing the first edition of this document, far more than can be credited here. Julie Allen and Anne Birkholz were both instrumental in beating the Summary section into shape. Karen Folger reviewed the entire document; she was a great reviewer. Anne Birkholz reviewed numerous sections and versions of both the first and second edition; I was fortunate to have such competent friends. More than 35 people reviewed the various updates that were made to the second edition.

My sister Karen Austin was tenacious in reviewing the entire document, checking grammar, punctuation, content, citations, the whole shebang. Karen was awesome as an editor, and you couldn’t ask for a finer sister. I only wish that she had lived to see this document printed.

Many people reviewed particular floods. Harold Werner, the national parks’ retired wildlife ecologist, reviewed the wildlife section. Jane Allen and Dave Humphrey reviewed the section on American Indians. Sarah Campe pointed out the need to write the document (well, rewrite it) so that it could be understood by those who didn’t have a national park background. She also provided insightful comments on the water demand and supply sections.

Tony Caprio was an insightful reviewer. Tony also had the skill set to fix errors detected in maps that were acquired from historic and other secondary sources. That was very handy.

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Charisse Sydoriak reviewed the entire document, looking at it from the perspective of an NPS manager. Colleen Bathe also reviewed the entire document, looking at it from the perspective of a Sequoia Natural History Association manager.

Dave Graber, retired NPS regional chief scientist, reviewed various sections of this document, especially those dealing with wildlife and ecological relationships. Dave makes a great reviewer; he knows his subject matter, and he quickly finds the weakness in an argument. In addition to being a reviewer, Dave provided guidance and encouragement in the writing and publication of this document.

Bob Meadows is a walking encyclopedia of weather knowledge on the national parks. He was especially helpful in reviewing this document and straightening out long-festering data errors.

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Terry Ommen reviewed the entire document, looking at it through the lens of an experienced writer. Rob Hansen was a spectacular reviewer; he was particularly critical in reviewing the Tulare Lake sections. Nobody knows Tulare Lake and its history like Rob does.

Bill Tweed, the national parks’ chief park naturalist from 1996–2006, has been a singular inspiration for this document; he built the foundation for it in so many ways. In many years, he was the force that motivated the park staff to produce the superintendent’s annual report. He has researched numerous papers on weather and climate over the years (several of which are shamelessly plagiarized in this document). Bill was a reviewer and consultant on many sections of this document. And most of all, he has preached the need to be aware of the world and the weather around us.

And lest I forget, my wife, Shauna Austin, put up with my spending over five years of evenings and weekends working on this document. I so appreciate her tolerance and forbearance during that period. I could not have prepared this document without her support.

Many individuals and agencies have assisted with the preparation of this document. Including their names in this Acknowledgement section does not imply that they have reviewed this entire document or blessed it. Furthermore, any errors that remain in this document are the sole responsibility of the author.
Background Material

Human Perspective
By its nature, this document illustrates floods and droughts as seen through the lens of humans. Their stories reflect the human dimension, what people thought was worth recording.

Furthermore, the occurrence of floods has changed with infrastructure (dams, levee maintenance, more flood-resistant bridges, etc.). For example, as dams were constructed, flooding became less frequent downstream. When levees weren’t maintained (or when a conduit became plugged), flooding could become much worse even though the hydrology didn’t change.

Peer Review Process
After long consideration, the National Park Service declined to conduct a formal peer review of this document. However, the first edition of this document did undergo an extended period of informal peer review, lasting from January 2011 through October 2012. During that period, it was reviewed, in part or whole, by nearly 70 individuals and agencies. The reviewers were selected based on interest and expertise. Some of the reviewers were directly involved in the collection, analysis, or reporting of some of the information used in this document. The author served as the peer review manager and maintains the records of the review process. The second edition went through a similar review process, lasting from November 2014 through April 2015. Approximately 35 individuals reviewed the second edition, in part or whole. This review focused primarily on new material added since the first edition.

Citation of Source Material
This document was designed to tell the story of water in the Tulare Lake Basin. Our basin has experienced a lot of floods and droughts during the last 2,000 years. A rich trove of stories and data has survived to tell the story of those events. The Literature Cited section of this document contains over 1,800 source citations documenting those stories and data.

A National Park Service reviewer observed that there are “uncounted source citations” missing from this document. While the comment may be a bit melodramatic, that does reflect on the intentional design of this document. It was not meant to be a scientific treatise of the subject matter. (See the Purpose section of this document.)

If this had been a scientific treatise, it would have been necessary to strictly adhere to the rule of only stating what could be backed directly by citation. While theoretically possible, that would have been an editorial challenge. This document is dense with facts. In addition to 30 figures and 105 tables, it contains 3,000 or so paragraphs.

Many of those paragraphs were composed using information from multiple sources. Some paragraphs have four or more sources. The challenge of citing those sources each time that they were used would have been made more difficult because they are so intertwined. The same source publication can be the source for dozens of sentences and tables that are scattered throughout this document. Individually calling out those citations would have made the document significantly less readable, so that generally wasn’t done.

Facts that can readily be found by searching the Internet have not always been cited. However, care has generally been taken to cite the hundreds of obscure sources that cannot be found with a simple online search.

Facts that are contained within the records management and archive systems of Sequoia and Kings Canyon National Parks have not been formally cited; these have only been identified by informal reference within the body of the document. Personal communications have also not been formally cited; these have only been identified by informal reference.

In order to avoid interfering with the flow of the document, the formal citations are contained as endnotes in a Literature Cited section at the end of the document rather than as footnotes.

Reliability of Source Material
This document uses a mixture of hard data and personal observations to tell the story of floods and droughts in the Tulare Lake Basin. Hard data alone would be insufficient. In part, that is a reflection of the nature of the
document’s design. (See the Purpose section of this document.) Personal observations are necessary to provide the human dimension of the story.

The heavy reliance on personal observations also reflects that there is limited data available to tell the complete story of floods and droughts in our basin, especially as we look further back in our history. In a general sense, data can be used to give the general outline of what has happened. Then the story has to be filled in by the observations of people who were there to witness the events.

However, this brings up a problem: Personal observations, even those of sworn eye witnesses, are not as reliable as those of machines. Most of the time people probably get most of their stories right, but sometimes they get parts of their stories wrong.

So the reader is cautioned to take these stories with a grain of salt. Wherever possible, stories used in this document have been cross-checked against other stories and available data.

Disclaimer Regarding Subject Matter Expertise

This is not meant to be a technical document or scientific treatise. (See the Purpose section of this document.) The author is neither a trained historian nor a physical scientist, and is not recognized by any entity as such. To use NPS jargon, this work cannot be ascribed to “a professional in the field of inquiry.”

Note about Completeness

This is not intended to be a complete document. By its design, it is a collection of the documentation that we have been able to find, brought together in a single place. This is meant to be a reference source for others to turn to.

There is a lot more source material out there. You could make a life’s work out of this; it’s hard to know when to stop tracking down loose ends. As Oscar Wilde said: books are never finished, they are merely abandoned.

But we now have a moderately complete listing of the major floods and droughts that have occurred in the Tulare Lake Basin since about 1850. We also know some of the big floods and droughts back as far as about the year A.D. 212. In general, the farther back that you look, the bigger the flood or drought has to be in order to be detectable.

National Park Service Involvement

This document was not commissioned or authorized by the National Park Service. It was researched and written as a volunteer effort by the author, primarily on weekends and evenings.

However, Sequoia and Kings National Parks did support the preparation of this document in various significant ways. The national parks provided full access to its historical files; that proved invaluable. Following NPS national policy, the national parks also allowed incidental use of government office equipment (copiers, printers, and computers) during non-duty hours.

This document is focused on Sequoia and Kings Canyon National Parks; as shown in Figure 4, those parks occupy the headwaters of the Tulare Lake Basin. The National Park Service considered publishing this document in its technical report series. However, that proved to be infeasible. The primary obstacles were:

- This document was designed as more of a historical than a traditional technical paper. (See the Purpose section of this document.) That made it less than an ideal fit for the NPS technical report series.
- If this document were a traditional technical paper to be included in the NPS technical report series, it would have to strictly adhere to the rule of only stating what can be backed directly by citation. However, this document intentionally does not cite all of the sources that it is based on. (See the Citation of Source Material section of this document.)
- This document identifies some of the flood risks in the Tulare Lake Basin that are not mitigated by the four federal reservoirs. This is a factual condition that is based on facts and has been thoroughly reviewed. However, if the NPS were to publish this document, that might be mistaken as implying that the national parks thought that local authorities should be acting to mitigate those risks. That would have the appearance of the parks criticizing neighboring agencies, something that generally isn’t done.

Therefore, the National Park Service chose not to publish or endorse this document.
Maps of the Tulare Lake Basin

Figure 2 illustrates California’s various water basins, including the Tulare Lake Basin.

Figure 2. Map of California’s water basins (aka hydrologic regions).
Source: California Department of Water Resources
Figure 3. Map of San Joaquin Valley.
Source: ECORP Consulting, report prepared for EPA.
Figure 4 is the best map that we have found of the Tulare Lake Basin. It is useful for providing an overview of the basin and its major features. However, it was prepared by others and contained some errors. The labeling errors have been corrected. The other errors could not be corrected without access to the GIS source data. In using this map, the reader should be aware of the following:

- The map doesn’t represent conditions at any particular point in time. River locations have changed over time. Lakes have dried up. Reservoirs have been constructed. Cities have grown up. In general, the map shows streams, rivers, and reservoirs as they are today.
- Tulare Lake is shown more or less at its high stand in 1862 at elevation of 216 feet. However, the lake is not precisely located. It is generally too far east and too close to Hanford.
- Mill Creek is shown as a significant stream course. It is actually less than a seasonal stream.
- Jerry Slough has not carried floodwaters since 1952. Its course is not accurately mapped.

Figure 4. Map of Tulare Lake Basin.
Source: Wikimedia Commons with corrections by Tony Caprio
Figure 5 shows the natural communities of the Tulare Lake Basin as they were in about 1850.
Figure 6. Map of Sequoia and Kings Canyon National Parks.
General Flood and Drought Notes

Basins, Watersheds and Deltas

Description and Identification of Basins

A "drainage basin" is defined as a part of the surface of the earth that is occupied by a drainage system, which consists of a surface stream or a body of impounded surface water together with all tributary surface streams and bodies of impounded surface water. An example is the Kaweah River Basin, the area drained by the Kaweah River. The term "watershed" can be used to describe the same feature.4

The Central Valley is divided into a northern portion (the Sacramento Valley) and a southern portion. The southern portion, extending from the Sacramento–San Joaquin Delta to the Tehachapi Mountains, is known as the San Joaquin Valley. The San Joaquin Valley is about twice as large as the Sacramento Valley. The San Joaquin Valley covers about 31,800 square miles. The floor of the valley is about 10,000 square miles.5

The San Joaquin River emerges from the Sierra in the middle of its namesake valley, and then turns north toward the San Francisco Bay (see Figure 3). Sometimes the term "San Joaquin Valley" is used to describe only that portion of the valley occupied by the San Joaquin River or only the valley floor. In this document, the term "San Joaquin Valley" is used to describe all of the San Joaquin Valley from the Sacramento–San Joaquin Delta to the Tehachapi, and from the crest of the Sierra to the Coast Ranges.

The California Department of Water Resources (DWR) has divided the state into 10 surface water basins or major hydrologic regions (see Figure 2). The San Joaquin Valley includes two of those surface water basins:

1. The San Joaquin River Basin (about 15,600 square miles)6 drains the northern half of the San Joaquin Valley. The San Joaquin River Basin contains the entire drainage area of the San Joaquin River and its tributaries. It extends from the Sacramento–San Joaquin Delta and the Cosumnes River in the north to the southern reaches of the San Joaquin River Basin, encompassing the area from Sacramento County (including the southeast corner of the county itself) to Madera County (and portions of Fresno County).

2. The Tulare Lake Basin comprises the southern half of the San Joaquin Valley. It ranges from the southern limit of the San Joaquin River Basin near Fresno to the crest of the Tehachapi Mountains.

The area of the Tulare Lake Basin has been variously reported. The USGS HUC boundary description for that basin identifies its size as about 16,200 square miles.7 The 2007 EPA report on the Tulare Lake Basin prepared by ECORP Consulting used the USGS HUC boundary and measured it as about 16,400 square miles using GIS technology.8

The northern boundary of the Tulare Lake Basin is not well defined in the western part of the valley because of the low gradient and alteration of the natural hydrography. The USGS HUC boundary does not include the Panoche Creek drainage. The DWR Water Plan Update 1993 said that it was 16,520 square miles. The DWR Water Plan Updates 2009 and 2013 have both stated that the area is 17,050 square miles, but they include the Panoche Creek drainage.9

The term "Tulare Lake Basin" is also used in two ways. It is generally used to describe the entire Tulare Lake Basin. But at other times, the term is used to describe the Tulare Lakebed (790 square miles). This document adopts the language of the California Water Plan, using "Tulare Lake Basin" to refer to the greater watershed and "Tulare Lakebed" to refer to the lakebed itself.10

Although DWR uses the nomenclature "Tulare Lake Basin," other sources use the shortened form, "Tulare Basin."

National Park Watersheds and Rivers

Sequoia and Kings Canyon National Parks are drained by the following five rivers:

1. San Joaquin River (South Fork)
2. Kings River (Middle Fork and South Fork)
3. Kaweah River (North Fork, Marble Fork, Middle Fork, East Fork, and South Fork)
4. Tule River (North Fork)
5. Kern River (North Fork)
**Description and Identification of Deltas**

When rivers flow west out of the Sierra onto the valley floor, they lose energy. When this occurs, they lose their ability to carry sediment and they often form fan-shaped deltas. In each flood, they deposit more sediment onto their delta. Or rather, that is what used to happen before the federal reservoirs were built on each of the major rivers: Kings, Kaweah, Tule, and Kern. Those reservoirs now function as sediment traps, intercepting much of the sediment load before the rivers can deliver it to their deltas.

For the bigger rivers, their deltas can be quite large, covering many square miles in area, and stretching out far across the valley floor. The Kings River Delta begins about where the town of Kingsburg is located; Visalia sits atop the Kaweah Delta.

The USACE estimates that the various channels of the Kaweah Delta (below McKay’s Point) are able to absorb and distribute a flow of up to 5,500 cfs of water before flooding begins to occur. Therefore, when the discharge below Terminus Dam (including the flow from Dry Creek) exceeds that amount, then flooding in the Visalia area can be expected. Prior to construction of the dam, some sort of June flooding occurred locally every few years on average. The older homes in Visalia were built with floor levels several feet above ground level because of this routine flooding. Few houses had basements.

Water on the valley floor in the San Joaquin Valley is generally trying to get from higher elevations (near Bakersfield) to lower elevations (the San Francisco Bay). The deltas of the Kern and Kings Rivers that extend out from the foothills can be thought of as a series of ridges or impediments to this flow. These ridges act as sills or dams, creating lakes on the valley floor.

The streams that flow east from the Coast Ranges also form deltas. By an odd coincidence, the delta formed by the east-flowing Arroyo Pasajero meets the much bigger delta formed by the west-flowing Kings River. The resulting sill (like a window sill or a broad saddle) has an elevation of 207 feet. Historically, that sill served to regulate the elevation of Tulare Lake in very wet years.

Arroyo Pasajero sits atop its delta. Kerry Arroues says that sometimes it sends water south into Tulare Lake, but sometimes it sends water north into the Fresno Slough and the San Joaquin River Basin.

Likewise, the Kings River sits atop its delta. Sometimes it sends water south into Tulare Lake, and sometimes it sends water north into the San Joaquin River Basin. This is discussed more fully in the section of this document on Pine Flat Dam.

The Sacramento–San Joaquin Delta (aka Bay-Delta watershed or simply the Delta) is a river delta and estuary formed by the confluence of the Sacramento and San Joaquin Rivers. It discharges into Suisun Bay, the upper arm of San Francisco Bay. In its natural state, the Delta was a large freshwater marsh, consisting of many shallow channels and sloughs surrounding low islands of peat and tule. Since the mid-19th century, most of the region has been gradually claimed for agriculture. The Delta consists of approximately 57 reclaimed islands and tracts, surrounded by 1,100 miles of levees that border 700 miles of waterways. The total area of the Delta, including both land and water, is about 1,150 square miles. Wind erosion and oxidation have led to widespread subsidence on the central Delta islands.

**Southern Sierra**

The Sierra Nevada runs from Lassen Peak in the north to Tehachapi Pass in the south. The Southern Sierra has been defined in different ways. In this document, that term is generally used as described by Bill Tweed to include all of the Sierra drained by the Kings, Kaweah, Tule, and Kern Rivers (the portion in the Tulare Lake Basin). North of that, the Central Sierra runs to Donner Pass, which coincides with the northern end of the continuous high alpine along with its barren peaks. Beyond that is the Northern Sierra.\(^\text{11}\)

**Elevations**

The elevations given in the literature about Tulare Lake levels must be treated carefully and do not necessarily represent what the elevations would be today with the current sea level reference datum. Some of the elevations are derived from surveys in the 1800s, and the sea level datum is usually not specified.

Sea level fluctuates from hour to hour and place to place. The sea level “datum” is the elevation that surveyors of a particular time and place use as their reference for a zero elevation. It is their vertical control point. When California was first settled, the state was on its own to establish a standard for what was the sea level reference datum. (Imagine the state engineer going to San Francisco Bay and taking the average of the tides.)
In any case, the early elevations of Tulare Lake and the surrounding area were made using the California State Engineering Department datum. By 1907, it was known that those elevations had to be reduced by 4.2 feet to get to the sea level datum as established by the U.S. Geological Survey (USGS). By 1929, a standard sea level datum had been established across North America. The current sea level datum for North America (NAVD 88) was set in 1991.

S.T. Harding reconstructed Tulare Lake levels in 1949. Harding determined that the elevations from the 1800s that were used by C.E. Grunsky in his graph of Tulare Lake levels should be reduced by 4.2 feet to conform to the sea level datum that was being used in 1949.

It’s critical to know the sea level datum associated with any given elevation. Without knowing the datum that was used when measuring a particular elevation, the reader can’t be sure what any given elevation really means.

An example of the confusion created when an author doesn’t provide this information is that some literature continues to report Tulare Lake’s high stand in 1862 as 220 feet. Those authors are still using W.H. Hall’s original measurements based on a sea level reference datum from the late 1800s; they just aren’t saying so. If they had used a sea level datum from the early 1900s or later, they would have stated the lake’s elevation as 216 feet. Sea level had been considered some 4 feet higher in Hall’s day, so the elevation for Tulare Lake also appeared to be 4 feet higher.

Likewise, some authors still report the Tulare Lake sill as being at elevation 211 feet. That is because they’re also using the sea-level datum from the late 1800s. If they had used a sea level datum from the early 1900s or later, they would have stated the sill’s elevation as 207 feet. Only in California could you have a lake with an elevation of 207 feet seemingly draining uphill over a sill with an elevation of 211 feet.

Conditions have changed significantly in the six decades since Harding’s time. The most dramatic change is the land subsidence that has occurred along the western side of the Tulare Lake Basin, particularly in Kern, Kings, and Fresno Counties. It has been speculated that some areas in those counties have now subsided upwards of 50 or 60 feet. See the section of this document on Land Subsidence for a more detailed discussion of this topic.

However, for the sake of consistency, most major studies of Tulare Lake continue to use Harding’s reconstructions of Tulare Lake level elevations, but note that they are doing this. This document also uses Harding’s reconstructed elevations.

The last comprehensive leveling surveys of the valley ended in 1970. Those were done as part of a USGS land subsidence research project. There have been no USGS general elevation surveys made in the Tulare Lake Basin since the 1960s. The data from those surveys are the basis for the USGS maps dated 1971. Large portions of the basin have subsided by significant amounts in the decades since the 1960s. (See the section of this document that describes Land Subsidence.) There are no general elevation maps that reflect the current — and ever changing — elevation of the Tulare Lake Basin.
Overview and Terminology

What Constitutes a Flood

There is a surprising variety in what constitutes a flood. This document contains a definition of what constitutes a drought. However, it does not have an all-encompassing definition of a flood; that has proved too messy a concept to define.

USGS broadly defines a flood as an overflow or inundation that comes from a river or other body of water and causes or threatens damage. That is almost like saying that a weed is any plant growing where you don’t want it; you’ll know it when you see it.

Some of our floods are obvious: a river overflows its banks or a downpour overwhelms a city’s drainage system. At the other extreme, some of our floods have two components: hydrologic and socioeconomic. Society decides what their tolerance is for natural processes and where they are willing to let a river flow. Some of our floods are the result of water appearing at the wrong place at the wrong time. They’re an inconvenience.

As one example, farmers wanted to drain Tulare Lake and keep it dry so that the lakebed could be used for agricultural purposes. They viewed Tulare Lake as an inconvenience, a nuisance to be prevented. Their viewpoint has prevailed. As a result, society has defined the presence of excess water in the lakebed as a flood. It’s an odd situation, but it effectively represents society’s values. Water managers go to great efforts to minimize this type of flood.

It’s helpful to look at lakebed flooding from the point of view of the farmers there. Runoff was below average in both 1970 and 1971, bordering on drought conditions. No significant storm or flood event happened in either year. Despite this, flooding occurred in the Tulare Lakebed in both 1970 and 1971. How was that possible? By looking at Figure 16, you can see that this flooding was left over from the big 1969 flood.

Lakebed flooding is a social construct; it is counted based on the number of growing seasons that are missed. The lakebed was flooded for three growing seasons: 1969, 1970, and 1971. Therefore, this is counted as three floods from the perspective of the lakebed farmers, even though the flood event occurred only once. (Something similar happened in the lakebed in 1982–84 and 1997–99. In each of those cases, lakebed flooding continued into a non-flood year.)

From the perspective of the natural environment, events like this would not be a flood at all. Tulare Lake was a natural body of water that society is trying to prevent from reclaiming its lakebed. Whenever the lake reforms naturally, we call that a flood.

So 1971 was a year with no storm event. It experienced near-drought conditions. And yet it is counted as a flood year. This can be a bit counter-intuitive, even mind-bending. That is one of the reasons why it proved so difficult to come up with a single definition for what constitutes a flood.

What Constitutes a Debris Flow

This document also addresses large-scale debris flows and landslide dams, both of which are associated with flood events. A debris flow is a moving mass of loose soil, rock, debris, and water that travels down a slope under the influence of gravity. It can carry material ranging in size from clay to exceptionally large boulders, and may contain a large amount of woody debris such as logs. Debris flows are typically associated with periods of heavy rainfall or rapid snowmelt and tend to worsen the effects of flooding that often accompany these events. Large-scale debris flows usually occur in small, steep stream channels and are often mistaken for floods. In fact, debris flows and flash floods often occur simultaneously in the same area.

What Constitutes "Normal"

Runoff is a reflection of precipitation. Figure 18 on page 111 show how widely runoff varies from year to year. Wet years and dry years commonly alternate, at least to some extent. The Tulare Lake Basin doesn’t have normal conditions in the sense of a statistical average. What is reliable about our climate is its extreme and relentless variability. That is our real normal; that is the lesson of Figure 18.

Most hydrologists avoid using the term normal since the average runoff can change depending on what period of record is used to define “normal.” For example, as shown in Table 2 of the 2007 EPA report on the Tulare Lake Basin prepared by ECORP Consulting — the long-term (1894–2001) Tulare Lake Basin runoff is about 10% less than the 50-year average (1962–2001). Some water managers in the Tulare Lake Basin like to use the period since 1962 after all the four major federal reservoirs in the basin were constructed.

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Figure 11 on page 46 illustrates the relative frequency of different water year types in the San Joaquin River Basin. It is based on the San Joaquin Valley Water Year Index shown in Figure 10 on page 45.

As Bill Tweed has written, the Tulare Lake Basin occupies the highly variable frontier between the wet winter climate of the Pacific Northwest and the often very dry winter climate of Southern California. We live on the unreliable southern edge of the winter storm track. Some years that storm track comes far enough south to include us, but some years it does not.

The 30-year mean (the statistical average) is what weathermen tend to call “normal,” but that does not make it normal in the sense that we can count on it. As shown in Figure 11, about half of our years are wetter than average and half are drier; that is just how the underlying index was set up. Other than that, the categories in this graph are somewhat arbitrary. They are standards used by the California Department of Water Resources to track and report on droughts over the past 114 water years (1901–2014). For more detail on these categories, see Table 3 and the associated text.

One of the lessons that we can learn from these figures is that most of our years are far from normal. As shown in Figure 11, only 36% of our years have been in the two categories that bracket the statistical average (Above Normal and Below Normal). The rest have been either much wetter or much drier. This is even more vividly illustrated in Figure 10.

There is no long-term pattern to our mixture of wet years and dry years. A team led by Dave Meko used tree-rings to reconstruct the flow on the San Joaquin River and its major tributaries for 1113 years (900–2012). The flow reconstructions contained no strong, regular, statistically significant cycles over their full lengths.

**Types and Duration of Floods**

There are three very different types of floods that occur in the Tulare Lake Basin:

1. Sharp, high flood peaks of short duration and comparatively small volumes, typically lasting a day or two. In this type of flood, elevated peak volumes and river heights damage developed areas along a river. These are our typical rain-floods. These are the type of floods that result in damage to national park infrastructure, and the ones more likely to be recorded in park monthly and annual reports. Dams are designed to offer those who live downstream a relatively high level of protection from this type of flooding event. For an example of a typical flood of this type, see Figure 29 on page 367 which graphs the discharge of the January 1997 flood.

2. Sustained high levels of discharge, typically lasting a month or more. These are snowmelt floods that typically occur between March or April and June. Such prolonged discharges inundate the valley floor. Dams are not designed to completely contain this type of flooding event. These floods don’t necessarily cause any damage in the Sierra or even in the deltas below. The floods themselves may not be recorded in national park reports, but there may be a mention of high precipitation, late springs, or difficulty accessing the higher elevations. These floods are primarily marked by the delivery of large volumes of water to the Tulare Lakebed. Residents of Corcoran (in the Tulare Lakebed) have a very different view of such floods than do residents of Three Rivers (above Lake Kaweah).

3. Cyclic mega floods. Floods in this category can be extraordinarily large. This category of flooding occurs approximately every 200 years and can result in a series of flooding events lasting up to 10 years. (See the section of this document that describes the [California megafloods](#)).

The above categorization of floods is based on duration, how long different types of floods last. It is also possible to distinguish between regional and localized flooding. This is addressed in the section of this document that describes the [Causes of Flooding](#).

**When Do Floods Occur**

A look at Figure 18 on page 111 will show how widely runoff varies from year to year. Wet years and dry years commonly alternate. The Tulare Lake Basin doesn’t have normal conditions in the sense of a statistical average. What is reliable about our climate is its extreme and relentless variability. That is our real normal; that is the lesson of Figure 18.

Floods are amazingly commonplace in our area. A look at the Table of Contents or Figure 25 on page 163 will show just how commonplace; this document describes what we know about approximately 188 floods that have occurred during the last 2,000 years.
Floods occur at all manner of times. They occur in wet years, and they occur during multi-year droughts. They occur during the winter wet season, and they occur during the summer dry season. When they occur varies so widely because there are such a variety of causes for floods.

**Causes of Flooding**

Sustained high levels of discharge (as opposed to high flood crests) result from an unusually deep snowpack. Megafloods that occur with a quasi-periodicity of approximately 200 years have their own unique set of causes (see the section of this document that describes the California megafloods).

Other high flood crests — whether local or widespread — typically result from one of six different conditions in the Tulare Lake Basin:

1. **Heavy winter rainstorms.** This is the most common cause of major flooding events. It results from heavy, relatively warm rains during November–February. Writing in 1900, John Muir described this type of flood quite well:

   > The Sierra Rivers are flooded every spring by the melting of the snow as regularly as the famous old Nile...Strange to say, the greatest floods occur in winter, when one would suppose all the wild waters would be muffled and chained in frost and snow...But at rare intervals warm rains and warm winds invade the mountains and push back the snow line from 2000 feet to 8000, or even higher, and then come the big floods.

   When weather conditions are right, the Southern Sierra can wring amazing amounts of water out of passing storms. Precipitation events producing 10 to 20 (or many more) inches of water equivalent are possible in the region. When such superstorms are unusually warm — and particularly if they drop heavy rain on an existing snowpack and melt much of it — our main rivers can rise to peak flows in excess of 50,000 cfs. Such extreme events have occurred infrequently in historic times. Some of the more impressive winter rain-floods occurred in 1861–62, 1867, 1955–56, 1966, 1969, 1982–83, and 1997.

2. **Spring snowmelt.** These floods take place at the onset of hot weather after a wet winter has built up a larger-than-average snowpack in the mountains. Such events typically occur several times a decade, and have relatively gentle peaks compared to the winter floods. The May–June 1850 flood may have been one of the largest such floods to occur in historic times. Bigger spring snowmelts have probably occurred since that time, but a large portion of the floodwaters have been diverted out of the rivers since the late 1800s.

3. **Remnants of Pacific hurricanes.** In this document, the terms “hurricane,” ”cyclone,” and “typhoon” are used somewhat interchangeably. Floods resulting from Pacific hurricanes are relatively rare events that occur between June and October. (The Tulare Lake Basin receives moistures from a variety of tropical storm systems, especially during the winter. However, Pacific hurricanes are a category unto themselves; something we seldom see, and never in the winter.) Because the waters off California’s coast are so cool, hurricanes always degrade before they make landfall. Only one such storm has ever been recorded coming ashore in California as a tropical storm (September 25, 1939 near Long Beach.) When remnants of hurricanes do come ashore, they can be intense and deliver a huge amount of precipitation. They cover a broader area and last longer than most other summer storms do. They usually make landfall in Southern California or in Mexico, they rarely come as far north as the San Joaquin Valley. The only examples of Pacific hurricanes causing floods in the Tulare Lake Basin that we are aware of were those of 1918 (unnamed), 1932 (unnamed), 1972 (Gwen), 1976 (Kathleen), 1978 (Norman), 1982 (Olivia), 1982 (Sergio), 1998 (Isis), and 2002 (Huko). For a typical example, see Figure 28 on page 341 which show Hurricane Olivia coming up the Pacific coast, recurving, and then coming ashore as a tropical depression. Typhoon Melor in 2009 was in a special category. The remains of Melor remained in the western Pacific, but water vapor from that typhoon moved across the ocean basin via an atmospheric river, a recently discovered phenomenon. Although never a hurricane, the cyclic storm that struck the Buena Vista Lake Basin in February 1978 behaved in a manner similar to a hurricane. Likewise, the cyclic storm that collapsed the Interstate 5 bridges near Coalinga in March 1995 behaved in a manner similar to a hurricane. Hurricanes and cyclic storms have the ability to punch through a mountain barrier and deliver extremely large rainfalls to the rain shadow on the lee side of those mountains. For example, the 1932 hurricane delivered rain at an average rate of 6 inches an hour over the entire Cameron Creek watershed near Tehachapi, causing catastrophic flooding.

4. **High intensity non-tropical storms.** This is somewhat of a miscellaneous category: With their high intensity, storms such as these are often described as cloudbursts. They frequently cause flash floods and overwhelm drainage systems. They are often relatively localized, although they can cover a wider area when associated with a frontal system. Cloudburst storms are often relatively brief, but sometimes last as long as three hours. The intensity of these storms is quite high, and they can produce enough precipitation to result in peak flows equal to or somewhat greater than those of a general flood-producing rainstorm. Flooding from cloudbursts is characterized by high peak flows, short duration of floodflow, and small volume of
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... runoffs. These high intensity storms are most common from late spring to early fall, but they can occur at any time of the year. Examples of such flood-related storm events that we are aware of were those of 1898, 1913, 1937, 1941, 1945, 1951, 1965, 1972, 1976, 1984, 1986, 1995, 1996, 1997, 2001, 2002, 2003, 2005, 2006, 2007, and 2008. There were surely many others, but it is easy to miss them in the record; they just don’t get systematically recorded like other floods.

5. **Monsoonal moisture.** The North American Monsoon is associated with a high pressure ridge that moves northward during the summer months and a thermal low (a trough of low pressure which develops from intense surface heating over the Mexican Plateau and the Desert Southwest of the U.S.) The monsoon typically develops in the U.S. between July and mid-September each year. The upper-level high pressure ridge may be coincident with a surface high. The clockwise flow around the upper-level pattern brings monsoonal moisture into Central California, including the Southern Sierra. The surface high brings lower-level moisture into the Desert Southwest, including the desert areas of Southern California. Pulses of low-level moisture are transported primarily from the Gulf of California and the eastern Pacific. Upper-level moisture is also transported into the region, mainly from the Gulf of Mexico by easterly winds aloft. The low-level moisture largely impacts Arizona and Sonora. The upper-level moisture can be transported as far as California. Depending on the position of the jet stream, the high pressure shifts from being centered over northern Mexico to over the Colorado Plateau or the Great Basin. When the high pressure is centered in the northern position, it causes south to southeasterly winds aloft to move tropical moisture into southwestern California and the Sierra. When this moisture-laden air reaches the Sierra, it is forced upwards and forms thunderstorms or convective cells which can be responsible for intense periods of rainfall. Examples of floods in the Tulare Lake Basin that were caused by monsoonal moisture were those of August 1983, August 1984, September 1997, August 2003, July 2008, and July 2011.

6. **Landslide dam failures.** These are localized events and, therefore, often go unrecorded. Some of the ones that we know of occurred in 1861–62, 1867–68, 1982–83 and 2002. When the landslide dams and the blocked streams are both small (such as in the 1982–83 and 2002 floods) there is relatively little damage. However, when a large landslide dam blocks a large river, the results can be highly destructive. The largest historic dam failures that we know of all occurred during a nine-day period in December 1867 and resulted in spectacular floods. Other large dam failures that may have occurred prehistorically are the Little Kern dam, the series of slides on the west slope of Moro Rock, and the enormous slide on the north side of Dennison Peak. The local canyons of the Kaweah show geologic evidence of dozens of massive landslides. Some of those undoubtedly produced severe flooding downstream similar to those that occurred in December 1867. Sierra-wide floods are common, even the norm. If one river floods, then usually the other rivers in the vicinity flood. However, floods resulting from landslide dam failures are different. Such localized flooding events generally don’t have much of an impact farther downstream on the main rivers. Because of these small-scale events, places such as Cedar Grove can be expected to have somewhat more large landslide dam flooding events than a downstream location such as Pine Flat would experience.

**Flash Floods**

A flash flood results from heavy rainfall within a short period of time, usually less than 6 hours, causing water to rise and fall quite rapidly. Typically a flash flood is associated with a thunderstorm or the remnant of a hurricane.

**El Niño/La Niña—Southern Oscillation**

The El Niño/La Niña–Southern Oscillation (ENSO) is a periodic climate pattern that occurs across the tropical Pacific Ocean at two- to seven-year intervals and lasts nine months to two years. This is the most intense short-term perturbation of the Earth’s climate system. These events have a strong impact on the continents around the tropical Pacific, and some climatic influence on half of the planet.

The El Niño/La Niña–Southern Oscillation has two phases. El Niño, the warm phase, is characterized by a warming of the ocean surface from the coasts of Peru and Ecuador to the center of the equatorial Pacific. La Niña, the cool phase, is characterized by a cooling of the surface waters in the equatorial Pacific. Nineteenth-century Peruvian sailors named the warm northerly current off their coast “El Niño” because it was most noticeable around Christmas.

In popular usage, the El Niño/La Niña–Southern Oscillation is sometimes called just “El Niño.” That usage, rather confusingly, lumps together the warm oceanic phase, El Niño, with the cool phase, La Niña.

The causes of — and relationship between — El Niño and La Niña events are not fully understood. Sometimes these phases alternate, but often they do not. El Niño events occur roughly twice as often as La Niña events.
During the last several decades, the number of El Niño events has increased while the number of La Niña events has decreased. Studies suggest that this variation is most likely linked to global climate change.

While El Niño events only occur in the tropics, their impacts are felt in many parts of the world. This happens because the location of the huge mass of warm water causes the location of the jet stream, or storm track, to shift. As a consequence, some regions are warmer or colder, or wetter or dryer, than average.

El Niño does not actually create any storms over California or anywhere else. It simply shifts the usual jet stream patterns so that some areas can be more susceptible to storm formation. It isn’t possible to say whether a particular weather event during a strong El Niño / La Niña winter would have occurred anyway. The El Niño conditions just made the storm more likely.

El Niño / La Niña events are capable of causing extreme weather such as floods and droughts in many regions of the world. Bill Tweed explained their effect on the Tulare Lake Basin in one of his columns for the Visalia Times-Delta. In our area, strong El Niño winters are usually wet, but moderate and weak events can be wet or dry. Moderate La Niña events are often dry, but can be wet. Strong La Niña events, as in the winter of 2010–11, can be quite wet.

Not all El Niño events have the same strength or location, and consequently their impacts can vary significantly. In general, the larger the area and the greater the warming of the eastern Pacific’s equatorial waters, the greater the impact on other regions.

Although a strong El Niño event is often associated with above-average levels of precipitation in California, it can have a very different effect in other parts of the world. For example, the El Niño events of 1876–77 and 1918–19 were associated with major droughts elsewhere in the world. These resulted in devastating famines and the deaths of millions of people.


There has not been a strong El Niño event since the winter of 1997–98, only two weak and two moderate events during that period. Bill Patzert is a climatologist at the NASA Jet Propulsion Lab in Pasadena. He says one reason for the lull in strong El Niños is another cycle known as the Pacific Decadal Oscillation, or PDO. It’s like El Niño but farther north and much less capricious.

The PDO takes a long time and a lot of momentum to switch phases. Patzert says the cool or negative phase of the PDO can be thought of as an “El Niño repellent.” When the PDO is in its negative phase, it affects the jet stream. When that happens, the big winter storms tend to get detoured around California. The PDO has been stuck in this negative phase since about 1999. The PDO shows signs that it may be on the verge of making a big switch into its warm or positive phase, which would encourage El Niño conditions. But that is far from certain.

Meteorologist Jan Null analyzed the relationship of the 22 El Niño events that have occurred since 1950 with precipitation in the seven climate regions of California. He found no correlation between weak and moderate El Niño events and precipitation for any region. However, strong El Niño events were correlated with increased precipitation for every region of the state, strongest and most frequently for Southern California. Tulare Lake Basin floods that were apparently associated with strong El Niño events include the floods of 1876, 1906, 1918, 1931, 1958, 1982–83, and 1998. The El Niño events of 1876–77, 1905–06, 1918–19, 1982–83, and 1997–98 were particularly strong. The El Niño event that developed during the winter of 1932–33 occurred during the extended drought of 1918–34 and resulted in a remarkable period of snow in the national parks.

Just as El Niño events can cause floods, so can La Niña events. Tulare Lake Basin floods that were apparently associated with La Niña events include the floods of 1911, 1916, 1924, 1938, 1955–56, 1976, and December 2010. The La Niña events of 1955–56, 1975–76, and 2010–11 were particularly strong.
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Weak El Niños have been associated with some of the driest winters of the 20th century, including the winters of 1976–77 and 1977–78. However, weak El Niño events are poor predictors of droughts, just as they are poor predictors of above-average winters. There just doesn't seem to be any relationship.

At the opposite end of the scale was the winter of 1997–98 which brought nearly double the state's average precipitation, and triggered massive mudslides and flooding throughout California. Patzert says he thinks that year is when Californians began to form the assumption that El Niño equals rain. But there are only a few strong El Niños like 1997–98 in every century.

Jan Null analyzed the relationship of El Niño and La Niña to California flood damage from 1949 until 1997. He found that occasionally — but only occasionally — a strong El Niño event means disastrous flooding for California. It is just as likely that California will have significant flooding in a non-El Niño year. Of the 10 costliest flood years in California since 1950, one occurred during a strong La Niña winter (December 19–27, 1955) and two occurred during strong El Niño winters (December 1982 – March 1983 and December 1997 – April 1998).

Just as an El Niño doesn't guarantee rain or flooding, the absence of one does not necessarily mean no rain. The bottom line is that California can get wet during a strong El Niño winter, but not always. El Niño is not the only thing happening in the atmosphere; other patterns can either enhance or detract from its overall impact.

For more detail about ENSO, PDO, and similar issues, see:

**Atmospheric Rivers**

At the heart of California's water supply are atmospheric rivers — long ribbons of moisture that transport huge amounts of water vapor from the tropics toward the poles. When atmospheric rivers move inland and strike mountains, the air rises and cools, creating heavy rainfall. Atmospheric rivers are the source of nearly half of California's precipitation, and they cause the large majority of the region's major serious floods.

Atmospheric rivers occur primarily between October and February; they supply the water behind many of our heavy winter rainstorms. Norm Miller agreed that atmospheric rivers are the primary cause of flooding. He said that extratropical cyclones, when stacked up in the Pacific also lead to major floods.

Atmospheric rivers are responsible for most of the horizontal transport of water vapor outside of the tropics. They are slightly more prevalent during years when there is no El Niño.

Approximately 20 atmospheric river events occur each year along the West Coast. The zone of most frequent occurrence for these events runs from Northern California into British Columbia. Farther north and south, atmospheric river events become less common. The Tulare Lake Basin falls on the meteorological boundary between wet and dry.

A cloud bulging above a mountain crest can hold several million pounds of water, but that water is divided into droplets too small to overcome the atmospheric updraft lifting them. The moisture has to freeze into ice or snow to be large enough to fall back to Earth. Tiny water droplets can remain stubbornly liquid at temperatures as low as 20 or 30 degrees below zero, perched on a thermodynamic cliff. The droplets need an instigator that tips them over the edge, into becoming ice; they need some microscopic flotsam floating in the air.

Atmospheric rivers vary widely in how much moisture they produce. Some drop 25% of their water; others, just 15%. That is a huge difference. Research by Kimberly Prather and Marty Ralph at the Scripps Institute of Oceanography suggests that the difference results from dust transported by high altitude winds from the Taklamakan Desert in far western China. When that type of dust is present, it appears to cause intense precipitation to fall out of the water-choked clouds. We can get upwards of 40% more moisture out of an atmospheric river event when those clouds contain Asian dust to seed the formation of ice crystals.

The nontechnical term “Pineapple Express” is popularly used to describe the meteorological phenomenon that causes moisture to be drawn from the Pacific Ocean near Hawaii and transported to the West Coast with firehose-like ferocity. The Pineapple Express is a subset of atmospheric rivers, distinguished primarily by the source of the water vapor and the strength of the southwesterly trending vapor-transport atmospheric river extending toward the West Coast. About 30% of atmospheric rivers fall into the Pineapple Express category.
Atmospheric rivers are embedded within much broader atmospheric storms referred to technically as extratropical cyclones. Extratropical cyclones are the winter-time analogue to hurricanes, but have a much different structure. Atmospheric rivers are the business end of extratropical cyclones because where an atmospheric river hits the mountains, it can create extreme precipitation, flooding, and high winds. In terms of impacts, an atmospheric river is to the broader extratropical cyclone it is embedded within, as the hurricane eyewall is to the broader hurricane of which it is a part.

That description of how an atmospheric river is embedded within an extratropical cyclone is based on the website for the USGS’s Multi Hazards Demonstration Project. However, Norm Miller said this is incorrect. He said that atmospheric rivers are not cyclonic they have low vorticity. Cyclones help to provide a pathway for atmospheric rivers.

According to NOAA, atmospheric rivers average 250–375 miles wide. Norm Miller said it is better to think of them as being on the order of about 600 miles (1000 km) wide. Mike Dettinger said that they tend to move through an area in ways that result in the dumping of precipitation on an area broader than the size of the river. That is, an atmospheric river might be only 200 miles across, but as it moves through an area, its footprint is likely to get smeared out to be wider than that. While that is the way that they usually work, not every storm does that. Atmospheric rivers are known for stuttering or stalling for periods of time ranging from a few hours to almost a day as they pass over California. Those areas that are under the place where they stall get extra doses of precipitation compared to areas elsewhere. That is, a broad part of the state (say, half of it) may get dumped on by an atmospheric river, but the area beneath where it stalls gets even more.

The storms of 1861–62 arguably caused the most widespread and intense flooding that the West Coast has experienced in historic times. The atmospheric mechanisms behind those storms are unknown. However, they were likely the result of an intense atmospheric river, or a series of atmospheric rivers.

Although atmospheric rivers occur in the mountains in the winter, they usually produce heavy rains rather than heavy snowfalls. That is because atmospheric rivers are usually warmer than most other storms. However, the word “usually” means that not all atmospheric rivers live up to that expectation. Some atmospheric rivers are cool and produce snow, especially at higher elevations.

The following table lists five of the largest atmospheric river floods that have impacted the Tulare Lake Basin:

1. November, 1861 – January 1862
2. February 11–24, 1986
5. October 13–14, 2009

Atmospheric rivers are generally viewed as flood producers. But they are also an important part of the state’s water supply. They contribute about 40% of California’s annual precipitation. When the frequency of atmospheric rivers decreases, this can contribute to droughts. In contrast, an atmospheric river may have the power to end a drought.

Just one or two such storms can help replenish the water system during dry spells. Persistent droughts often end as a result of the arrival of an especially wet month or a few very large storms. A 2013 study by Mike Dettinger analyzed how drought events ended along the West Coast over the 60-year period 1950–2010. Atmospheric river events broke up 60%–74% of all persistent droughts in the Pacific Northwest and about 33%–40% of all droughts in California. The remaining droughts in California were mostly broken up by rainfall resulting from local low-pressure systems.

Preparing for the Next Big Flood

At one time, the USACE defined the “intermediate regional flood” as the flood with a recurrence interval of 100 years. That agency did a study in 1974 which calculated what the intermediate regional flood would be on the lower Kaweah River, taking into account the operation of Terminus Dam. (USACE no longer uses the term “intermediate regional flood.” They now use frequency events such as a flood with a recurrence interval of 100 years.)

The 1974 USACE study said that the intermediate regional flood on the valley floor, with Terminus Dam in operation, would be of somewhat lesser magnitude than the November 1950 flood on the Kaweah River.
Damage in the 1950 flood was characterized as extensive. Many houses and automobiles were swept away and roads and bridges were extensively damaged. Flooding extended from 2–4 miles wide from Woodlake to Visalia. Wide areas of cotton, pasture, and grain were inundated along the St. Johns River and Cross Creek and south of Visalia along various distributary channels and canals. Floodwaters extended 3–4 miles wide to the northwest of Visalia.

Mill Creek overflowed in Visalia, resulting in extensive flooding of the business section. The areas flooded in November 1950 would resemble the calculated intermediate regional flood, except that flooding in Visalia would be more widespread under intermediate regional flood conditions than in 1950 as a result of freeway construction along Highway 198.

Risk management agencies have different terms for projected floods that have a recurrence interval of 100 years. FEMA’s term for such an event is the “base flood.” The base flood is the national standard used by the National Flood Insurance Program for the purposes of requiring the purchase of flood insurance and regulating new development. Base flood elevations are typically shown on flood insurance rate maps.

The Tulare Lake Basin has experienced several floods during historic times that equaled or exceeded a flood with a recurrence interval of 100 years (aka, the intermediate regional flood or base flood).

The 1974 USACE study also laid out a bigger flood: the standard project flood. The USACE formerly defined the “standard project flood” as the flood that may be expected from the most severe combination of meteorological and hydrological conditions that are considered reasonably characteristic of the geographical area in which the drainage basin is located, excluding extremely rare combinations. This is a rare event, but one that could reasonably be expected to occur.

The 1974 study found that a general rain-flood (possibly augmented by melting snow) or cloudbursts would create the most severe flood conditions in the study area. Therefore, the standard project flood is that which can be expected from a standard project rainstorm, either general rain or cloudburst, centered over the Kaweah drainage basin, taking into account the operation of Terminus Dam and ground saturation conditions.

The December 1955 flood was the largest and most damaging rain-flood known to have occurred in northwestern Tulare County in the 20th century prior to the completion of Terminus Dam. The Kaweah drainage has experienced several floods larger than the 1955 flood. The December 6, 1966 flood was a bigger flood than the 1955 flood; it occurred after Terminus Dam was in operation. The 1861–62 and the December 1867 floods were significantly larger floods than even the 1966 flood.

As shown in Figure 7, most of northwest Tulare County, with the exception of a few island areas and high ground, is subject to flooding under standard project flood conditions. Large areas would experience sheet flooding. Floodwaters would back up and spread out on the uphill side of barriers such as roads, railroads, and canal embankments. Portions of Orosi, East Orosi, Cutler, Goshen, Farmersville, Exeter, Visalia, and Woodlake would be subject to inundation under standard project flood conditions. A standard project flood, with Terminus Dam in operation, would be of somewhat greater magnitude than the December 1955 flood was without Terminus Dam.
Figure 7 is the USACE’s estimate of how much of northwest Tulare County would be flooded during the standard project flood. The study area extended from Highway 99 on the west to Highway 137 and Lindsay on the south. This flood projection assumes that Terminus Dam is in operation.

The standard project flood is a rare event, but one that could reasonably be expected to occur. It is bigger than the 1966 flood. Think of it like the 1861–62 flood or the 1867–68 flood, our two biggest floods in historic times.

The USACE no longer uses the term “standard project flood” when assessing flood risk potential. But the 1974 study gives a good illustration of what the effects of a very large flood could be like, even with Terminus Dam in operation. It is a reminder of how important downstream levees are.

By some measures, the 1861–62 flood was the greatest flood of historic times. It wasn’t a fluke and it is only prudent that we plan for a future flood of similar or greater magnitude.

The USGS Multi Hazards Demonstration Project (MHDP) applies science to improve the resiliency of communities in Southern California in their response to a variety of major natural hazards. The MHDP assembled experts from a number of agencies to design a large, but scientifically plausible, hypothetical storm scenario that would provide emergency responders, resource managers, and the public a realistic assessment of what is historically possible. One of the MDHP’s full scenarios, called ARkStorm, addresses massive West Coast storms analogous to...
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those that devastated California in 1861–62. Storms of this magnitude are projected to become more frequent and intense as a result of climate change.

The ARkStorm scenario is patterned after the 1861–62 historical events, but uses modeling methods and data from large storms in 1969 and 1986. The ARkStorm draws heat and moisture from the tropical Pacific, forming a series of atmospheric rivers that approach the ferocity of hurricanes and then slam into the West Coast over several weeks.

The website for the USGS’s Multi Hazards Demonstration Project warns that an ARkStorm is plausible, perhaps inevitable. The 1861–62 storm was not a freak event, was not the last time that California will experience such a severe storm, and was not the worst case. The geologic record shows that six mega-storms more severe than the 1861–62 have struck California in the last 1800 years (see the section of this document that describes the California megafloods), and there is no reason to believe that similar events won’t occur again. In the Tulare Lake Basin, the 1867–68 flood was bigger than the 1861–62 flood on all four of our major rivers. With the right alignment of conditions, a single intense atmospheric river hitting the Sierra east of Sacramento could bring devastation to the Central Valley. An independent panel wrote in October 2007 to the California Department of Water Resources: California’s Central Valley faces significant flood risks. Many experts feel that the Central Valley is the next big disaster waiting to happen. This fast-growing region in the country’s most populous state, the Central Valley encompasses the floodplains of two major rivers — the Sacramento and the San Joaquin — as well as additional rivers and tributaries that drain the Sierra Nevada. Expanding urban centers lie in floodplains where flooding could result in extensive loss of life and billions in damages.

Water Year and Runoff Terminology
Because we live in a Mediterranean climate where nearly all of our precipitation comes during the winter months, terms (metrics) have been created to measure this yearly accumulation of water as a whole, rather than arbitrarily splitting it on December 31. The two most common terms are:

- **Water year.** The term “water year” refers to the twelve-month period beginning October 1 in one year and ending September 30 of the following year. The water year is designated by the calendar year in which it ends. Therefore, water year 2014 would cover the 12-month period from October 1, 2013 through September 30, 2014. This term is commonly used to measure runoff, precipitation, and other water-related metrics.

- **Weather year.** The term “weather year” refers to the twelve-month period beginning July 1 in one year and ending June 30 of the following year. The weather year is designated by the calendar year in which it ends. Therefore, weather year 2013 would cover the 12-month period from July 1, 2013 through June 30, 2014. This term is commonly used to measure annual precipitation.

By default, this document uses the term “year” to refer to calendar year. It generally makes clear when referring to water year, fiscal year, etc.

Much of the water storage in the Sierra is in the form of snow, so there is considerable interest in what the runoff will be during the April–July snowmelt period. The April–July period is a focus of agricultural interests, but is also relevant for management of montane meadows. (Snowpack on the April 1 and May 1 snow surveys is typically used as a predictor of when meadows will dry out enough for stock grazing.) The amount of snowmelt runoff varies by year and watershed, but constitutes roughly two-thirds of total annual runoff for the Tulare Lake Basin.

The term “annual runoff” is sometimes used to describe all of the runoff over the October–September water year. But at other times, the term is used to describe only the portion that occurs during the April–July snowmelt period.

**Acre-foot Water Measurement**
An acre-foot of water is enough water to cover an acre of land one-foot deep. It is enough to provide a 12- to 18-month water supply for an average family in the Tulare Lake Basin. Water flowing at a steady rate of 1 acre-foot per day is equivalent to: approximately 0.504 cubic foot per second, 226 gallons per minute, 43,560 cubic feet of water per day, or about 326,000 gallons per day.
Measurements of Flows and Runoff

The flow and runoff measurements provided in this document were generally obtained from publicly available sources. The measurements of annual runoff for the four main rivers in the Tulare Lake Basin generally came from the California Data Exchange Center (CDEC) using the full natural flow sensor (#65).^59

Regardless of source, the measurements were generally based upon correlations, calculations, or estimates based upon spot measurements (for late 19th century and early 20th century figures) and continuous measurements or daily records that have been adjusted for diversions and dams to produce the unimpaired runoff record. Unimpaired flow represents the flow that would occur if there were not any diversions or reservoir regulation upstream of the gage.

Although unimpaired flow is sometimes referred to as the full natural flow, the unimpaired flow does not reflect fully natural conditions since it does not account for changes in natural watershed flow characteristics that have occurred in the past 150 years due to land use alterations and vegetation conversion. It is assumed for convenience, however, that the cumulative effects of those alterations on the seasonal runoff from the upland ecosystem are relatively minor and the unimpaired runoff is a satisfactory representation of natural upland runoff.

Measurements of Peak Floodflows

The peak flow in a river during a flood can be measured in a variety of ways, among them:

1. **Instantaneous flow at the peak moment of the flood.** Measurements of this type typically come from a manually read gaging station.
2. **Peak hourly flow.** This is the average flow for the peak hour of the flood. This type of measurement typically comes from an automatic gaging station.
3. **Peak daily flow.** This is the average hourly flow for the peak day of the flood. This typically comes from an automatic gaging station. It is calculated by averaging the 24 hours of the peak day (midnight to midnight). The term “peak daily flow” suggests that this is a measurement of total daily flow when that really isn’t the case. To avoid that confusion, this document generally refers to this measurement using the nontechnical term “peak average daily flow.”

When you see a reference saying that a flood peaked on a particular day, it is usually referring to one of the first two measurements.

Every flood is different, and there is a wide range of variation. However, in a ballpark sense, the peak hourly flow measurement on the big rivers in the Tulare Lake Basin is often about 150% greater than the peak average daily flow. As an example, the peak hourly flow for the Kern River in the 2002 flood was 26,500 cfs while the peak average daily flow (the average hour for the peak day) was 10,306 cfs.

Flood Rate and Flood-risk terminology

One way of describing the size of a particular flood is by describing the risk of a flood that big (or bigger) occurring in any given year. It used to be common practice for the USACE and other risk management agencies to express the risk of floods using exceedence interval terminology (e.g., 50-year flood, 100-year flood, etc.). This is also commonly called a flood’s recurrence interval or return interval.

Most of those agencies have now adopted the practice of expressing risk in terms of exceedence frequency, or flood likelihood (e.g., a 20% chance of a flood of a certain size or larger occurring in any given year).

The public still likes to think of a flood in terms of its recurrence interval. If you know the exceedence frequency for a particular flood, you can calculate its recurrence interval using the following conversion formula:

\[
\left( \frac{1}{\text{Exceedence Frequency}} \right) \times 100 = \text{Recurrence Interval}
\]

For example, the flood that occurred on the Kings River in December 1955 had an exceedence frequency of 1%. The above formula can be used to calculate its recurrence interval:

\[
\left( \frac{1}{1\%} \right) \times 100 = 100
\]

A flood with a recurrence interval of 100 is commonly referred to as a 100-year flood. A flood of that magnitude has a 1% chance of being equaled or exceeded in any given year.
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Table 2 presents the two types of terminologies for a selection of different size floods.

<table>
<thead>
<tr>
<th>Flood Recurrence Interval</th>
<th>Likelihood of Flood</th>
</tr>
</thead>
<tbody>
<tr>
<td>10 years</td>
<td>10% chance/year</td>
</tr>
<tr>
<td>20 years</td>
<td>5% chance/year</td>
</tr>
<tr>
<td>50 years</td>
<td>2% chance/year</td>
</tr>
<tr>
<td>100 years</td>
<td>1% chance/year</td>
</tr>
<tr>
<td>200 years</td>
<td>0.5% chance/year</td>
</tr>
<tr>
<td>500 years</td>
<td>0.2% chance/year</td>
</tr>
</tbody>
</table>

Calculation of the risk of different size floods occurring is done by looking at the history of past floods for a given watershed. The results of those calculations — the predictions of the likelihood of future floods — are a series of flood frequency curves. In our area, those calculations are generally done by the Sacramento District of the USACE.

Since a flood with a recurrence interval of 100 is commonly referred to as a 100-year flood, there is a temptation to think that such a flood has a 100% chance of occurring in the next 100 years, but nothing is guaranteed in nature. There is actually only a 63.4% probability of a 100-year (or bigger) flood occurring in the next 100 years. (That is just how probability works out.)

The probability, $P_e$, that a certain-size flood occurring during any period will exceed the 100-year flood threshold can be calculated using the formula $P_e = \left[1 - \frac{1}{T}\right]^n$, where $T$ is the return period of a given storm threshold (e.g., 50-year, 100-year, etc.), and $n$ is the number of years that you’re interested in. So for a 100-year flood, the formula would be $1 - \left[1 - \frac{1}{100}\right]^{100} = 63.4\%$. That is almost enough to make you feel like you’re back in high school again.

Some caution is necessary when using flood frequency curves, since they are derived from historic data. One of the assumptions used in building flood frequency curves is that the probability distribution function is stationary, meaning that the mean, standard deviation and max/min values are not increasing or decreasing over time. If temperatures are changing and precipitation cycles are being altered, then the probability distribution is also changing. The simplest implication of this is that not all of the historical data can be considered valid as input when building the frequency curves. In an era of changing climate, the past cannot necessarily be used to predict the future.

There are two other flood-risk terms that are used by dam designers:
- probable maximum flood
- level of flood protection

The probable maximum flood is the flood that may be expected from the most severe combination of critical meteorological and hydrologic conditions that are reasonably possible in a particular drainage area. This is a very unlikely event, much less than a 500-year flood. Dams are designed to safely pass the probable maximum flood without being overtopped. This is especially critical for earthen dams like Terminus, Lake Success, and Lake Isabella. If a dam were overtopped, this could lead to failure of the dam and catastrophic consequences for the people downstream.

The other dam-related flood-risk term is “level of flood protection (aka design protection level).” This is the level of protection provided by a reservoir for people and property downstream of that reservoir. The level of flood protection is related to the storage capacity of a reservoir.

Authorized flood-control reservoirs are designed to provide a particular flood-control pool (aka flood control space or flood management reservation space). That flood-control pool is used to store high inflows from a flood event so that flows downstream of a dam do not exceed the stated channel capacity. Hydrologists manage this flood-control pool to temporarily store the rain-flood runoff which would otherwise pass by a dam.

The goal is to keep flows downstream of the dam within their stated channel capacity so that flooding conditions are avoided. The maximum size flood that a dam can control is termed a dam's "level of flood protection." The maximum flow that the channel downstream can handle without causing flooding conditions is termed the "maximum objective flow."
The term “level of flood protection” does not imply or guarantee that communities downstream of a reservoir have any such level of flood protection. A reservoir’s estimated “level of flood protection” is based on the results of a probabilistic computer model. The model assumes that a certain set of conditions are present. If those conditions aren’t present during a real flood, then the maximum objective flow may be exceeded, and flooding could result.

The models typically use a fairly simple set of assumptions about conditions downstream of a reservoir. To assess the level of flood protection for communities downstream of a reservoir would require a full risk and uncertainty assessment of downstream hydrologic, hydraulic, geotechnical, and economic conditions.

A reservoir’s level of flood protection can change over time. An example is Terminus Dam which forms Lake Kaweah. When originally constructed in 1962, Lake Kaweah’s storage capacity was estimated to be sufficient to provide a 60-year level of flood protection downstream. But as sediment accumulated in the reservoir, the level of protection had decreased by 1978 to only a 46-year level of flood protection. When fuse gates increased the flood-control pool size of the reservoir in 2004, the level of protection increased to a 70-year level of flood protection (see Table 8).

As described above, the flood-control pool of a reservoir is used to manage rain-floods. However, the entire gross-pool capacity of the reservoir is theoretically available to control snowmelt floods.

**Landslides and Landslide Dams**

Landslides sometimes dam rivers and streams. In the Sierra, dams are occasionally formed by rockslides and debris flows as well.

The USGS has urged local risk management agencies to prepare for a return of a flood as big as the 1861–62 flood. We definitely don’t have plans in place for dealing with such an event.

However, the Tulare Lake Basin experienced a flood even bigger than the 1861–62 flood just six years later: the 1867–68 flood. It was a bigger flood on all four of our major rivers. In addition to being a major flood, the storm that brought on that flood was a soaking rain that lasted for upwards of six weeks. Much of the Sierra foothill and montane zones consist of steep hillsides of unconsolidated debris slopes. When these hillsides are soaked to depth, landslides can be triggered.

In a short intense storm like the December 2010 storm, we get many relatively small landslides and debris flows. But in an extended event like the 1867–68 storm, we get cataclysmic landslides. In the past, some of these have formed landslide dams across our major rivers that were up to 400 feet high.

When dams such as those fail, the results downstream can be catastrophic. For example, the residents of Bakersfield woke on New Year’s Day, 1868, to a 200-foot-high flood coming out of the Kern Canyon.

In the 1867–68 storm, the landslide dams on the Kaweah and Kern held the flooding rivers back long enough for the residents downstream to react and get out of the floodplain. In contrast, the dams on the San Joaquin River and Mill Flat Creek presented a less clear signal downstream, partly because those events happened at night.

The residents of Old Kernville and Weldon had about 24 hours’ notice because the Kern River stopped running. They were able to evacuate their towns before the river submerged them under about 50 feet of water. The residents of Millerton, the county seat of Fresno at the time, weren’t aware of what was happening. The disintegrating remnants of one or more landslide dams hit their town just before midnight on Christmas Eve, 1867, destroying it. That is why Fresno is now the county seat.

Just as USGS is urging us to prepare for a return of a storm similar to the 1861–62 flood, it might be prudent to prepare for a return of a storm similar to the 1867–68 flood, complete with landslides and landslide dams. This is especially true in high landslide hazard zones.
Long-term Temperature Changes

Popular perception is that North American temperatures have been relatively stable for the last thousand years or so and have recently begun an unprecedented climb in recent decades.

However Tom Swetnam and other researchers have found evidence that temperature anomalies corresponding to Europe’s Medieval Warm Period (approximately 950–1250) and Little Ice Age (approximately 1450–1850) occurred in the Tulare Lake Basin. (The Medieval Period aka the Middle Ages lasted from the 5th to the 15th century.)

These two different views of long-term temperature changes have been presented by others as the Battle of the Graphs:

![Battle of the graphs](image)

Figure 8. Comparison of two temperature reconstruction graphs.

Although the concept of the lower graph in Figure 8 may be correct in general concept, it exaggerates and greatly oversimplifies the situation. For one thing, it is based just on European conditions and probably overstates the magnitude of the temperature anomalies even there, especially during the Medieval Warm Period. Figure 9 is a more reliable and peer-reviewed presentation of the available data.

The Earth has warmed by about 1.4 degrees since 1880; that is one measure of global warming. The biggest change in California is not a global-warming-related increase in temperature. The biggest change stems from the extreme makeover we have done to much of our state. As a result, the average temperature in the San Joaquin Valley has increased more than 5 degrees since 1950. (These measurements were made prior to the recent warming that California experienced in 2013 and 2014.)

That increase has been caused by urban agricultural heat islands. The San Joaquin Valley is a prime example of this. Before the Central Valley Project, the California State Water Project, and many other smaller projects — the average rainfall here was 5 inches a year. It literally was a desert: the San Joaquin Valley Desert. Now the valley is heavily irrigated, and it is a major agricultural area.

But when you take a desert, which the San Joaquin Valley essentially is, and you make it wet, it starts to absorb heat. So now heat waves are longer, they are more intense, and they are more frequent. So the direct impact of such changes to the land by man has had a bigger impact in California than global warming.\(^{63, 64}\)
The temperatures in Figure 9 are based on reconstructions using multiple climate proxy records. They all show that the Northern Hemisphere experienced significant warming during the Medieval Warm Period (approximately 950–1250) as well as significant cooling during the Little Ice Age (approximately 1450–1850). However, the current global warming period that we’re experiencing is unprecedented in the past 1,300 years.

The year 2012 continued this trend in rising temperatures. For example, there were 3,215 daily high-temperature records set or tied during June. Among those records, 1,748 of them were for temperatures of 100 degrees or higher. July 2012 was the warmest month on record for the contiguous U.S., breaking the record set in July 1936. The warm July temperatures contributed to the warmest 12-month period that the nation has experienced since record-keeping began in 1895.

This was just one of a series of temperature records that has been recently set and broken.

Calendar year 2014 was remarkably warm in California, the West, the contiguous U.S., and for the Earth as a whole. This fits within a context of a long-term warming trend that has been going on for several centuries and has been accelerating in recent decades.

Globally, 2014 was the warmest year since record-keeping began in 1880, breaking the record set in 2010. Seven out of 12 months tied or topped previous monthly global temperature records. The year was 1.1 degrees above the 20th-century average. It was the 38th year in a row with global temperatures above the 20th-century average.

2014 set the new global temperature record in the absence of an El Niño, a phenomenon which raises global temperature. Many of the previous hottest years on record have occurred during El Niño years, including 2010 and 2005, which now share the record for the second hottest year.

2014 was the latest in a series of warm years, in a series of warm decades. February 1985 was the last month when global temperature fell below the 20th century monthly average, making December 2014 the 358th consecutive month (29.8 years) where the combined global land and ocean surface temperature was above average.

The 10 warmest years in the instrumental record, with the exception of 1998, have now occurred since 2000. Since 1880, Earth’s average surface temperature has warmed by about 1.4 degrees. The majority of that
warming has occurred in the past three decades. Each of the last three decades has been much warmer than the decade before.

In the contiguous U.S., 2014 was 0.5 degree above the 20th-century average. 2014 was the 18th year in a row with U.S. temperatures above the 20th-century average.

The 2014 heat in the U.S. was most pronounced in the West. Both water year 2014 and calendar year 2014 were the hottest in California since record-keeping began in 1895. California’s average temperature for calendar year 2014 was 61.5 degrees; breaking the record set in 1934 by an impressive 1.8 degrees. This was 4.1 degrees higher than the 20th century average.

In 2012, Adrian Das and Nate Stephenson at the USGS field station at Sequoia at Kings Canyon National Parks performed an analysis of long-term historical temperature records in the vicinity of the national parks. They analyzed data from the Lemon Cove, Ash Mountain, Independence, Bishop Airport, Grant Grove, and Lodgepole weather stations.

Their analysis showed that the long-term temperature records for the above six stations, taken as a whole, have generally followed patterns seen statewide: there has been an increase over the periods of record.

Laura Edwards and Kelly Redmond prepared an assessment of the climate for the Sierra Nevada Network parks (Yosemite, Sequoia, Kings Canyon, and Devils Postpile) in 2011. Among other things, they looked at trends in annual temperature for the five counties that encompass that area (Tuolumne, Mariposa, Madera, Fresno, and Tulare). In all cases, mean temperatures have been increasing since the mid-1970s. The maximum temperatures have risen a small amount in the past 75 years (since the mid-1930s), but the minimum (nighttime) temperatures have risen fairly dramatically. This behavior is seen elsewhere in California and throughout much of the Western U.S. Mean temperatures have been and continue to be on the rise primarily because minimum temperatures are increasing; although in recent years maximum temperatures show some small increases as well.

Peter Vorster said that average temperature change is not the best indicator if you are interested in the impact of temperature on water supply. The change in minimum temperatures, especially in the late spring and summer, has arguably had the greatest impact in California. The retreat of the glaciers in the Little Ice Age and the rising snowline in the 20th century and earlier spring snowmelt have been the result of increasing nighttime temperatures in spring and summer. Peter said that there are many studies, temperature records, and historical accounts to support this relationship.

Focusing on the Sierra Nevada, a 2004 study by Knowles and Cayan modeled the future possible reduction in snowpack for the 21st century. They demonstrated that the historical trends found in a 2005 study by Mote and others are projected to continue through the end of the century. Depending on the climate change scenario used, they predicted a 60%–80% reduction in April 1 snow water equivalent in 2070–2099 compared to the 1961–1990 period.
California Snow Conditions during the Little Ice Age

In 1542, Juan Rodríguez Cabrillo led the first European expedition to explore what is now the West Coast of the United States. The Cabrillo expedition sailed out of the port of Navidad, Guatemala (near modern day Manzanillo) on June 24, 1542. He sailed as far north as the Monterey Peninsula. Returning south on November 18 of that year, he noted the snow-capped Santa Lucia Mountains in southern Monterey County. That is not a place where we expect to see such conditions today.

Sebastian Vizcaino led an expedition that departed Acapulco on May 5, 1602, and arrived in the bay that he named Monterey on December 14 of that year. That expedition occurred during the height of the Little Ice Age (approximately 1450–1850).

Vizcaino spent a very cold Christmas in the area. He recorded that on Christmas Day, 1602, the mountains near the port were covered with snow and that on New Year’s morning the water holes were frozen to the depth of a palm. The expedition encountered no American Indians, but did find a deserted village. Vizcaino speculated that the inhabitants had taken refuge in the interior to escape the biting cold. (Today, the average low temperature in Monterey in December and January is a much more pleasant 43°.)

Vizcaino wanted to attract colonists to his new bay. So although he was numb with cold, he wrote a glowing report in which he said the area’s climate was like that of Seville’s. So untrue was the picture that he painted, that when Captain Gaspar de Portolá and Father Juan Crespi arrived with their colonizing party 167 years later, they failed to recognize the fabled port. Food was so scarce that they were reduced to eating seagulls and pelicans. After snow began to cover the hills on November 30, 1769, the survivors decided to return to San Diego.

Zenas Leonard was a fur trapper, a mountain man. He left St. Louis on April 24, 1831, with a fur trapping company. He returned to civilization some 4½ years later on August 29, 1835. He left us a highly readable account of his adventures: the Narrative of the Adventures of Zenas Leonard.

In July of 1833, Leonard joined with the Joseph Walker party in exploring the unknown country from the Great Salt Lake to the Pacific Ocean. With great difficulty, they traveled down the Humboldt River, crossed the Sierra in October 1833, followed the San Joaquin River downstream, and arrived in the vicinity of Half Moon Bay on November 20, 1833.

Their route down the Humboldt River took them to Carson Lake and the vicinity of present-day Carson City, Nevada. How they actually crossed the Sierra and reached the San Joaquin River is less clear. They are traditionally thought to have crossed north of the Merced River, becoming the first whites to discover Yosemite Valley. However, that route would have been well south of the Carson City area, and that would seem inconsistent with Leonard’s account.

Furthermore, Leonard described the section of the Sierra that the Walker Party passed through. Based on his research, George Durkee is convinced that the area Leonard described wasn’t Yosemite. Scott Stine’s research indicates that the Walker Party’s probable route was well north of Yosemite: over Ebbetts Pass, and generally along the route of present-day Highway 4. In any case, Leonard described encountering a lot of old and consolidated snow as they crossed the Sierra in October 1833:

> In some of these ravines where the snow is drifted from the peaks, it never entirely melts, and may be found at this season of the year, from ten to one hundred feet deep. From appearance it never melts on the top, but in warm weather the heap sinks by that part melting which lays next the ground. This day’s travel was very severe on our horses, as they had not a particle to eat...but the most of the distance we this day traveled, we had to encounter hills, rocks and deep snows. The snow in most of the hollows we this day passed through, looks as if it had remained here all summer, as eight or ten inches from the top it was packed close and firm — the top being loose and light, having fell only a day or two previous.

The Walker Party encountered snow that was persisting from year to year. They were crossing the Sierra near the end of the Little Ice Age. There is no longer persistent snow in that area; conditions have changed dramatically in the last 180 years.

The glaciers in the Sierra have advanced and retreated with each of the various ice ages. Scott Stine said that the Sierran glaciers reached their maximum extent during the Little Ice Age about 1850.
What Constitutes a Drought

Most people think of a drought as a period of unusually dry weather that persists long enough to cause problems such as crop damage and water supply shortages. But because dry conditions develop for different reasons, there is more than one definition of drought. Drought can be caused not only by a lack of precipitation, but by a lack of reserve supply (or imports), high temperatures, overuse, and overpopulation.

In 1985, two researchers, Donald Wilhite and Michael Glantz, uncovered more than 150 published definitions of drought. In an effort to bring some order to measuring drought, they grouped the various definitions of droughts into four basic approaches: meteorological, hydrological, agricultural and socioeconomic. The first three categories track drought as a physical phenomenon. The last category dealt with drought as a supply and demand problem.

The descriptions of these four categories have changed and broadened somewhat since Wilhite and Glantz first described them. But they are more or less as follows:

- **Meteorological drought.** Period of below-average precipitation. Meteorological drought is usually defined based on the degree of dryness (in comparison to the statistical average) and the duration of the dry period. Typically this is two or more successive years of less than average precipitation.

- **Hydrological drought.** Period of below-average runoff and water supplies. This is when the available water supply from all sources (streams, imports, reservoirs, and groundwater) falls below the statistical average. A hydrological drought usually occurs following a period of extended precipitation shortfall (a meteorological drought) that impact water supply (streamflow, imports, reservoir levels, and groundwater), potentially resulting in significant societal impacts. Human activities, such as drawdown of reservoirs or groundwater overdrafts, can worsen hydrological droughts.

- **Agricultural drought.** Agricultural drought links various characteristics of meteorological or hydrological drought to agricultural impacts, focusing on precipitation shortages, soil water deficits, reduced ground water or reservoir levels needed for irrigation, and so forth.

- **Socioeconomic drought.** Wilhite and Glantz described socioeconomic droughts as those when demand for water exceeds the available supply. Others have since started using this term to describe agriculture and other sector impacts due to meteorological or hydrological droughts. Used this way, socioeconomic drought links various characteristics of meteorological or hydrological droughts to socioeconomic impacts, focusing on precipitation shortages, soil water deficits, reduced reservoir levels or groundwater needed for irrigation, etc. This includes adverse effects on crop production or agricultural land ecology due to restricted water use. Socioeconomic drought is when the available water supply fails to meet our needs (the amount of water we choose to apply). Examples of this kind of drought include too little irrigation water at affordable prices or when low river flow forces hydroelectric power plant operators to reduce energy production sooner than normal. One could also say that our society’s water use has imposed a socioeconomic drought on the environment. For example, diversions from the Delta have imposed critical drought conditions (measured by alteration of unimpaired runoff out of the Delta) in more than 50% of the years or 80% of the years just on the San Joaquin River inflow to the Delta. The result of this environmental drought, which in turn is partially responsible for the decline in abundance of critical fish species, then leads to cutbacks in Delta exports or diversions from the San Joaquin River which in turn exacerbates the socioeconomic drought described herein.

Traditionally, the term “drought” has most commonly been used in the sense of a meteorological drought, a period of significantly less than average precipitation. That is generally how the term drought is used in this document. Typically this condition has to last for at least two years before it is recognized as a drought. This type of drought ends when precipitation returns to average or above-average conditions for at least one year or preferably two. The 1976–77 drought is an example of this type of drought.

The socioeconomic category of drought often has a political component to it. It is used to describe conditions when, for various reasons, we don’t have all the water that we feel we are entitled to. In this kind of drought, 1) available water supply fails to meet our needs (the amount of water we choose to apply), and 2) we perceive that water that is rightfully ours is being used somewhere else. If “they” would just let us have our water, we would be better able to meet our needs.

The Tulare Lake Basin has had a lot of the socioeconomic type of drought in recent years. That results in large part because our basin relies on a great deal of supplemental water imported from the San Joaquin River (via the Friant-Kern Canal) and from the Delta via the state and federal canals. Reduced imported supplies can stimulate a socioeconomic drought even when precipitation and runoff in the Tulare Lake Basin is not in a meteorological drought. It is more challenging to mark the end of this type of drought. Precipitation can return.
to average or even above-average conditions, but there still isn’t enough water to meet our needs (the amount of water we choose to apply). The latter part of the 2007–09 drought is an example of this type of drought.

In the 2007–09 drought, water years 2007 and 2008 really were drought years by the traditional definition, they constituted both a meteorological and a hydrological drought. The runoff was so low in those years that the state’s water year index rated those years as critically dry. The 2007–09 drought was California’s first drought for which a statewide proclamation of drought emergency was issued.

That turned out to be critical. When precipitation returned to near-average or above-average, it was hard politically for the governor to declare an end to the drought. There clearly wasn’t enough water to go around, there wasn’t enough to meet our needs. It wasn’t until March 30, 2011, after an incredibly wet winter, that the state of emergency was finally rescinded. That was long after the end of the meteorological drought. The 2007–09 drought had morphed from the traditional meteorological type of drought (below-average precipitation) into the socioeconomic type (we’re entitled to more water than we’re getting).

Unlike floods, droughts are not clearly defined. Identifying periods of drought is a matter of subjective interpretation, even in retrospect. The droughts noted in this document are generally those that meet both of the following criteria:

- recognized at a statewide level or identified in a peer-reviewed scientific publication
- appear to have impacted the Tulare Lake Basin in some way, such as being reflected in the annual flows of the rivers within that basin

Most water users in California are cushioned to some degree from drought. We have developed a number of conveyance and storage sources and have invested in redundant systems in many areas. However, it is generally not cost-effective to provide this level of secondary storage for rural areas. As a result, rural areas have to put a greater reliance on rain for their principal water supply. That leaves them very vulnerable to drought.

Impacts of drought are typically felt first by those most reliant on annual rainfall — ranchers engaged in dryland grazing, rural residents relying on wells in low-yield rock formations, or small water systems lacking a reliable water source. (Criteria used to identify regional and state drought conditions generally do not address such localized impacts.)

It comes down to the cost of developing secondary sources for such rural areas and the memory of those that are now alive. John Steinbeck grew up in the Salinas Valley, just west of the Tulare Lake Basin. He captured this view well in his greatest novel, East of Eden:

> I have spoken of the rich years when the rainfall was plentiful. But there were dry years too, and they put a terror on the valley. The water came in a thirty-year cycle. There would be five or six wet and wonderful years when there might be nineteen to twenty-five inches of rain, and the land would shout with grass. Then would come six or seven pretty good years of twelve to sixteen inches of rain. And then the dry years would come, and sometimes there would be only seven or eight inches of rain. The land dried up and the grass headed out miserably a few inches high and great bare scabby places appeared in the valley. The live oaks got a crusty look and the sagebrush was gray. The land cracked and the springs dried up and the cattle listlessly nibbled dry twigs. Then the farmers and the ranchers would be filled with disgust for the Salinas Valley. The cows would grow thin and sometimes starve to death. People would have to haul water in barrels to their farms just for drinking. Some families would sell out for nearly nothing and move away. And it never failed that during the dry years the people forgot about the rich years, and during the wet years they lost all memory of the dry years. It was always that way.

East of Eden was published in 1952, but Steinbeck’s message is still resonant today. Human memories are short; we forget our rainfall and drought history all too fast.

**Measurements of Drought**

There is no universal definition of when a drought begins or ends. There are many measurements of precipitation, runoff, and soil moisture in the Central Valley. To various degrees, these measurements can be used as indicators or proxies for drought conditions in the Tulare Lake Basin.

Runoff is an indicator of precipitation since about 24% of total precipitation appears in the runoff of our streams and rivers. (See the section of this document that discusses Where does precipitation wind up?) Based on an
analysis of the nine modern droughts shown in Table 22, it appears that droughts in the Tulare Lake Basin tend to begin when the combined annual runoff for our four major rivers is less than 75% of the long-term average for two years in a row. They tend to end when the annual runoff is at least 75% of the long-term average for two years. That’s not a hard and fast rule, but it seems to generally fit with how we identify our droughts.

The State of Washington is one of the few Western states with a state statutory definition or process for defining or declaring drought. They defines a drought condition as when water supply for an area is below 75% of normal and the water shortage is likely to create undue hardships for various water uses and users. That seems consistent with our experience in the Tulare Lake Basin.

An indicator of California’s hydrology and the annual surface water supplies is the amount of water that flows into the state’s major rivers. For the central portions of California, the Sacramento River basin and San Joaquin River basin indices have been used for many years to evaluate the amount of available surface water:

- the Sacramento Valley Water Year Index
- the San Joaquin Valley Water Year Index

Many decisions about annual water requirements for the Delta are based on these indices, as are the amounts of surface water supplies available to many agricultural and urban regions of the state.

Those indices measure unimpaired natural runoff from 1906 to the present for the Sacramento River Basin and from 1901 to the present for the San Joaquin River Basin. As part of those indices, DWR categorizes each year compared to average runoff.

The San Joaquin Valley Water Year Index is based on the combined unimpaired flows of four rivers: the Stanislaus, Tuolumne, Merced, and San Joaquin Rivers. The index for the current year is not a measure of the unimpaired runoff in that year. Because this index is designed to reflect drought conditions, it considers the flow from the previous water year as well as the flow from the current water year. It is an index of the current year’s runoff, seasonal precipitation prior to snowmelt, and the previous year’s index. The formula for the index is 0.6 * current April–July runoff forecast (in million acre-feet) + 0.2 * current October–March runoff in (million acre-feet) + 0.2 * previous water year’s index.

For the combined flow of those four rivers, the hydrologic index classification (water year type) is grouped as:

- wet (equal to or greater than 3.8 million acre-feet)
- above normal (between 3.1 and 3.8 million acre-feet)
- below normal (between 2.5 and 3.1 million acre-feet)
- dry (between 2.1 and 2.5 million acre-feet)
- critical (less than 2.1 million acre-feet)

Table 3 summarizes the number of years in each of the hydrologic index classifications illustrated in Figure 10. For example, the San Joaquin River Basin has experienced 21 critically dry years during the past 114 water years (1901–2014).

<table>
<thead>
<tr>
<th>Water Year Classification</th>
<th>Number of Years</th>
<th>Percent</th>
</tr>
</thead>
<tbody>
<tr>
<td>Wet</td>
<td>37</td>
<td>32%</td>
</tr>
<tr>
<td>Above normal</td>
<td>22</td>
<td>19%</td>
</tr>
<tr>
<td>Below normal</td>
<td>18</td>
<td>16%</td>
</tr>
<tr>
<td>Dry</td>
<td>16</td>
<td>14%</td>
</tr>
<tr>
<td>Critical</td>
<td>21</td>
<td>18%</td>
</tr>
<tr>
<td>Total</td>
<td>114</td>
<td>100%</td>
</tr>
</tbody>
</table>
Figure 10 illustrates the San Joaquin Valley Water Year Index during the past 114 water years: 1901–2014. That index is a measure of runoff in the San Joaquin River Basin.

Figure 10. San Joaquin Valley Water Year Index for past 114 years: 1901–2014.
Source: California Department of Water Resources
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Figure 11 summarizes the relative frequency of different water year types in the San Joaquin River Basin. It is based on the San Joaquin Valley Water Year Index shown in Figure 10.

![Frequency of Different Water Year Types](image)

**Figure 11.** Frequency of different water year types in the San Joaquin River Basin (1901–2014).

The San Joaquin Valley Water Year Index is based on the runoff of rivers in the San Joaquin River Basin, and reflects drought conditions in that basin. The Tulare Lake Basin is the next basin to the south and there is frequently a similarity between our droughts and those that occur in the San Joaquin River Basin.

However, there have been multiple occasions when our two basins have experienced different drought conditions. The San Joaquin Valley Water Year Index does not always represent drought conditions in the Tulare Lake Basin. The index can be used with care as a proxy for the Tulare Lake Basin droughts. Think of it as just one more piece of information. Wherever possible, it should be used in conjunction with some other proxy such as the runoff of the major rivers in our basin.

Another proxy frequently used in this document is the flow of the upper San Joaquin River at the inflow to Millerton Lake. Peter Vorster did a quick and dirty comparison of the unimpaired runoff records for the Upper San Joaquin and the Kern. It appears that the droughts are correlated enough to use it as a proxy. The very wet years on the Kern are generally much wetter than on the San Joaquin River in terms of % of average.

Dry years happen periodically; sometimes dry conditions persist over multiple years, eventually resulting in sufficient impacts for these dry conditions to be termed a drought. It is useful to identify which individual years are dry years. However, from a human or environmental perspective, a single dry year does not constitute a drought. Droughts occur slowly, over a multi-year period. That is generally how this document treats droughts, as multi-year events.

Droughts in the San Joaquin River Basin seem to occur when DWR’s San Joaquin Valley Water Year Index is dry or critical for two years in a row. That index is frequently — but not always — applicable to droughts that occur in the Tulare Lake Basin.

Researchers often focus on estimating the dryness of individual years. This can be thought of as single-year droughts as opposed to the multi-year droughts discussed above.

The Palmer Drought Severity Index (PDSI) is a widely used indicator of long-term (meteorological) drought severity. PDSI is based on a supply-and-demand model of soil moisture. Supply is precipitation and stored soil moisture. Demand is the potential evapotranspiration, the amount of water needed to recharge the soil, plus the runoff needed to keep rivers, lakes, and reservoirs at an average level.

The PDSI serves as an estimate of soil moisture deficiency and roughly correlates with drought severity. It is the most commonly used index for drought monitoring and research. It has been widely used in tree-ring-based reconstructions of past droughts in North America and other regions.
The term “climatic water deficit” (CWD) is the amount of water by which potential evapotranspiration exceeds actual evapotranspiration. CWD effectively integrates the combined effects of solar radiation, evapotranspiration, and air temperature on watershed conditions given available soil moisture derived from precipitation. CWD can be thought of as the amount of additional water that would have evaporated or transpired had it been present in the soils given the temperature forcing. In a Mediterranean climate like that of the Tulare Lake Basin, CWD can be thought of as a surrogate for water demand based on irrigation needs. Changes in CWD effectively quantify the supplemental amount of water needed to maintain current vegetation cover, whether natural vegetation or agricultural crops. PDSI is a quick and dirty way of trying to get at CWD. PDSI is a biologically meaningful measure of drought severity, but CWD is even more biologically meaningful.

The PDSI combines temperature, precipitation, evaporation, transpiration, soil runoff, and soil recharge data for a given region to produce a single number that indicates drought conditions. As shown in Table 4, the index uses a 0 as normal, and drought is shown in terms of negative numbers; for example, values less than negative 4 indicate extreme drought. Palmer’s algorithm is also used to describe wet spells, using corresponding positive numbers.

<table>
<thead>
<tr>
<th>Palmer Index</th>
<th>Classification</th>
</tr>
</thead>
<tbody>
<tr>
<td>4.0 or more</td>
<td>extremely wet</td>
</tr>
<tr>
<td>3.0 to 3.99</td>
<td>very wet</td>
</tr>
<tr>
<td>2.0 to 2.99</td>
<td>moderately wet</td>
</tr>
<tr>
<td>1.0 to 1.99</td>
<td>slightly wet</td>
</tr>
<tr>
<td>0.5 to 0.99</td>
<td>incipient wet spell</td>
</tr>
<tr>
<td>0.49 to -0.49</td>
<td>near normal</td>
</tr>
<tr>
<td>-0.5 to -0.99</td>
<td>incipient dry spell</td>
</tr>
<tr>
<td>-1.0 to -1.99</td>
<td>mild drought</td>
</tr>
<tr>
<td>-2.0 to -2.99</td>
<td>moderate drought</td>
</tr>
<tr>
<td>-3.0 to -3.99</td>
<td>severe drought</td>
</tr>
<tr>
<td>-4.0 or less</td>
<td>extreme drought</td>
</tr>
</tbody>
</table>

In 1999, the U.S. Drought Monitor replaced the PDSI as the nation’s drought indicator. The Drought Monitor uses a lot of different tools to assess drought. The Drought Monitor categorizes drought into five levels of severity:
- abnormally dry (category D0, corresponding to a PDSI between -1.0 and -1.9)
- moderate drought (D1, PDSI between -2.0 and -2.9)
- severe drought (D2, PDSI between -3.0 and -3.9)
- extreme drought (D3, PDSI between -4.0 and -4.9)
- exceptional drought (D4, PDSI between -5.0 and -5.9)

Relationship of Temperature and Drought

In a study published in 2015, a team of Stanford researchers led by Noah Diffenbaugh examined the role that temperature has played in California droughts. Higher temperatures reduce snowfall and increase snowmelt; they decrease soil moisture, increase evaporation, and intensify our dry season. All of these accentuate the effects of below-normal precipitation.

The Stanford study took advantage of a recently released trove of monthly precipitation, temperature and drought data for California for the 119-year period 1896–2014. Using that observed dataset, the scientists calculated the probability of drought years occurring in different temperature and precipitation conditions. The Stanford researchers used the Palmer Modified Drought Index (PMDI), a variant of PDSI, as their primary drought indicator. They calculated PMDI for each of the past 119 years, with a year defined as August 1 – July 31. (This is a slightly different definition than is used for water year or weather year.)

California’s average precipitation has not been declining over this period. What has been changing is an increase in temperature, especially in recent decades (see the section of this document that describes Long-term Temperature Changes).

The Stanford team analyzed California’s 119-year historic precipitation and temperature record from 1896–2014. They identified the most severe drought years as those with a negative PMDI anomaly exceeding -1.0 Standard Deviations, 1-SD drought for short. There have been 20 such drought years during this 119-year period.
Figure 12 shows the results when the Stanford research team used PMDI to calculate statewide drought severity for the past 120 years.

Figure 12. California drought severity (PMDI index) for past 119 years: 1896–2014.
Source: Noah Diffenbaugh, School of Earth, Energy & Environmental Sciences, Stanford University
The researchers found that years with 1-SD droughts have occurred approximately twice as often in the past two decades as in the preceding century:

- 14 events in 1896–1994 = 14% of years
- 6 events in 1995–2014 = 30% of the years

Most 1-SD droughts have occurred when conditions were both dry (precipitation less than the long-term average) and warm (temperature above the long-term average). This includes:

- 15 of the 20 1-SD drought years during 1896–2014
- all 6 of the 1-SD drought years that occurred during 1995–2014

Similarly, dry years were much more likely to produce a 1-SD drought if they occurred in warm years. Years that were both warm and dry were about twice as likely to produce a severe drought as years that were cool and dry. For the 63 warm years during 1896–2014:

- 5 of the 31 cool-dry years (16%) produced a 1-SD drought.
- 15 of the 32 warm dry years (47%) produced a 1-SD drought.

The probability of dry years occurring during a warm year has been greater in the past two decades than in the preceding century. There has been more than a doubling of the frequency of warm-dry years in California:

- During 1896–1994, only 41% of the years were warm-dry years.
- During 1995–2014, 91% of the years were warm-dry years.

The probability of dry years producing 1-SD droughts has been approximately twice as great in the past two decades as in the preceding century.

- During 1896–1994, only 27% of dry years resulted in 1-SD droughts.
- During 1995–2014, 55% of dry years resulted in 1-SD droughts.

These increases in drought risk have occurred despite a lack of substantial change in the occurrence of low or moderately low precipitation years. Dry years are not becoming more frequent. Dave Meko found the same thing when he reconstructed precipitation for the San Joaquin River Basin. The average of recomputed flows for the period 900–2012 is very similar to the average of actual flows since record-keeping began.\(^{89}\)

The observed doubling in the frequency of years with 1-SD droughts and the observed doubling in the probability of dry years producing 1-SD droughts are both due to the rising temperatures in the state. There has been nearly a doubling of the frequency of warm years in California:

- During 1896–1994, 45% of years were warm years.
- During 1995–2014, 80% of years were warm years.

There are far more warm years, so dry years are routinely occurring during warm years. Those are the conditions that have produced the most severe droughts in the past.

The team's results "strongly suggest that global warming is already increasing the probability of conditions that have historically created high-impact drought in California."

The team assessed climate simulations to see what this likely means for California's future:

- It appears that the situation is set to get worse. The team predicted that a continuing rise in global temperatures will greatly increase the probability of warm and dry conditions occurring together. Historically those have been the most severe droughts in California. This means that both drought frequency – and the potential intensity of those droughts which do occur – will likely increase as temperatures continue to rise.
- Essentially all years are likely to be warm – or extremely warm – in California by the middle of the 21st century. Nearly every year that has low precipitation will also have warm temperatures.
- More frequent warm years also increase the likelihood of multiyear droughts in the future. According to the team's analyses, the 2012–15+ drought is one of the longest consecutive periods in the historical record during which conditions were both severely dry and severely warm. The climate models also indicate that such conditions will become even more common if global warming continues in the future, as the state enters a regime in which there is a nearly 100% risk that every year is warmer than conditions experienced during the 20th century.
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Cedar Grove Flooding

Challenge of Modeling Flows

The impetus for preparing this document was a project to replace the bridge over the South Fork of the Kings River in Cedar Grove. (That bridge would eventually be replaced beginning in the fall of 2010. The new bridge opened for use in the spring of 2013.) To design the replacement bridge, the national parks needed to know the magnitude of floods on that reach of the river. Unfortunately, there are no stream gages along the portions of the Middle and South Forks of the Kings River that are within the national parks.

The nearest long-term gaging station is far downstream at Pine Flat Dam on the mainstem of the Kings River. (Fragmentary data are available from a collection of gages on the South Fork Kings below the park. See the section of this document that describes the Stream Gages on the South Fork Kings.)

Therefore, in 2006, the Federal Highway Administration (FHWA) estimated the probable floodflows by modeling the hydrology of the Cedar Grove drainage basin. The results of the peak flow discharge computations are shown in Table 5.  

<table>
<thead>
<tr>
<th>Flood Recurrence Interval</th>
<th>Likelihood of Flood</th>
<th>Peak Discharge</th>
</tr>
</thead>
<tbody>
<tr>
<td>50 years</td>
<td>2% chance/year</td>
<td>13,300 cfs</td>
</tr>
<tr>
<td>100 years</td>
<td>1% chance/year</td>
<td>18,500 cfs</td>
</tr>
</tbody>
</table>

Those peak flow discharges are calculations of how big the model projected a flood event would likely be. Bridges like the Cedar Grove Bridge are typically designed for a 50-year-flood event because that is a comparatively rare occurrence. At the same time that the model was being constructed, the national parks were working to rebuild the actual flood history for that reach of the South Fork Kings. Our goal was to ensure that the new bridge would span the 50-year flood event. We wanted to stop interfering with the natural processes of this designated Wild and Scenic River. To our surprise, we discovered that floods of this magnitude (floods that would cover an area equivalent to the 50-year-flood predicted by the model) were a relatively frequent event.

According to the Environmental Assessment for the Cedar Grove Bridge,


Obviously a flow event that occurs every eight years on average (nine times in 70 years) is significantly different from an actual 50-year flow event. (Using more complete data, we would later learn that Cedar Grove has experienced these modeled 50-year flow events at least 13 times in the past 70 years, an average of once every five years.) How to explain the huge discrepancy between the modeled flows and the observed flows? FHWA used a reliable model, so presumably the difference resulted from one or more of the three model inputs:

1. The size of the drainage basin: approximately 357 square miles.
2. A U.S. Geological Survey (USGS) regional regression equation for the Sierra that predicts the frequency of floods for given rainfall values.

One possible source of error was the USGS regression equation. It was developed to reflect average conditions across the Sierra; it does not necessarily reflect the actual conditions of the South Fork Kings. Another likely source of error was the NOAA precipitation data. This part of the Sierra has very few precipitation gages. That is a serious problem because some of our big floods are caused by localized storm events.

When the national parks started the bridge design process, it would have been very helpful to have had a description of the flood history of the South Fork Kings; but nothing like this document existed. The parks were always playing catch-up. The model results were completed while reconstructing the flood history was still a work in progress. That meant that the actual flood history was not on hand to validate the model results: a serious shortcoming.

Completeness of Flood History

We have a very incomplete record of Cedar Grove floods. Many of the floods occur in the winter when no one is present in Cedar Grove to observe the event. Just as important, the national parks have no records.
management system in place for routinely recording flood events when they are observed. For example, we know about the 1937 flood in Cedar Grove only because the U.S. Forest Service (USFS) was designing a new bridge and recorded the high-water mark on their design drawings. We are very lucky that those drawings survived.

All that we know of the 1943 and 1945 floods at Cedar Grove is what we can infer from conditions some 50 miles downstream at Piedra. (We didn’t have that information until June 2011, well after construction of the new Cedar Grove Bridge had begun.)

We initially learned about the 1950 flood only because it swept away a major bridge downstream in Reedley, and because of an NPS Washington Office memo documenting the damage in Cedar Grove and elsewhere in the national parks. We were very lucky that memo survived. We know about the 1955 flood in Cedar Grove only because Wayne Alcorn flew into that area after the flood and photographed the damage. (Helicopters weren’t that common in 1955. Landing one in Cedar Grove brings to mind the opening scene of the 1972–1983 CBS M*A*S*H* television series.) We are very lucky that those photographs could be found.

All that we know of the 1966 flood at Cedar Grove is what we can infer from conditions downstream at Pine Flat Dam. (We didn't have the data to calculate the flood recurrence intervals in Table 30 until March 2011, well after construction of the new Cedar Grove Bridge had begun.)

We know about the 1969 flood on the South Fork Kings because it swept away a major trail bridge, and because we were fortunate to record Jim Harvey's recollections of the aftermath of that flood before he retired.

We know about the 1978, 1982, and 1983 floods in Cedar Grove only because Jerry Torres (Kings Canyon National Park's trails supervisor at the time) was there to observe them; we have no surviving documentation. CalTrans and Jeff Manley (a national park employee) flew into Cedar Grove after the 1997 flood, took lots of photographs to document the event, and yet none of those photographs survived. Thanks to Jerry Torres, we now know that there are some other surviving photographs of damage from that event, but we have yet to find them.

This incomplete record of floods suggests that Cedar Grove may have experienced more "50-year floods" in the last 70 years than the nine that we could document at the time that the environmental assessment was printed. For example, it now appears that the 1943, 1945, 1963, and 1980 floods should have been included in the list of modeled (as opposed to actual) 50-year floods. By compiling this document, the expectation is that the next project team won’t have to start from scratch. They will at least have a general idea of the flood history of the rivers within the Tulare Lake Basin.

**Summary of Flood History**

Pine Flat Dam on the Kings River was completed in 1954. The history of floods on the South Fork Kings prior to 1954 can be inferred from sources such as the California Water Plan and historical records. Some of the bigger floods from that earlier period were 1861–62, 1867, 1890, 1893, 1906, 1914, 1937, and 1950. Several of those would probably have qualified as actual (as opposed to modeled) 50-year floods. At least two (the 1861–62 and the December 1867 floods) appear to have qualified as 100-year flood events. The 100-year floodplain has been mapped in Cedar Grove, at least in general outline. The Cedar Grove Lodge appears to lie just within it, and the Cedar Grove Visitor Center appears to be just outside.

With completion of the Pine Flat Dam, much more reliable gaging data became available for the Kings River. In the absence of flow data specific to the South Fork Kings, it is not unreasonable to expect that the floodflows in Cedar Grove are roughly proportional to those downstream at Pine Flat. If that is the case, then Cedar Grove has probably experienced two actual (as opposed to modeled) 50-year or bigger floods since the dam was completed: December 23, 1955 and December 6, 1966. See Table 30 for the actual ratings (the flood recurrence intervals) of these and other floods.

**Stream Gages on the South Fork Kings**

USGS operated a recording stream gage on the South Fork Kings from 1950–1957. It was located 0.3 mile downstream from Grizzly Falls, across the highway from the big rock deposit that resembles a terminal moraine. It was installed on the downstream side of a big boulder that may have been a glacial erratic. Evidence of that stream gage is still visible today. The bottom section of the 3-foot-wide corrugated metal stilling well protrudes from the streambed alongside that big boulder (photograph on file in the national parks). There is a bolt sticking out of that boulder that was either an elevation pin or part of the attachment for the catwalk to the stilling well.
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Either because of good planning or incredible good luck, that gage began operation on November 16, 1950, just as the second and biggest of the 1950 storms struck. It continued to operate through September 30, 1957. The data from November 16–30, 1950 is available in a report that covered the 1950 flood. The data from December 1, 1950 – September 30, 1957 is available online from the USGS. This gage is known in USGS parlance as USGS 11212500 SF Kings River near Cedar Grove CA.

There has never been any electricity in this section of the canyon. Alternative power technology was rather primitive in the 1950s. This gage was probably operated by a hand-wound mechanical timer. This is the only recording stream gage ever installed on this reach of the Kings. The USFS is considering reestablishing a recording stream gage on this stretch of the river.

In addition to this stream gage, there have been two pairs of crest-stage gages, two staff gages, and a cable car on the South Fork Kings. USGS has a website that explains the various types of gages.

Crest-stage gages provide information on the highest flow since the gage was last visited. They are similar in concept to a min/max thermometer. A pair of crest-stage gages are located in the South Fork Kings on separate boulders where the old USGS stream gage was located (multiple photographs on file in the national parks). Another pair of crest-stage gages is located on the Grizzly Creek culvert, one on each end (multiple photographs on file in the national parks).

The lower pair of crest-stage gages was operated by USGS through December 6, 1966. Presumably those gages were installed in October 1957 when the stream gage was removed. The data from the crest-stage gages are available online from USGS. For the sake of continuity, the data were reported in the USGS database using the same database code and name as the old USGS stream gage USGS 11212500 (SF Kings River near Cedar Grove CA).

The Grizzly Creek station was operated from April 12, 1960 – May 20, 1973. Those data are also available online from USGS. The station is known in the USGS database as USGS 11212450 (Grizzly Creek near Cedar Grove CA).

All four of these crest-stage gages are still in place and functional to varying degrees. Bill Templin has data that he has collected from them in recent years.

USGS also installed a cable car over the South Fork Kings. That cable car was located about ½ mile downstream from Grizzly Falls. The exact location of this former cable car installation is unknown.

Staff gages are located on the Lower South Fork Bridge (the lowermost bridge across the South Fork Kings within the national park). The gage numbers are painted on both ends of the center bridge pier. Floodwaters gradually erode away the paint, making the gage more difficult to read. Bill Templin last painted the downstream numbers in 2011. The upstream gage is the less accurate of the two. That gage is affected by the hydraulic jump when the flow hits the edge of the pier, making it difficult to accurately read the numbers. Although the staff gage has traditionally been located on this bridge, this is not the ideal bridge for that purpose. The problem with using this bridge for discharge measurements is that the bridge crosses the river at a severe angle, and a good discharge measurement needs to be made perpendicular to the flow line. Ideally the staff gage should be placed on a bridge that crosses the river perpendicular to the flow.

Bill Templin rehabbed both of the staff gages on the lower Cedar Grove Bridge in 2012. He installed ceramic plates on the downstream gage and repainted the upstream gage.

There are three bridges over the South Fork Kings in the national park. In 2013, the Federal Highway Administration (FHWA) installed stream gages on the Lower and Upper South Fork Bridges that can be read remotely. The middle bridge (the new Cedar Grove Bridge) is unsuitable for a stream gage because of the riprap that has been placed around the bridge piers.

As part of the California Cooperative Snow Survey Program, the Kings River Water Authority installed a stream gage on the Upper South Fork Bridge in 2014 that can be read remotely.
Tulare Lake and other Valley Lakes

Geologic History of the Tulare Lake Basin
The Sierra Nevada batholith began intruding into the overlying rocks near the end of the Jurassic period (roughly 145 million years ago). A general regional uplift at this time was followed by a long period of erosion, during which large areas of granite were exposed and the metamorphic rocks were left as roof pendants. There were subsequent elevations of the region, the greatest being at the close of the Pliocene era (about 2.6 million years ago), when the entire Sierra Nevada was uplifted and tilted to form the present range, and the present cycle of erosion was initiated.

The San Joaquin Valley was once part of a large, ocean-covered basin. About 50 million years ago, parts of the current San Joaquin Valley, particularly north of present-day Coalinga, rose above sea level for the first time and the Coast Ranges were uplifted. By five million years ago, the seaways connecting the valley with the Pacific Ocean had all closed, leaving the San Joaquin Valley an isolated inland sea.

About 700,000 years ago, the continued filling of the valley by sediment from the Sierra and Coast Range mountains left Lake Corcoran (aka Lake Clyde) as the last remaining widespread ancient lake. It extended about 250 miles from present-day Bakersfield to Stockton. Lake Corcoran was connected to the Salinas River, which drained into Monterey Bay.

The Corcoran Clay was deposited in Lake Corcoran between about 600,000–800,000 years ago during the Pleistocene. The Long Valley Caldera eruption occurred about 747,000 years ago near present-day Mammoth Lakes. Tephra from that eruption (known as the Bishop Tuff) is associated with the Corcoran Clay.

The Corcoran Clay underlies 6,600 square miles of the San Joaquin Valley. It extends from Stanislaus County in the north, to Kern County in the south. In the Visalia area, that clay layer extends from near Highway 99, west toward the Coast Ranges. The top of the Corcoran Clay is up to 900 feet deep, and it is up to 200 feet thick. It separates a deeper aquifer of high-quality water from a shallower aquifer that in some places is lower quality.

The younger sediments of the foothills in the Kaweah River Basin include old terrace deposits and recent valley fill alluvium. Terrace deposits occur on remnants of elevated benches along both sides of the valley from Terminus Dam upstream to Three Rivers. The deposits are generally less than 20 feet thick and are so deeply weathered that they are relatively impervious. Included granite boulders are crumbly, and the feldspar-rich sand matrix has been thoroughly decomposed. These terrace deposits were probably formed early in the Pleistocene (up to 2.6 million years ago). The recent valley fill deposits are unconsolidated sands and gravels averaging about 20 feet deep at the location of Terminus Dam. This alluvium increases in depth downstream from Terminus Dam and extends westerly out over the river delta as alluvial fan deposits.

The San Joaquin Valley Basin has accumulated up to six vertical miles of unconsolidated marine and continental sediment. The top 2,000 feet of these sediments consist of continental deposits that generally contain freshwater. By the end of the Pleistocene, roughly 12,000 years ago, the valley was completely filled except for three depressions which were occupied by Tulare, Kern, and Buena Vista Lakes.

General Notes on Tulare Lake
At the time of Euro-American settlement, there were five generally recognized valley lakes. In order, stretching upstream from near present-day Lemoore to Bakersfield, those lakes were:

1. Summit Lake
2. Tulare Lake
3. Goose Lake
4. Buena Vista Lake
5. Kern Lake

These lakes were the anchors of a wetland complex of over 400,000 acres (see Figure 5). That complex constituted the largest single wetland in California. It connected with the wetlands that fringed the San Joaquin River, making a continuous wetland all the way to the Sacramento–San Joaquin Delta.

Those lakes supported an extensive fringing tule marsh. The tules (also known as bulrushes) grew in very dense stands. The plants were up to 20 feet tall and the stems were 2–3 inches in diameter. The tule-bed around
Floods and Droughts in the Tulare Lake Basin
General Flood and Drought Notes

Tulare Lake was typically a hundred yards or more wide. However, at the south end of the lake, the tule-bed extended outward from the shore for about 15 miles.

Floating islands of tules, many large enough to support the weight of several people, are reported to have drifted windblown across Tulare Lake’s surface. Where the rivers and streams entered the lake, dense thickets of willows and buttonbush (Cephalanthus occidentalis) crowded the shore.\(^{104}\)

John W. Audubon (one of John James Audubon’s sons) described his encounter with the tules when he traveled through the San Joaquin Valley in November 1849.

> Following down the San Joaquin southwest and west, we came to the river of the lakes, and stood off northwest (its general course) for nearly two days, but were so impeded in our progress by the bull-rushes that we turned aside to a clump of trees, where we expected to find water and grass; but not succeeding, returned to the river, about eight miles, and with great difficulty reached the edge of it for water at dusk cold, tired, and regretting our lost time. We resolved, nevertheless, to steer off from the rushes next day. This is the locality from which, I suppose, the valley takes its name, ”tulare“ meaning ”rush,” this plant taking here the place of all others.

> There is no trail but that of wild horses and elk, all terminating at some water-hole, not a sign of civilization, not the track of a white man to be seen, and sometimes the loneliness and solitude seem unending.

The Kings, Kaweah, and Kern Rivers all divide into distributaries in their lower reaches, spreading out to form a ”perfect swamp” according to early travelers.\(^{105}\)

Another traveler passing through the San Joaquin Valley during that period observed:\(^{106}\)

> The San Joaquin? What, that is the end of the World.

Tulare Lake fluctuated in size depending largely on the amount of runoff coming from the Sierra. In very wet years, it could grow to at least 790 square miles. That is over four times larger than Lake Tahoe (192 square miles), easily qualifying it as the largest freshwater lake west of the Mississippi. One author described what it was like to be on Tulare Lake in the mid-1870s:\(^{107}\)

> At this time Tulare Lake was beautiful to look at. It was nearly always as smooth as glass and the water as clear as crystal. One could take a small boat and row out far from shore where the sensation would be that of swinging in midair. During any hour of the day the outlook was a pleasing one. Far to the east the outlines of the Sierra Nevada Mountains could be seen rising above a pale blue mist that hung over the vast stretch of lowland. To the south the Tehachapi Range loomed up, and on the west the Coast Range looked as close as if you could put your hand on it. But how deceiving it all was, for the nearest mountain was at least 25 miles away.

It is possible that Tulare Lake was large enough to induce some lake-effect snow. Gary Sanger said that lake-effect snow typically occurs immediately downwind of the shoreline as the airmass transitions from being over the relatively warm water surface to the cooler land. The main question is whether the moving airmass would have enough time to pick up moisture from the Tulare Lake surface. The strongest winds likely would be moving too fast to pick up appreciable moisture; slower winds would have time to pick up more moisture. Gary said that if there was lake-effect snow, it would probably have fallen on the valley floor in Tulare County as well as the downwind Sierra.

Although classified as a freshwater lake, Tulare Lake tended to be more brackish and alkaline along its margins and when it became very shallow, particularly so in dry years. Orlando Barton described pioneers driving their wagons \(\frac{1}{2}–\frac{3}{4}\) mile into the lake in order to fill their water tanks with fresh, cold water from the inflow of the rivers.\(^{108}\) Sounds like the old-time equivalent of a visit from the Culligan Man.

Various sources purport that Tulare Lake was at one time or another upwards of 80 or 100 miles long.\(^{109}\) Those numbers are improbably large and belong in the realm of legend. John C. Fremont reported that the lake was about 60 miles long when he explored it in the winter of 1845–1846. S.T. Harding concluded that this 60-mile figure probably included flooded areas outside of the actual lakebed, such as the Fresno Slough area.
George H. Goddard published a map of the State of California in 1857 that showed Tulare Lake as being about 40 miles long. The Goddard map was produced by order of the Surveyor General of California and may be one of the more reliable source documents from that time period. According to the 1892 Thompson Historic Atlas Map, the lake had a length of 44 miles during the early years of American occupation. During the time of highest water, the lake’s maximum width was about 22 miles. The shoreline was quite irregular.

See Figure 4 for an overview map of Tulare Lake and the entire basin. At its maximum size, the eastern shore of Tulare Lake was within five miles of present-day Highway 99 near Pixley and the south shore was near the state historical monument at Allensworth. The western boundary was at present-day Kettleman City. Lemoore marks the approximate northern boundary of the lake.

The outline of the lake as shown in Figure 4 is only a general approximation at best. Frank Latta researched and mapped Tulare Lake in much more precise detail (see Figure 13).
Figure 13. Historic map of Tulare Lake.
Source: Frank Latta
Legend for Figure 13

The following explanatory notes were provided largely by Frank Latta. Elevations have been adjusted for consistency with the current sea level reference datum (see the section of this document that describes Elevations). Note that the stated elevation for a community often represent a relatively high point in that community; much of the community may be located at a lower elevation.

1. Lanare, elevation 207 feet
2. Riverdale, elevation 223 feet
3. Lemoore, elevation 230 feet
4. Stratford, elevation 203 feet
5. Santa Rosa Rancheria, elevation 218 feet
6. Guernsey, elevation 223 feet
7. Waukena, elevation 226 feet
8. Corcoran, elevation 206 feet (as measured at the train depot)
9. Alpaugh, elevation 213 feet
10. Atwell’s Island, elevation 212 feet
11. Skull Island, elevation 212 feet
12. Gordon’s Point at 190 foot contour
13. Orton’s Point at 196 foot contour
14. Kettleman City, elevation 253 feet
15. White River
16. Deer Creek
17. Tule River
18. Cameron Creek
19. Cross Creek
20. Kings River
21. Summit Lake, elevation 207 feet
22. Natural channel through Sand Ridge connecting the north and south Tulare Lakes (aka Táché and Ton Táché lakes)
23. Entrance of Kern River floodwaters from Kern, Buena Vista, and Goose Lakes
24. Lost Hills
25. Cox and Clark adobe
26. (Dotted line) Approximate 221-foot elevation contour. According to Latta, some pioneers said that Tulare Lake reached this elevation in the 1852–53 and 1862–63 floods. This is greater than the generally recognized elevation for the lake’s high stand.
27. Approximate 216-foot elevation contour. This is the generally agreed upon maximum elevation of the lake during the floods of 1852–53, 1862–63, and 1867–68. S.T. Harding calculated the lake’s elevation after these three floods to be 215.5, 216.0, and 215.4 feet respectively.
28. Lowest point in the bottom of the Tulare Lakebed, elevation 179.1 feet.
29. Terrapin Bay
30. Township lines, 6 miles apart
The annual elevation of the water in Tulare Lake is available for each year since 1850. It is as if a permanent
gaging station had been established in the deepest part of the lake, two years before Nathaniel Vise and the
first settlers came to the Four Creeks Country that would later be known as the Kaweah Delta. There clearly was
no such gaging station back then. And yet we are blessed with this wonderful dataset. The two people most
instrumental in its creation were:

- C.E. Grunsky, a civil engineer who first examined the Tulare Lake area in the 1870s
- S.T. Harding, a professor of irrigation at UC Berkeley for 35 years and a long-time consultant to the Tulare
Lake Basin Water Storage District

The very involved story of how Grunsky and Harding pieced together the Tulare Lake elevation dataset is told
elsewhere.113 But to greatly oversimplify their work, they talked to all the watermasters and collected all the
gaging data. They dedicated many years to studying how the lake worked and built an elaborate model of the
lake. From this they were able to fill in the gaps and calculate the lake elevations for those years when the
gaging data weren’t available.

The highest lake stage was attained three times during the historic period:

- After the 1852–53 flood raised the lake to elevation 215.5 feet
- After the 1861–62 flood raised the lake to 216 feet
- After the 1867–68 flood raised the lake to 215.4 feet

The above elevation figures are based on research done by S.T. Harding.114

The lowest elevation of the lakebed is 179.1 feet. C.E. Grunsky estimated that the bottom of the lakebed — the
area that had roughly this elevation — covered about 100 square miles.

The Kings, Kaweah, and Tule Rivers historically flowed into Tulare Lake, while the Kern River flowed into Kern
and Buena Vista Lakes (which occasionally overflowed to Tulare Lake). These four rivers formed expansive, low-
gradient, fan-shaped deltas near the lakes that were covered by vegetation. The fans formed by the Kings and
Kern Rivers extended far out into the valley. On the western side of the Tulare Lake Basin, coalesced fans
originating from the Coast Ranges are comparatively short and steep.

Surface waters were periodically exchanged between the San Joaquin River Basin and Tulare Lake Basin through
a complex of slough channels. Some of the channels branching off the mainstem of the San Joaquin River near
Firebaugh extended southward, and eventually formed a deep slough channel about 40 miles long and 250 feet
wide. This feature (Fresno Slough, aka Fresno Swamp) eventually branched into smaller channels 8 to 10 miles
from the San Joaquin River, which became intricate and ramified as they entered Tulare Lake, completing the
surface connection. A large bar at the mouth of the slough (on the Tulare Lake side) prevented water exchange
between Tulare Lake and the San Joaquin River except during periods of high flows.115 Fresno Slough merges
with the San Joaquin River at Mendota Pool near the present-day city of Mendota.

Flow in the Fresno Slough system was generally from south to north, bringing in seasonally high water from a
Kings River distributary, groundwater, and the occasional overflow from Tulare Lake. (A distributary is a branch
of a river that flows away from the mainstem. Distributaries are common on deltas. The Kings, Kaweah, and
Kern Rivers all have distributaries.)

Eyewitness reports variously describe flows in the Fresno Slough system at different times as both south from
the San Joaquin toward Tulare Lake, as well as north from Tulare Lake into the San Joaquin. C.E. Grunsky
believed that Lieutenant George Derby had crossed the delta of the Kings River in May 1850. Based on Derby’s
account, Grunsky concluded that the water in the Fresno Slough was flowing from the Kings River Delta north
toward the San Joaquin River and that part of the Kings River was flowing south to Tulare Lake.116

The Kings River is the largest of the four rivers in the Tulare Lake Basin; it was historically the largest source of
supply for Tulare Lake. At the time of EuroAmerican settlement, most of the water of the Kings River flowed into
Tulare Lake via what is now called the South Fork Kings, along the south side of the delta. Since then, a system
has developed that routes a portion of the Kings River floodwaters along the north side of the delta. That is a
heavily engineered system, but it was made possible by huge new sloughs eroded during the 1861–62 and
1867–68 floods. See the section of this document on Pine Flat Dam for a discussion of the system that has been
developed to manage Kings River waters on its delta.
Tulare Lake was divided into two parts by a sand ridge about 12 feet high which extended across the lake. The ridge extends from present-day Alpaugh to Kettleman City. It consists of Atwell's Island, Skull Island and Dudley Ridge, all more or less connected.\textsuperscript{117} The ridge varies in width, but tends to be about 100 yards wide.\textsuperscript{118}

During much of the year, the prevailing winds across Tulare Lake are from the north and northwest. In the distant geologic past, that sand ridge was formed when those winds piled up a ridge of sandy material along the southeast side of the lakebed during a period when the lake was not present.\textsuperscript{119}

A slough connected the two parts of Tulare Lake, passing through Sand Ridge. During very high-flow years, this slough served (and still serves on occasion) as part of the extension of Kern River. Kern River floodwaters last reached Sand Ridge in 1983.

At moderately high-water levels, Sand Ridge was mostly submerged with parts of it forming islands. The biggest island was about two miles wide by nine miles long and was known variously to early settlers as Root Island, Hog Island, or Atwell's Island. Today that island is the site of Alpaugh.\textsuperscript{120}

At relatively lower water levels, Sand Ridge divided Tulare Lake into two separate lakes. To the Yokuts Indians who lived there, the southern lake was known as Ton Taché, in contrast to the northern lake which was known as Taché. As the water level lowered farther, the southern lake would become little more than a marsh, having drained completely into the northern lake. The elevation of the lowest point in the Ton Taché lakebed is 204 feet.\textsuperscript{121}

Today, the southern lake, Ton Taché, might still be considered to exist but in a highly modified form. Tulare Lake water storage districts and irrigators know it as the South Flood Area and use it to store floodwaters. It includes all three of the Hacienda Reservoirs that are south of Sand Ridge as well as the South Wilbur Flood Area which is north of Sand Ridge.

Ton Taché was more than just the overflow for the northern part of Tulare Lake. In addition to being fed by floodwaters from the Kern River, Ton Taché also received the runoff from Deer Creek, White River, and Poso Creek.\textsuperscript{122, 123}

Tulare Lake sits in a natural depression. But it is also dammed on the north by the meeting of the Kings River Delta and the much smaller Arroyo Pasajero Delta. In this document, the broad feature formed by the meeting of these two deltas is referred to as the Tulare Lake sill or the delta sill. Other documents have described this feature by a variety of terms, including a ridge.\textsuperscript{124}

The Kings River Delta was formed by the outflow of glacial meltwater along the ancient Kings River during the period leading up to the end of the last glaciation (the Wisconsin) about 10,000 years ago. It has been described by one author as an alluvial fan dam.\textsuperscript{125}

The Arroyo Pasajero Delta is an immense and very broad landscape feature, some 450 square miles in size.\textsuperscript{126} Like other deltas along the Coast Ranges, it is relatively young. The bulk of the Arroyo Pasajero Delta has formed in the last 10,000 years, and deposition is still active. To begin to grasp its immensity, drive Highway 198 west from Lemoore. Virtually everything you see between you and I-5 is the Arroyo Pasajero Delta.

As explained earlier, all the water in the Kings River used to flow into Tulare Lake via what is now called the South Fork Kings, along the south side of the delta. Some Kings River water still flows into the lakebed during larger flood events.

On those occasions when the lake used to fill, the lake would overflow the delta sill and flow north toward the Fresno Slough and thence into the San Joaquin River.

The two-way channel where Tulare Lake spills north over the delta sill (and where the Kings River flows south into Tulare Lake) is about 15 feet across at the bottom and 60 feet across at the top.\textsuperscript{127} The elevation of this point on the delta sill is 207 feet. When Tulare Lake was full to that elevation, it was about 28 feet deep at its deepest point (elevation 207 - 179 feet).

If only it were so simple, but it’s not. The Tulare Lake sill was broad and densely vegetated with tules. This greatly reduced the rate at which water flowed out of the lake. C.H. Lee found that although some outflow started when the lake reached a depth of 28 feet (elevation 207 feet), significant outflow didn’t really start until the lake reached a depth of 31 feet (elevation 210 feet).
There is one further complication. Left to their own devices, rivers like the Kings, Kaweah, and Kern have moved back and forth across their alluvial fans. At the time of Euro-American arrival in the region (1840s), the Kings was flowing down the south side of its delta and into Tulare Lake.

In earlier geologic time, that river would have been flowing across other parts of its fan and contributing much less water to the lake. For example, the Kings River appears to have been flowing northward from about 21,000–15,000 B.P. (Before Present). This was during the recessional phase of the Tioga glaciation in the Late Pleistocene, when lakes east of the Sierra were reaching their post-glacial high stands. The Buena Vista Basin has deposits that record the runoff from that time period, but the Tulare Lakebed has no comparable record. So apparently the Kings River was flowing north across its fan during that period, contributing no water to Tulare Lake.128

Peter Vorster suspects that some Kings River water could have periodically flowed north during other high-water periods in the recent geologic past prior to settlement. This is suggested by the geomorphology of the alluvial fans on the Sierra rivers in the late Pleistocene and Holocene when large amounts of water and sediment flowed out of the Sierra.

As Euro-Americans have settled and developed these alluvial fans, we have fought to constrain the rivers and prevent their continued migration.

At the time of settlement, most of the water in Tulare Lake came from the Kings and Kaweah rivers. The Tule River flowed into the lake, but provided less volume than the Kaweah. In wet years, Kern River water entered the lake from the south. In high runoff years (perhaps every 5–10 years on average), Tulare Lake overflowed the delta sill and connected through the Fresno Slough to the San Joaquin River. From there, the water flowed on to San Francisco Bay.

Summit Lake is a tiny lake set in the fan-shaped delta of the Kings River. It sits in the throat of the channel at the north end of Tulare Lake. Whenever there was outflow from Tulare Lake, those waters flowed north through Summit Lake. Likewise, whenever Kings River waters flow into Tulare Lake, that water flows south through Summit Lake.

Summit Lake is located west of present-day Lemoore. This overflow point in Tulare Lake is immediately west of the intersection of Eight Ave and the south end of 26¼ Ave. In recent years, Summit Lake has been reduced to a circular alkaline flat. Imagine a highly reflective white disc. Apparently it is used as a landmark by pilots landing at the nearby Lemoore Naval Air Station (NAS).

When Tulare Lake overtopped the delta sill, water would flow northerly in a well-defined channel toward the Fresno Slough and thence into the San Joaquin River. At the lake's highest stage, about six feet of water flowed in a broad expanse northerly over the delta sill.

The sill has an elevation of 207 feet. At its highest stage, Tulare Lake had an elevation of 216 feet. For comparison, the highest point in the city of Corcoran, which was built within the lakebed, is the train depot: elevation 206 feet. Emergency levees have to be constructed to protect Corcoran when lake levels approach about 190 feet. Stratford, with a nominal elevation of 203 feet, has a similar problem with lake flooding.

In 1857, the California State Legislature granted a private company the right to construct a canal from the San Joaquin River to Tulare Lake, and then on to Buena Vista and Kern Lakes. That canal would have drained those three lakes and would have carried boats of up to 80 tons. The company was also given the exclusive right to reclaim all the swamp and overflow lands in the huge area represented by the lakebed of those three lakes and all the surrounding wetlands. That was one of the grandest opportunities ever given to enterprise in California, second only to the grants for the construction of the Pacific railroads. Tulare County bitterly fought back as soon as it learned of the act, but the State Supreme Court ruled that the Legislature could not revoke the franchise that it had granted. The company struggled to build the canal but never succeeded in that grand enterprise.129

Tulare Lake lies in the rain shadow of the Coast Ranges and is normally protected from large Pacific storms. The average annual rainfall in Tulare Lakebed is six inches.130 Water comes into the lakebed from the various tributary rivers and creeks. It leaves via a variety of ways:

- Natural outflow. When the elevation of the lake is higher than the elevation of the lowest point on the Tulare Lake sill (elevation 207 feet), the water begins to flow out of the lake. That last occurred in 1878. Since
1878, the Tulare Lake Basin has functioned largely as a closed basin, an inland sink without a regular outlet to the ocean.

- **Evaporation.** S.T. Harding calculated gross evaporation from Tulare Lake at 4.6 feet per year. Subsequent measurements showed that actual evaporation is somewhat higher: 5.2 feet per year. That is the primary way that water leaves the lake.\(^{131}\)
- **Absorption into the ground.** That is pretty minor. For example, during the 11 years between 1906–16, Harding calculated that only 4% of the water flowing into the lake was absorbed into the ground.
- **Used for irrigation within the lakebed.** This varies but is relatively minor in a big runoff year.
- **Pumped out of the lakebed.** This is generally minor.

There is much misinformation and outdated information out there about Tulare Lake. However, there are at least three good publications that are based on exhaustive searches of the literature, both published and unpublished, including gray literature:

- The Bay Institute. 1998. *From the Sierra to the Sea: The Ecological History of the San Francisco Bay-Delta Watershed*.\(^{132}\)
- United States Bureau of Reclamation. 1970. *A Summary of Hydrologic Data for the Test Case on Acreage Limitation in Tulare Lake*.\(^{133}\)
Floods and Droughts in the Tulare Lake Basin
General Flood and Drought Notes

General Notes on Kern, Buena Vista, and Goose Lakes

The three historic lakes in the south end of the Tulare Lake Basin are:

- Kern Lake (south of Bakersfield)
- Buena Vista Lake (just east of Taft)
- Goose Lake (southeast of the junction of Highway 46 and Interstate 5)

One source said that the five lakes (Summit, Tulare, Buena Vista, Kern, and Goose) were commonly joined in very wet years, but that is misleading. It would be more accurate to say that those lakes were set within a huge wetland complex. In very wet years, it would have been a challenge to go cross-country by foot between any two of those lakes; dry ground was probably quite scarce. However, it would have been easy to go among the lakes by boat if you had known how to navigate through the maze of interconnecting and frequently changing sloughs and waterways. One source estimated that during high water, the five lakes had a combined shoreline totaling 2,100 miles.\textsuperscript{135}

Before the 1861–62 flood, the Kern River channel ran where the Kern Island Canal now runs in Bakersfield: by the Beale Library (between Chester and Union Ave) on its way to Kern Lake. That flood shifted the river to the west. The new channel began at Gordon’s Ferry (just north of present-day Bakersfield College) and passed through what is now Old River and into the Las Palomas slough system between Kern Lake and Buena Vista Lake on its way to Tulare Lake (see Figure 14).\textsuperscript{136, 137} Not only did that new channel bypass Kern Lake, but one source said that it also bypassed Buena Vista Lake, meaning that those lakes would only get water during years with very high runoff. In any case, the river would shift even farther northwest in the 1867–68 flood.

The Sinks of the Tejón was the first Butterfield Overland Mail stage stop north of Fort Tejon. It was located at the intersection of present-day David and Wheeler Ridge Roads, roughly 10 miles northeast of where Interstate 5 and Highway 99 diverge.

The winter of 1861–62 was a high water year on all the rivers in the Tulare Lake Basin. When the Kern River came out of the canyon that winter, it created one vast sea of water from Gordon’s Ferry to the Sinks of the Tejón. Kern Lake was located in the southeast corner of that huge sheet of water.\textsuperscript{138}

Estimates of the size of Kern Lake at full pool vary a good bit, but it was somewhere between 9,000–13,000 acres. It was roughly triangular in shape, some 9 miles long (measured along its northern edge) and 4 miles wide in the middle. It would be possible to determine its exact size rather precisely using the Natural Resources Conservation Services (NRCS) soils map for southwest Kern County. Tejón Creek flowed into the southeast end of Kern Lake in a channel two feet deep and ten feet wide.\textsuperscript{139} Tejon Ranch Headquarters were located some miles south of that point.

Although it is convenient to talk about the Kern River as if it were a regular river channel flowing into Kern, Buena Vista, and Tulare Lakes, that is not exactly how it worked. The mainstem channel of the Kern River effectively terminated not too far after emerging from its canyon near Bakersfield. After that, the waters of the Kern passed through an ever-changing network of sloughs, some larger than others, across its very sandy alluvial fan.

For example, the soils map shows that the Kern flowed south to Kern Lake via a number of finger sloughs. The connection from Kern Lake west to Buena Vista Lake was via the 7-mile-long Connecting Slough. It was much the same throughout the area north to Tulare Lake; this was one vast wetland. The oft-cited comparison to the Everglades isn’t all that farfetched.

Prior to 1861, the Kern River flowed directly through Kern Lake in high-water years. The 1861–62 flood rerouted the Kern River to the west, through the Las Palomas slough system, bypassing Kern Lake. In subsequent years, Kern Lake received waters from the Las Palomas slough system.

The 1867–68 flood moved the river channel even farther northwest to its present location, ending in the Buena Vista Slough (aka Kern River Flood Channel), a few miles north of Buena Vista Lake and entered that lake from the northwest as shown in Figure 14. This reduced the floodflows into Kern Lake still farther.
Figure 14. 1880 map of Goose, Buena Vista, and Kern Lakes.  
Source: Charles Lux with modifications by Tony Caprio
The Kern Lakebed is located about 10 miles northwest of where Interstate 5 and Highway 99 diverge. The interstate passes through the western portion of the lakebed. The lakebed is immediately northeast of the small community of Lakeview, but that community no longer has a lake view.

The Kern Lakebed is largely or entirely owned by the J.G. Boswell Co. Virtually all of the acreage has been converted to cropland except for an 83-acre tract on the south edge of the lakebed. For many years, it contained the last remnant of Kern Lake. (That pond was known locally as Gator Pond, named for a small alligator that supposedly lived there in the 1930s.) From 1984–1995, the J.G. Boswell Co. partnered with The Nature Conservancy, allowing them to manage this valuable tract as the Kern Lake Preserve. Recent accounts are sketchy, but the tract has apparently since been dewatered and it no longer appears to be a functioning wetland.

Buena Vista Lake is immediately west of Kern Lake. Estimates of the size of Buena Vista Lake at full pool vary a good bit, but it was somewhere between 25,000–50,000 acres, some 8 miles wide by 12 miles long. Both Buena Vista and Kern Lakes were 290 feet in elevation at full pool.

In 1973, two manmade lakes (998-acre Lake Webb, and 86-acre Lake Evans) were created within the Buena Vista Lakebed for irrigation and recreation purposes. Kern County Parks and Recreation manages those two lakes as the Buena Vista Aquatic Recreation Area. The rest of the lake bottom is farmed, part of the J.G. Boswell Co. cropland.

Waterworks have been constructed to allow the choice of whether to impound Kern River floodflows in Buena Vista Lakebed or pass them through to the Tulare Lakebed. In the 1952 flood, 232,000 acre-feet of Kern River floodwaters were stored in the Buena Vista Lakebed. The water storage district has apparently never chosen to use that lakebed for such purposes in subsequent floods, particularly in the big 1969 flood. That decision would eventually be ruled on in a landmark decision by the U.S. Supreme Court. For a more complete description of this event, see the section of this document that describes the 1969 flood.

The Kern River is traditionally considered to terminate in Buena Vista Lake, but that is somewhat misleading. In low-water years, the waters of the Kern never reached the lake. However, in high-water years, the lake would overflow or spill.

When Buena Vista Lake overflowed, water then flowed northwest through Buena Vista Slough to Tulare Lake (see Figure 14). About midway to Tulare Lake, Buena Vista Slough passed through what Miller and Lux called the Buttonwillow Swamp along the west side of Buttonwillow Ridge. At the northern toe of Buttonwillow Ridge, Jerry Slough (aka Bull Slough or the northern extension of Goose Lake Slough) joined with Buena Vista Slough and these waters then passed through Sand Ridge to enter the main body of Tulare Lake. During high-water periods, there was a smaller portion of Tulare Lake on the south side of Sand Ridge, known to the American Indians as Ton Taché.

But that was only one of two ways that Kern floodwaters found their way into Tulare Lake. The other way was via a natural overflow flood channel, a distributary that began much farther upstream (see Figure 14). In high-water years, the Kern would overflow its banks just west of Bakersfield, a few hundred yards east of the present-day Stockdale Bridge, at a place where the historic wooden Bellevue Weir was built across the river. That weir is now gone, but there is a rock spillway across the river at the same location, which slightly raises the level of the river upstream. See the 1950 flood for an account of private interests working to contain flooding in this flood channel.

This flood channel (Goose Lake Slough) flowed 14 miles west through what is today the Rosedale neighborhood until, just before it joined the Buttonwillow Swamp, it was deflected by the east end of a gentle rise known as Buttonwillow Ridge. Here the slough channel (now called Jerry Slough) veered 14 miles northwest to Goose Lake. Jerry Slough formed a network offorking and rejoining channels; it was a challenge to navigate.

Goose Lake is the widest point on what has been termed the Jerry Slough Delta. That landform is located in the northwest to southeast-trending trough of low-lying land between Buttonwillow Ridge (southwest of Goose Lake) and Semitropic Ridge (northeast of Goose Lake).

Goose Lake is a slight depression, averaging a little more than a mile in diameter. At full pool, Goose Lake and Jerry Slough formed a single body of water, elevation 250 feet, up to 20 miles long, and 1–4 miles wide.
Today, the Goose Lake Bottom is divided into quadrants by various canals and channels. The southwest quadrant of the lake bottom is a nearly permanent body of shallow water surrounded by emergent marsh vegetation. This “surge pond” is the one quadrant of the lake bottom that most often contains water. Were Goose Lake to fill to a surface elevation of 235 feet, Goose Lake would measure approximately 3 miles long by 1–1½ miles wide and would cover approximately 1,600 acres. When duck club lands in the Goose Lake Bottom are flooded to an elevation of 237 feet, water covers approximately 922 acres of wetlands inundated to a depth of 6–8 inches. This is in addition to the 1,600 acres in the surge pond.¹⁴²

Today, Interstate 5 runs along the crest of Buttonwillow Ridge. The ridge begins near Stockdale Highway and peters out a few miles south of Highway 46, just a short distance downstream from Goose Lake. The ridge separates Buttonwillow Swamp from Jerry Slough.

Where Buttonwillow Ridge peters out, the waters of Buttonwillow Swamp and Jerry Slough come back together. The slough channel north of Goose Lake and Buttonwillow Ridge is known as Bull Slough.

The actual location of Bull Slough is somewhat uncertain; this is a nomenclature issue. In 1880, Buena Vista Slough was the main waterway and Bull Slough was apparently only the northern portion of the short slough that connected Goose Lake and Buena Vista Slough.¹⁴³ Possibly the name Bull Slough later came to apply to a larger area extending farther to the north. In any case, Goose Lake Slough, Jerry Slough, and Bull Slough have not carried Kern River floodwaters since 1983.

In the 1800s, Henry Miller (of Miller & Lux fame) set about draining Buttonwillow Swamp under the provisions of the Swamp and Overflow Act. He accomplished this by building a huge levee along the west side of the swamp, which created a manmade channel (known today as the Lokern Flood Channel) running along the toe of the Elk Hills, McKittrick and Belridge alluviums between Buena Vista Lake and Highway 46. That confined the Kern River to a narrow channel, thus drying up a lot of the wetland. However, in high-water years, the Kern would still overflow into the Goose Lake system and enter Miller’s empire through the back door.

To solve that problem, Miller built another huge levee across the narrowest spot in the Goose Lake system, which happened to be at the west edge (downstream side) of the Goose Lake depression. That levee connected Buttonwillow Ridge to the Semitropic Ridge, right at the toe of Semitropic Ridge, running northwest along the toe, all the way to the present-day Kern National Wildlife Refuge, thereby holding this “back door” water against the Semitropic Ridge. These massive earthmoving projects effectively reclaimed Buttonwillow Swamp so that it was no longer a wetland.

The Goose Lake system was left unaltered, except that when water does come down the Kern, it is held slightly deeper than it would otherwise be in Goose Lake. The last significant floodwaters to come down Jerry Slough were in the fall of 1951 and the spring of 1952. Goose Lake has been filled three or four times since 1952, when there was a need in some high-water years to divert water to anyplace they could find a spot that would take it.

During high-flow events, Kern River water continues flowing to the north, passing near the west side of present-day Kern National Wildlife Refuge, crossing through Sand Ridge between Alpaugh and Dudley Ridge, and enters the south end of Tulare Lake.
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General Notes on Bravo Lake
In addition to the five generally recognized valley lakes, there was reported to have been a sixth natural lake in the valley: Bravo Lake (aka Wood Lake). This lake was located near the upper end of the Kaweah Delta in the Kaweah River Swamp. It was immediately north of the St. Johns River, about two miles west of McKay’s Point (Township 17 S, Range 27 E, Section 31). Today Bravo Lake is in the southeast edge of the city of Woodlake.

Relatively little is known about Bravo Lake compared to the five better-known valley lakes. Assuming that reports are correct, the lake was originally a natural feature, but was converted into a reservoir in the 1870s. We have found no written description of what the natural lake looked like or how it was formed. However, we can make reasonable inferences based on its location on the Kaweah Delta and old maps. There really is very little hard documentation to go on, so this should all be read cautiously.

The earliest reference to the lake that we have found was as a landmark for the small community of Stringtown that was caught up in the 1867–68 flood. Stringtown was a settlement of five families described as living in a line south of present-day Woodlake along the Kaweah River, east of Bravo Lake. (This reference to Stringtown’s location is somewhat unclear. Bravo Lake was located adjacent to the St. Johns River.) Although Stringtown may not have been much of a town, this suggests that Bravo Lake may have existed prior to the 1867–68 flood.

The lake does not appear to have occupied a natural depression or to have been on a stream course. Judging from available data, the contours under Bravo Lake appear to slope toward the southwest, generally toward the St. Johns River. The lake appears to have been a floodplain feature associated with the St. Johns River. The south edge of the lake generally paralleled the north bank of the St. Johns River. That suggests that the lake may have been formed by the natural levee created on the river’s north bank during floods.

Levees are commonly thought of as man-made, but they can also be natural. When a river floods over its banks, the water spreads out, slows down, and deposits its load of sediment. The coarsest sediment is dropped first as the river no longer has the energy to carry it. This coarse material forms a natural embankment (levee) along the edge of the river channel.

It is tempting to think that Bravo Lake was created primarily during the great floods of 1861–62 and/or 1867–68. Those were the two floods that created the St. Johns River. Both of those floods deposited large amounts of silt and debris on the Kaweah Delta.

Ed Reynolds, an early Tulare County pioneer, recalled how Bravo Lake got its name. According to Reynolds, a fight took place near the lake in 1870 between Tom Fowler and “Swamp John” Asbill. Each man had a rooting section. Many of Fowler’s supporters yelled “Bravo, Bravo”, and that is how the lake apparently got its name. This suggests that the lake may have had no generally accepted name prior to 1870.

In 1872 the Wutchumna Ditch Company organized and commenced the construction of an irrigation system which eventually consisted of about forty miles of main and branch ditches. The water was taken from the Kaweah just above McKay’s Point. Bravo Lake, situated near the intake of the canal, was used as a storage reservoir for floodwaters so that a supply was maintained throughout the year. Water from Wutchumna Ditch entered on the east side of Bravo Lake and exited on the west. The ditch system was largely completed by 1877.

Joe Childress was the manager of the Wutchumna Water Company for many years. While he has heard that a natural lake once existed where the reservoir is today, he has never seen any records to substantiate that. The company has no surviving engineering records to document how the reservoir was constructed in the 1870s. Joe speculated that the original (natural) Bravo Lake may have been a shallow ponding basin on the edge of the St. Johns. The ditch company may then have built a low levee around that to raise the water level and convert it into a reservoir.

The Tulare County Times reported that Bravo Lake was full of water in July of 1886. Presumably that was a reference to the use of the lake as a reservoir for Kaweah River water delivered via the Wutchumna Ditch. The 1892 Thompson Historic Atlas Map labeled the lake as “Bravo Lake (Reservoir)” and showed it as connected to Wutchumna Ditch on both the east and west sides (map on file in the national parks). A photograph of the lake taken in about 1900 shows it to have a relatively natural shoreline with a scattering of small islands (photograph on file in the national parks).
The original (natural) Bravo Lake had no creek or river flowing into it. The nearest cross drainage of any size is Antelope Creek which is about half a mile west of the lake. The lake would have received water only when the St. Johns overflowed its banks. Perhaps the Kaweah Delta contained other lakes of this type, but Bravo Lake is the only one of this type that we know of.

Our understanding of the hydrologic operation of the natural Bravo Lake is based largely on supposition. We have found no records about when the St. Johns overflowed in the vicinity of Bravo Lake in the 19th century. The floods of 1861–62 and 1867–68 almost surely inundated this area, quite possibly creating the lake. It is also possible that the St. Johns River overflowed its banks in the floods of 1872, 1874, 1875, 1876, and 1877, putting water into Bravo Lake. In the 1877 flood, the levee on the south bank the St. Johns failed farther downstream, causing flooding in Visalia.

Once Bravo Lake began functioning as a reservoir, it presumably got most of its water from the Wutchumna Ditch. As man-made levees were raised around the reservoir, it would have been completely cut off from the St. Johns' floodwaters.

The community of Woodlake was founded in 1912 by Gilbert F. Stevenson, a wealthy land developer from Southern California. He planned for the town to be a resort community with the lake the center of a "Mecca of prosperity." Stevenson constructed a large levee around the lake. That is the levee that we see today.

Stevenson had big plans for the lake. The lake level was to be raised substantially. Several islands were to be constructed within the lake which would be serviced by pleasure craft. On these islands would be elaborate dance pavilions, bathhouses, and restaurants. A narrow gauge railroad would run around the levee for scenic excursions (one source said that the train was intended to connect the islands).

Stevenson also planned to develop several parks outside the levee, featuring shade trees, fountains, walkways, baseball diamonds, etc. The Tulare County Times reported in 1913 that "Steps are now being taken to make the lake, formerly Bravo and now Woodlake, the most popular picnic ground in the county." Stevenson's numerous financial commitments resulted in his downfall during the Great Depression; his plans for the resort and the lake were never realized. The Wutchumna Water Company took over control and operation of Bravo Lake and its irrigation system.

The USGS 1928 Lemon Cove quad sheet labeled the lake as Wood Lake, maintaining the name that Stevenson had apparently applied to it. However, the 1952 Woodlake quad had switched to using Bravo Lake for the lake's name.

American Indians and Early Exploration of the Tulare Lake Basin

The Yokuts Indians were the dominant American Indian group on the San Joaquin Valley floor and in the adjacent foothills; their population throughout the region at the time of European contact was at least as high as 40,000 and probably much higher. At least 19,000 Yokuts lived in the Tulare Lake Basin or visited it seasonally.

In the Kaweah River Basin, the present-day town of Three Rivers (junction with South Fork Kaweah) marked a distinct linguistic and cultural boundary between the Penutian-speaking Yokuts and the Shoshonean-speaking Monachi. The Monachi group (the Balwisha or Patwisha) lived to the east of there.

The Wukchumni, a sub-group of the Yokuts Indians, lived to the west. Their outpost village of hotnú nyu was apparently located at Slick Rock, about one mile west of Three Rivers.

The American Indians who lived in the lower elevations of the Tulare Lake Basin are believed by some to have had a population density exceeding that of any group in North America exclusive of Mexico. That is remarkable since these were non-agricultural people. However, these population density estimates may be a bit of an exaggeration.

What can be said is that population densities in the Tulare Lake Basin were high compared to those of non-farming aborigines in other parts of North America. The highest population densities by far — six to seven people per square mile — occurred in the lush stream delta and delta foothill areas along the Kings and Kaweah Rivers. The total population along the lower Kaweah was approximately 3,800.

Deserting Spanish soldiers from the mission at San Diego are believed to be the first Europeans to enter the San Joaquin Valley. Don Pedro Fages, Lieutenant of Catalanian Volunteers, came in search of them in the fall of
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1772, entering the valley by way of Tejón Pass. Fages was the first person to make a written report of the San Joaquin Valley. In his report, Fages referred to an American Indian village located on the valley’s southernmost lake, calling that village “Buena Vista,” hence the name that we now use for that lake.

In 1773, Commander Tagus came into the Tulare Lake Basin looking for army deserters. He traveled far and wide in the region and found a large lake that he named Laguna de las Tules, but did not find any deserters.

Missionary explorer Father Garcés was the first to explore into what is now Tulare County. His purpose in exploring this area was to find native converts to Christianity. He entered the valley by way of Tejón Pass in 1776 and traveled northerly, possibly as far as the Kaweah Delta.

In November 1805, Father Juan Martín journeyed into the Tules because the natives of the region, through neophytes at Mission San Miguel (near Paso Robles), had expressed a desire that he visit them. (Some accounts say that this trip occurred in 1804.) Because of a disagreement with the Spanish Governor on establishment of missions in the Tules, his expedition was without official sanction. Father Martín visited the rancherías of Bubal and Sumtache on Tulare Lake.

At the time of Father Martin’s visit, the rancheria of Bubal was located on Atwell’s Island. Bubal was one of the largest Yokuts villages, with a population of about 1,300. Martin estimated that there were at least 4,000 American Indians living around the lake.

The Spanish began a period of active exploration into the interior valleys in 1806. These expeditions were primarily for discovering potential mission sites, capturing or punishing runaway neophytes, or bringing converts to the padres of the coastal missions.

The first of these expeditions, which the governor called “civilizing missions,” left Mission Santa Barbara on July 19, 1806 under the leadership of Lieutenant Francisco Ruiz and Father é María Zalvidea. This party visited Buena Vista Lake, Tulare Lake, and the Kings River country in the vicinity of present-day Kingsburg. On August 4th they were in an oak forest on the Kaweah Delta. The party probably camped in or near what is now Mooney Grove. Father Zalvidea was impressed with the oak grove and believed it would be a suitable mission site.

The Moraga-Muñoz expedition left Mission San Juan Bautista (near present-day Hollister) on September 21, 1806. It traveled inland to enter the San Joaquin Valley in the vicinity of San Luis Creek, thence generally north as far as the Calaveras River. From the Calaveras River, the party proceeded southeasterly and on October 14-16 was exploring upstream and downstream on the Kings River from a camp in the vicinity of present-day Sanger or Centerville. On October 18, scouts reached the great oak forest in the Kaweah Delta.

On October 20 the party explored east until they reached what we now call the Kaweah River, but the Spanish knew as Rio San Gabriel. Father Muñoz recorded that:

The river is known as the San Gabriel. It divides into two branches, one of which we call the San Miguel, and the latter sends its water into several other branches. [A mission in this place], in case the King, our Lord, whom God protect, grants its establishment, could have available pine and redwood timber and fine lands for crops.

On October 21, they explored to the San Pedro [Tule River], but found that it was dry.

The Moraga party remained in the vicinity of the oak forest on the Kaweah Delta for a week. On October 26 they traveled southeasterly to the Tule River and out of the valley by way of Tejón Pass.

There is a legend among descendants of Wukchumni and Padwisha Indians that a great battle took place between a Spanish mining expedition and American Indians on the North Fork of the Kaweah River in an area inside present-day Sequoia National Park. The legend says that this battle took place about 1811 and resulted in a total victory for the American Indians.

In 1814, Master Sgt. Don Juan de Ortega and Father Juan Cabot went to the Tulare Lake area. They left Mission San Miguel (near Paso Robles) on October 2. They first visited the rancherías of Bubal and Sumtache, each with a population of about 700. The primary reason for the expedition was to pacify Sumtache and establish peaceful relations between that village and Bubal.
After achieving that, they went to the Rio San Gabriel (Kaweah River) which they described as being very full and of good water even though it was early October. From the ford of the San Gabriel, they marched three leagues (nine miles) to the rancheria of the oaks, Telame. They identified this as the only place in the area suitable for founding a mission or a presidio. From there, they visited other rancherias along the river Reyes (Kings River).\textsuperscript{168, 169}

Ortega and Cabot visited the region again in 1815. Their party left Mission San Miguel on October 5. Ortega reported great famine when they visited Telame on October 11. They explored around present-day Visalia and up the Kaweah River to about Lemon Cove.\textsuperscript{170, 171}

The Telamni were a sub-group of the Yokuts; they lived in the vicinity of present-day Visalia and Goshen. Their principal village was Telame, the largest rancheria in the Tules. It was located immediately to the northeast of Visalia. It was in an immense oak forest on the Kaweah Delta, a league (three miles) from the Sierra. With reference to the lower Kaweah River region, the degradation of the American Indian population, which began essentially with the first Spanish contact, was characterized by Sherburne Cook:\textsuperscript{172}

\[\text{Telame} \text{ had originally been a very large village but the disturbances caused by the Spanish expeditions had substantially destroyed it. The heavy mortality and great famine mentioned by Ortega were undoubtedly due to the continuous state of fugitivism [fleeing the Spanish], severe exposure to the weather, and inability to gather and store the customary stocks of food such as acorns and fish. No specific epidemic was recorded [but] no fulminating epidemic was necessary to produce the mortality. Starvation, exposure, and respiratory diseases would be quite adequate.}\]

In 1816 Father Luis Antonio Martinez visited Buena Vista Lake.

In 1819 Lieutenant Jose Maria Estudillo led a military force into the oak grove on the Kaweah Delta.\textsuperscript{173} The Kaweah drainage above present-day Visalia was inhabited by the Gawia, Yokod, and Wukchamni. The Wukchamni were by far the most numerous, and Estudillo left an excellent account of them. In addition to being a competent field commander, Estudillo appears to have been a scholar and a gentleman. His report on the Wukchamni village of Chischa is the most complete and accurate left us by any of the Spanish explorers. Estudillo was the first white man to see that village. It was located at or just above Lemon Cove. The village was crescentic in shape and large enough to have covered eight city blocks. The population was at least 1,250.\textsuperscript{174}

The last important Spanish expedition of the pioneering period was that of Estudillo. Subsequent expeditions were purely punitive military raids and campaigns against American Indians, and it was generally concluded that presidios as well as missions were needed if the San Joaquin Valley were ever to be colonized.

The Spanish in their distant coastal missions had a difficult time converting the American Indians who lived in the Tulare Lake Basin. Many became accomplished horse thieves and occasionally raided for horses. In his report of 1818–1819, Father President Mariano Payeras said “the Tulare Indians are inconstant. Today they come, tomorrow they are gone, not on foot as they came, but on horseback [and] having crossed the Tulare Valley and the mountains that surround it, they kill the horses and eat them.” Father Payeras referred to Telame as “a republic of hell and a diabolical union of apostates.”\textsuperscript{175}

Acorns are one of the food items most associated with the American Indians who lived in the foothills and the Sierra. Archeological evidence indicates that a major switch to acorns pounded in rounded mortars with pestles happened about 3800 B.C. in the San Francisco Bay Area, by 2600 B.C. in the Central Valley, and at 1000 A.D. in the Sierra.\textsuperscript{176}

The American Indians who lived around the valley lakes had a very different lifestyle than those who lived in the foothills. Frank Latta gave a vivid description of what life was like for those who lived in the vicinity of Goose Lake:\textsuperscript{177}

\textit{The Tuhoumne Yokuts were on Kern River...and on Buena Vista, Jerry, Goose Lake, and Bull Sloughs, from the eastern portion of the Elk Hills past Goose Lake and Adobe Holes toward Tulare Lake. Except for an occasional antelope surround, or a ground squirrel smoke-out on the West Side, theirs was strictly a goose, duck, mudhen, swan, blue heron, egret, pelican, lake, slough, swamp-and-overflow culture; water and mosquitoes, willows and mosquitoes, tules and mosquitoes everywhere; tule boats, tule bags, tule skip-rings, and other tule equipment — and mosquitoes; tule houses, tule sunshades, tule windbreaks, piled up tules for sails on tule boats; tule clothing — caps, capes, hoods, parkas and skirts; tule mattresses, tule mats, tule blankets, pounded tule-fibre disposable diapers for babies, tule baby...}
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cradles, tule fuel, tule blinds for hunting, tule-seed mush, tule-root bread, tule baskets, tule shrouds, tule rope, tule string, tule elk, beaver, sea and freshwater otter, tules, tules, tules — and mosquitos; seal, raccoon; waterfowl and fish in myriads; more tules, tules, tules — and mosquitos.

Malaria was unintentionally introduced into the San Joaquin Valley from Oregon in early 1833 by a party of beaver trappers from the Hudson’s Bay Company. More than 20,000 American Indians died from the disease that spring. These included Yokuts, Chumash, Miwok, and others.178

By 1853, the population of Yokuts in the Tulare Lake Basin had been reduced to no more than 1,100. The survivors were located primarily in the foothills, on the eastern shore of Tulare Lake, and among the timber on the upper Kaweah Delta. In 1876, the Tule River Indian Reservation was created east of Porterville where 1,200 people were taken from various aboriginal groups. A few Yokuts managed to stay outside the reservation, but even that small aboriginal population proved incompatible with many of the American settlers, who outnumbered them by three to one as early as 1860, becoming the new dominant culture in the basin. Most of the remaining Yokuts were eventually rounded up and taken to the Santa Rosa Rancheria near Lemoore.179, 180

In 1806, Father Pedro Muñoz and Second Lieutenant Don Gabriel Moraga visited San Joaquin Valley as far south as the Kings River. They noted salmon and beaver. Moraga changed the name from Valle de los Tulares to the San Joaquin Valley.

Jedediah Strong Smith was probably the first American to visit the Kaweah Delta area. Searching for a mythical river that supposedly flowed from the Rocky Mountains to the Pacific Ocean, he went south from the Great Salt Lake, picked up the old Spanish Trail to the Colorado River, and after crossing the Mojave Desert, eventually reached Mission San Gabriel late in 1826.181

Smith and his party were regarded with suspicion by the Spanish and were ordered in January 1827 to leave California by the same route by which they arrived. However, the party turned north from San Bernardino, crossed Cajón Pass into the desert, then over the Tehachapis into the San Joaquin Valley, probably by the same route Father Garcés had taken over Tejón Pass in 1776. Smith’s group traveled up the east side of the valley, trapping beaver on the Kern, Tule, Kaweah, and Kings Rivers as they went. Smith later reported this to be the best beaver country he had ever seen.182

Thomas L. “Peg-Leg” Smith, called El Cojo (the crippled one) by the Spanish, was another early visitor to the valley. He came in about 1830 and gained a reputation as a horse thief. He is believed to have encouraged thievery among the Indians.

Hudson’s Bay Company trappers were active in the region in the early 1830s. Ewing Young, with Kit Carson as a member of his group, encountered them as competitors as his party trapped throughout the valley in 1832.

As described in the section on California Snow Conditions during the Little Ice Age, Zenas Leonard and the Walker Party crossed the Sierra in October 1833. They followed the San Joaquin River downstream, and arrived in the vicinity of Half Moon Bay on November 20, 1833. They traveled down the coast to Monterey, the capital of Mexican California. They spent the winter living with the secular Spanish.

On February 14, 1834, the Walker Party departed from their camp about 40 miles east of San Juan Bautista, heading for the San Joaquin Valley, working their way south along the foothills, and crossed the Tulare Lake Basin. They are believed to have then entered the Kern River Basin and left the valley by what is now known as Walker Pass on about March 15, 1834. At an elevation of 5,250 feet, this is much lower and more snow-free than any other pass across the Sierras. Walker led another party across this pass, heading west, in 1843. John C. Fremont named it Walker’s Pass in 1845.

John C. Fremont, with Kit Carson as guide, passed through the Kaweah Delta area on his second expedition. After an arduous winter crossing of the Sierra, he traveled south from Sutter’s Fort on March 24, 1844. On April 9, he camped in the Kaweah Delta area, probably on the banks of the Kaweah.183

California became a United States territory in 1848, and the Army sent Engineer Officer Lieutenant George H. Derby into the Tulare Lake Basin in 1850 to select a site for a military post to protect the San Luis Obispo area from Indian raids. With reference to the Kaweah Delta, Derby said:184
the only point in the whole valley... at all suitable for a military post was a small portion of the interval land contained by the [distributary] creeks of the River Frances [Kaweah River]. The land is excellent for cultivation, well timbered and an abundance of excellent building material may be found close at hand. The country is 8 miles in length by 6 miles in width between these branches; it is a beautiful, smooth, level plain covered with clover of many different kinds and high grass and shaded by one continuous growth of oaks of a larger and finer variety than I have ever seen in the country.

Lieutenant Derby pointed out that this location would be central to the passes to the west, to Tejón Pass, Walker Pass, and directly on the road from the mines to the south.

The first attempt at settlement in the Kaweah Delta area is generally credited to John Wood who came to California early in 1850 in the John Hudgins party. Wood left that party in Los Angeles and went on to the northern mines. In the fall of 1850 Wood led a party of 14 or 15 men from Mariposa into the Kaweah Delta area.

They settled east of present-day Visalia, built a cabin of oak logs, and began to prepare the land for cultivation. Wood chose a poor time to attempt colonization, for the Indians throughout the valley had become alarmed at the encroachment of white settlers on their land. The local tribe gave Wood’s group ten days to leave the area. When they were still there on the tenth day, December 13, the Indians attacked, killing Wood and most of the settlers.

Visalia was founded in 1852. The first citizens of that community built a stout stockade against the American Indians, but despite many alarms there was never any significant trouble with the American Indians.

The first settler in the Lake Kaweah area, and the first American to see what is now Sequoia National Park, was Hale D. Tharp, a Michigan cattleman who had successfully mined gold near Placerville. He came into the Kaweah Delta area in 1856 in search of a place to settle his family. On the way to California he married a widow who had several children. Tharp built a crude log cabin on the west side of Horse Creek a few hundred yards upstream from its confluence with the Kaweah River. That location is north of the present-day Horse Creek Campground.

Tharp befriended the Indians and got along admirably with them. In reciprocation, they shared their knowledge of the country. They told him of the high mountains where there were trees so large it took 25 men clasping hands to encircle them, and of lush mountain meadows that were green all summer. Because 1858 was a dry year and Tharp was a cattleman, he decided to investigate the stories the American Indians had told him about the perennial meadows and big trees in the high mountains (see the section of this document that describes the 1855–1861 drought).

Wildlife in and around Tulare Lake

Thanks to the writings of James Carson and others, we have a pretty good idea of the wildlife that lived in and around Tulare Lake in the middle of the 19th century. For example, Carson wrote that beavers, mink, and river otters were all present.

Rob Hansen cautions that these early accounts need to be viewed critically. For example, Carson wrote that muskrats were also present. However, muskrats were almost certainly not present in the Tulare Lake system until 1943. So it’s worth examining what evidence we have that beavers, mink, and river otters were actually present in the Tulare Lake system in the 19th century.

Some have questioned whether beavers were really present in the San Joaquin Valley, but Carson wasn’t the first or the only one to report them. Father Pedro Muñoz and Second Lieutenant Don Gabriel Moraga observed beavers when they explored the central portion of the San Joaquin Valley in 1806. Felipe Santiago García reported seeing many beavers in all the rivers of the Tulare Lake Basin.

The Spanish were aware of the beavers in the San Joaquin Valley, but they generally didn’t exploit the valley’s many resources. All of California, including the San Joaquin Valley, was Spanish territory. The Spanish settlements were largely on the coast, so that left the beavers in the San Joaquin Valley as fair game for the English and American trappers to exploit. The history of the early beaver trappers in California and the Southwest has been well researched by Dr. Robert Cleland and others.

The first and most famous of the trappers was Jedediah Strong Smith. When Smith and his band of trappers arrived in California, they came by a route far south of the Sierra. After crossing the Great Basin from the Great...
Salt Lake, they struggled up the Mojave River, finally reaching the haven of Mission San Gabriel (southeast of present-day Pasadena) at the end of November 1826.

They stayed there for the next two months, learning much about California. Harrison G. Rogers, (Smith’s clerk) wrote on December 26, 1826, that there were supposed to be plenty of beavers at both Buena Vista and Tulare Lakes. When Mexican Governor Echeandia learned of Smith’s presence, he ordered Smith to immediately leave by the way that he had come.

Smith had no intention of abandoning his search for new beaver country. So, although he departed San Gabriel as ordered on January 18; instead of going back by the Mojave, he headed northward over the Tehachapis and entered the San Joaquin Valley.

There has been speculation about which lakes and streams Smith would have trapped as his expedition wandered north through the Tulare Lake Basin. It does seem reasonable that he would have checked to see if Buena Vista and Tulare Lakes really did have the large number of beavers that he had been told.

On the other hand, the map of his expedition shows Smith’s route touching Tulare Lake at the south and gradually bearing away from it as he went northward. That suggests that he might not have hunted beavers on the Kern, Tule, or Kaweah Rivers or on the other small streams that flow out of the Sierra before he reached the Kings. Smith wrote a letter to General William Clark, Superintendent of Indian Affairs, saying that he began his spring hunt there. By the end of April, Smith had trapped his way north through the San Joaquin to the American River and his horses were packing great bundles of beavers.

On May 20, 1827, Smith and two of his group began an attempt to cross the Sierra, following generally the route of present-day Highway 4. They reached the crest near Ebbetts Pass eight days later, becoming the first white people to cross a Sierrapass.

Smith returned to the San Joaquin by the Mojave route in 1828, rejoining his party. After some difficulties with the Spanish authorities, they were allowed to travel to San Francisco. There they sold 1,568 pounds of beaver pelts and 10 otter skins for $3,940. Smith's party spent the next few months trapping beavers in the lower tributaries of the San Joaquin.

Peter Skene Ogden led a 60-man trapping expedition for the Hudson's Bay Company in 1829–30; he was one of the first English trappers to explore California. His expedition trapped down the Colorado River nearly to the Gulf of California, then worked their way trapping up the San Joaquin Valley.

After Ogden came into the San Joaquin Valley, his trapping route took him along what he called the South Branch of the Bonaventura. The Bonaventura is the stream now known as the San Joaquin River, so his route up the “South Branch” presumably took him along the Kern River through Kern, Buena Vista, and Tulare Lakes.

Ogden is the only English trapper to venture so far south. After Ogden’s expedition, Hudson’s Bay Company trappers, working from their post at Fort Vancouver, generally trapped only as far south as the San Joaquin River. Most or all the trappers in the Tulare Lake Basin after Ogden were Americans. According to Donald Tappe, trappers working for the Hudson's Bay Company took beaver furs as far south as Buena Vista Lake, but they usually considered it unprofitable to work farther south than the shores of Tulare Lake.

Ogden’s English trappers weren't the only trapping party to enter the San Joaquin Valley in 1830. Also in that year, an American trapping party led by Ewing Young which included Kit Carson traveled from Taos, New Mexico to Mission San Gabriel. After crossing into the San Joaquin Valley, Young’s party initially trapped beavers along the Kern River. Following Ogden's trail, they caught up with his group in a few days.

The two groups trapped the San Joaquin Valley together for 10 days until they came to the Sacramento–San Joaquin Delta. Although Ogden had collected a thousand skins by then, he generally found that the California waters were the poorest in furs of any area that he had yet explored.

In 1833, Zenas Leonard recorded beavers living in the San Joaquin River and being traded out of Monterey. Leonard did not say where those beavers were trapped or by whom.

Stephen Hall Meek recorded that his fur trapping party pitched their camp for the winter on the shore of Tulare Lake in December 1833.
Clearly beavers occupied the lower elevations of the San Joaquin River and Tulare Lake Basins in the early 19th century. If beavers behaved in those basins as they did in the Rockies and elsewhere, they would go upriver as far as they found suitable habitat. That habitat existed in the form of cottonwood and aspen; the only question is whether beavers made use of it.

The various subspecies of American beaver that live in the Rocky Mountains have adapted to occupy virtually all areas of suitable habitat, even in the alpine. For example, the subspecies that lives in Colorado (*Castor Canadensis concisor*) lives throughout that state in suitable habitat, although it is most abundant in the subalpine zone. Many of the high alpine ponds and meadows that exist in Colorado today are the work of generations upon generations of beavers.

The subspecies of beaver that lives in our state is the California Golden beaver (*Castor Canadensis subauratus*). Beaver have been considered native to Northern California’s Klamath and Pit River watersheds and much of the Central Valley, but not to the Sierra. Current wildlife management policies in California and Nevada continue to cite early 20th-century zoologists Joseph Grinnell and Donald Tappe, who concluded that beaver were not historically extant at elevations above 1,000 feet on the western slope, nor on the eastern slope, of the Sierra.

There is no obvious biogeographical barrier that would have prevented beaver from migrating up the tributaries of the Sacramento and San Joaquin Rivers into the many low gradient streams of the western slope of the Sierra. Beaver have occurred historically in every other North American mountain range from the Arctic Circle to northern Mexico. Preferred food sources for beaver (aspen, cottonwood, willow, and mountain alder) are widely available in the Sierra. These facts raise questions as to why beaver would not have been historically extant in the Sierra. In addition, the widespread presence of incised, rapidly eroding streams in Sierran montane meadows, which were historically low gradient and meandering, raises the question as to whether beaver might have stabilized these streams in the past.

In effect, Grinnell and Tappe were saying that they had no evidence of beavers living above the 1,000 foot elevation in the Sierra. They didn’t present any biological argument that our subspecies was incapable of colonizing suitable habitat above that elevation limit. Information has since become available documenting that our subspecies did dwell well above that elevation, making use of the habitat that they found.

Much of that evidence comes from areas in the Central Sierra such as the upper reaches of the Carson River. For example, Donald Tappe recorded an eyewitness who said that beaver were plentiful on the upper part of the Carson River and its tributaries in Alpine County until 1892 when they fell victim to heavy trapping. After the January 1986 flood, several old beaver dams dating to the early 1800s were re-exposed at elevation 5,400 feet in the Feather River Basin.

Accounts also survive of beavers at relatively high elevations in the Tulare Lake Basin including the area that is now Sequoia and Kings Canyon National Parks.

A 14-man party led by Ewing Young which included Colonel Jonathan J. Warner trapped the Kings River in the fall of 1832 “up to and some distance into the mountains and then passed on to the San Joaquin River, trapped that river down to canoe navigation in the foothills…”

Young and Colonel Warner are believed to have trapped the North Fork of the Kings River up to about present-day Courtright Reservoir (elevation 8,170 feet). They then crossed the divide to the South Fork of the San Joaquin River. A logical route would have been via Hell for Sure Pass, taking them into present-day Kings Canyon National Park.

Aspen Meadow (elevation 8,206 feet) and Blaney Meadows on the upper reaches of the San Joaquin have what would seem to be quite suitable beaver habitat. As Young’s party progressed down the San Joaquin, they came upon the trail of another trapping group. When they caught up with that group, it turned out to be a Hudson’s Bay Company party led by Michel Laframboise.

Earle Williams similarly interpreted accounts of Colonel Warner’s expedition, stating that “Warner had been trapping fur-bearing animals at the headwaters of the Kings River about the same time that the Walker party was descending the Merced River.”
Floods and Droughts in the Tulare Lake Basin
General Flood and Drought Notes

Donald Tappe recorded an eyewitness account from a retired game warden in 1940, who stated that beavers were “apparently not uncommon on the upper part of the Kings River” until 1882–83.221

On the Kern River, Roy De Voe, a native of the lower Kern Canyon, recalled that he had seen “very old beaver sign” at Lower Funston Meadow (elevation 6,480 feet) in 1946. De Voe also reported that his friend Kenny Keelor trapped the Kern River for beavers around 1900, making his camp at the mouth of Rattlesnake Creek (elevation 6,585 feet) until they were largely trapped out by about 1910–1914.222

The presence of Beaver Canyon Creek, tributary to the lower Kern River just east of Delonegha Hot Springs, is also consistent with the Kern River Basin having historically supported a population of native beavers.

In 2012, Charles James and Richard Lanman found evidence that beaver were present historically in the Sierra. They reevaluated historical records of occurrence by reliable observers, as well as new sources of indirect evidence including newspaper accounts, geographical place names, American Indian ethnographic information, and assessments of habitat suitability. They concluded that beaver historically occurred throughout some, if not most, of the Sierra and on its eastern slope.223

Officially, mink records extend only as far south as Fresno County. We have only James Carson’s account that mink were in the Tulare Lake area in the 19th century. Because all the wetlands were interconnected, it seems logical that beavers, mink and river otters were able to move about from one end of the Tulare Lake Basin to the other.

There are a number of well-documented mink sightings from Kaweah Oaks Preserve in the 1980s and possibly later. These include numerous photographs taken from a blind set up on the edge of Deep Creek, a mink that entered that blind to take food from the photographer, and a road-kill mink on Highway 198 next to the preserve. Rob Hansen recalled an anecdotal reference from another observer who saw a mink on the 3¼ mile stretch of the Kaweah River channel between McKay’s Point and Highway 245 (Road 212).

Kirk Stiltz recalled observing mink twice in the Kaweah River. The first time was in about 1979 just below the Slicky swimming hole in Three Rivers. The other time was in about 1986 just above where Highway 245 (Road 212) crosses the river south of Woodlake.

There are four records of mink in Sequoia and Kings Canyon National Parks. All of these were by reasonably reliable observers, but none were backed up with photographs or specimens. Therefore, the national parks consider the current status of mink in the park to be unconfirmed. The last reported mink observation in the parks was in 1994 near Bearpaw Meadow (elevation about 7,700 feet).

The presence of river otters in the Tulare Lake Basin ecosystem has also been questioned, but there have been observations in addition to those of James Carson.

Frank Latta recorded that the Yokuts who lived in the vicinity of Goose Lake hunted both beavers and otters.224

In addition, a group of sailors arrived at the south end of the Tulare Lake Basin in 1853 and began trapping beavers and otters from Kern and Buena Vista Lakes (see the section of this document that describes the 1852–53 floods.)

Joseph Dixon interviewed J.W.B. Rice who lived on the Kaweah River, four miles northeast of Lemon Cove. This would have been in the vicinity of present-day Terminus Dam. Rice reported having trapped three river otters in the Kaweah near his place and knew of others being taken on that stream.225

There are four reported observations of river otters in the national parks:226
1. William Colby and Poly Kanawyer reported river otter in the Kings Canyon at about 5,000 feet elevation in about 1910. That would have been on the Middle Fork Kings River, about four miles upstream from Tehipite Valley.
2. Poly Kanawyer also reported river otter in Simpson Meadow (on the Middle Fork Kings River) in about 1910.
3. Ray Walls, an electrical foreman, reported seeing a river otter on the South Fork Kaweah River near Ladybug Camp in March 1941.
4. CCC educational adviser McDonald of the Maxon Ranch CCC Camp also reported seeing a river otter near Ladybug Camp one month later in April 1941. Presumably this was the same animal that Walls had seen.
The fur-bearing population in the San Joaquin Valley was rapidly hunted to near extinction, so no fur post was ever established here. James "Grizzly" Adams noted that a few beavers survived in secluded spots up to the late 1850s.227 Walter Fry reported beavers at Ash Mountain (on the Kaweah River) in the national parks in 1920.

The California Department of Fish and Game (CDFG) reintroduced beavers to the upper tributaries of the South and North Forks of the Kern River between 1949–52.228

According to James Carson, deer (both "red and black-tailed") were present in large numbers in the general vicinity of Tulare Lake in the middle of the 19th century.229 Apparently "red deer" was a reference to California mule deer. What we think of as red deer today are native to Europe. The California mule deer that lived in the mountains moved down to the area around Tulare Lake in the winter.230

The reference to black-tailed deer is a more interesting question. Today Columbian black-tailed deer are found largely in Northern California. They range south through the Monterey Bay and Big Sur areas to Ragged Point where they are replaced by California mule deer. That is some 150 miles west of Tulare Lake.

John W. Audubon recorded that his party killed a black-tailed deer just north of the Mission of San Fernando when they traveled through that area in 1849.231 That mission is located on the north side of Los Angeles. So Audubon was apparently describing the Santa Clarita area. That is some 150 miles south of Tulare Lake.

There is no official record of Columbian black-tailed deer ever having lived in the San Joaquin Valley. However, perhaps Carson was right that a population of Columbian black-tailed deer was living in the Tulare Lake area in the 1850s. We really don't know what was present in that ecosystem before it was radically altered.

Coyotes, elk, and pronghorn were also present in very large numbers in the general vicinity of Tulare Lake in the middle of the 19th century.232

The following excerpt from a 1904 article in the Bakersfield Daily Californian entitled "The Phantom Antelope" provides a sense of the historical extent of the grassland-wetland ecosystem along Cross Creek just north of Visalia and offers some notion of the wildlife populations encountered by early settlers in this part of the San Joaquin Valley.233

*In the early (eighteen) fifties the plains between Kings River on the north and the Four Creeks timber (Kaweah River in vicinity of Visalia) on the south were the ranging ground of vast herds of antelope. A stream of water known as the Elbow swamp, caused by the spreading out of a branch of the Kaweah River that waters the Visalia country, split the plain between Kings River and the Four Creek timber into about two equal portions, debauching into Tulare Lake about midway between the two above named points. A level plain of about ten or twelve miles in width lay on each side north and south of Cross Creek and from the point where it left the Elbow swamp it was about eighteen or twenty miles in length to the point where it emptied into Tulare Lake. A narrow fringe of willows grew along its banks, with occasional bare breaks, and all along its banks on either side it was a favorite watering place for the herds of antelope that ranged upon the plain, which was covered at the time I speak of, with a luxuriant growth of grass extending all the way from the Sierra Nevada low hills, a distance of almost thirty to forty miles to Tulare Lake. Along the banks of this creek was an ideal hunting ground and it was the principal source from which the larders of the settlers at and around the vicinity of Visalia were supplied with flesh; and extremely palatable and juicy flesh it was, for nothing in the shape of meat can excel for flavor and excellence the roasted ribs of a fat antelope or a steak from one of his hind quarters. They were usually at their best in June and a prime buck at that stage carried globes of tallow on his kidneys that would rival the fattest of our south-down sheep at their best. It was customary for a hunting party to encamp in one of the depressions easily found along the creek banks, and then spread out a mile or more apart and await the advent of a band of antelope at an accustomed watering place three or four men could secure a large supply of game in a day’s hunt. I have been one of a party of five that killed ninety antelope in one day in this manner.*

Pronghorn (antelope) meat became cheap and abundant in markets in the 1850s.234 Pronghorn and elk populations were quickly decimated by market hunters in the days before game laws. Pronghorn and elk were essentially eliminated by 1870.

Black and grizzly bears were also present. (The last grizzly bear in the state was killed in 1922 at Horse Corral Meadow in what is today Giant Sequoia National Monument.)
Carson and others reported gray wolves around Tulare Lake. Some have suggested that Carson was mistaking coyotes for gray wolves. It isn’t universally accepted that gray wolves were even native to California. However, a review of early pioneer diaries shows that gray wolves were most likely present throughout California, including in the Central Valley. Joseph Grinnell concluded that “unquestionably wolves ranged regularly over the northern one-fourth of the State and south along the Sierra Nevada to Inyo County at least.”

Gray wolves had a significant presence in the San Joaquin Valley based on accounts left by early visitors. John C. Fremont noted in 1844 that he saw “wolves frequently during the day — prowling about for the young antelope, which cannot run very fast.”

John W. Audubon reported that gray wolves were very numerous when he traveled through the San Joaquin Valley in November 1849. He said that “their long, lonely howl at night ... tell the melancholy truth all too plainly, of the long, long distance from home and friends.” The wolves were so bold at night that Audubon had “several pieces of meat and a fine goose stolen from over (his) tent door.” He assumed that the wolves preyed on the elk that were abundant in the area.

Hale Tharp was the first Euro-American settler in the Kaweah canyons. He recalled that wolves were very plentiful in that area when he arrived in 1856. He saw six wolves when he took a trip to Log Meadow in what is now Sequoia National Park in the spring of 1861.

Resident populations of gray wolves are generally thought to have been extirpated from California sometime in the 1800s. Gray wolves seen or trapped in the state in the late 1800s and early 1900s are generally presumed to have been wanderers from Oregon and Nevada. However, wolves apparently held on in parts of the Southern Sierra into the early 20th century. In describing the principal animals of Sequoia National Park, the superintendent’s annual report for 1900 listed both coyote and “black wolf.”

On September 25, 1908, Charlie Howard killed a wolf at Wolverton in Sequoia National Park. At the time, Howard was slaughtering beef for a troop of soldiers, and the wolf came up within 50 yards of his camp in broad daylight and was eating some of the beef offal. Someone, almost certainly Walter Fry, inspected the carcass. (The event happened in his ranger district, and he was a very curious naturalist.) Fry later reported that the wolf was a large male in fairly good condition, but quite old, as evidenced by badly worn teeth. Guy Hopping, a long-time national park ranger and former superintendent of General Grant National Park, reported seeing and hearing a wolf in the Roaring River country of what is now Kings Canyon National Park in the summer of 1912. He described the howl as deep, like that of a big old hound. Those are the last two reliable records of gray wolves in the Tulare Lake Basin.

Walter Fry served Sequoia National Park for 25 years as its chief ranger, superintendent, judge, and naturalist. He concluded from his research that the gray wolf was native to the park. The last known specimen of a native gray wolf to be collected in California was killed by a government trapper in Lassen County in 1924. The few wolves seen in California since then are presumed to have been captive-bred wolves released by humans. One of the few examples of such a record is a wolf shot and killed while raiding a chicken coop near Woodlake in 1962. That wolf turned out to be an animal of Asian descent that was apparently an escaped pet.

On Dec. 28, 2011, a male gray wolf with a GPS collar crossed the state line from Oregon into Lassen County, becoming the first gray wolf known to live wild in California since 1924.

Mountain lions, bobcats, and gray foxes were all present around Tulare Lake. Carson reported that ocelots were also present. Whether that is a reliable account or not is uncertain. Much has changed since the 1850s. Although the ocelot’s present-day range comes close to California’s border in Mexico, there have been no confirmed sightings in the state during recent times. The ocelot’s historic range probably included at least Southern California. Harold Werner, the national parks’ former wildlife ecologist, sees no reason that it couldn’t have extended into the Tulare Lake Basin. So it’s tempting to think that Carson was correct and a few ocelots were living in the area around Tulare Lake. Unfortunately, there was nothing like a thorough inventory done of that ecosystem before it was radically altered.

Today, jaguars are present in Mexico and occasionally venture into the timbered mountains of southern Arizona and New Mexico. But jaguars inhabited a much larger area of the U.S. during the early settlement days of California. They were found in the Colorado Desert (think Palm Springs area), on islands in the delta of the
Colorado River, and in the Cuyamaca Mountains of San Diego County. They were reported from the South Coast Ranges as far north as Monterey and San Francisco up to at least 1826.\textsuperscript{247,248} James Capen (Grizzly) Adams recorded vivid accounts of his encounters with jaguars on the south side of Tejon Pass in the Tehachapis in the summer of 1855.\textsuperscript{249} The last known jaguar in California was killed in Palm Springs in 1860.\textsuperscript{250}

Felipe Santiago Garcia recorded in 1807 that wild horses and cattle were present in large numbers.\textsuperscript{251} Those feral animals were descended from stock escaped and stolen from Spanish settlements along the coast. Some accounts from early settlers said that they didn’t bother to hunt wild game because of the abundance of wild cattle.

Carson also reported that red foxes were present, but that seems suspect. The red fox is thought to be a relative newcomer to the Tulare Lake ecosystem. Non-native red foxes were introduced into the lowlands of California beginning in the late 1800s for fur farming and fox hunting.\textsuperscript{252} Today the native Sierra Nevada red fox lives at relatively high elevation (e.g., the Sonora Pass region and within the northern edge of Yosemite National Park. Based on that, it is plausible that this species was once present further south in the Tulare Lake Basin, at least at higher elevations. We have no way of knowing whether this species once came down to lower elevations. Perhaps Carson’s report was correct, or he may have been mistaken.

Carson reported that large numbers of gulls and band-tailed pigeons lived around the lake. He didn’t say whether those populations were transient or resident. In more recent times, gulls have not nested at the lake. Similarly, band-tailed pigeons today are primarily a bird of forested parts of the Sierra and foothills where blue oaks and California bay grow. At the time when Carson was writing, the Kaweah Delta was a forested area with an abundance of valley oaks. With our 21st-century perspective, it’s hard for us to imagine just how different the ecosystem was back then. It may well be that large flocks of band-tailed pigeons resided on the delta or at least visited the area seasonally.

John Xantus has been described as an insufferably arrogant, hypersensitive, difficult, jealous, generous Hungarian. He enlisted in the U.S. Army in 1855. At his first station, Fort Riley, Kansas Territory, he came under the influence of Dr. William A. Hammond, an Army surgeon, who taught the immigrant private how to collect and prepare specimens for Spencer Baird at the Smithsonian Institution. Xantus’s achievements as a tireless, all-round collector soon gained him a transfer to the medical department and promotion to the equivalent rank of sergeant.

Baird arranged for Xantus to be transferred to Fort Tejon as a hospital steward, with additional duties as librarian and baker. Stationed there for 20 months, Xantus was indefatigable in collecting insects, reptiles, and mammals as well as birds. Thanks to Xantus, we know a lot about the birds that were present in Southern California at the time, including species such as parrots and imperial woodpeckers. The imperial woodpecker is the largest woodpecker in the world and a close relative of the ivory-billed woodpecker. John James Audubon said the magpie-jay was present in the wooded areas of Northern California, so it may have been present in suitable habitat (think oaks and bay laurel) in Southern California as well.\textsuperscript{253} Today these species are associated with more tropical climates such as Costa Rica or high mountains in Mexico.

The Tulare Lake ecosystem was a significant stop for hundreds of thousands of ducks, geese, sandhill cranes, swans, curlews, snipe, and other birds migrating along the Pacific Flyway. Colonel Andrew Grayson described the density of waterfowl in the fall of 1853:

\begin{quote}
On October 31 our surveying operations brought us to the main Kern River. Here we found any quantity of elk and waterfowl, and such a place for hunters I never saw! The mallard duck abounded, but of every description of waterfowl my pen could scarcely describe the numbers or the excitement they would create in the breast of a sportsman. Your ears are confused with the many sounds — the quacking of the mallard, the soft and delicate whistles of the baldpate or teal, the underground-like notes of the rail or marsh hen, the flute-like notes of the wild geese and brant, the wild rantings of the heron, not to forget the bugle-like notes of the whooping crane and swan and a thousand other birds mixing their songs together — creates that indescribable sensation of pleasure that can only be felt by one fond of nature in its wildest and most beautiful form.\textsuperscript{254}
\end{quote}

John W. Audubon recorded that the Hutchinson’s goose (aka brant) was abundant in the San Joaquin Valley when he traveled through there in November 1849.\textsuperscript{255}
In the above account, Grayson was almost certainly referring to the call of sandhill cranes. Whooping cranes were native to the Central Valley, but there is no evidence that they were ever present this far south in the valley. John James Audubon reported whooping cranes from “upper California northward.” There were several reports of whooping cranes from the wetlands around Yuba City. Ornithologist Lyman Belding saw a flock in April 1841 near Gridley in Butte County and another flock in April 1884 over the tules on Butte Creek in Sutter County.

The following account describes what conditions were like in the mid-1870s:

The surface of the water was nearly always covered with some sort of water fowl. Geese, ducks, pelicans, snipes, mudhens, cranes and other birds were there by the millions at their own season of the year. Men went out with their blunderbusses and killed ducks by the thousands and they never seemed to grow less. The waters were filled with fish that could be caught with bare hook. During the hunting season parties left all parts of the state to go shooting on Tulare Lake. Arks were built and keepers engaged to take care of them. Hundreds of sportsmen made their annual pilgrimage to this big body of water.

The lake and surrounding wetlands were a major waterfowl hunting area until at least the mid-1880s. (For example, a wagonload of swans was sold on Main Street of Visalia on the morning of January 7, 1886. The birds, weighing 18–20 pounds each, brought $1 each.)

Sequoia National Park was created in 1890. Walter Fry recorded the changes that he observed in bird life between 1906 and 1931 and then summarized these in a report. Because of Sequoia National Park’s location, it was never a significant breeding ground for waterfowl. However, Fry said that it served as a “splendid refuge” for waterfowl and shorebirds during the winter months. Until about 1930, it was a common sight in the park to see many ducks, geese, swans, and other such fowl during the autumn, winter, and spring, some of which remained in the park throughout the year. But by about 1930, most of those species were seldom seen, and when they were seen, they were few in number. Fry reported that the rapid decrease in water- and shore-birds started about 1909 and continued through the time of his report (1931).

Fry attributed the decline of these birds to four principal causes:
1. the settlement and drying up of their breeding grounds
2. the length of the open hunting season and bag limit
3. the increase in the number of hunters
4. disease

Fry provided a table in his report comparing the number of birds present by family in 1906 and 1931. That table is reproduced in part in Table 6:

<table>
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<th>Family</th>
<th>1906</th>
<th>1931</th>
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</thead>
<tbody>
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<td>1</td>
</tr>
<tr>
<td>Loons</td>
<td>1</td>
<td>1</td>
</tr>
<tr>
<td>Cormorants</td>
<td>1</td>
<td>1</td>
</tr>
<tr>
<td>Ducks, geese, swans</td>
<td>17</td>
<td>2</td>
</tr>
<tr>
<td>Herons, egrets, bitterns</td>
<td>4</td>
<td>2</td>
</tr>
<tr>
<td>Rails and coots</td>
<td>2</td>
<td>1</td>
</tr>
<tr>
<td>Stilts</td>
<td>1</td>
<td>0</td>
</tr>
<tr>
<td>Snipes and sandpipers</td>
<td>2</td>
<td>1</td>
</tr>
<tr>
<td>Plovers</td>
<td>2</td>
<td>1</td>
</tr>
<tr>
<td>Vultures</td>
<td>2</td>
<td>1</td>
</tr>
<tr>
<td>Hawks and eagles</td>
<td>12</td>
<td>10</td>
</tr>
<tr>
<td>Cuckoos</td>
<td>2</td>
<td>1</td>
</tr>
<tr>
<td>Flycatchers</td>
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<td>6</td>
</tr>
<tr>
<td>Crows, jays, magpies</td>
<td>6</td>
<td>5</td>
</tr>
<tr>
<td>Blackbirds, orioles, meadowlarks*</td>
<td>5</td>
<td>3</td>
</tr>
<tr>
<td>Shrikes</td>
<td>1</td>
<td>0</td>
</tr>
<tr>
<td>Wood warblers</td>
<td>11</td>
<td>10</td>
</tr>
</tbody>
</table>

*3 species were lost from this family during this period, but the western meadowlark was added.
This represents a loss of 30 species in 16 families during this 25-year period. One new species was added: the western meadowlark. Fry observed that the losses were all among the water birds and migratory species. The period that Fry documented (1906–31) was generally the period during which the Tulare Lake ecosystem was being lost and the five valley lakes were drying up.

We understand the magnitude of the loss of Sequoia National Park bird life during this time, but we are unable to describe it in detail. We have a very incomplete list of which species were lost. In general we can only identify the loss at the family level. The tundra swan is the one notable exception: Fry recorded that it used to roost near Potwisha where the Marble and Middle Forks of the Kaweah River join.

We also lack a census of the number of birds by species. It is apparent from Fry’s description that waterfowl and shorebirds used to visit the park regularly during the early decades of the park’s existence. Since then, the park has remained largely undeveloped, but adjacent lands have changed radically. It is a reminder that our national parks do not exist as islands.

American white pelicans were migrants to Tulare Lake and also bred there periodically until at least 1942. Waterfowl returned in spectacular numbers when Tulare Lake had its last great reappearance in 1937–46, and there was abundant breeding by waterfowl, colonial water birds (grebes, cormorants, herons, egrets, and ibises), stilts, avocets, and terns in the South Flood Area during the floods of 1982–83 and 1997. While exciting to see, these bird congregations were ephemeral and moved on once the floodwaters receded.

During wet years, Tulare Lake was the terminus of the Western Hemisphere’s southernmost (chinook) salmon run. The Kings River supported both spring and fall runs of salmon. On November 2, 1819, Spanish Lieutenant José María Estudillo observed Tachi tribesmen catching salmon and other fish in the Kings River by means of hand nets:

> This they did before my very eyes, with great agility, diving quickly and staying under the water so long that I prayed.

Others left accounts of American Indians spearing fish in Tulare Lake and elsewhere in the Central Valley. Yokuts also set basket traps in the shallow waters to catch fish and eels. They dried and smoked large quantities of fish, prizeing salmon above all other. Yokuts tribesmen built tule balsa (rafts) and fished in Tulare Lake. These craft were reminiscent of the reed boats used on Lake Titicaca in Bolivia. The tule balsas were up to 50 feet long, could stay out on the lake for days, held up to a dozen people, and often had mud fireplaces for cooking.

In addition to the occasional spring salmon run, Tulare Lake had small populations of white sturgeon and steelhead trout and a very large population of “lake trout.” The lake trout lived year-round in Tulare Lake and in the larger tributary rivers. A very wet winter such as 1861–62 would allow them to come up the smaller streams, rising as high as Antelope Valley on the Kaweah Delta (vicinity of present-day Elderwood). The lake trout was a fine, white-fleshed fish that grew to 30 pounds and was appreciated for its taste. Despite its name, it wasn’t a salmonid; it was most likely the Sacramento pikeminnow (aka Sacramento squawfish or pik). Sacramento pikeminnow are still present in the larger rivers of the Tulare Lake Basin.

With the coming of American settlers, Tulare Lake became an important commercial fishery, shipping tons of fish to San Francisco each year. The fishery included lake trout, chinook salmon, Sacramento perch, and white catfish (after the introduction of that fish in 1873). Freshwater mussels (aka lake clams) were abundant. In addition to white catfish, a number of other exotic fishes were introduced into Tulare Lake and other lakes within the basin.

Tulare Lake was also known for its population of western pond turtles (locally called terrapin). Harold Werner, the national parks’ former wildlife ecologist, recalled reading that the turtles were once so abundant that a roar was created when sunning turtles were disturbed and took flight into the water. Those turtles were the source of a regional favorite. They were caught in seines and shipped live in sacks to San Francisco. There they were relished in terrapin soup and other delicacies.

The last category of animal that might have been present in the Tulare Lake ecosystem was marine mammals. This is not as farfetched a proposition as it might sound. During high-water periods, Tulare Lake was connected to San Francisco Bay by the San Joaquin and Kings Rivers. That was how chinook salmon entered the lake, so it is plausible that marine mammals could have used the same route. If they did do this, they would have
encountered apparently suitable habitat. Tulare Lake was brackish, rather like an estuary. There was a food supply including an abundance of freshwater mussels and a wide variety of fish.

Frank Latta reported that the Yokuts who lived around the lakes harvested both seals and sea otters. He also observed that Spanish expeditions reported seals and sea otters 150 leagues (375 miles) upstream from San Francisco Bay. That measurement was presumably made along the Sacramento River. The same distance along the length of the San Joaquin would encompass the Tulare Lake ecosystem. Marine mammals, at least sea lions, are still occasionally observed coming up the San Joaquin River. In February 2004, a male California sea lion came up the river and canals as far as Henry Miller Road north of Los Banos. He just kept going upstream until he ran out of water. At that point, he was about 65 miles from San Francisco Bay and only 100 miles from Tulare Lake. When a California Highway Patrol car arrived, the animal lumbered over, jumped up on the trunk and lay down. In April 2014, a sea lion pup was found near Modesto close to the boundary of the San Joaquin River National Wildlife Refuge.

"Wildlife" includes animals of all sizes, even mosquitoes. California has multiple species of Anopheles mosquitoes, anyone of which can transmit malaria. Although those mosquitoes are native to the state, malaria is not. Malaria was unintentionally introduced into the San Joaquin Valley from Oregon in early 1833 by a party of beaver trappers from the Hudson’s Bay Company. More than 20,000 American Indians died from the disease that spring. These included Yokuts, Chumash, Miwok, and others.

Malaria remained epidemic in the Central Valley from 1833 until the late 1800s. Case recognition and treatment, combined with successful mosquito abatement, essentially eliminated malaria as a major health concern in California by the early 1900s. In the past, malaria was endemic throughout much of the continental United States. More than 600,000 cases occurred during 1914. During the 1940s, a combination of improved socioeconomic conditions, water management, vector control efforts, and case management were successful in vastly reducing locally transmitted cases of malaria in the U.S.

In the Southern U.S., we got rid of malaria in the early 1900s through a concerted federal effort to drain the areas where mosquitoes were breeding. But that didn’t seem to be as necessary in the San Joaquin Valley where the canals served to drain the wetlands. We never got rid of all the Anopheles mosquitoes in California. Drying up the lakes and wetlands helped by greatly reducing the number of mosquitoes. But the Anopheles mosquitoes are still out there. On very infrequent occasions, there are still small, localized outbreaks of malaria in California, transmitted by these species of mosquitoes.

In its pristine state, the Tulare Lake Basin was like a wheel of water, with Tulare Lake as the hub and all the Sierra streams as spokes in the wheel. Once Tulare Lake (and the other four valley lakes) had been dried up, disintegration of this remarkably complex system was sealed with the damming of the four main rivers. The functioning infrastructure of this formerly biodiverse ecoregion was so badly broken that it resulted in the loss of most of the wetland habitat and nearly all of the biological connectivity between the watersheds in the high country and the lowland floodplains. Water-dependent habitats on the adjacent land, particularly on the Kaweah Delta and other riparian corridors, were also significantly degraded during the ensuing decades.

The loss of the Tulare Lake ecosystem affected even protected areas like the national parks. For example, we speculate that the relict populations of beavers, mink, and river otters that hung on in the park became isolated from populations elsewhere in the parks and the basin. Their numbers gradually declined, and some of those species may now be extinct in the parks. We also speculate that a similar problem occurred with many populations of fish, birds, and other animals.

Why is there no lake in the Tulare Lakebed today?

One major cause is that we’re using a lot more water, primarily for agricultural purposes. We’re taking that water from our rivers before the water reaches the historic lakebed. In a big picture sense, there is not enough water to sustain both Tulare Lake and the needs of people. Society has found what it considers to be a better use for the water: serving the needs of people, rather than serving the needs of a natural resource.

Table 7 details both total runoff and inflow to Tulare Lake for the 19 largest runoff years that we are aware of. The total runoff shown in the center column is based on the data behind Figure 18 on page 111. For the source of the runoff data in the middle column, see the section of this document that addresses Measurements of Flows and Runoff. The inflow shown in the right-hand column is based on data covered under individual floods. Diversion of river water for irrigation began on the Kings and Kaweah Deltas in about the 1870s. By the end of the 19th century, there was a large network of canals diverting water out of the rivers in order to support extensive irrigated agriculture.
The story of how irrigation came to the Kings River service area is particularly well known. In the 1850s, there was virtually no settlement in the region, and the empty plains were considered a worthless desert. The valley floor wasn’t really a desert but a seasonal grassland. Those grasses grew lushly when adequately watered by rain, a sign of the land’s incredible fertility and growing conditions. That began to change in the 1850s in what is now known as the Centerville Bottoms where a few very small ditches were dug near the river. Other ditches tapped into the Kings between 1863–66 to bring water into the Centerville area. The first canal of substance was the Fresno Canal in 1871; it was an enormous success.

Dozens of canals were built across the plains over the next 30 years, bringing water to the immense, previously uncultivated prairie and allowing it to be converted into farms. Most major canal construction was completed by 1900. By then, the Kings River — historically Tulare Lake’s most important source of water — was irrigating over a million acres, more land than any other stream in the world except the Nile and Indus Rivers.

At the time of EuroAmerican settlement, most of the water of the Kings River used to flow into Tulare Lake via what is now called the South Fork system, along the south side of its delta. From 1861–1884, various floods and man-made ditches began making a channel for the Kings River along the north side of its delta. The key constriction was the Zalda Canal, a ditch constructed in 1872 and enlarged by two subsequent floods. But the majority of floodwaters still flowed along the south side of the delta and into Tulare Lake.

The 1916 flood significantly opened the Zalda Canal, and thereafter it became the main channel for the Kings River. This reach is now known as the North Fork of the Kings or the Kings River North Channel. In the 1916 flood, this channel is said to have discharged 60% of the Kings runoff into the San Joaquin River and thence to San Francisco Bay. These changes in the flow of the Kings River soon left the farmers in the Tulare Lakebed without sufficient water to irrigate the reclaimed grain land, forcing them to sink deep wells for their irrigation water. The combination of all the canals plus the rerouting of the Kings River is the primary reason that there is generally no lake in the Tulare Lakebed.

Table 7 illustrates how dramatically inflows to Tulare Lake decreased as a result of those diversions. The large runoffs that used to sustain Tulare Lake continued to come. But after the turn of the 19th century, farmers were very successful in diverting most of those waters from the lakebed onto their irrigated lands. When the four federal reservoirs began operation during the 1954–61 period, they had an important, but relatively less noticeable effect on the amount of inflow to Tulare Lake.

<table>
<thead>
<tr>
<th>Water Year</th>
<th>Total runoff of 4 major rivers¹ (acre-feet)</th>
<th>Total inflow to Tulare Lake² (acre-feet)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1853</td>
<td>3</td>
<td>5,096,000</td>
</tr>
<tr>
<td>1862</td>
<td>3</td>
<td>6,290,000</td>
</tr>
<tr>
<td>1868</td>
<td>3</td>
<td>5,360,000</td>
</tr>
<tr>
<td>1906</td>
<td>7,195,240</td>
<td>1,530,000</td>
</tr>
<tr>
<td>1909</td>
<td>5,689,840</td>
<td>1,175,000</td>
</tr>
<tr>
<td>1916</td>
<td>6,512,710</td>
<td>1,041,700</td>
</tr>
<tr>
<td>1938</td>
<td>5,773,470</td>
<td>126,000</td>
</tr>
<tr>
<td>1952</td>
<td>5,375,050</td>
<td>583,000</td>
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<tr>
<td>1967</td>
<td>6,253,344</td>
<td>94,300</td>
</tr>
<tr>
<td>1969</td>
<td>8,379,585</td>
<td>1,155,000</td>
</tr>
<tr>
<td>1978</td>
<td>6,078,925</td>
<td></td>
</tr>
<tr>
<td>1980</td>
<td>5,821,879</td>
<td></td>
</tr>
<tr>
<td>1982</td>
<td>5,201,438</td>
<td></td>
</tr>
<tr>
<td>1983</td>
<td>8,746,222</td>
<td>1,069,000</td>
</tr>
<tr>
<td>1986</td>
<td>5,692,766</td>
<td></td>
</tr>
<tr>
<td>1995</td>
<td>5,814,847</td>
<td></td>
</tr>
<tr>
<td>1997</td>
<td>4,931,557</td>
<td></td>
</tr>
<tr>
<td>1998</td>
<td>5,990,549</td>
<td></td>
</tr>
<tr>
<td>2011</td>
<td>5,910,342</td>
<td></td>
</tr>
</tbody>
</table>

¹This is the total runoff of the Kings, Kaweah, Tule and Kern Rivers.
²See Table 18 for examples of exports that have been made in recent decades to keep floodwaters out of the Tulare Lakebed.
³No runoff data is available for these rivers prior to 1894.
⁴There were some relatively small inflows to Tulare Lake in each of these years, but no measurements are available.
Floods and Droughts in the Tulare Lake Basin
General Flood and Drought Notes

But there is another reason that the lake isn’t in the lakebed, at least in most years. That also involves society’s values: the value that it places on the lakebed itself. Society used to place a high value on Tulare Lake as a resource for all the food that the lake and its associated wetlands produced. However, today the resource that society values is the irrigated agricultural crops, improvements, and the towns that occupy the lakebed. Therefore, society marshals its resources to try to prevent the lake from returning to its lakebed. And when the lake does return in varying degrees, people strive to minimize the damage that it causes. Some people view Tulare Lake as an inconvenience, a nuisance to be prevented. Society has defined the presence of excess water in the lakebed as a flood.

A variety of steps have been taken in recent decades to keep water out of the Tulare Lakebed:
- Encouraging users below the four federal reservoirs to take all the water that they can productively use or store in retention basins so that the reservoirs can be drawn down in anticipation of an oncoming flood.
- Coordinating the operation of the federal reservoirs to keep floodwaters out of the lakebed. This has included engaging the assistance of private interests and of PG&E in this complex effort.\(^{276}\)
- Installing temporary sandbag (or sack concrete) barriers on the spillway of the federal dams, thereby allowing the reservoirs to operate as much as 5½ feet above full pool level; thus keeping this water from flowing down to the Tulare Lakebed. See Figure 4 for a map showing how the federal reservoirs sit upstream of the lakebed. At Terminus Dam, the need for such temporary measures was eliminated when fuse gates were installed in 2004.
- Blocking the Kern River at Sand Ridge, causing a huge holding pond to form at the south end of Tulare Lake. This holding area has since been further developed and is now known as the South Flood Area.
- Diverting Kings River floodwaters to the San Joaquin River in order to minimize flooding in the Tulare Lakebed. This is done using the North Fork / Fresno Slough / James Bypass channel. The Fresno Slough Bypass (now known as the James Bypass) began operation in 1872. The capacity of the associated system has since been increased several times. Prior to about 1872, all of the Kings River water flowed into Tulare Lake. See the section of this document on Pine Flat Dam for a more detailed description of the James Bypass. Water that is sent through this system winds up in San Francisco Bay; it is essentially a loss from the point of view of Tulare Lake Basin water users. The transfer of water through this system was greatly reduced. Even so, diversions through this system have occurred in 38% of the years since the dam was completed.\(^{277}\)
- Diverting Kern River floodwaters into the California Aqueduct rather than into Buena Vista and/or Tulare Lakes. This is done using the Kern River Intertie and Cross Valley Canal. Once the water enters the California Aqueduct, it is pumped over the Tehachapi Mountains and sent to the Los Angeles area. The Kern River Intertie was completed in 1977.\(^{278}\) Prior to that, a big flood on the Kern would first fill Buena Vista and/or Goose Lakes, and then spill into Tulare Lake. The term "beneficial use" refers to a reasonable quantity of water applied to a non-wasteful use. Transferring Kern River floodwaters over the Tehachapis to the Los Angeles area has now been determined to be a beneficial use.\(^{279}\)
- Transferring water from the Kings, St Johns, and Tule Rivers into the Friant-Kern Canal in order to minimize flooding in the Tulare Lakebed. This is done by using pumps at the point where each of those rivers cross the canal. Once the river water enters the canal, it flows by gravity to the canal’s terminus near Bakersfield. There it is emptied into the Kern River. The water is then routed to the Los Angeles area using the Kern River Intertie and Cross Valley Canal as described above. A combined total of over 472,000 acre-feet of floodwaters was pumped into the canal during the years 1978, 1980, 1982, 1983, 1986, 1995, 1997, 1998 and 2006. (The total amount may have been a good bit more than this; records are incomplete.) Transfers may have been made in later years as well. Including all sources (four rivers), exports to the LA area have occurred in 30% of the years since the Kern River Intertie began operation in 1977.\(^{280}\)

Despite all of the above efforts, floodwaters still make it to the Tulare Lakebed on occasion, especially in heavy runoff years (see Figure 16). To assist in the reclamation of the lakebed, over 20 reclamation districts were formed under California general reclamation district laws between about 1896 and 1925. The reclamation districts have built levee systems which divide the lakebed into cells or sumps. As floodwaters come into the lakebed, the sumps are filled, more or less in order. The first four cells (the South Wilbur Flood Area and the three Hacienda Reservoirs) are devoted to holding floodwater; they are never planted in crops. This use of lakebed levees minimizes the damage and allows the remaining portions of the lakebed to be used for agricultural purposes. In huge runoff years, emergency levees still have to be constructed within the lakebed to protect the towns of Corcoran and Stratford.
Role of Floods in Maintaining Tulare Lake

Our first-hand knowledge of Tulare Lake dates back over 150 years to the middle of the 19th century. Floods would abruptly raise the level of the lake after which it would gradually shrink during the drought or non-flood years that followed (see Figure 15).

As described earlier, Tulare Lake and the other four valley lakes were not landlocked, an inland sink, the way that we think of them today. They were the anchors of a wetland complex of over 400,000 acres (see Figure 5). That complex connected with the wetlands that fringed the San Joaquin River, making a continuous wetland all the way to the Sacramento–San Joaquin Delta.

The flood cycle of the Tulare Lake Basin was critical in maintaining that ecosystem; the floods provided sufficient water storage to keep the lake going through the drought or non-flood years. Once Tulare Lake and the other four lakes had been dried up, the last remnants of the ecosystem totally disintegrated with the damming of the four main rivers. The functioning infrastructure of this formerly biodiverse ecoregion was so badly broken that it resulted in the loss of most of the wetland habitat and nearly all of the biological connectivity between the watersheds in the high country and the lowland floodplains.

Floods — with the water that they brought — created a marvelous ecosystem in the Tulare Lake Basin. Reminders of that ecosystem survive in disjointed preserves in the valley, in the foothills, and in the Sierra. The framework of the hydrologic system that powered that ecosystem still exists today. On occasion, flooding can recreate a portion of Tulare Lake. The last significant reappearances of the lake were brought on by the floods of 1982–83 and 1997.

But just adding water to the Tulare Lakebed is not enough to recreate the complex ecosystem that once existed. The associated habitat is highly degraded, and the ability of the ecosystem to provide connections among the various river and stream courses in the Tulare Lake Basin has largely been lost. See the section of this document that describes Wildlife in and around Tulare Lake for a discussion of how waterfowl have responded when high-water years have rewatered portions of the Tulare Lakebed. Birds and other wildlife are still attracted to water wherever they can find it. Preservation and restoration of lost and degraded wetland habitat is being pursued by various landowners and conservation groups.

Chronology of Tulare Lake

A popular perception is that Tulare Lake was relatively stable before agricultural diversions began. Perhaps reflecting this mythology, one source said that when the Spanish first visited Tulare Lake in 1772, it was about 50 miles long and 35 miles wide. No source was given for this measurement, so it should probably be attributed to legend or wishful thinking. What information we do have on lake levels prior to 1844 indicates that it was not constant, but varied as a function of runoff and perhaps other climatic factors.

Annie Mitchell wrote that American Indians said that Tulare Lake went dry about 1825.

John C. Fremont (along with his scout, Kit Carson) led two government expeditions through the San Joaquin Valley. Carson was a good choice because he had traveled from Taos to the San Joaquin Valley in 1830 on a trapping expedition. On his first expedition in 1844, Fremont explored the east base of the Sierra as far south as present-day Bridgeport. Short on supplies, Fremont then decided to make the first ever mid-winter crossing of the Sierra.

In late January, the party turned west and started pushing their way up the East Fork of the Carson River. By February 6, conditions were appalling: they were lost, out of food, and the stock was in poor shape. Quoting from Fremont’s diary for that day:

Two Indians joined our party here, and one of them, an old man, immediately began to harangue us, saying that ourselves and animals would perish in the snow; and that if we would go back he would show us another and a better way across the mountain. He spoke in a very loud voice, and there was a singular repetition of phrases and arrangement of words, which rendered his speech striking, and not unmusical.

We had now begun to understand some words, and, with the aid of signs, easily comprehended the old man’s simple ideas. “Rock upon rock — rock upon rock — snow upon snow — snow upon snow” said he; “even if you get over the snow, you will not be able to get down from the mountains.” He made us the sign of precipices, and showed us how the feet of the horses would slip, and throw them off from the narrow trails which led along their sides. Our Chinook, who comprehended even more readily than
ourselves and believed our situation hopeless, covered his head with his blanket, and began to weep and lament. "I wanted to see the whites," said he; "I came away from my own people to see the whites, and I wouldn’t care to die among them; but here" — and he looked around into the cold night and gloomy forest, and, drawing his blanket over his head, began again to lament.

Seated around the tree, the fire illuminating the rocks and the tall bolls of the pines round about, and the old Indian haranguing, we presented a group of very serious faces....

Fremont decided to continue on up the Carson River, and the party crossed what would later be known as Carson Pass on February 14. Traveling through deep snow and blizzard conditions, the expedition reached Sutter’s Fort (in present-day Sacramento) on March 6, 1844. Only 33 of their 67 horses and mules survived the passage, most of the rest had been eaten. Continuing south, Fremont reached the Kings River on April 8, and noted that most of the flow of that river was going into Tulare Lake. This might imply that Fremont thought that a small amount of Kings River water was flowing north.

He also observed that Tulare Lake was overflowing into the San Joaquin River, indicating that the lake was sufficiently high to overtop the delta sill that dams it. This is the earliest reliable measurement of the size of Tulare Lake. In 1844, the southern lake (Ton Taché) was nearly as extensive as the northern lake (Taché). A slough connected the two lakes, passing through Sand Ridge. During very high-flow years, this slough served (and still serves) as part of the extension of the Kern River.

Fremont returned to the San Joaquin Valley on his next expedition in the winter of 1845–1846. This time, he split his party and led one portion west up the Truckee River and crossed the Sierra at Donner Pass, arriving at Sutter’s Fort on December 8, 1845. From there, he headed south to rendezvous with the rest of his expedition at the river that he knew as the Lake Fork of the Tulare River. We know that river today as the Kings River. Fremont’s group arrived at the rendezvous first.

When he didn’t find the rest of his group at the rendezvous spot, Fremont led his portion of the expedition (16 men on horseback herding a number of cattle) up the Kings River in search of them. What began as a search would develop into a mid-winter exploration of the Sierra.

After a one-day rest, they began their trek east on December 24, driving their cattle with them. They rode through the oak woodlands along what was likely the mainstem and North Fork of the Kings River for several days until they started to climb higher.

They worked their way up through the oak and conifer forests and some “extremely large” trees. The historian Francis Farquhar interpreted those trees to be the McKinley Sequoia Grove, seven miles west of present-day Wishon Reservoir in the North Fork Kings River Basin. The party eventually reached the 11,000 foot elevation, coming out onto a bare granite ridge that divided the North Fork of the Kings and the South Fork of the San Joaquin River. One of the few places on the divide where they could have driven cattle to this elevation was in the vicinity of Hell for Sure Pass at the west boundary of present-day Kings Canyon National Park. It was a beautiful day to be in the Sierra, the weather comfortably warm.

The next day, December 31, Fremont’s party headed back down to the San Joaquin Valley. They had almost waited too late to begin their return. The weather quickly turned bad as a big snowstorm blew in. Fremont’s decision to check out the Sierra had been an incredibly rash act, and the storm almost cut off their escape from the mountains.

The old year went out and the new year came in, rough as the country.

They soon had to abandon their cattle and had difficulty getting themselves out from the snow. But within a few days they returned to the valley floor, completing their grand scouting adventure. On January 4, 1846 they returned to the Kings River, camping at the east end of present-day Pine Flat Reservoir.

After returning to the valley, the expedition engaged in more mundane explorations. Among other accomplishments, they mapped Tulare Lake as being about 60 miles long. If that accurately represents the amount of water in the winter of 1845–1846, it would suggest that the lake was close to being full. However,
S.T. Harding concluded that Fremont’s length measurement probably included flooded areas outside of the actual lakebed, such as the Fresno Slough area.

Lieutenant George H. Derby of the U.S. Army’s Topographical Engineers visited the Tulare Lake area in May 1850. By then, the southern lake (Ton Taché) was essentially dry, having been drained by the slough that passed through Sand Ridge. The remains of that southern lake formed a tule swamp 10 miles wide and 15 miles from north to south.

Derby reported that the gradual receding of the water was distinctly marked by a ridge of decayed tules upon its shore, and that he had been informed, and had no reason to disbelieve, that 10 years previous it had been nearly as extensive a sheet of water as the northern lake. That seems plausible. Fremont’s measurement of the lake suggests that it was full or nearly so when he visited it in 1844, six years before Derby’s visit.

One source said that Tulare Lake measured 570 square miles in 1849, but the reliability of that measurement is questionable. There were no known surveyors in the area at the time. Possibly this is a reference to the lake’s size in 1850 when Derby measured it.

On May 9, 1850, Derby came to the Kern River, which was discharging into Buena Vista Lake by two separate mouths. At that time, Buena Vista Lake was 10 miles long and 4–6 miles wide.

After leaving the Kern, Derby crossed the Tule, Kaweah, and Kings Rivers, heading north. He then turned west to explore the huge wetland complex between Tulare Lake and the San Joaquin River (i.e., the Fresno Slough system). 1850 was a very heavy runoff year, the heaviest in the memory of the American Indians who lived on the Kings River. Derby discovered that the entire flow of the Kings and a significant portion of the San Joaquin were flowing toward Tulare Lake.

As noted in the section of this document on General Notes on Tulare Lake, C.E. Grunsky believed that Derby may have been mistaken about the San Joaquin River flowing south. Some of the San Joaquin River downstream of Gravelly Ford could flow to the southwest in sloughs during flood periods but would eventually be intercepted by the Fresno Slough. Grunsky concluded that the water in the Fresno Slough was flowing from the Kings River Delta north toward the San Joaquin River, and that part of the Kings River was flowing south to Tulare Lake.

As the peak of the flood was approaching, the lake was still not overflowing the delta sill. Tulare Lake had been about one foot above the delta sill prior to the flood (elevation 208 – 207 feet). But because the sill was densely vegetated with tules, significant outflow didn’t really start until the lake reached an elevation of 210 feet. The flood would eventually raise the lake to a maximum elevation of 211.5 feet, sending water flowing back toward the Sacramento–San Joaquin Delta. (Derby was lucky. The waters were rising rapidly while he was in the area and his party barely escaped entrapment.)

In 1851 and 1852, the lake remained almost brim-full, nearly to the level of the delta sill that serves to regulate its maximum height. However, the flood of 1852–53 raised Tulare Lake by 11.5 feet. At this point, the lake had a depth of about 37 feet at its deepest point and a maximum elevation of 215.5 feet. Tulare Lake would reach this size only twice more: in 1862 and 1867. Over the next eight years (1853–61), the lake dropped 16 feet. This became the pattern for the lake over subsequent decades. Floods would abruptly raise the level of the lake after which it would gradually shrink.

While Tulare Lake was at this high stage during the 1852–53 flood, some sailors jumped ship in San Francisco and stole a whaleboat. They hoisted the sail and headed inland. Taking advantage of the prevailing winds, they sailed south up the San Joaquin River, through the Fresno Slough, and entered Tulare Lake. This was the first of six documented trips between that lake and San Francisco Bay to occur in historic times. (The other five trips were in 1868, 1938, 1966, 1969, and 1983.)

John M. Barker lived on a cattle ranch on the Kings River near Tulare Lake. One morning in the winter of 1857, he and a neighbor started out on horseback to search for some horses that had strayed. They skirted the shores of Tulare Lake between Cross Creek and the Kings River (apparently just west of present-day Corcoran). For a couple of miles from the shore, the waters in the shallows were covered with burnt tules and other refuse matter unfit for use by man or beast.

They knew that their horses would not drink from the lake (presumably because it was brackish and alkaline), but there were sloughs and water in depressions outside of the lake, where the water was clear and fit for use.
Floods and Droughts in the Tulare Lake Basin

General Flood and Drought Notes

They headed to one of those waterholes in order to look for tracks of their missing stock. As several of them were shod, they knew if they found shod tracks that they would be on the right trail.

Barker dismounted and walked to the edge of the water. Just as he reached it, a massive earthquake struck. The lake commenced to roar like the ocean in a storm, and the cowboys rode as fast as they could to get away from there. They returned the next day and found that the lake had run up on the land for about three miles. Fish were stranded in every direction and could have been gathered by the wagon-load.

Apparently the earthquake that Barker experienced was the Southern California Earthquake of 1857 (aka the Fort Tejon Earthquake). At magnitude 7.9, this was the most powerful earthquake to hit Southern California in historic times.

Tulare Lake had been at a very high stage after the 1852–53 flood, the second highest ever recorded. After 1853, there was a gradual shrinkage of the lake until the fall of 1861. Over those eight years, the lake dropped about 13 feet in elevation.

There were multiple causes of this:
- The maximum elevation in 1853 (215.5 feet) was higher than the elevation of the Tulare Lake sill (207 feet). The water above this elevation simply flowed out of the lake and connected through the Fresno Slough to the San Joaquin River, and from there it flowed on to San Francisco Bay.
- Normal evaporation in our hot valley summers (averaging 5.2 feet per year).
- Eight years with only low or average runoff and no floods.
- Two years of drought (1856 and 1857).
- Diversion for irrigation was just getting underway (negligible).

Gordon’s Ferry (aka Gale’s Ferry) was located just north of present-day Bakersfield College. The Sinks of the Tejón was the first Butterfield Overland Mail stop north of Fort Tejon. It was located at the intersection of present-day David and Wheeler Ridge Roads, roughly 10 miles northeast of where Interstate 5 and Highway 99 diverge. When the Kern River came out of its canyon in the winter of 1861–62, it created one vast sea of water from Gordon’s Ferry to the Sinks of the Tejón. Kern Lake was located in the southeast corner of that huge sheet of water. Buena Vista Lake backed up to within 12 miles of Fort Tejon.

In the summer of 1861, Tulare Lake reached a low of 200.3 feet. The 1861–62 flood raised the lake by 15.7 feet to elevation 216, the highest that the lake has been during historic times. At elevations above 207 feet, the lake over-topped the lowest point on the Tulare Lake sill. At the lake’s highest stage, about 9 feet of water flowed in a broad expanse northerly over this sill (elevation 216 - 207 feet). From there, the water flowed into the Fresno Slough and the San Joaquin River. At the height of this flood, the lake was about 37 feet deep at the deepest point (elevation 216 - 179 feet). The surface area increased from about 350 square miles in 1861 to about 790 square miles in July 1862.

S.T. Harding estimated that 6,290,000 acre-feet of water flowed into Tulare Lake in the single season 1861–62. For comparison, that is 3.9 times greater than the combined current capacity of all four of the federal reservoirs in the Tulare Lake Basin.

Before the 1861–62 flood, the Kern River channel ran where the Kern Island Canal now runs in Bakersfield: by the Beale Library (between Chester and Union Ave) on its way to Kern Lake. The flood shifted the river to the west. The new channel began at Gordon’s Ferry (just north of present-day Bakersfield College) and passed through what is now Old River and into the Las Palomas slough system between Kern Lake and Buena Vista Lake on its way to Tulare Lake. Not only did that new channel bypass Kern Lake, but one source said that it also bypassed Buena Vista Lake, meaning that those lakes would only get water during years with very high runoff. In any case, the river would shift even farther northwest in the 1867–68 flood.

Bill Tweed calculated that when all the tributaries of the San Joaquin River were swollen with snowmelt, the total flow of that river as it approached the Sacramento–San Joaquin Delta could exceed 100,000 cfs. At such times, the river spread out for miles across the flat valley floor. That would presumably describe the condition that existed in the flood of 1861–62.

Tulare Lake gradually declined in elevation after the 1861–62 flood. In the summer of 1867, the lake level was elevation 200.7 feet. However, the 1867–68 flood raised it by 14.7 feet, bringing it back to a maximum
elevation of 215.4 feet. At the height of the flood, Tulare Lake was almost 37 feet deep at the deepest point. The lake has not been this deep since (see Figure 15).

In 1868, Richard Smith loaded a 16-foot scow with a one-ton cargo of honey and made the 170 mile journey from Tulare Lake to San Francisco Bay. That remains the only recorded commercial trip ever made between Tulare Lake and San Francisco Bay to occur in historic times. (There were five non-commercial trips: 1852, 1938, 1966, 1969, and 1983.) Apparently Smith was able to make the return trip back through the tules to Tulare Lake.

In 1872, the Fresno Slough Bypass (now known as the James Bypass) began operation. The Kings River Handbook says that the James Bypass was “developed” in 1912–14. This may mean that the bypass was further improved at that point. In any case, this channel works with the North Fork to route a portion of the Kings River floodwaters to the San Joaquin River. Prior to about 1872, most of the Kings flowed into Tulare Lake. See the section of this document on Pine Flat Dam for a more complete description of the North Fork / James Bypass.

White catfish were introduced to Tulare Lake in 1873. This is just one of several exotic fish that would be introduced to the lake. See the section of this document on Wildlife in Tulare Lake, for a discussion of the lake fishery.

The landlocked form of Atlantic salmon (Salmo sebago, S. salar sebago, or S. salar ouananiche) was introduced to Tulare Lake in about 1878.

From 1854 to 1872 the lake changed very little in area. Almost due west from Bakersfield there was a shrinking, but otherwise its area remained about the same. It was about these years that irrigation started in the valleys around Visalia and Bakersfield and the shrinking became very rapid. The rivers were tapped in several places and the water that would have gone into Tulare Lake was spread out over the dry pastures and cotton fields. The shrinking was most marked from 1872 to 1875. The southern end of the lake contracted and became somewhat in the form of a creek. It narrowed until it was not more than a mile wide and had drawn up from the southern end at least 15 miles.

The 1878 flood filled Tulare Lake to elevation 207.5 feet, causing it to spill over the delta sill and into Fresno Slough and the San Joaquin River for the last time. That was the last natural overflow of the lake; Tulare Lake has never filled again. Since 1878, the Tulare Lake Basin has functioned largely as a closed basin, an inland sink without a regular outlet to the ocean.

A number of sailboats and at least two steamboats plied the lake in the 1870s and 1880s. The Mose Andross was a 50-foot long, side-wheel steamboat that A.J. Atwell built and operated from 1875 until 1879. This is the same Atwell who owned the lumber mill that began operation at Atwell Grove in 1879. The Mose Andross was built primarily to service Atwell’s farming interests at Atwell’s Island (site of present-day Alpaugh), but it also served as general transport during the years that it operated.

The Mose Andross was flat-bottom, so it could pull in almost anywhere. However, there were six regular landings that it serviced in addition to Atwell’s Island (that was the original spelling). Those landings were located around the lake at the following points (see Figure 13):

- Cox and Clark ranch (a double adobe, 3 miles south of present-day Kettleman City. Located where El Camino Viejo from San Pedro to San Antonio (present-day Los Angeles to East Oakland) ran closest to the lake.)
- Gordon’s Point (6 miles north of Kettleman City)
- Dan Rhoades ranch (an adobe, south of Orton Point, near present-day Lemoore)
- Buzzards Roost Landing (immediately south of present-day Waukena)
- Near the Artesia Schoolhouse (at the mouth of Cross Creek, south of present-day Waukena)
- Creighton Ranch (at the mouth of the Tule River)

The Mose Andross was used as much for pleasure trips as for freighting. The following announcement appeared in the Tulare Times in 1875:

EXCURSION ON THE LAKE: There is to be a May Day excursion on Tulare Lake and a dance on a barge in the evening. We acknowledge receipt of complimentary tickets from Captain Atwell, owner of the lake steamer to be used on this occasion.
Gustav Eisen, the Swedish natural scientist, visited Tulare Lake in 1878, and recounted his adventure 20 years later:

In 1878 I crossed Tulare Lake on a steamboat. This was a regular packet that ran between Hanford and a small town on the west side of the lake. (Presumably he was referring to Kettleman City.) The distance across was about thirty miles. There were one or two other steamboats running on the lake at the time. Sailboats were numerous and altogether Tulare Lake was of considerable use to the commerce of the region.

On the occasion that I made my trip across the lake we were all treated to a surprise. When we were about twelve miles from Hanford, and almost out of sight of land, the boat ran over the ruins of an old ranch. We could look down through ten or twelve feet of clear water and see the fence posts of an old pig sty. There was also the foundation of a house and several metal utensils scattered about. Nobody on the boat knew whether the ranch had been on an island that had sunk from sight or whether it had been on the mainland during some previous dry year. It was a mystery.

That ranch presumably became established during the low-water years of either 1857–61 or 1865–67. Each of those dry periods ended with a dramatic flood, causing the lake to rapidly rise by 10 feet or more. The owners of the ranch Eisen saw must have been stunned when that occurred. The lakebed floods can arrive with surprising swiftness and virtually no warning. See the section of this document that describes the 1867–68 flood for an account of several hog camps that were caught up in that flood. The onset of flooding in the lakebed swept in abruptly on Christmas Eve, 1867, catching people by surprise. They had to beat a hurried retreat, the rising waters on their heels the whole way.

The schooner Water Witch (formerly the Alcatraz) was brought from San Francisco to Tulare Lake by "Eating" Smith in 1878. (Smith earned his nickname for his big appetite.) The ownership of the Water Witch changed hands twice after its arrival at the lake. First Smith traded it to the McCoy brothers for some cows. The McCoys used the Water Witch for two seasons of harvesting turtles, sending as many as 300 dozen to San Francisco in one season. They then sold it to Captain Thomas J. Conley in 1880. Conley was described as living "near the notorious Work's adobe" near the South Fork of the Kaweah.

In 1880, at the time of the Creighton Survey, Tulare Lake had dropped to an elevation of 200 feet; its surface covering only 445 square miles.

In 1881, Thomas Conley patented 80 acres of land on the west side of the South Fork, near where the Shoshone Inn is today. That would have been right across the South Fork from Work's adobe, apparently a well-known landmark of the day. Presumably Conley Creek is named for Thomas Conley or his family.

In 1881, Captain James W.A. Wright of Hanford met Conley in Slapjack Canyon on the road to Mineral King. Conley had just cut new masts for his ship. The following May, Conley and Wright would embark on a six-day excursion to map Tulare Lake in detail.

Mussel Slough is located along the western edge of present-day Hanford. In 1881–82, Tulare Lake reached what was then considered an unusually low stage, about elevation 192 feet. At that point, the lake margin laid bare an area near the mouth of Mussel Slough which was covered with the broken stumps of long submerged trees (illustration on file in the national parks). C.E. Grunsky made a number of visits to that area while doing a study of the water resources of the San Joaquin Valley during the years 1881–88.

At that location, he could see the location of an old channel entering the lake from the northeast. He deduced that this was the former channel of Mussel Slough during a protracted period in which the lake was at or below its then low stage. Some of the stumps had a diameter of about four feet. Their dimensions and position indicated that they were the remnants of a grove of willows which had reached mature growth along the bank of the watercourse and the margin of the lake.

Grunsky concluded that low lake stages with conditions favorable to the growth of those willows must have been continuous for a period of some 40–50 years or perhaps much longer. Therefore, at sometime in the past before the arrival of the white man, the lake had been at or below that elevation for a long period of time. After the lake rose to a stage high enough to drown the willows, it remained at or above that stage for 50–100 or more years, keeping the stumps submerged until their discovery. The long period of persistently light or moderate
rainfall favoring the growth of the willows was followed by a long period in which the frequency of fairly wet winters kept the lake at fairly high stages, culminating with the very high waters of 1853, 1862, and 1868.

Grunsky presented his findings to a meeting of the American Meteorological Society in June 1930. At the time of his talk, the San Joaquin Valley was undergoing a severe drought. Tulare Lake had been repeatedly dry during the preceding 30 years. The prevailing assumption was that the lakebed would remain generally dry forever more. In his talk, Grunsky made the point that we needed to take a longer term view toward climate change. Tulare Lake had been dry for an extended period in the not too distant past, and he expected that it might come back to life in the foreseeable future.

In 1878, Tulare Lake had filled to elevation 207.5 feet and spilled over the delta sill. However, in 1883, just five years later, the lake had decreased to just nine feet deep (elevation 188 - 179 feet). While that was remarked upon as the lowest elevation in memory, the lake continued to dwindle in size over the next 16 years.

The lake shrank from the south to the north. In 1882 the southern border of the lake left Kern County altogether.

In May of that year, Captains James W.A. Wright (a former Civil War officer) and T.J. Conley (a man with sea-faring experience) made a six-day excursion across and around the lake in the Water Witch, making careful measurements and soundings of the lake. They found the greatest depth to be 22 feet, in a comparatively narrow depression like a river channel, from the mouth of Kings River on the north to Terrapin Bay on the south (see Figure 13). Other than this channel, the rest of the lake had a maximum depth of eight feet or less. The area of the lake was then 417 square miles.

Wright left a detailed and highly readable account of their six-day mapping adventure. The Water Witch sailed the lake for a few more months until capsizing in a severe storm about three miles southeast of the mouth of the Kings River in the winter of 1882. That location was just south of present-day Stratford.

The rapid drying up of Tulare Lake was written up in the New York Times in 1884. A few years earlier, the lake had been 33 miles long by 21 miles wide. By 1884, it had shrunk to 15 miles long with an average width of 8 miles.

As Tulare Lake shrank, the fishing technology changed. By 1887, large scale, land-based seining had become possible. Seines were taken out into the lake about a mile and drawn to shore by horse-powered windlass. There were at least five seines running, making two hauls each day with up to 2,200 pounds per haul. Each haul included 500–1200 pounds of perch in addition to catfish and lake trout. The lake trout were plentiful and ranged in size from 2–20 pounds, occasionally as large as 30 pounds.

As Tulare Lake shrank, it changed shape and configuration. The lake had been described as having the shape of an oyster when it was at full pool. By 1888, the remnants of the lake had become almost circular in shape.

As the lakes shrank, the alkalinity rose. The ecosystem started to go into a tailspin. 1888 seems to have been the pivotal year. The fishing (or seining) was apparently terrific that year as the ecosystem crashed. Over 133,600 pounds of fish from Tulare Lake were shipped to San Francisco in one ten-week period in the fall of 1888. By the end of that year, the catfish, lake trout, pond turtles, mussels, and clams had reportedly died out of all three lakes (Tulare, Kern, and Buena Vista) due to the increasing alkalinity.

The New York Times reported in the summer of 1889 that Tulare Lake had dropped to less than three feet deep at its deepest part. Elevation that year was 183.5 feet.

An editorial in the Visalia Weekly Delta in 1889 captured the current thinking of the day:

Tulare Lake, from present appearances, will soon have to be erased from the maps of the state of California. As a geographical fact it exists today as a "lake" by courtesy only — for it is not a lake.

By the time that Kings County was formed in 1893, the lake had shrunk to about 220 square miles.

The New York Times reported in 1898 that the lake had dried up completely. This was the first time that had happened in historic times.
Floods and Droughts in the Tulare Lake Basin
General Flood and Drought Notes

The drying up of the lake was best described in an August 1898 issue of The San Francisco Call. That newspaper article was essentially an obituary for the lake.\(^ {314} \)

The lake had gone from full-pool to bone-dry in just 20 years (1878–98). There were apparently five causes for the lake drying out during this period:

1. Diversion of river water for irrigation; this began on the Kings and Kaweah Deltas in about the 1870s. The effectiveness of these diversions is dramatically illustrated in Table 7 on page 81. These diversions were the primary reason that Tulare Lake dried up.
2. The Fresno Slough Bypass (now known as the James Bypass) began operation in 1872. Prior to about 1872, all of the Kings River water flowed into Tulare Lake.
3. The general absence of winters with heavy precipitation. The 20-year period from 1878–98 had very few winters that were excessively rainy. There were three water years with moderately heavy runoff (1884, 1886, and 1890) as measured by inflows to the Tulare Lakebed. But those weren’t nearly enough to make up for the deficit in the other years. This was in contrast with the 1860s which had storm systems that put the greater portion of the San Joaquin and Sacramento Valleys under water several times.\(^ {315} \)
4. Only one major flood during this 20-year period: 1890.
5. At least six years of drought during this 20-year period: 1879, 1882, 1887, 1888, 1897 and 1898. These were part of four droughts that occurred during the late 1800s: 1873–79, 1882–83, 1887–89, and 1898–1900.

Gustav Eisen was a Swedish natural scientist and a member of the California Academy of Sciences. He provided a short article for the San Francisco Call issue on Tulare Lake giving his observations about the reasons for the lake drying up. Eisen didn’t have any data or claim subject matter expertise. However, he thought that the principal cause of the drying-up of the lake was the use of the waters of the tributary rivers for irrigation purposes. In addition, he recognized that the general absence of excessively rainy winters since 1874 had a good deal to do with it.\(^ {316} \)

As the lake shrank, agriculture moved in. By May 1895, there was 50,000 acres of grain growing in what had been the Tulare Lakebed. The lake was essentially not to be seen.

The Tulare Lakebed dried completely for the first time in August 1898 and remained dry through 1900. The 1901 flood brought the lake back to life, if only modestly. After that flood, the lake was about six feet deep at the deepest point (elevation 185.5 - 179 feet).

Tulare Lake was virtually dry when Hobart Whitley visited it in 1905. However, the high runoff of the 1906 flood brought the lake back. That flood left the lake about 12 feet deep at the deepest point, submerging 300 square miles. (This compares to 790 square miles at its maximum in 1862 and 1868.) As a relative measure of the volume of the runoff, that was the biggest increase in the lake’s depth since the 1890 flood. Combined runoff of the four rivers in the Tulare Lake Basin during water year 1906 was 7,195,240 acre-feet, the third-largest runoff of record (only 1983 and 1969 were larger). The total floodwater entering the lake in 1906 was about 1,530,000 acre-feet. This inflow exceeds that of any year since that time.\(^ {317} \)

Many levees were constructed in the Tulare Lakebed between 1903–1905 when lake levels were low. Unfortunately, those levees were light and poorly constructed. As a result, they failed when the high flows of 1906 entered the lake. The failure of those levees resulted in large financial losses, as almost 175,000 acres of wheat and barley had been planted that year. Most of that land was flooded before the crops could be harvested.

The lake continued to rise with the floods of 1907 and 1909, and then gradually receded for the next seven years.

By September 1914, the lake had dropped to an elevation of 180.0 feet, less than one foot deep. Avian botulism became a problem in the lake that year. Corcoran was incorporated in 1914, the year that the 1912–13 drought ended.

In 1907, a massive levee had been built around four sides of Tulare Lake, attempting to constrain it to a fraction of its full natural size. Ripley’s Believe It or Not featured the “Square Lake” in its syndicated cartoon. The lake was now harnessed, the lakebed declared safe for growing orchards. However, the 1916 flood brought the lake back to life and put an end to those hopes, at least temporarily.
At the time of EuroAmerican settlement, most of the water of the Kings River used to flow into Tulare Lake via what is now called the South Fork system, along the south side of its delta.

From 1861–1884, various floods and man-made ditches began making a channel for the Kings River along the north side of its delta. The key constriction was the Zalda Canal, a ditch constructed in 1872 and enlarged by two subsequent floods. But the majority of floodwaters still flowed along the south side of the delta and into Tulare Lake.

The 1916 flood significantly opened the Zalda Canal, and thereafter it became the main channel for the Kings River. This reach is now known as the North Fork of the Kings or the Kings River North Channel. In the 1916 flood, this channel is said to have discharged 60% of the Kings runoff into the San Joaquin River and thence to San Francisco Bay. These changes in the flow of the Kings River soon left the farmers in the Tulare Lakebed without sufficient water to irrigate the reclaimed grain land, forcing them to sink deep wells for their irrigation water.\(^{318}\)

The lakebed again went completely dry on April 30, 1919.

Floodwaters from the Kings and Kaweah arrived in the lakebed in May 1922. A total of 23,680 acres of lakebed cropland was inundated that year. A little more water (from both the Kings and the Kaweah) was added from heavy rains during the winter of 1923–24, but the lakebed was completely dry again early in 1924. Thanks in part to the extended drought of 1918–34, it would be 13 years before the lake would come back to life.

The lake did get some inflows from both the Kings (1923, 1927, 1932, and 1935) and the Kaweah (1923 and 1932) during the drought years. However, most of the quantities during those years were small, and S.T. Harding believed that much of that water was quickly absorbed by the soil or used directly for irrigation of crops growing in the lakebed. Tulare Lake would not reappear as a large lake until 1937.

The 1937 flood was a major flood, bringing an end to the drought years. The Kern River sent floodwaters into Tulare Lake for the first time since 1916. Tulare Lake reappeared on February 7, 1937 for the first time since 1924. The lake rose to an elevation of 191.9 feet and would stay at roughly that elevation for nine years. American white pelicans, waterfowl, and shorebirds reappeared almost instantly and in incredible numbers.

Ward B. Minturn was a prominent Fresno businessman and one of the most prominent field ornithologists in the San Joaquin Valley. He made at least 43 visits to Tulare Lake between 1937–54. Thanks to Rob Hansen's research, we have a wonderful collection of Minturn's field notes. An entry from his field notes of October 16, 1937 gives an idea of how quickly the bird life of Tulare Lake responded when the water returned:

> After being dry for several years, Tulare Lake is back. Main body is confined by levees in an area 6 miles by 8 miles. Quite a sea! Today I saw there one of the greatest bird sights of my experience. Great masses of white pelicans so thick on the levees that those in the center could not spread their wings to rise until those on the edges had taken flight! Pelicans flying, pelicans swimming, pelicans feeding so thick in nearby fields as to look at a distance like great snowbanks. Truly a sight I shall never forget. How many? I wish I knew. Possibly 40,000 to 50,000. I did not know there were so many left in the U.S.A.

(To some extent, this resilience still exists. Just add water and almost instantly an enormous number of waterfowl will appear. Prime examples are the floods of 1982–83 and the flood of 1997.)

In February and March, 1938, heavy storms flooded the San Joaquin Valley. When the elevation of Tulare Lake reached 192 feet, one of the main levees in the lakebed broke and the lake spilled over 49 square miles of land. The lake continued rising, eventually cresting at 195 feet. By June, 135,600 acres of the lakebed was underwater. That was the maximum acreage flooded since the 1906 flood. Tulare Lake has not been this big since.\(^{319}\)

While the high lake levels of 1938 were a disaster for the lakebed farmers, others saw opportunity. Near the height of the flood, Frank Latta and three boys took a 15-foot homemade motor boat from Bakersfield to San Francisco.\(^{320}, 321\) This was the third of six documented trips between Tulare Lake and San Francisco Bay to occur in historic times. (The other five trips were in 1852, 1868, 1966, 1969, and 1983.) It seems like all the trips after 1868 must have encountered impediments of one type or another; water last flowed out of Tulare Lake in 1878.
When World War II started, the Navy needed a place in the southern San Joaquin Valley where seaplanes could land in an emergency. The realization of this need occurred after 11 seaplanes from Hawaii arrived over Alameda Naval Air Station (NAS) from Hawaii only to find the entire region fogged in. Diverted south to Los Angeles, the pilots, low on fuel, were surprised to find themselves flying over a body of water that they knew nothing about — Tulare Lake. They landed without incident.

In January 1942, the Navy leased 3,000 acres of the Tulare Lakebed for an emergency seaplane landing base (Tulare Lake Outlying Field) and other purposes. In addition to a radio building and lighted buoys for navigation, the facility consisted of about two tents and six sailors. It was located about 10 miles south of Stratford. The only activity after that first emergency landing was a monthly seaplane visit from the commanding officer stationed at Alameda and a few low-level practice torpedo runs. In at least a conceptual sense, Tulare Lake Outlying Field might be thought of as the precursor of Naval Air Station Lemoore. Regrettably, NAS Lemoore does not have facilities for seaplane landings.

In 1943, enough runoff made it to the valley to raise the level of Tulare Lake to near the top of the lakebed levees. Wave action caused levee breaks and the flooding of 28,000 acres. Those levee breaks increased the size of Tulare Lake from 46,000 acres to 74,000 acres. By summer, 100,000 acres would be flooded.

1945 was a big flood year in the Tulare Lake Basin. It was also the first year of use of the new works, built by the USACE, to keep the Kings River out of the Tulare Lakebed. They did not work quite as designed. A break in that bypass occurred on February 3, 1945, about 20 miles south of Hanford at the height of the flood. Some ranchers were driven from their homes on the east side of the bypass and considerable grain was flooded on the west side.322

The J.G. Boswell Co. bought the Cousins Ranch in 1946. At that time, the ranch was still under water as a result of the 1938 flood.323

Thanks in large part to the extended drought of 1947–50, Tulare Lake again went completely dry on July 17, 1946.

The lake reappeared briefly four years later:
- November 19, 1950 – March 10, 1951 (maximum elevation 184.8 feet)

The winter of 1951–52 brought near-record snows to the Southern Sierra. The addition of rain to this snowpack caused Tulare Lake to reappear on January 19, 1952. That was the biggest episode of lake flooding (both by height and duration) between the early 1940s and 1969. See the section of this document that describes the 1952 flood for the considerable efforts that were made to store the floodwaters of the Kings and Kern Rivers before they reached the Tulare Lakebed. Despite these efforts, the 1952 flood still raised Tulare Lake by 15.5 feet to a maximum elevation of 194.6 feet, flooding 72,700 acres. The lake has never been this high since, although the 1969 flood would come close.

Tulare Lake again went dry late in 1953. It would reappear briefly on three occasions over the next 16 years
- December 23, 1955 – April 21, 1956 (maximum elevation 187.4 feet)
- March 31, 1958 – August 15, 1958 (maximum elevation 187.9 feet)
- December 6, 1966 – August 9 1967 (maximum elevation 183.1 feet)

Bill Cooper recalled that somebody made the trip to San Francisco in a motorboat in 1966.324 This was the fourth of six documented trips between Tulare Lake and San Francisco Bay to occur in historic times.

Tulare Lake next reappeared on January 20, 1969. By the end of March, 125 square miles (80,000 acres of farmland) had been inundated. The total estimated lakebed inflow in 1969 was about 1.155 million acre-feet. This is the second biggest flood since the federal reservoirs were completed (both by volume and by area flooded); only the 1983 flood was bigger.

In 1969, 960,000 acre-feet were impounded in the Tulare Lakebed, inundating 88,700 acres (139 square miles), significantly more than was flooded in 1952. The J.G. Boswell Co. had more land flooded in the Tulare Lakebed than any other landowner (almost 50,000 of the total 88,700 acres). Although huge, the 139 square miles inundated in 1969 was just 18% of the 790 square miles that Tulare Lake used to cover when it was at full pool.
All of the flow from the Kings was diverted into the San Joaquin River in 1969, at least during the January and February floods. The Kaweah and Tule Rivers both contributed significant flows to the lake. Even ephemeral Deer Creek (which fed the historic Ton Taché Lakebed near present-day Alpaugh) was flowing into the lake just south of Sand Ridge that spring. However, the majority of the inflows to the Tulare Lakebed came from the Kern River.

Historically, the Kern would fill Buena Vista Lake before spilling over into Tulare Lake. However, in 1969, a giant dike protected two-thirds of Buena Vista Lake from being filled. When the other third of the lake filled, the Kern then spilled or passed through to Tulare Lake. The decision to keep the remainder of Buena Vista Lake dry was not appreciated by those downstream in the Tulare Lakebed.

In 1952, the Tulare Lake Basin Water Storage District had stored Kern River floodwaters in Buena Vista Lake. That was presumably possible because the J.G. Boswell Co., which had a long-term agricultural lease for the Buena Vista Lakebed, was willing to have its land flooded. In any case, no such water storage was allowed in 1969. That created hard feelings among some who were being impacted by the flooding that was occurring in the Tulare Lakebed in 1969. Emotions ran high as did financial losses.

The decision to pass through the Kern River floodwaters was challenged in court. But in the meantime, the floodwaters continued to come. As a result, about 222,000 acre-feet of Kern River water flowed into the Tulare Lakebed. (The majority of the Buena Vista Lakebed remained dry, safe behind its giant levee.) For a more complete description of that event, see the section of this document that describes the 1969 flood.

On May 8, 1969, the USACE received approval for a half-million-dollar project to throw up levees to connect the separated segments of Sand Ridge, south of the current Tulare Lake, creating a gigantic holding pond capable of containing 100,000 acre-feet of Kern River floodwater. Along with the South Wilbur Flood Area (located north of Sand Ridge), that is the area known today by Tulare Lake water storage districts and irrigators as the South Flood Area.

On June 24, 1969, Tulare Lake reached its highest modern level at 192.5 feet. (The last time that the lake had been this high was in 1952.) An emergency levee was hurriedly built just west of the Corcoran Airport. Tulare Lake was deep enough to cause significant erosion to that levee.

When the lake came back, it brought an abundance of crayfish. Mo Basham’s family lived in Corcoran at the time, and her father recognized that this was a natural resource not to be wasted. For the next two years, he organized the neighborhood kids to go on crawdad hunts along the edge of the lake. The kids would bring back hundreds at a time, and their families would eat them just like lobster. To read Mo’s description of these hunts, see the section of this document that describes the 1969 flood.

In 1969, two fathers and their sons took advantage of the high water to boat from Bakersfield through Buena Vista Lake and Tulare Lake to San Francisco Bay. For a few more details on this adventure, see the section of this document that describes the 1969 flood. This was the fifth of six documented trips between Tulare Lake and San Francisco Bay to occur in historic times.

Minor flooding occurred in the Tulare Lakebed in 1970, 1971, 1973, 1978, 1980 and 1982. This was followed by the major flood of 1983. In that year, the lake rose to 191.44 feet and flooded a slightly larger area than in 1969. Bill Tweed recalled that Tulare Lake was so big in the summer of 1983 that you could see it from the High Sierra, shining through the valley haze. To see it was like seeing a ghost, a relic of another time.

In 1983, Bill Cooper and John A. Sweetser, Sr. kayaked from the banks of the Kern River just outside of downtown Bakersfield all the way to Richmond Marina on the shores of San Francisco Bay. This was the sixth documented trip between Tulare Lake and San Francisco Bay to occur in historic times. (The other five trips were in 1852, 1868, 1938, 1966, and 1969.)

In a 2014 interview, Bill recalled how he and John did it. Bill was not an experienced kayaker; he had never been in a kayak before. They first scouted to Tulare Lake which was the hardest section, and decided they could make it. Then they threw sleeping bags and a couple tents in their kayaks and put in.

They didn’t expect their trip to be a big deal, but it made the national news. Radio and TV news reported on their progress. They were followed by an airplane for part of their route. Dave Graber, retired NPS regional chief scientist, recalled that their trip was written up in The Fresno Bee. Their goal was just to go across Tulare Lake.
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However, it caught the attention of the news media, so they decided to try to make it all the way to San Francisco Bay.

Bill and John thought they could make it to the south end of Tulare Lake in one day, but they barely made it to Buttonwillow in the first day. Ken Wedel landed his plane and took them up to scout their route. He then returned and re-provisioned them by plane that first day, giving them an air drop of water jugs. After that, they resupplied at farm houses and stores along the way at places like Firebaugh. It took them a full day to cross Tulare Lake, staying close to the western levee. It took them 12 solid days of work, sunrise to sunset, to make it to the Richmond Marina.

Bill said that he and John are apparently the last two people to make it through to San Francisco Bay. He heard that some guys from Reedley tried to get through in canoes in some year after 1983, but didn’t make it; the wind gave them a hard time in the Fresno Slough.

After two years of flooding (1982 and 1983), cotton growers decided to drain their lands, and also save the lake towns of Corcoran, Stratford, and Alpaugh in the process. They proposed to pump the excess water over the top of the Tulare Lake sill. The water would then flow into the North Fork of the Kings River, and from there to the San Joaquin River and the Sacramento–San Joaquin Delta.

The Tulare Lake Irrigation District applied for a permit to pump the excess water over the top of the Tulare Lake sill. It appears that there was considerable opposition to granting this permit. Under an emergency proclamation issued by the USACE during the spring of 1983, reclamation districts and land companies remade the channel along some 29 miles of the Kings River to dewater the lake and drain the water north into the Sacramento–San Joaquin Delta region.

A series of pumps were installed with a total lift of 43 feet. The project was designed to remove approximately 2,000 acre-feet of water per day from the lakebed. Pumping began on October 7, 1983 and continued intermittently until the program was terminated on January 19, 1984. Only about 90,000 acre-feet was pumped northward over the Tulare Lake sill under that program. Pumping was stopped earlier than scheduled due to concern that white bass might be transferred from the Tulare Lakebed into the San Joaquin River. The lakebed would not be fully drained until water year 1985.329

Exotic white bass had been illegally introduced into Lake Kaweah by fisherman during the 1970s. Large numbers escaped into waters downstream of Lake Kaweah during the record 1982–83 flood runoff. A large population became established in the flooded Tulare Lakebed and connecting waterways. Lake Kaweah and downstream waters of the Tulare Lakebed were treated with rotenone in the fall of 1987. This was one of the largest such chemical treatments ever carried out in the U.S., and certainly California's largest. The cost of the project was about $9.7 million.330 Apparently, a complete kill of white bass was achieved; they were completely eliminated from Lake Kaweah and from throughout the rest of the Tulare Lake Basin.331

Minor flooding occurred in the Tulare Lakebed in 1985 and 1986. Tulare Lake had appeared in 1937–46 and 1951–53. However, since the damming of the Kings (1954) and the Kaweah (1962), only very wet years have seen water return to the lakebed in significant amounts. The lake occasionally reappears during unusually wet years, as it did in 1969, 1983 and 1998.

In order to minimize flooding in the Tulare Lakebed, a combined total of over 472,000 acre-feet of floodwaters was pumped into the Friant-Kern Canal during the years 1978, 1980, 1982, 1983, 1986, 1995, 1997, 1998 and 2006. Once those floodwaters reached the Bakersfield area, they were emptied into the Kern River and then routed via the Kern River Intertie (a structure that was completed in 1977) into the southbound California Aqueduct. That water then made its way to the Los Angeles area rather than being available for use or groundwater recharge within the Tulare Lake Basin.

Additional amounts of water from the Kings and Kern Rivers have been diverted out of the basin in many years. See the section of this document that describes the Groundwater Overdraft for a discussion of those diversions.

Parts of the Tulare Lakebed still become periodically inundated during major flood events (see Figure 16). However, only remnants of the historic wetland area in the lakebed remain, confined primarily to privately owned waterfowl hunting clubs, former agricultural ground that has been enrolled in wetland reserve programs, and the Pixley and Kern National Wildlife Refuges (although neither one of those refuges is in the actual lakebed).
Figure 15 illustrates how the elevation of Tulare Lake has varied for the 120 years for which we have data: 1850–1969. Data since 1969 are only available from the J.G. Boswell Co., and that has proved impossible to obtain.

Source: Data from USBR which obtained it from USACE which compiled it from a variety of sources.
Figure 16. Portion of the Tulare Lakebed flooded each year 1954–99.
Source: Tulare Lake Basin Water Storage District
Federal Dams and Reservoirs

Friant Dam
This 319-foot-high concrete dam forms Millerton Lake on the San Joaquin River north of Fresno. The watershed drainage area is 1,675 square miles. The dam was built by USBR and completed in 1942; it is part of the Central Valley Project. Storage capacity is 520,528 acre-feet. The San Joaquin River produces a long-term average annual flow (measured at Friant Dam) of about 1.8 million acre-feet per year. Although Friant Dam is not in the Tulare Lake Basin, much of the water stored in that lake is exported to our basin.

Pine Flat Dam
This 429-foot-high concrete dam is on the Kings River at river mile 95 a few miles above Piedra, about 28 miles northeast of Fresno. Regulation of discharges from Pine Flat Reservoir began December 4, 1951, although the dam was not completed until 1954. The watershed drainage area of the Kings as measured from above the dam is 1,545 square miles. A gaging station for the Kings is incorporated into the bridge immediately downstream of the dam. Flow at that gage is above the confluence of Mill and Hughes Creeks. Unimpaired flow (full natural flow) is calculated for that gage and is apparently reported as CDEC station KGF Kings R-Pine Flat Dam.

The drainage area as measured from above the USGS gage at Piedra (USGS 11222000 Kings R A Piedra CA) is 1,693 square miles. This gage was located at Winton Park in Piedra and was operated from October 1895 – September 1959. Discharge data after 1951 was affected by operation of Pine Flat Dam.

The Kings River Water Association (KRWA) and USACE have continued to calculate the unimpaired flow (full natural flow) of the Kings at the location of the former gage. This calculated number is intended to make the record from 1954 to present match the largely unobstructed observations recorded from 1895-1954. This blended gage data is apparently reported as CDEC station KGP Kings Pre-Project Piedra. Data from this gage site best represents full runoff for the Kings River Basin because it is below the confluence with Mill and Hughes Creeks.

Pine Flat Dam was built with a gross-pool capacity of 1,000,000 acre-feet. That amount of gross storage can hold 60% of the 121-year average runoff (1894–2014) for the Kings River. That gross storage percentage is important primarily from the standpoint of irrigation, not flood control. Pine Flat Reservoir has a gross-pool capacity of 1,000,000 acre-feet, of which 475,000 acre-feet is reserved for a flood-control pool. That leaves 525,000 acre-feet available for a conservation pool in the winter.

Pine Flat Dam has a Dam Safety Action Classification (DSAC) rating of 4: Priority (Marginally Safe). The dam safety characteristics associated with that rating are: Inadequate with low risk. The dam is operated to reduce floodflows to a downstream objective release of 4,750 cfs below Crescent Weir. Dam operators also try to minimize floodflows into the Tulare Lakebed.

Although the maximum objective flow of Pine Flat Dam is 4,750 cfs below Crescent Weir, that location is many miles downstream of the dam, several miles west of Highway 99 (see Figure 17). There are many creeks that enter the Kings River prior to that point. There are also many diversion canals upstream of Crescent Weir to divert water out of the river. Wayne Johnson is chief of the Water Management Section in the Sacramento District of the USACE. He recalled that in the past, releases of 10,000 cfs from the dam have occasionally resulted in flows less than 4,750 cfs at Crescent Weir.

Pine Flat Reservoir is more formally (but perhaps less commonly) known as Boone Lake. At full pool, the reservoir covers about 5,956 acres and extends 20 miles back from the dam with 67 miles of shoreline. Its gross pool elevation is 951.5 feet.

Under the Kings River Fisheries Management Program, Kings River Water Association member units have agreed to maintain a minimum reservoir storage of not less than 100,000 acre-feet. The purpose is to maintain a pool of cool water for use in the reservoir and downstream fisheries under many, although possibly not all, critically dry conditions.

The Kings River divides in its lower reaches, about 60 miles below the dam (see Figure 17). The distributary point is located north of Lemoore, about 1½ miles above Highway 41. (A distributary is a branch of a river that flows away from the mainstem. They are common on deltas. The Kings, Kaweah, and Kern Rivers all have...
The southerly channel (generally known as the South Fork system) flows southeasterly and southerly into the Tulare Lakebed. The South Fork system is known by a variety of names in different stretches. It begins as the Clark’s Fork. Farther along, it empties into the South Fork of the Kings which turns south toward Stratford and terminates in the Tulare Lakebed. However, since that lakebed is normally dry, the river has been extended another 10 miles in the South Fork Canal, which intersects the Tule River Canal at a point 12 miles west of Corcoran. It’s an ignominious end for a fine river. USACE lists the capacity of the South Fork Kings as 3,200 cfs. The northerly channel (generally known as the North Fork system) is also known by a variety of names in different stretches. It begins as the North Fork of the Kings. (Not to be confused with the river of the same name that originates in the High Sierra.) At the Crescent Weir, the North Fork empties into the Fresno Slough which later empties into the 12-mile-long James Bypass channel. The James Bypass merges with the San Joaquin River at Mendota Pool near the city of Mendota. The Kings River Delta begins about where Kingsburg is located. At the time of EuroAmerican settlement, most of the water of the Kings River used to flow into Tulare Lake via what is now called the South Fork system, along the south side of the Kings River Delta. Only limited amounts of Kings River rain and snowmelt floodwaters typically reached across the north side of the river’s extensive delta to the Fresno Slough. Since then, a system has developed that routes a portion of the Kings River floodwaters along the north side of the delta. The 1861–62 flood on the Kings River began the formation of Cole Slough, cutting the head of that slough. The slough was named for William T. Cole, who dug the irrigation ditch that the Kings River enlarged to form the slough. The 1867–68 flood completed the formation of Cole Slough. From Cole Slough, the floodwaters flowed on through Murphy Slough, possibly also formed in the 1861–62 and/or 1867–68 floods. These two floods created the conditions necessary to start moving a significant portion of the flow of the Kings River from the south side of its delta to the north. In 1872, a ditch was constructed on the north side of the Kings River Delta that became known as the Zalda Canal. Floods in 1878 and 1884 substantially enlarged that canal for about four miles, enabling connections to be made with other channels. One source said that it was the 1879 flood (not the 1878 flood) that enlarged the Zalda Canal, but that was apparently an error. After those floods, the Kings River began sending a portion of its floodwaters along the north side of its delta via this channel. Starting in 1909, reclamation districts enlarged sections of the channel and constructed levees along it. The 1916 flood significantly opened the Zalda Canal, and thereafter it became the main channel for the Kings River. This reach is now known as the North Fork of the Kings or the Kings River North Channel. In the 1916 flood, this channel is said to have discharged 60% of the Kings runoff into the San Joaquin River and thence to San Francisco Bay. These changes in the flow of the Kings River soon left the farmers in the Tulare Lakebed without sufficient water to irrigate the reclaimed grain land, forcing them to sink deep wells for their irrigation water. This reach is now known as the North Fork of the Kings or the Kings River North Channel. The USACE increased the capacity of this channel from 3,500 cfs to 5,500 cfs just prior to the onset of the January 1969 flood. The North Fork Channel now has a rated capacity of 6,300 cfs. The James Bypass is a manmade channel. The Fresno Slough Bypass (now known as the James Bypass) began operation in 1872. The bypass reached its present configuration during the years 1912–14. It was built to bypass the meanders of a portion of the Fresno Slough and is sometimes still referred to as the Fresno Slough Bypass. The USACE lists the capacity of the Fresno Slough and James Bypass as 4,750 cfs. However, flows up to 6,000 cfs have passed through this reach. A USGS stream gage (USGS 11253500 James Bypass (Fresno Slough) NR San Joaquin CA) (a water-stage recorder) was maintained on this stretch from October 1, 1947 – September 30, 2009. The small original Fresno Slough channel (the part that has been bypassed) meanders for nearly 15 miles, from southeast of San Joaquin to north of Tranquility. It is no longer a part of Kings River operations. A large portion of the old slough is now managed as a wetlands area by the California Department of Fish and Game. (This agency became the California Department of Fish and Wildlife on January 2013.)
Total channel capacity in the lower reaches of the Kings is 7,950 cfs (3,200 cfs for the South Fork system and 4,750 cfs for the North Fork system). For comparison, the flood-of-record on the Kings River is 112,000 cfs, set on January 3, 1997.

Under typical flood operations, the first 4,750 cfs of flood release water from Pine Flat is directed through the North Fork / Fresno Slough / James Bypass channel to the San Joaquin River and, ultimately, San Francisco Bay. When the capacity of that channel has been reached, floodwater is sent into the Kings River South system up to its published channel capacity of 3,200 cfs. Flow in excess of 7,950 cfs is supposed to be divided equally between the two channels. In practice, the stage of the San Joaquin River during large floods may affect how water is divided between the two channels.

Figure 17. Map of Lower Kings River features.
Source: Kings River Handbook
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Terminus Dam
This 250-foot-high\textsuperscript{347} earthen dam forms Lake Kaweah on the Kaweah River. The watershed drainage area of the Kaweah as measured from above the dam is about 561 square miles.\textsuperscript{348, 349, 350} The Dry Creek watershed drainage area immediately below Terminus Dam is about 80 square miles.\textsuperscript{351} The watershed of the Kaweah as measured from above McKay’s Point is about 647 square miles.\textsuperscript{352, 353} The entire watershed drainage area of the Kaweah as measured from above the Tulare Lakebed is about 952 square miles.\textsuperscript{354}

One USACE report said that the dam was closed on November 1, 1961.\textsuperscript{355} However, most sources say that the dam was completed and storage began in February 1962.\textsuperscript{356}

Terminus Dam was built with a gross-pool capacity of 149,600 acre-feet (often rounded to 150,000). We haven’t been able to find a record of how much of this was reserved for the flood-control pool, but it may have been 148,600 acre-feet.

The dam is operated to reduce floodflows so that the Kaweah doesn’t exceed its rated channel capacity (5,500 cfs) three miles downstream at McKay’s Point. Since Dry Creek enters the Kaweah below Terminus Dam, releases have to take the flow of that creek into consideration. The floodflow as measured at McKay’s Point is considered the maximum objective flow for the dam. Dam operators also try to minimize floodflows into the Tulare Lakebed.\textsuperscript{357}

When originally constructed, the lake’s level at full pool (its gross pool elevation) was 694 feet. That provided a storage capacity of 150,000 acre-feet, which was estimated to be sufficient to provide a 60-year level of flood protection downstream.\textsuperscript{358, 359}

A reservoir’s gross-pool capacity is defined by its area-capacity curve. That curve is a graph showing the relation between the surface area of the water in the reservoir and the corresponding volume.

A reservoir’s rated gross-pool capacity changes through time. The statement of a reservoir’s gross-pool capacity is based on three factors:

- The volume of the reservoir as of the date of the most recent area-capacity curve.
- The assumed long-term annual rate at which sediment is accumulating in the reservoir.
- The year that the reservoir’s capacity is being projected to.

A reservoir’s area-capacity curve may be revised periodically for various reasons. This revision can be triggered by new information about the rate of sediment accumulation or by a change in the elevation of the pool level.

The Kaweah River, upstream of Lake Kaweah, has an unusually high potential to erode and carry sediment. That is due to its gradient and the type of soil that it is eroding.

The Kaweah is the only river in the U.S. that drops 10,000 feet in less than 100 miles. Measured over its entire length, the Kaweah is the steepest river in the U.S. It drops 10,826 feet in 76.5 miles, a gradient of 142 feet per mile. Individual reaches of rivers can be much steeper. The Marble Fork of the Kaweah has a gradient of 559 feet per mile, dropping 8,549 feet in just 15.3 miles.

From a geomorphic standpoint, a key metric is the drop from the headwaters to the range front. By that metric, the Kaweah is also the steepest river in the U.S. When measured from its headwaters to the range front at Terminus Dam, the Kaweah drops 10,505 feet in 37.5 miles, a gradient of 280 feet per mile.\textsuperscript{360}

Lower elevation soils are more erodible than higher elevation soils. Compared to rivers such as the Merced, the various tributaries of the Kaweah have many more miles in which to erode these lower elevation soils. In addition, the geology of the lower elevation of the Kaweah River Basin consists of sedimentary or metamorphic rock that is far more erodible than the granite of basins such as that of the Merced. This combination of high-energy streams and erodible soils give the Kaweah the ability to carry a high sediment load relative to other Sierra streams.\textsuperscript{361}

The Kaweah’s sediment load used to be delivered to the Kaweah Delta. See the section of this document on Description and Identification of Deltas. But once Terminus Dam was closed in 1962, Lake Kaweah became a sediment trap. There has never been a sediment gage on the Kaweah, so the USACE had to estimate what the sediment load was when they were designing Terminus Dam. One source said that the initial design was based
on an estimated long-term sediment yield for the watershed upstream of Lake Kaweah of 150 acre-feet a year. We haven’t been able to find any documentation to clarify this.

Once the dam was in place, the sediment load could be measured through periodic surveys of the lakebed during the summer when the reservoir was drawn down. Detailed surveys were made in 1961, 1967, 1977, and 1988. Additional samples of bank and bed materials were analyzed and a field reconnaissance was done in 1988 and 1989.\textsuperscript{362, 363}

The reports on the four reservoir surveys (1961, 1967, 1977, and 1988) have apparently all been lost. There have not been any general surveys of the reservoir bed since 1988.

The 1967 and subsequent surveys showed that the quantity of deposition in the reservoir bed was quite high. USACE studies attributed that largely to the 1966 flood, a storm that had a recurrence interval in excess of 100 years.\textsuperscript{364}

In the lake’s first 17 years, it received an average of 474 acre-feet a year of sediment.\textsuperscript{365} As noted above, that was the largely the result of an unusually large flood event in 1966. However, as noted elsewhere, the Kaweah has been unusually quiet of late. As illustrated in Table 30, the Kaweah hasn’t seen any 20-year or larger floods in over 40 years.

The USACE revised Lake Kaweah’s area capacity curve significantly downward in 1978 as a result of the 1977 survey. They documented this in a 1978 report, but apparently all copies of that report have been lost. Fortunately, Allen Wilson of the Kaweah Delta Water Conservation District (KDWCD) has a clear recall of some of that report’s main findings. There are also allusions to that 1978 report (and to the 1977 survey) in various USACE reports. Apparently Lake Kaweah’s area capacity curve was not revised again until after the fuse gates were installed in 2004.

The new area capacity curve set the lake’s gross-pool capacity at 143,200 acre-feet (generally rounded to 143,000 acre-feet). Of this total, 142,000 acre-feet was reserved for the flood-control pool.\textsuperscript{366} That was sufficient to provide only a 46-year level of protection.\textsuperscript{367, 368}

The USACE studied the situation in 1989 and concluded that sediment yield for the watershed upstream of the reservoir should be approximately 100 acre-feet per year.\textsuperscript{369} All copies of that report have apparently been lost.

Lake Kaweah’s flood-control pool is small compared with the drainage area tributary to the lake. Because of this, the lake provides a relatively low level of protection from rain-floods (see Table 8). This is also illustrated by the 1997 flood. That had a recurrence interval of only 14 years for the Kaweah (see Table 30). Even so, Lake Kaweah filled and emptied twice during that flood.\textsuperscript{370}

Fuse gates were installed in the spillway in 2004. This raised the lake level at full pool 21 feet (from elevation 694 to 715 feet). When the Lake Kaweah Enlargement Project was originally designed in the 1990s, it was estimated that this would increase the lake’s storage capacity to an estimated 183,300 acre-feet (generally rounded to 183,000). As the project neared completion, this estimate was revised to 185,630 acre-feet (generally rounded to 185,600).\textsuperscript{371} There have been no new measurements of sedimentation rates since 1988.

As illustrated in Table 8, the level of flood protection for downstream communities has varied since the dam was built. When originally constructed in 1962, Lake Kaweah’s storage capacity was estimated to be sufficient to provide a 60-year level of flood protection downstream. That is, it could catch that level of rain-flood while keeping flows downstream of the dam within stated channel capacity / maximum objective flow. However, protection decreased significantly after 1966 due to an accumulation of sedimentation. When fuse gates increased the flood-control pool size of the reservoir in 2004, the level of protection increased to a 70-year level of flood protection.\textsuperscript{372, 373}

<table>
<thead>
<tr>
<th>Year</th>
<th>Elevation of lake at full pool</th>
<th>Gross-pool capacity (acre-feet)</th>
<th>Level of flood protection</th>
</tr>
</thead>
<tbody>
<tr>
<td>1962</td>
<td>694</td>
<td>150,000</td>
<td>60-year</td>
</tr>
<tr>
<td>1978</td>
<td>694</td>
<td>143,200</td>
<td>46-year</td>
</tr>
<tr>
<td>2004</td>
<td>715</td>
<td>185,600</td>
<td>70-year</td>
</tr>
</tbody>
</table>
Wayne Johnson is chief of the Water Management Section in the Sacramento District of the USACE. He said that in calculating the 70-year level of flood protection, the USACE estimated that the actual storage in the reservoir was probably less than 185,600 acre-feet due to sediment that has occurred within the lake area since the sediment survey that was completed in 1977. (Records from the sediment survey completed in 1988 have apparently been lost.)

If Lake Kaweah’s current gross storage capacity were 185,600 acre-feet, that could hold 44% of the 121-year average runoff (1894–2011) for the Kaweah River. That gross storage percentage is important primarily from the standpoint of irrigation, not flood control.

In 2008, the USACE conducted a one-day inspection of Terminus Dam, part of a series of safety reviews on its dams that looked at their performance histories, ages, construction and the geology around them. Although no history of problems came up at Terminus, questions were raised about whether seepage was an issue and whether any faults in the area could affect it.

Because of the potential risks to Visalia and other communities if the dam were to fail, the USACE chose in 2012 to be conservative and gave Terminus a Dam Safety Action Classification (DSAC) rating of 2: Urgent (Unsafe or Potentially Unsafe). The dam safety characteristics associated with that rating are: Failure initiation foreseen or very high risk.

This rating was put on the dam until a more comprehensive study could be done. However, no reduced pool restriction limits were placed on the reservoir. The safety study is now largely complete. Apparently that study found that the dam sits on a relatively quiet seismic zone that does not pose a high risk for earthquakes. Seepage in the ground and abutment on the sides of the dam appears to be higher than expected. However, that may be the result of when fuse gates increased the gross pool capacity of the reservoir in 2004.

The USACE is currently preparing the Baseline Risk Assessment report detailing the safety findings for the dam. When that report is released, the USACE will determine if the DSAC rating should be changed. In any case, no safety retrofit is likely for Terminus Dam because the concerns about seepage and the risk of an earthquake were never high.

At full pool, Lake Kaweah covers about 2,154 acres and extends 6 miles back from the dam. Its gross pool elevation is 715.0 feet.

Lake Kaweah has a gross-pool capacity of 185,600 acre-feet, of which 184,600 acre-feet is reserved for a flood-control pool.

Prior to the Lake Kaweah Enlargement Project, Lake Kaweah had to be kept practically dry each winter. The winter conservation pool was only 1,000 acre-feet (143,000-142,000 acre-feet). The rest of the reservoir was kept dry for the flood-control pool of 142,000 acre-feet.

The Lake Kaweah Enlargement Project Water Control Plan established a more flexible plan for winter operations. It established a conditional winter rain-flood storage pool of 12,000 acre-feet. Wayne Johnson said that this value may be reduced based on a rain-flood variable (aka rain parameter). This parameter is based on the precipitation that has occurred in the Kaweah River Basin above the dam. The wetter the basin is, the greater the flood pool requirement is, and so the lower the water conservation pool must be. In other words, if the basin is wet, the lake must be lower. If the basin is dry, the lake can be higher, up to 12,000 acre-feet. This conditional winter rain-flood storage pool has essentially no effect on the level of flood protection.

The flood-of-record on the Kaweah is 105,000 cfs, set on December 6, 1966. The dam with fuse gates is designed to withstand a flood of 300,000 cfs. The first fuse gate is designed to tip at 190,000 cfs.


The Kaweah River divides in its lower reaches. Major floods used to periodically result in the relocation of the distributary point where those channels divide. However, since the December 1867–68 flood, the channels have divided at McKay’s Point, about a mile northwest of present-day Lemon Cove and three miles below Terminus Dam. A variety of structures have been built at this location since 1870 in an attempt to control the Kaweah and...
split the flow between the two channels. For a more complete description of McKay’s Point and the diversion structures that have been built there, see the section of this document that describes the 1867–68 flood.

The southerly channel, known as the Lower Kaweah River, flows to the southeast. That channel ends east of Visalia, just north of the Ivanhoe turnoff (Road 156/158) on Highway 198. From there, it feeds distributaries on the south side of the Kaweah Delta, principally Mill Creek, Packwood Creek, and Cameron Creek.

The northerly channel, known as the St. Johns River, flows along the north side of Visalia. It merges with Cross Creek northwest of the city, a few miles upstream of Highway 99. The Shipp Cut was made in 1854, a small drain ditch from the Kaweah River Swamp near Rocky Ford (north of present-day Kaweah Oaks Preserve) west to Canoe Creek. The 1861–62 flood cut a new channel along the northern border of the swamp. Shipp Cut and a section of Canoe Creek were enlarged by the floodwaters and became a part of this new channel, and finally a connection was established with the Cross Creek channel, creating what we now know as the St. Johns River. The 1867–68 flood further enlarged the St. Johns and eroded a new head for that river about a mile farther upstream, farther into the swamp.

The USACE considers the total channel capacity in the lower reaches of the Kaweah / St. Johns (below McKay’s Point) to be 5,500 cfs. Above this, flooding occurs in Visalia and other delta towns. This compares to the flood-of-record on the Kaweah River of 105,000 cfs that occurred during the December 1966 flood.

The weir at McKay’s Point is used to send the majority of the Kaweah River floodwaters around the north side of Visalia through the St. Johns River. A levee on the south side of that channel protects the city, but has failed in a number of floods. See the section of this document that describes the St. Johns Levee — Condition in Recent Years, for a summary of the challenges that the city and county face in maintaining that levee and keeping it from failing.

Figure 7 on page 33 illustrates the standard project flood for the lower Kaweah and northwest Tulare County. The USACE formerly defined the “standard project flood” as the flood that may be expected from the most severe combination of meteorological and hydrological conditions that are considered reasonably characteristic of the geographical area in which the drainage basin is located, excluding extremely rare combinations. This is a rare event, but one that could reasonably be expected to occur. A standard project flood, with Terminus Dam in operation, would be of somewhat greater magnitude than the December 1955 flood was without Terminus Dam. For a more complete description of this flood, see the section of this document that describes Preparing for the Next Big Flood.

Success Dam

This 142-foot-high earthen dam is on the Tule River. The watershed drainage area of the Tule as measured from above the dam is 393 square miles. The dam was completed and storage began in November 1961. It is located off Highway 190 between Porterville and Springville.

Success Dam was built with a gross-pool capacity (aka usable storage capacity) of 85,400 acre-feet. USACE has since revised the area-capacity curve for the reservoir. The reservoir is now rated as having a gross-pool capacity of 82,291 acre-feet (generally rounded to 82,300). That amount of gross storage can hold 60% of the 121-year average runoff (1894–2014) for the Tule River. That gross storage percentage is important primarily from the standpoint of irrigation, not flood control.

Success Reservoir has a gross-pool capacity of 82,300 acre-feet, of which 75,760 acre-feet is reserved for a flood-control pool. That leaves 6,540 acre-feet available for a conservation pool in the winter.

The dam is operated to reduce floodflows in order to achieve a maximum objective flow immediately downstream of the dam to 3,200 cfs. Dam operators also try to minimize floodflows into the Tulare Lakebed.

Success Dam is capable of providing a downstream level of protection of 100-200 years. That is, it can catch that level of rain-flood while keeping flows downstream of the dam within stated channel capacity / maximum objective flow.

The above level of protection estimate has not been reviewed and approved by appropriate USACE personnel and represents a draft value only. It is based on the existing dam design. The estimate reflects only hydrology, not seismic or geotechnical risk. It assumes that the full flood-control pool is available at the beginning of the flood. There are no recent data available on sedimentation rate or on the amount of sedimentation that has occurred in the reservoir.
At full pool, Lake Success covers 2,477 acres and extends 35 miles back from the dam. Its gross pool elevation is 652.5 feet.\(^{287}\)

In the science fiction novel “Lucifer’s Hammer,” fragments of a comet hit the lake and destroy the dam. Although less romantic, the earthen dam has been found to have seepage problems and to be at risk of failure in the event of an earthquake. After studying alternative solutions, the USACE chose a preferred solution: constructing a 350-foot extension downstream and the replacement of the dam’s core. However, the USACE did further study and announced in April 2012 that the risk of catastrophic failure was not as great as formerly thought.

Success Dam has a Dam Safety Action Classification (DSAC) rating of 2: Urgent (Unsafe or Potentially Unsafe). The dam safety characteristics associated with that rating are: Failure initiation foreseen or very high risk.\(^{388}\)

The full pool elevation of Lake Success is 652.5 feet, equivalent to 82,291 acre-feet. In September 2004, the USACE set interim restrictions on the maximum amount of water that the lake could hold, lowering the maximum amount of water that the lake could hold to 40,900 acre-feet. In late 2006, the USACE further lowered the maximum water level to 620 feet elevation (equivalent to 29,183 acre-feet). This pool restriction was felt necessary because at the time:

1. The USACE thought that there was a large risk due to earthquakes.
2. They thought that there was a bigger seepage risk with water running through the earthen dam.

In 2009, the pool restriction was increased to 41,000 acre-feet. In April, 2012, the USACE came to the conclusion that there was not as much seismic risk as they had thought. When the dam was evaluated a few years earlier, concerns were raised that sandy soil beneath the dam might be so prevalent that if a sizable earthquake occurred, the soil would settle, causing the dam above it to settle lower than its current height. That could cause large amounts of water to spill over the dam. Depending on the volume and speed at which the water spilled out, it could further eat away at the dam and make for a worse flood.

As a result of the reevaluation, the official pool restriction was increased to 640 feet elevation (equivalent to 56,084 acre-feet), but encroachment was allowed to 645 feet elevation (65,473 acre-feet). Sensors were monitoring seepage rates in and under the dam as the amount of water in the lake increased, and so far no problems have been found. The USACE still has to look at the risks. The previously identified risks have not been eliminated; they have just come down.\(^{389}\) The USACE was reevaluating the water level (i.e., the pool restriction and allowable encroachment level) while they evaluate the preferred structural solution for the dam.

This pool restriction largely impacted recreation and irrigation users. It did not affect the dam’s ability to control floods. The flood-control pool remained unrestricted and the dam’s ability to provide a downstream level of protection remained unaffected.

Studies apparently showed that the seismic risk was not as significant as originally thought because the fault under Success isn’t active. There is apparently still concern about the spillway not being able to stop water overtopping it. The USACE removed the 645 feet elevation (65,473 acre-feet) pool restriction on April 11, 2014.

The USACE is currently preparing the Baseline Risk Assessment report detailing the safety findings for the dam. When that report is released, the USACE will determine if the DSAC rating should be changed and if they are going to go forward with any type of construction to expand the dam’s capacity.\(^{390}\) This document assumes that Lake Success is now rated as having a gross-pool capacity of 82,300 acre-feet without pool restriction.
Isabella Dam
This 185-foot-high earthen dam is on the Upper Kern River, just below where the North Fork and South Fork of the Kern merge. That is about 34 miles northeast of the city of Bakersfield. The watershed drainage area of the Kern as measured from above the dam is 2,074 square miles. The watershed drainage area as measured from above the First Point of Measurement gage (located just upstream of the city limits of Bakersfield) is 2,407 square miles. The entire watershed drainage area of the Kern as measured from above the Buena Vista Lakebed is 3,612 square miles.

The dam began operation on April 15, 1954. Storage in the reservoir prior to that date was negligible. However, the USACE apparently reports Lake Success as officially beginning operation in 1953.

Isabella Dam was built with a gross-pool capacity of 568,100 acre-feet. That amount of gross storage can hold 79% of the 121-year average runoff (1894–2014) for the Kern River. That gross storage percentage is important primarily from the standpoint of irrigation, not flood control.

Isabella Reservoir has a gross-pool capacity of 568,100 acre-feet, of which 169,760 acre-feet is reserved for a flood-control pool. That leaves 398,340 acre-feet available for a conservation pool in the winter.

The dam is operated to reduce floodflows in order to achieve a maximum objective flow at the First Point of Measurement gage (located just upstream of the city limits of Bakersfield) to 4,600 cfs. Dam operators also try to minimize floodflows into the Tulare Lakebed.

Isabella Dam is capable of providing a downstream level of protection of 50–100 years. That is, it can catch that level of rain-flood while keeping flows downstream of the dam within stated channel capacity / maximum objective flow.

The above level of protection estimate has not been reviewed and approved by appropriate USACE personnel and represents a draft value only. It is based on the existing dam design. The estimate reflects only hydrology, not seismic or geotechnical risk. It assumes that the full flood-control pool is available at the beginning of the flood. There are no recent data available on sedimentation rate or on the amount of sedimentation that has occurred in the reservoir. There are several uncertainties involved in modeling flow downstream of the dam.

At full pool, Lake Isabella Reservoir covers about 11,500 acres, making it one of the larger reservoirs (by surface area) in California. It is nearly twice as large as the reservoir formed by Pine Flat Dam.

At full pool, Lake Isabella’s gross pool elevation is 2,605.5 feet.

The USACE is conducting a wide-ranging study of the main and auxiliary dams at Lake Isabella. The study was undertaken as a result of seismic concerns as well as water seepage detected in 2006. Scientists looked at core samples of the rock below the dams and dug several deep trenches to look for movement along the Kern Canyon Fault, which runs directly under the western edge of the auxiliary dam. Once thought to be inactive, the fault is now believed to be capable of causing a 7.5 magnitude earthquake, large enough to rupture the auxiliary dam.

The studies may lead to the need to repair the dams. It could take 10–15 years before the anticipated repairs can be completed. In the meantime, the USACE has reduced the amount of water that can be safely stored in the reservoir. As originally designed and constructed, the lake’s capacity was 568,000 acre-feet. However, until the dam is once again certified safe to hold that volume of water, the fill-limit has been set at 360,000 acre-feet, about 63% of the total capacity of the reservoir. That amount of gross storage can hold 50% of the 121-year average runoff (1894–2014) for the Kern River.

This 360,000 acre-feet pool restriction largely impacts recreation and irrigation users. It does not affect the dam’s ability to control floods. The flood-control pool remains unrestricted and the dam’s ability to provide a downstream level of protection remains unaffected.

Isabella Dam has a Dam Safety Action Classification (DSAC) rating of 1: Urgent and Compelling (Unsafe). The dam safety characteristics associated with that rating are: Critically near failure or Extreme high risk. There is only one other USACE-operated dam in Central California that has a DSAC rating of 1; that is Martis Creek Lake near Truckee.
Floods and Droughts in the Tulare Lake Basin

General Flood and Drought Notes

Historically, the Kern River divided in its lower reaches. The distributary point was located just west of Bakersfield, a few hundred yards east of the present-day Stockdale Bridge, at a place where the historic wooden Bellevue Weir was built across the river. That weir is now gone, but there is a rock spillway across the river at the same location.

At that point, the Kern split into two parallel channels, both of which flowed eventually north toward Tulare Lake. The main channel flowed along the west side of the San Joaquin Valley, connecting with Buena Vista Lake and Buttonwillow Swamp. Between Buena Vista Lake and Tulare Lake, this channel was known as Buena Vista Slough. The other channel, known as Goose Lake Slough, flowed along the east side of the valley. Interstate 5 occupies the high ground between those two channels.

Goose Lake Slough typically only carried water during flood periods. Bull Slough was the northern extension of Goose Lake Slough and was located north of Goose Lake. Goose Lake Slough and Bull Slough have not carried Kern River floodwaters since 1983. For a more complete description of these channels and their associated wetlands, see the section of this document: General Notes on Kern, Buena Vista, and Goose Lakes.

With the completion of Isabella Dam in 1954, the risk of Kern River water reaching the Tulare Lakebed was greatly reduced. During floods, all four federal reservoirs coordinate their operations to minimize inflows to the lakebed. Still, that has proved insufficient in a major flood such as 1969.

Now it is possible to route a large portion of Kern River floodwaters into the California Aqueduct rather than into Buena Vista and/or Tulare Lakes. This is done using the Kern River Intertie and Cross Valley Canal. Once the water enters the California Aqueduct, it is pumped over the Tehachapi Mountains and sent to the Los Angeles area. The Kern River Intertie is located just north of the Taft Highway (Highway 199). It was completed in 1977 and has a capacity of 3,500 cfs.

Storage Capacity in the Tulare Lake Basin

As shown in Table 9, the total storage capacity of the four southern federal reservoirs (Pine Flat, the expanded Kaweah, Success, and the reduced Isabella) is 1,627,900 acre-feet. If and when Isabella is restored to full capacity, the total storage capacity of these four reservoirs will be about 1,835,900 acre-feet.

<table>
<thead>
<tr>
<th>Reservoir</th>
<th>Original Capacity (acre-feet)</th>
<th>Current Capacity (acre-feet)</th>
<th>Planned Capacity (acre-feet)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Pine Flat</td>
<td>1,000,000</td>
<td>1,000,000</td>
<td>1,000,000</td>
</tr>
<tr>
<td>Kaweah</td>
<td>150,000</td>
<td>185,600</td>
<td>185,600</td>
</tr>
<tr>
<td>Success</td>
<td>82,300</td>
<td>82,300</td>
<td>82,300</td>
</tr>
<tr>
<td>Isabella</td>
<td>568,300</td>
<td>360,000</td>
<td>568,000</td>
</tr>
<tr>
<td>Total</td>
<td>1,800,300</td>
<td>1,627,900</td>
<td>1,835,900</td>
</tr>
</tbody>
</table>

The 1,627,900 acre-feet in current capacity is based on pool restriction at Lake Isabella that largely impact recreation and irrigation users. These restrictions do not affect the ability of those dams to control floods. The full 1,835,900 acre-feet is available for flood control purposes without restriction, and the dams’ ability to provide a downstream level of protection remains unaffected.

People who live below the reservoirs tend to think that they’re safe, that the reservoirs are so big that they can catch and hold the floodwaters of the biggest events. Those reservoirs have been very effective at protecting downstream communities since their construction, but they do have their limits. For comparison:

- The 1,627,900 acre-feet in combined current capacity can hold 55% of the 121-year average runoff (1894–2014) of the four rivers.
- The combined runoff of the four major rivers in the Tulare Lake Basin in 1983 was 8,746,222 acre-feet (see Table 83 and Figure 18). For comparison, that is 5.4 times the combined current capacity of the federal reservoirs on those four rivers.
Major State and Federal Canals

California State Water Project

The California State Water Project, commonly known as the State Water Project (SWP), is the world’s largest publicly built and operated water and power development and conveyance system. The primary purpose of the SWP was to provide water for arid Southern California, whose local water resources were insufficient to sustain that region’s growth. The SWP is operated by the California Department of Water Resources. The two major feature of the SWP are the Oroville Dam on the Feather River and the 702-mile-long California Aqueduct.\(^{402}\)

The bond measure that funded the SWP was the largest in the nation’s history, almost equal to California’s entire state budget for 1959. As required by the state constitution, the legislature submitted the measure to voters for their approval. Governor Edmund G. (Pat) Brown believed that the SWP was essential for California’s future growth and economic prosperity and campaigned throughout the state for the measure. The bond measure had strong support in the San Joaquin Valley. However, Sacramento Valley voters, fearing that their water would be contracted away to users in the south, generally opposed the measure.

In the November 1960 election that sent John F. Kennedy to the White House, California voters approved the SWP by a margin of less than three-tenths of 1% of the 5.8 million ballots cast, the narrowest election in the state’s history. All northern counties except for the recently flooded Yuba and Butte Counties (the future site of the Oroville Dam) voted no. However, there was sufficient Southern California support to provide the margin of victory. The SWP vote highlighted the north-south divide that would dominate California water politics for the next quarter century.\(^{403}\)

Construction on the SWP began in 1961. The San Luis Canal is considered a shared asset of the State Water Project and the federal Central Valley Project (CVP). It was completed in 1968.

The project supplies a portion of the drinking water to agencies that serve more than 26 million people.\(^{404}\) In the first 20 years of the SWP, the majority of deliveries were for agriculture; in most of the last 20 years, the majority of the deliveries have been for urban use.\(^{405}\)

DWR attempts to provide farms and cities with all of the water each summer that they buy under contract. In some summers, however, DWR is only able to allocate a certain percentage of the water deliveries from the SWP that farms and cities have under contract. See the section of this document on the Role of the Endangered Species Act in Reducing Delta Exports for a discussion of some of the reasons that DWR is not able to allocate full deliveries.

Water users must pay for all the SWP water that they contract for, even if DWR is unable to deliver it. SWP contractors pay all of their SWP contract costs, regardless of whether any water is received, plus pay additional costs for any water received (transportation costs, etc.).

According to DWR’s water portfolio data, the Tulare Lake Basin received an average of 1.125 million acre-feet per year from the State Water Project during the 13-year period 1998–2010, That represents 9% of our basin’s total water supplies (excluding 9.8 million acre-feet of direct precipitation as shown in Table 12) during that period. See Figure 21 for a discussion of our basin’s various sources of water supplies.

SWP exports from the Delta are highly variable but average 2.607 million acre-feet.\(^{406}\) SWP urban and agricultural contractors received full requested deliveries in all years from 1969–89, excepting 1977. These contractors received 100% of their requested contractual amounts in only 6 years (1995–99 and 2006) during the 23-year period from 1990–2014.
Table 10 shows the allocations from the State Water Project from 1968–2015.

Table 10. Water deliveries from the State Water Project for 48 years (1968–2015).

<table>
<thead>
<tr>
<th>Calendar Year</th>
<th>Allocation for Agriculture</th>
<th>Allocation for Urban</th>
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<tr>
<td>2015*</td>
<td>20%</td>
<td>20%</td>
</tr>
</tbody>
</table>

Source: DWR press releases.⁴⁰⁷, ⁴⁰⁸

* Tentative water supply allocation

Table 10 covers total SWP deliveries. Only 42% of those deliveries on average are sent to the Tulare Lake Basin (based on total deliveries for water years 1998–2013).
Central Valley Project

The Central Valley Project (CVP) is a federal water project operated by USBR. It was devised in 1933 in order to provide irrigation and urban water to much of the Central Valley – by regulating and storing water in reservoirs in the northern half of the state, and transporting it to the San Joaquin Valley and its surroundings by means of a series of canals, aqueducts, and pump plants, some shared with the SWP. Two of the major features of the Central Valley Project are the Friant–Kern Canal and the Delta–Mendota Canal.

The Friant–Kern Canal is a 152-mile aqueduct that was built by USBR and completed in 1951; it is part of the Central Valley Project. The canal’s purpose is to convey water to augment local surface and groundwater supplies in Fresno, Tulare, and Kern Counties. It begins at Millerton Lake on the San Joaquin River and flows south along the eastern edge of the San Joaquin Valley, ending at the Kern River near Bakersfield. The canal’s initial capacity is 5,000 cfs, gradually decreasing to 2,000 cfs at its terminus.

The Delta–Mendota Canal is a 117-mile aqueduct that was built by USBR and completed in 1951; it is part of the Central Valley Project. The canal’s purpose is to replace water in the San Joaquin River that is diverted into Madera Canal and Friant-Kern Canal at Friant Dam. The canal begins near Tracy and runs south, parallel to the California Aqueduct for most of its journey, but diverges to the east after passing San Luis Reservoir. The canal ends at Mendota Pool on the San Joaquin River near the town of Mendota.

CVP south-of-Delta agricultural contractors received full requested deliveries in all years from 1978–89. These contractors received 100% of their requested contracted supply amounts in only 3 years (1995, 1998, and 2006) during the 23-year period from 1990–2014, and 75% or better in only 8 of those years.
Table 11 shows the allocations for the federal Central Valley Project from 1998–2015.

<table>
<thead>
<tr>
<th>Calendar Year</th>
<th>Water Supply Allocation for Year</th>
<th>North of Delta</th>
<th>South of Delta</th>
<th>Friant' Class 1</th>
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<td>Ag Urban Refuge</td>
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<td>Friant East Side</td>
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<td>100% 55%</td>
<td></td>
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<td>0% 25%</td>
<td>?</td>
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<td>?</td>
<td>0%</td>
</tr>
</tbody>
</table>

Sources: USBR’s Summary of Water Supply Allocations website and press release, and DWR.

1 Friant Division contractors’ water supply is delivered from Millerton Lake on the upper San Joaquin River. The first 800,000 acre-feet of water supply is considered Class 1, and the next 1.4 million acre-feet is considered Class 2. Class 2 deliveries are not shown because reliable data for total Class 2 deliveries made in wetter years after 1998 is not readily available.

Table 11 covers all deliveries from the Central Valley Project. Only a portion of those deliveries are made within the Tulare Lake Basin. According to DWR’s water portfolio data, the Tulare Lake Basin received an average of 2.031 million acre-feet per year from the Central Valley Project during the 13-year period 1998–2010, That represents 16% of our basin’s total water supplies (excluding 9.8 million acre-feet of direct precipitation as shown in Table 12) during that period. See Figure 21 for a discussion of our basin’s various sources of water supplies.
Precipitation and Runoff

Variation in Runoff over Past 122 Years: 1894–2015

Figure 18 illustrates how unimpaired flow (full natural flow) runoff has varied for each of the four major river drainages in the Tulare Lake Basin over the last 122 years. Conditions for 2015 are projected. The average runoff of the four rivers during the period 1894–2014 was 2,941,237 acre-feet.

Source: Data compiled from DWR California Data Exchange Center and USBR which obtained it from USACE and 2015 projections from NOAA River Forecast Center.
Runoff Reconstructions

- Christopher Earle and Harold Fritts used tree-rings to reconstruct annual runoff for the Sacramento River Basin for 1560–1980.\textsuperscript{416} Dave Meko and others used tree-ring data to push that reconstruction back to A.D. 869.\textsuperscript{417}

- Lisa Graumlich used tree-ring data from subalpine conifers in the Southern Sierra to reconstruct temperature and precipitation back to A.D. 800.\textsuperscript{418} That study found that summer temperatures were warmer than late twentieth-century values from about 1100–1375, corresponding to the Medieval Warm Period. It also found a period of cold temperatures from approximately 1450–1850, corresponding to the Little Ice Age. Precipitation during the 1,000± year record varied, but generally averaged less than twentieth-century levels regardless of the variation in temperature. Refer to the section of this document that describes the 1566–1602 drought for cautions about using the results from Graumlich’s reconstruction.

- D.A. Graybill and G.S. Funkhouser used tree-ring data to reconstruct the climate from 1100 to 1987 for the Southern Sierra Nevada and Owens Valley.\textsuperscript{419}

- Malcolm Hughes and Peter Brown used giant sequoia tree-rings to reconstruct the Palmer Drought Severity Indices for the period 101 B.C. to A.D. 1988.\textsuperscript{420} That study found that the period from roughly 1850 to 1950 had one of the lowest frequencies of drought of any one hundred year period in the 2089-year record. The 20th century had a below-average frequency of extreme droughts.

- Edward Cook and others used tree-rings to reconstruct the Palmer Drought Severity Index for North America.\textsuperscript{421, 422} Their studies found that the Sierra had long periods of exceptional and extended drought from the late 800s to about 1300; . The driest two periods in western North America were centered on the mid–1100s and the mid–1200s; those are both reflected in parts of the Sierra.

- Dave Meko and others used tree-rings to reconstruct the flow on the San Joaquin River and its major tributaries for 1113 years (900–2012).\textsuperscript{423}

- A 3,000-year record from giant sequoia trees in Sequoia National Park was reconstructed by Tom Swetnam, Tony Caprio, and others.\textsuperscript{424} Their study found that the western Sierra was droughty and often fiery during the Medieval Warm Period (i.e., from 950–1250). This time period had the most frequent fires in the 3,000 years studied. During that period, extensive fires burned through parts of Giant Forest at intervals of about 3–10 years. Any individual tree was probably in a fire about every 10–15 years.

Tree-rings can only be used to reconstruct the precipitation record for a few thousand years at most. Scott Mensing and others analyzed a set of sediment cores extracted from Pyramid Lake that had been deposited over the past 7,630 years.\textsuperscript{425} They used the ratio of moisture-loving Asteraceae to drought- and salt-tolerant Chenopodiaceae in that lakebed as a proxy for drought. This allowed the authors to reconstruct a drought record for the western Great Basin. Since Pyramid Lake gets most of its water from the Sierra, this drought record can — with caveats — be applied to the Sierra as well. The authors documented multiple droughts in the region that each lasted 150–200 years. Their work suggested that variable solar activity may well be the major factor in determining the hydrological state of the region.

There have been two important studies of sediments in the Tulare Lakebed.

- In 1999, Owen Kent Davis at the University of Arizona provided a record of late Quaternary climate for the Tulare Lake region based on the palynology (pollen study) of a depocenter core. A depocenter is that part of a basin where the greatest subsidence occurred. Davis’s study found that the vegetation of the southern San Joaquin Valley used to resemble that of the contemporary Great Basin, including abundant greasewood. He also found that giant sequoia was widespread along the Sierra Nevada streams draining into Tulare Lake prior to 9,000 year B.P. (Before Present). The end of Great Basin plant assemblages 7,000 B.P. coincided with increased charcoal (i.e., fire frequency in the woodland and grasslands). Davis’s study also included conclusions regarding relative lake levels throughout the Holocene.\textsuperscript{426}

- In 2006, Rob Negrini and his associates at CSU Bakersfield built on Davis’s results with improved constraints on elevations and ages of past lake levels from trench sites at higher elevations in the Tulare Lakebed.\textsuperscript{427}

Where does precipitation end up?
DWR maintains an annual water portfolio for each of the state’s water basins (aka hydrologic regions) in support of the California Water Plan Update and for other purposes.

PRISM is an analytical tool, a model, that uses point data, a digital elevation model, and other spatial datasets to generate estimates of precipitation and other climatic parameters. As part of the water portfolio, DWR uses PRISM data to estimate the total precipitation that falls in each water basin annually.
An analysis of DWR’s water portfolio data for water years 1998–2010 (the most recent reliable water portfolio data available) shows that the Tulare Lake Basin receives an average total precipitation of about 13.6 million acre-feet per year.

Most of that precipitation soaks into the ground. Some of that precipitation, one way or another, finds its way into local rivers. A majority of the water that flows into those rivers (about 81% or 2,294 thousand acre-feet) is delivered by conveyance infrastructure (mostly canals) to water users. Most of the remainder of the water in the rivers soaks into the ground.

Some of the precipitation that soaks into the ground, however it gets there, is later pumped back up by water users. That component of groundwater withdrawals is termed “groundwater natural recharge.” As defined in the California Water Plan Update 2013 glossary, natural recharge is the percolation to groundwater basins from precipitation falling on the land and from flows in rivers and streams.\(^2\)

Most of the water that soaks into the ground is put to use by plants naturally, without the use of conveyance infrastructure. That is, the water soaks into the ground, the roots draw it up, and it eventually evaporates from the surface of the plant leaves. This can be thought of as direct precipitation because it is delivered directly to plants (agricultural, urban, or environmental) without the use of conveyance infrastructure. The other two components of total precipitation are delivered to plants indirectly: either deliveries from local rivers or natural recharge, later to be pumped back up in the form of groundwater withdrawal.

Table 12 breaks out the three different components of total precipitation, to the best of our knowledge.

<table>
<thead>
<tr>
<th>Component (Where did the total precipitation wind up)</th>
<th>Average Water Delivery (thousand acre-feet)</th>
<th>Average Percent Component</th>
</tr>
</thead>
<tbody>
<tr>
<td>Deliveries from local rivers (that is, canals)</td>
<td>2,294</td>
<td>17%</td>
</tr>
<tr>
<td>Natural recharge. Precipitation that soaks into the ground, later to be pumped back up in the form of groundwater withdrawals(^1)</td>
<td>1,545</td>
<td>11%</td>
</tr>
<tr>
<td>Direct precipitation(^2)</td>
<td>9,771</td>
<td>72%</td>
</tr>
<tr>
<td>Total precipitation</td>
<td>13,611</td>
<td></td>
</tr>
</tbody>
</table>

\(^1\) This is an educated guess based on the best available data. 
\(^2\) Precipitation used by plants without man’s assistance. For simplification, this category also includes that portion of total precipitation which evaporates before plants can use it.

The 121-year average runoff (1894–2014) of the four major rivers in the Tulare Lake Basin (Kings, Kaweah, Tule, and Kern) is 2,941,237 acre-feet. About 325 thousand acre-feet of additional runoff on average comes from drainages other than the four major rivers.\(^3\) These include the White River, and Deer, Poso, Yokohl, Cottonwood, Dry, and Mill Creeks. That makes a total average runoff of about 3.3 million acre-feet per year. This is only about 24% of the 13.6 million acre-feet that falls as average total precipitation. Partly that is because some of the precipitation goes to naturally recharging the groundwater aquifer. However, most of the difference is due to water being vaporized through evapotranspiration (evaporation plus transpiration of water vapor from plants) before it can flow into the rivers and streams.

Most of the average 13.6 million acre-feet in total precipitation that falls in our basin is used by plants of one sort or another (agricultural, urban, or environmental). Once those plants pull the water into their roots and through the stems, it is vaporized from the leaf surfaces and tends to be lost from our basin.

Based on the water portfolio, about 4% of total precipitation (541 thousand acre-feet) is considered to have fallen on crops planted in developed irrigated land areas which can use the precipitation. That portion of total precipitation is termed “ag effective precipitation on irrigated lands” by the California Water Plan Update 2013.\(^4\) The rest fell elsewhere in the Tulare Lake Basin such as in the foothills or the national parks, or it fell at times when crops planted in developed irrigated lands could not use the precipitation. Direct precipitation that falls on non-irrigated areas (dryland farms, rangelands, forests, etc.) is particularly valuable to the plants that grow in those areas because they have no supplemental source of irrigation.
Sources of Water

Background about Water Demand and Water Supply

Water Demand

Many people can agree that water demand exceeds water supplies in the Tulare Lake Basin; that’s the core or our problem. We have insufficient water supplies to meet our demand. But what exactly do we mean when we refer to demand? The term “demand” is tricky. It can mean different things in different contexts, and it can mean different things to different people.

The California Water Plan Update 2013 uses the term “demand” in a variety of ways. These range from a measure of consumption (how much water we used) to a measure of desire (how much water would we use if it were available. The Update 2013 uses “demand” in at least eight different ways:

1. Applied water (the total amount of water actually used in a given year)
2. Dedicated and developed water supplies (sustainable component of applied water)
3. Net water use (demand)
4. Average year water demand (average amount of applied water)
5. Explanation for past changes in the amount of applied water (demand)
6. Temporal demand
7. Projecting future water demand relative to future water supplies under various scenarios
8. Water demand (how much water would we use if it were available)

1. **Applied water.** The California Water Plan Update 2013 glossary defines this term as "the total amount of water diverted from any source to meet the demands of water users without adjusting for water that is used up, returned to the developed supply or irrecoverable."^[431]

Applied water includes both consumptive use and reuse and return flows. Applied water includes all groundwater withdrawals, including groundwater overdraft. Total applied water includes four types of applied water: agriculture, urban, managed wetlands, and instream.

Examples of how the Update 2013 uses “applied water” and “demand” in a similar context include:
- The update refers to how much of the “applied water demand” in a given year was met by surface supplies versus groundwater supplies.
- The update says that aquifer conditions and groundwater levels change in response to varying supply, demand, and climate conditions. That use of “demand” apparently means something like withdrawals or need.
- There are similar references to “seasonal or short-term changes in groundwater demand” or “decreasing groundwater demand.” All of these uses are clearly referring to the amount of groundwater withdrawals, a component of applied water.

The average amount of applied water during the 13-year period 1998–2010, was 13.1 million acre-feet per year. The amount of applied water is fairly stable from year to year; it typically varies less than 5% from the 13.1 million acre-feet average. During the 13-year period 1998–2010, the amount of applied water varied from 12.3 million acre feet (2001) to 14.8 million acre-feet (2006).

2. **Dedicated and developed water supplies.** The Update 2013 glossary defines this term as “This represents water distributed among urban and agricultural uses, and which is used to protect and restore the environment or for storage in surface water and groundwater reservoirs. In any year, some of the dedicated supply includes water that is used multiple times (reuse) and water held in storage from previous years.”^[432]

Not all surface water supplies have been developed. For example, some floodwaters are lost from the system before they can be stored or used. In addition, much of the total precipitation in the basin is not captured in the dedicated and developed water supplies. The term dedicated and developed water supplies includes sustainable groundwater (groundwater that has been recharged from the surface), but it does not include groundwater overdraft. The term dedicated and developed water supplies is identical to applied water minus one component: groundwater overdraft. This is the sustainable component of applied water.

The Water Plan updates are in a transition period where they are talking more and more about the components of groundwater withdrawals. However, DWR does not yet have enough data to quantify groundwater with the same specificity that it quantifies surface water supplies.
Due to a relative lack of data, the water portfolios do not fully differentiate among the various components of groundwater as described in Table 15. Groundwater overdraft is lumped in with other groundwater withdrawals recharged by surface precipitation. The Update 2013 is built using the water portfolios, so the dedicated and developed water supplies includes all groundwater withdrawals, including groundwater overdraft. They do not call out what portion of the groundwater withdrawals represent groundwater overdraft. Therefore, Table TL-15 in Update 2013 treats all of groundwater withdrawals as part of dedicated and developed water supplies. However, a note in that table recognizes that a portion of those groundwater withdrawals are in fact groundwater overdrafts.

The average amount of dedicated and developed water consumed during the 13-year period 1998–2010 was 11.8 million acre-feet. It excludes the 1.2 million acre-feet of groundwater withdrawals that were not recharged. This is the amount of our groundwater overdraft.

3. Net water use (demand). The Update 2013 glossary defines this term as "For the California Water Plan water portfolios, this represents the amount of water needed in a water service area to meet all requirements. It includes the consumptive use of applied water, the irrecoverable water from the distribution system, and the outflow leaving the service area. It does not include reuse of water within a service area." The term net water use is identical to applied water minus two components: reuse and groundwater withdrawals recharged from surface applications.

That definition is a bit misleading. Net water use does not really represent the amount of water needed to meet all the requirements in a water service area. Such a definition better suits the term "applied water." Water users need all the water that they apply in order to meet their requirements. As detailed in Table 16 on page 121, the average net water use consumed during the 13-year period 1998–2010 was 8.2 million acre-feet. It excludes the average annual total 4.8 million acre-feet total of reused and recycled water. Net water use can be thought of as the amount of water that we have to work with, including groundwater overdraft. Applied water is what we do with that water; it is the gross or total amount of water that we use, including reused and recycled water.

The 4.8 million acre-feet of reused and recycled water represents 59% of the 8.2 million acre-feet in net water use. We apply that water once and then apply it again. The rest of the 8.2 million acre-feet is applied once, but then it evaporates and leaves our basin. The 4.8 million acre-feet of reused and recycled water has two components: 1.6 million acre-feet of reuse and 3.2 million acre-feet of groundwater withdrawals recharged from surface application. Together, that water represents 37% of all the water used in the Tulare Lake Basin.

4. Average year water demand. The Update 2013 glossary defines this term as "Demand for water under average hydrologic conditions for a specific level of development." Eric Osterling with the Kings River Conservation District said that average year water demand is simply an overall annual consumption average for a given area, unless “demand” is calculated specifically for a particular end user (e.g., agriculture, or even a specific crop type).

Average year water demand is calculated by averaging the applied water over a period of years. As shown in Table 14, the average water supply delivered by conveyance infrastructure for water years 1998–2010 in the Tulare Lake Basin was 13,073,000 acre-feet. This measure of water demand / water supply excludes the average 9.8 million acre-feet of direct precipitation.

The water plan updates sometimes refer to the average year water demand as the historical average water demand in order to contrast it with the projected future average water demand. As shown in the Update 2013, the historical average water demand (aka average year water demand) for 1998–2005 was 9,466 thousand acre-feet for agriculture and 676 thousand acre-feet for urban. As shown in Table 13, those figures would be 10,709 and 711 if calculated using DWR’s water portfolio data for 1998–2010.

"Demand" in this sense is all about applied water. The average year water demand is calculated by averaging the applied water over a period of years. The allocation figures in the water plan update make it hard to understand that because of how the dedicated wild and scenic flows are presented. That water flowing through those rivers isn’t really applied in the sense that agriculture or urban users apply water. It is just a dedicated flow that has to be provided for by law.

This water is dedicated to the wild and scenic rivers only while it is flowing through those rivers. Once the water passes through the designated segments, it is reclassified as “reuse of return flows” and becomes available for use downstream by agriculture and urban users.
Floods and Droughts in the Tulare Lake Basin
General Flood and Drought Notes

To clarify the issue, Table 13 presents the allocation of the available water supply both ways:

1. Showing the dedicated water as it flows through the designated wild and scenic river segments.
2. Showing total water use after downstream water users have applied the water that flowed through those wild and scenic rivers.

<table>
<thead>
<tr>
<th>Component</th>
<th>Average Year Water Demand (thousand acre-feet)</th>
<th>Percent Contribution</th>
</tr>
</thead>
<tbody>
<tr>
<td>Agriculture (nominal, excluding reuse of W&amp;S water)</td>
<td>10,709</td>
<td>82%</td>
</tr>
<tr>
<td>Urban (nominal, excluding reuse of W&amp;S water)</td>
<td>711</td>
<td>5%</td>
</tr>
<tr>
<td>Managed wetlands (refuges)</td>
<td>86</td>
<td>1%</td>
</tr>
<tr>
<td>Wild and scenic rivers — dedicated flow*</td>
<td>1,567</td>
<td>12%</td>
</tr>
<tr>
<td>Nominal application of water supply</td>
<td>13,073</td>
<td>100%</td>
</tr>
<tr>
<td>Agriculture (total, including reuse of W&amp;S water)</td>
<td>12,179</td>
<td>93%</td>
</tr>
<tr>
<td>Urban (total, including reuse of W&amp;S water)</td>
<td>808</td>
<td>6%</td>
</tr>
<tr>
<td>Managed wetlands (refuges)</td>
<td>86</td>
<td>1%</td>
</tr>
<tr>
<td>Total</td>
<td>13,073</td>
<td>100%</td>
</tr>
</tbody>
</table>

* Based on use of other river water, it is assumed that 94% of the wild and scenic river water is later used by agriculture and 6% by urban although this use is not tracked by the water portfolios.

The above analysis was done using DWR’s water portfolio data for 1998–2010, calculated using two methods. Similar analyses were done using the water portfolio data for 1998–2005. DWR in the 2009 Water Plan used data to calculate the agriculture/urban/environment breakdown for those years as 82% / 5% / 13% (using the upper method from Table 13). In contrast, Ellen Hanak and her coauthors used the same water portfolio data and reported the breakdown for those years as 98% / 4% / 1% (using the lower method from Table 13). It all comes down to how you treat the dedicated flows in the wild and scenic rivers.

The average amount of applied water during the 13-year period 1998–2010, was 13.1 million acre-feet per year. That is the average year water demand. It is also the water supply; we consumed what was supplied. The amount of applied water is fairly stable from year to year; it typically varies less than 5% from the 13.1 million acre-feet average. During the 13-year period 1998–2010, the amount of applied water varied from 12.3 million acre feet (2001) to 14.8 million acre-feet (2006).

There are a number of demand and supply terms that are used interchangeably, including:

- The (historic) average year water demand for the 13-year period 1998–2010 in the Tulare Lake Basin was 13.1 million acre-feet per year.
- The average amount of applied water for this 13-year period was 13.1 million acre-feet per year.
- The average water consumption for this 13-year period was 13.1 million acre-feet per year.
- The average water need or requirement (the amount of water we chose to apply) for this 13-year period was 13.1 million acre-feet per year.
- The average total water supply for this 13-year period was 13.1 million acre-feet per year.

5. Explanation for past changes in the amount of applied water (demand). The Water Plan updates sometimes use the term “demand” to explain how and why use the amount of applied water (or a component of applied water such as groundwater withdrawals) has changed in the past. Examples of how the Update 2013 describes changes in past demand include:

- “...significantly reduced the demand for groundwater for agricultural use.”
- “In order to meet the rapidly increasing demand for groundwater supplies during the 2007–09 period, the annual installation of new agricultural wells nearly tripled.”

6. Temporal demand. Eric Osterling said that demand fluctuates depending upon location, season, and use. That is temporal demand. Water demand models have monthly and sometimes daily time steps and adjust demand in any one particular area of the model based upon demand variables. Temporal demand information for the entire Central Valley and its sub-regions can be extracted from DWR’s C2VSim and USGS’s CVHM hydrologic models (see the section of this document on Summary of Groundwater Overdraft). DWR’s California Central Valley Groundwater-Surface Water Simulation (C2VSim) model is an integrated numerical model that simulates water movement through the linked land surface, groundwater and surface water flow systems in the
Central Valley. The C2VSim model contains monthly historical stream inflows, surface water diversions, precipitation, land use and crop acreages from October 1921 through September 2009. C2VSim dynamically calculates crop water demands, allocates contributions from precipitation, soil moisture and surface water diversions, and calculates the groundwater withdrawals required to meet the remaining demand. The model simulates the historical response of the Central Valley’s groundwater and surface water flow system to historical stresses, and can also be used to simulate the response to projected future stresses.

7. Projection of future demand. The Update 2013 projects the future demand (average year water demand for 2043–2050) under various scenarios compared to the historical 1998–2005 average year water demand. It then calculates unmet demand, which is the difference between projected future water demand and future water supply. Possibly DWR may have determined water demand in the past through surveys of end users. It appears that future water demand is now determined in part through IRWM plans.

8. Water demand. The Update 2013 glossary defines this term as “The desired quantity of water that would be used if the water were available and if a number of other factors, such as price, did not change. Demand is not static.” As worded, this definition of demand refers to the desires of water users, not to the amount of water that they use. Evelyn Tipton, DWR’s expert on the water portfolios, said that the water portfolios do not address water demand in the sense of users' desire or need. The portfolios address actual use (applied use, net use, etc.) in past years. Since water portfolios are developed for past years, “demand” is not applicable. If a demand wasn't met for some reason (cost, availability, whatever), the water was not used for that purpose and so does not appear in the portfolio. The purpose of the water portfolios is to quantify the amount of water that was used in a prior year and the sources of that water.

In normal usage, “demand” has an economic component to it. As price goes up, demand goes down. If the price of steak were to go up, people would tend to buy less of it regardless of availability. However, the Water Plan updates makes the assumption that price does not change when talking about water demand. Water demand, in this sense of the term, is how much water irrigators would use if it were available at a price similar to what they have been paying. Realistically, the answer would be a very large amount. Agriculture in the Tulare Lake Basin has a large land base and an essentially unlimited thirst. They can make productive use of as much water as they can get.

Viewed that way, water demand will almost always be higher than the actual dedicated and developed water supplies. Water users will want more water than society is able (or willing) to provide. Dedicated and developed water supplies can never equal this kind of water demand in any but the wettest years. But that isn’t quite how “water demand” has worked in practice. As the Update 2009 explained:

Water demand is more or less controlled by water supplies. Over the years, agricultural acreage has risen based on water supplies.

That is just a recognition that water demand (despite the definition) is essentially the same as applied water, the amount of water available.

Public perception of water demand. As shown above, the Update 2013 uses the term “demand” in at least eight different ways. Most of those boil down to the same meaning: applied water or some subset of applied water such as dedicated and developed water supplies. However, we need to keep in mind that the term water demand is sometimes used by the public and water users to mean some other unit, one that is larger than dedicated and developed water supplies.

When used in that way, the term water demand is more closely associated with how much water users want or need rather than with dedicated and developed water supplies. Because the term water demand is being used in different ways, some formally defined, and some not defined, it can be challenging to have clear communication on this issue. It is a bit confusing to have so many different metrics for demand. Wherever possible, this document uses terms other than “water demand” to minimize confusion.

What is our demand / supply problem?

Based on the above review of the many different ways that the term “demand” is used, we can revisit what many perceive as our core water problem:

demand exceeds supply
Floods and Droughts in the Tulare Lake Basin
General Flood and Drought Notes

There seem to be at least three possible interpretations of that statement:

1. Average year water demand exceeds average dedicated and developed water supplies.
2. A long-term groundwater overdraft condition exists. That is the same as #1 above. The term dedicated and developed water supplies is identical to applied water minus groundwater overdraft.
3. Dedicated and developed water supplies are a limiting factor for agriculture. If the amount of dedicated and developed water supplies were increased, agriculture would make productive use of that water.

Allocation of Surface Water Rights

Water is limited in California. For the last 100 years, California’s State Water Resources Control Board (SWRCB) and its predecessors have been responsible for allocating available water supplies to beneficial uses. A water right, granted by the state, bestows the power to divert a certain amount of water from a stream or river. California’s system of water rights is the state’s way of sharing limited water resources among many different users. California has a dual system of both riparian and appropriative surface water rights. Riparian land is land that touches a lake, river, stream, or creek. A riparian water right is a right to use the natural flow of water on riparian land. Riparian rights usually come with owning a parcel of land that is adjacent to a source of water. Riparian rights have a higher priority than appropriative rights. The priorities of riparian right holders generally carry equal weight; during a drought all share the shortage among themselves.

Appropriative water can generally be defined as water that is diverted for use on non-riparian land. Prior to 1914, there was no comprehensive permit system available to establish appropriative water rights in California, Senior water rights are those appropriative rights bestowed before 1914; junior water rights are those bestowed post–1914.

Senior water rights holders tend to comprise the same mix of users as junior water rights holders, with the major difference being that they simply started using the water first and established a legal right to it earlier. In times of shortage, state water law says if there is not enough water for all water right holders, holders of junior water rights will be curtailed before restrictions are imposed on holders of senior water rights. The curtailment of water use during periods of severe drought by holders of junior rights is intended to preserve limited water supplies for those with senior rights.

Water users are required to report their diversions to the SWRCB. Until recently, most senior water right holders (especially riparian water right holders) did not report their diversions despite the requirement to do so. During times of severe drought, the SWRCB has further restricted the extent to which diversions can be made.

In August 2014, Ted Grantham and his co-author analyzed the state water-rights database. Water experts had long known the amount of surface water granted by water rights far exceeds the state’s average supplies. Historically, the over-allocation has not raised much concern; in most years, there has been enough runoff of rain and snowmelt to go around, at least for withdrawals and consumptive uses by humans (although not for the needs of the environment).

But circumstances appear to be changing. Water year 2014 was our fourth driest water year in the Tulare Lake Basin since 1894 (see Table 23), and demands for water were at an all-time high. The huge gap between allocations and natural flows — coupled with great uncertainty over water-rights holders’ actual usage — was increasingly creating conflicts between water users and confusion for water managers trying to figure out whose supplies should be curtailed during a drought.

To understand where and to what degree California rivers have been claimed, the authors mapped all appropriative water rights recorded by the SWRCB. They quantified the maximum annual diversion volume of water rights for all rivers and streams and compared that data with estimates of water supply.

The authors of the Grantham study found that water rights exceed average supplies in more than half of the state’s large river basins. The results showed that water right allocations total 400 billion cubic meters, approximately five times the state’s mean annual runoff. In the state’s major river basins, water rights account for up to 1000% of natural surface water supplies, with the greatest degree of appropriation observed in tributaries to the Sacramento and San Joaquin Rivers and in coastal streams in Southern California. In the San Joaquin River, water rights exceed flows by as much as eightfold.

Not only are many rivers over-allocated, but the amount of water actually used by water-rights holders is poorly understood. Comparisons of allocations with water use suggested that in most of California, only a fraction of claimed water is being used. Statewide, appropriative water-rights claims for consumptive uses are about five times greater than average surface-water withdrawals.
The extent to which water rights have actually been over-allocated is uncertain. Eric Osterling with the Kings River Conservation District said that the issue is not as clear-cut as the Grantham study makes it appear. Water rights are reported with several numbers; almost never are they a simple single number. Permits and licenses must address all flow conditions from drought to flood. They must also address different runoff patterns from early to late. Often a permit or license will have a maximum instantaneous and maximum 30-day average rate of diversion in cfs or gpm. They will include season of diversion which can range from days to an entire year. All of those parameters are often capped by an annual maximum.

For the purposes of billing water rights fees, the SWRCB uses the maximum daily rate multiplied by the season of diversion and converts that to acre-feet. In the case of the Kings River, that exceeds the largest runoff period ever recorded by a large amount. However, the state-licensed annual maximum diversion for the Kings is approximately half of the total 1983 runoff of 4.5 million acre-feet (see Table 83 on page 346).

The numbers get even more inflated when you add the hydro generation licenses (which are non-consumptive) to the equations. One of the best example of this is in Imperial Irrigation District where there are about 10 hydro plants on the All American Canal. Each plant has its own license, so the same water runs through each plant. For illustration purposes, assume that each plant is licensed to 1,000,000 acre-feet per year. Based on licenses, this could be represented as if the 10 plants are using 10,000,000 acre-feet. In reality these plants aren’t consuming any water despite their licenses; their water use is non-consumptive. This example illustrates the highly questionable use of licensed values in any discussion of California water without excruciatingly detailed understanding of the specific license.

**Water Supply**

Based on DWR’s annual water portfolio data for water years 1998–2010 (the most recent reliable water portfolio data available), the Tulare Lake Basin uses (consumes) an average of about 22.9 million acre-feet of water per year (see Table 14). Of this, about 13.1 million acre-feet is delivered and applied via conveyance infrastructure, largely on the valley floor. The difference between the two numbers (9.8 million acre-feet) is the amount of direct precipitation that is used by plants without man’s assistance.

<table>
<thead>
<tr>
<th>Source of Water</th>
<th>Average Water Delivery (thousand acre-feet)</th>
<th>Average Percent Contribution</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td></td>
<td>Excluding Direct Precipitation</td>
</tr>
<tr>
<td>Deliveries from local rivers</td>
<td>2,294</td>
<td>18%</td>
</tr>
<tr>
<td>State Water Project deliveries</td>
<td>1,125</td>
<td>9%</td>
</tr>
<tr>
<td>Central Valley Project deliveries</td>
<td>2,031</td>
<td>16%</td>
</tr>
<tr>
<td>Total reuse of water</td>
<td>1,626</td>
<td>12%</td>
</tr>
<tr>
<td>recharge from surface application</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Groundwater withdrawals — recharge from surface application</td>
<td>3,206</td>
<td>25%</td>
</tr>
<tr>
<td>Other groundwater withdrawals</td>
<td>2,791</td>
<td>21%</td>
</tr>
<tr>
<td>Total applied water excluding change in surface water storage</td>
<td>13,073</td>
<td>100%</td>
</tr>
<tr>
<td>Direct precipitation*</td>
<td>9,771</td>
<td>100%</td>
</tr>
<tr>
<td>Total water use excluding change in surface water storage</td>
<td>22,845</td>
<td>100%</td>
</tr>
</tbody>
</table>

* Precipitation used by plants without man’s assistance. For simplification, this category also includes that portion of total precipitation which evaporates before plants can use it. As shown in Table 12, direct precipitation is the total precipitation in the basin after removing two categories:
1. Deliveries from local rivers (that is, canals).
2. Natural recharge. Precipitation that soaks into the ground, later to be pumped back up in the form of groundwater withdrawals. This is an educated guess based on the best available data.

The Tulare Lake Basin uses (consumes) more water than any other region of California — about 13.1 million acre-feet a year (excluding 9.8 million acre-feet of direct precipitation). Deliveries from local rivers are only able to provide 18% of that water, so we have a long history of looking to other sources to meet our needs (the amount of water we choose to apply). Delta imports and San Joaquin River diversion supply about 24% via the State Water Project and Central Valley Project.
Floods and Droughts in the Tulare Lake Basin
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Table 10 shows the allocations for the State Water project from 1968–2014. Table 11 shows the allocations for the federal Central Valley Project from 1998–2014. The water from these projects serves a wide variety of users in Central and Southern California. In drought years, the total amount of water deliveries is reduced. Agricultural users seldom get their full allotments of water because all the federal water delivery contracts add up to more water than exists, even in years without drought. That is because the water projects were developed when California had fewer people, and the infrastructure of the Central Valley Project was never completed.

These water cuts affect some water users more severely because of the way that their contracts were written. When droughts occur, farmers with the newest water delivery contracts are the first to face cutbacks. That is why farmers on the west side of the Tulare Lake Basin receive the least water during droughts. In recent years, there has been a campaign to convince others that state and federal politicians are diverting water that should rightfully be coming to farmers on the west side of the Tulare Lake Basin. This campaign is carried out in large part through signs and billboards. The signs typically say things like “Man-made Drought” and “Congress-created Dust Bowl.” This seems to be a statement that the sign-makers believe their water is being sent elsewhere, presumably for environmental or some similar low-value use.443

During the 2007–09 drought, there was considerable controversy around the role that environmental protections, particularly the Endangered Species Act, played in the reduced exports to south-of-Delta water users. However, analyses from the California Department of Water Resources and the Congressional Research Service showed that over three-quarters of the reductions in Delta exports in 2009 was due to drought conditions, and that less than a quarter of the reductions was due to environmental protections such as protecting endangered fish and maintaining Delta salinity standards.444 Westside farmers were the first to feel the impact of the reduced exports because of the way that their water delivery contracts were written.

In recent years, some studies have attempted to show when reductions in south-of-Delta exports (and increases in groundwater overdrafts) were a result of “regulatory droughts” versus “climactic droughts.”445 As the California Department of Water Resources and the Congressional Research Service analyses show, that is not an easy distinction to make.

As shown in Table 14, total groundwater withdrawals provide an average of about 6 million acre-feet of water per year in our basin. If you exclude direct precipitation, this accounts for 46% of all the water used. Table 15 breaks out the various components of the groundwater withdrawals.


<table>
<thead>
<tr>
<th>Source of Water</th>
<th>Average Water Delivery (thousand acre-feet)</th>
<th>Percent of Groundwater Deliveries</th>
<th>Average Percent Contribution</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td></td>
<td></td>
<td>Excluding Direct Precipitation</td>
</tr>
<tr>
<td>Groundwater withdrawals — recharged from surface application</td>
<td>3,206</td>
<td>53%</td>
<td>25%</td>
</tr>
<tr>
<td>Groundwater withdrawals — from natural recharge</td>
<td>1,545</td>
<td>26%</td>
<td>11%</td>
</tr>
<tr>
<td>Groundwater withdrawals — not recharged</td>
<td>1,246</td>
<td>21%</td>
<td>10%</td>
</tr>
<tr>
<td>Total groundwater withdrawals</td>
<td>5,998</td>
<td>100%</td>
<td>46%</td>
</tr>
<tr>
<td>Total applied water — excluding change in surface water storage</td>
<td>13,073</td>
<td>100%</td>
<td>57%</td>
</tr>
<tr>
<td>Direct precipitation</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Total water use — excluding change in surface water storage</td>
<td>22,845</td>
<td>100%</td>
<td></td>
</tr>
</tbody>
</table>

1 The components of the total groundwater withdrawals are not understood with precision. This estimate of natural recharge may be overstated somewhat.

2 Data presented in the California Water Plan Update 2013 for the years 2005–2010 suggest an average groundwater aquifer storage loss of about 1.2 million acre-feet per year.446 Other sources speculate the overdraft is somewhat larger.

Approximately half of the groundwater withdrawals consists of water that has been recharged from surface application. That water has been applied on the surface for agricultural purposes, soaked into the ground, and then has been pumped back up to be reapplied. In effect, it is recycled water; it is being applied twice.
The other half of the groundwater withdrawals comes from two sources:

- Approximately half is recharged from natural hydrologic processes. This is termed “groundwater natural recharge.” As defined in the California Water Plan Update 2013 glossary, groundwater natural recharge “represents the percolation to groundwater basins from precipitation falling on the land and from flows in rivers and streams.”

- The remaining portion (roughly 1.2 million acre feet per year) is the amount of our average groundwater overdraft, basin-wide. This groundwater overdraft supplements the natural amount of water available in our basin, the amount available from total precipitation, imports from north of our basin, and other sources. This is more fully described in Table 17.

As shown in Table 15, the groundwater overdraft is roughly 5% of all the water used in our basin. If you exclude direct precipitation (and only consider water delivered by conveyance infrastructure), then the groundwater overdraft is about 10% of all the water used in our basin.

The amount of groundwater withdrawals in the San Joaquin Valley is among the highest in the world according to Fresno engineer and water authority Ken Schmidt, who has worked on groundwater issues for more than four decades. Ken says the overdraft is staggering, somewhere between 1 and 2 million acre-feet of water per year.

Table 16 shows the components of reused and recycled water in the Tulare Lake Basin.

<table>
<thead>
<tr>
<th>Source of Water</th>
<th>Average Water Delivery (thousand acre-feet)</th>
<th>Average Percent Contribution</th>
<th>Excluding Direct Precipitation</th>
<th>Including Direct Precipitation</th>
</tr>
</thead>
<tbody>
<tr>
<td>Total reuse of water (reuse of return flows)</td>
<td>1,626</td>
<td>12%</td>
<td>7%</td>
<td></td>
</tr>
<tr>
<td>Groundwater withdrawals — recharged from surface application</td>
<td>3,206</td>
<td>25%</td>
<td>14%</td>
<td></td>
</tr>
<tr>
<td>Total reused and recycled water</td>
<td>4,832</td>
<td>37%</td>
<td>21%</td>
<td></td>
</tr>
<tr>
<td>Water applied only once</td>
<td>3,409</td>
<td>26%</td>
<td>15%</td>
<td></td>
</tr>
<tr>
<td>Net water use (amount of water actually consumed)</td>
<td>8,241</td>
<td>63%</td>
<td>36%</td>
<td></td>
</tr>
<tr>
<td>Net water use</td>
<td>8,241</td>
<td>63%</td>
<td>36%</td>
<td></td>
</tr>
<tr>
<td>Total reused and recycled water (this water applied a second time)</td>
<td>4,832</td>
<td>37%</td>
<td>21%</td>
<td></td>
</tr>
<tr>
<td>Total applied water — excluding change in surface water storage</td>
<td>13,073</td>
<td>100%</td>
<td>57%</td>
<td></td>
</tr>
<tr>
<td>Direct precipitation</td>
<td>9,771</td>
<td>43%</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Total water use excluding change in surface water storage</td>
<td>22,845</td>
<td>100%</td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

1As shown in Table 13, 1,567 thousand acre-feet (97%) of the “Total reuse of water (reuse of return flows)” is water that was dedicated to wild and scenic rivers. Once that water flows through those river segments, it is reclassified and made available for use by agriculture and urban users.

2Net water use is the total amount of water that is actually consumed. Of this 8,241 thousand acre-feet, 4,832 thousand acre feet (59%) is applied twice. So the total amount of applied water is 13,073 thousand acre-feet.

In Northern California and in the Sierra, the mean annual precipitation of more than 40 inches exceeds the annual evapotranspiration, and the resulting perennial water surplus eventually becomes runoff in streams. But perennial water deficiency characterizes all of Southern California and practically all the lowlands farther north. This water deficiency is a natural condition of the climate that is related to — but distinct from — water demands or water requirements of man and his crops and other activities.

The term “actual evapotranspiration” represents the amount of evapotranspiration that used to occur on the valley floor naturally when it was a desert, the San Joaquin Valley Desert. The term “potential evapotranspiration” describes the amount of evapotranspiration that occurs on much of the valley floor now, with an unnaturally high level of vegetation sustained by irrigation.

The vegetation of the desert is sparse and uses little water because water is deficient. When water supply increases, as in a desert irrigation project, evapotranspiration also increases. This is potential
evapotranspiration, as distinct from actual evapotranspiration. That is the situation we have created on irrigated lands in the San Joaquin Valley today; a situation in which potential evapotranspiration exceeds actual evapotranspiration.\textsuperscript{451}

The valley floor in the Tulare Lake Basin, where agricultural production is most intense, has an average water deficiency (precipitation minus potential evapotranspiration) under natural conditions of as much as 40 inches per year or more.\textsuperscript{452, 453} Thus, agricultural development in the valley is dependent on water from sources other than direct precipitation.

The supplemental water needed for agricultural production in this desert environment is obtained from four sources:

1. Streams and rivers that enter the valley from the surrounding mountain ranges, where there is a surplus of water. This surface water is diverted by canals to areas of farming.
2. Water imported from streams and rivers north of the Tulare Lake Basin.
3. Reuse and recycled water.
4. Groundwater, which is used primarily where surface-water supplies are not available or are not sufficient or dependable enough to support desired agricultural activities.

Table 17 shows where the 13.1 million acre-feet of water comes from that is applied in the Tulare Lake Basin in an average year. (Or 22.9 million acre-feet if you include the 9.8 million acre-feet of direct precipitation.) We receive 13.7 million acre-feet of total precipitation. We capture 3.8 million acre-feet of that and deliver it to water users (deliveries from local rivers plus groundwater withdrawals from natural recharge). Since that isn’t nearly enough to meet our needs (the amount of water we choose to apply) of 13.1 million acre-feet, we supplement it with 9.2 million acre-feet of water from other sources. If you exclude direct precipitation, those supplemental sources account for 71% of all the water applied in the Tulare Lake Basin in an average year.

What’s the big deal about direct precipitation, the 43% of our water supply that is not delivered by conveyance infrastructure? Why should we care? There are at least four reasons why we need to pay attention to direct precipitation when accounting for water supplies in our basin:

1. Including direct precipitation gives us the big picture of all our water sources. Knowing the full amount of water available (and used) in our basin provides context for the individual sources of water.
2. Including direct precipitation as a separate category in tables and charts allows for two different views of our water budget depending on your interest: the 13.1 million acre-feet that are delivered by conveyance infrastructure or the total 22.9 million acre-feet.
3. Direct precipitation that falls on non-irrigated areas (dryland farms, rangelands, forests, etc.) is particularly valuable to the plants that grow in those areas because they have no supplemental source of irrigation.
4. Direct precipitation represents a potential source of additional water in the form of dedicated and
developed water supplies; see the section of this document that describes the Potential for a Sustainable
Water Supply.

Figure 19 shows the amount of water delivered by conveyance infrastructure in the Tulare Lake Basin from
1998–2010. It does not include the direct precipitation (see Table 12 and Table 14). This is the amount of water
that we use (consume). It is the amount of applied water. The black line on the chart shows the relative amount
of precipitation in each year.

Although our precipitation varies dramatically, the total amount of water delivered by conveyance infrastructure
in the Tulare Lake Basin typically varies less than 5% from year to year. What does vary are the sources of that
water. The preference of water users is to use surface water, presumably because it is the cheapest and
highest-quality water.

Analysis of DWR’s water portfolio data for water years 1998–2010 shows that in an average year, about 18% of
our basin’s water supply deliveries (excluding 9.8 million acre-feet of direct precipitation) comes from local
rivers (2.3 million acre-feet). About 24% of our deliveries (3.2 million acre-feet) is imported from rivers to the
north of our basin via the State Water Project and Central Valley Project. We turn to our groundwater aquifer to
make up most of the rest of our needs (our average water consumption of 13.1 million acre-feet).

In drier years, we rely more heavily on our groundwater aquifer than we do in wetter years. As shown in Figure
10 on page 45, water year 2006 was classified as a wet year in the San Joaquin Valley. Figure 19 shows that
more water than average was delivered by rivers flowing down from our local mountains that year. Water year
2006 was also classified as a wet year in the Sacramento Valley; this resulted in relatively more water being
delivered to our basin through the Sacramento–San Joaquin Delta and via the State Water Project and Central
Valley Project.

That can be contrasted with water year 2008 which was classified as a critically dry year in the San Joaquin
Valley. It was the second year of the 2007–09 drought. In that year, less water was delivered by rivers flowing
down from our local mountains. Water year 2008 was also classified as a dry year in the Sacramento Valley;
this resulted in relatively less water being delivered to our basin from the Delta. When available dedicated and
developed surface water supplies are insufficient to meet our needs (the amount of water we choose to apply),
we generally turn to our groundwater reserves.

Figure 19 shows how we use groundwater withdrawals to supplement our surface water supplies in order to
meet our needs. Because of the drought conditions, we withdrew nearly twice as much groundwater in 2008
than in 2006 (8.4 million acre-feet compared with 4.3 million acre-feet).
Runoff is closely correlated with the amount of precipitation. The combined runoff of the four rivers in the Tulare Lake Basin during water year 1998 was 5,990,549 acre-feet, 204% of the 1894–2014 average. As shown in Figure 19, that allowed the groundwater aquifer to be recharged slightly that year. Such a recharge event rarely happens.

The combined runoff of the four rivers in the Tulare Lake Basin during water year 2011 was 5,910,342 acre-feet, 201% of average. This was nearly as large as the runoff in 1998. This may have allowed some recharge of the groundwater aquifer. The groundwater aquifer may have been recharged in 2011 as well, but the water portfolio for that year is not yet available.

Our needs usually exceed our available dedicated and developed water supplies. In effect, we are routinely in drought. (This is a socioeconomic drought, not a meteorological or hydrological drought. See the section of this document on What Constitutes a Drought for a description of the different types of droughts.) Years like 1998 when we break even or can recharge the groundwater aquifer are the rare exception. In most years we withdraw more water from the groundwater aquifer than is recharged naturally or through surface application. That is, we overdraw the groundwater aquifer in most years.

Figure 20 shows where the Tulare Lake Basin gets its water in different types of water years from 1998–2010. This chart represents all water sources including direct precipitation (see Table 12). As shown in Table 14, direct precipitation makes up about 43% of the water that we use in our basin in an average year. That precipitation is used by plants without man’s assistance.

Figure 20 shows the source of all 22.9 million acre-feet of water that are consumed in the Tulare Lake Basin in an average year. Figure 19 shows the source of only the 13.1 million acre-feet that are delivered by conveyance infrastructure. Those different views can be useful depending on your interest.

California water managers traditionally exclude direct precipitation when accounting for sources of water. They generally focus on those water sources that are delivered through conveyance infrastructure like canals and pumps.
Figure 21 shows where the Tulare Lake Basin gets the water that we use (consume) in an average year; it is a summary of the data shown in Figure 19 for the years 1998–2010. This is the amount of applied water. This chart shows only the 13.1 million acre-feet of water delivered in an average year by conveyance infrastructure. It does not include the 9.8 million acre-feet of direct precipitation (see Table 12 and Table 14).

When excluding direct precipitation from the picture, our largest water source by far is groundwater withdrawals; these provide 46% of our water. All the rivers in our basin combined provide only 18% of what we use. The State Water Project and Central Valley Project together make up about 24%. That water is imported from rivers to the north of our basin.

But in a serious drought year like 1977, we have turned to groundwater withdrawals for 78% to 82% of our water supply.\textsuperscript{454, 455} This is illustrated in Figure 22.
Floods and Droughts in the Tulare Lake Basin
General Flood and Drought Notes

Description of the Groundwater Aquifer

The valley floor component of the Central Valley, the groundwater aquifer, is a large structural trough filled with poorly permeable marine sediments that are overlain by coarser continental sediments. The aquifer that underlies the San Joaquin Valley stretches from the foothills of the Sierra Nevada to the Coast Ranges, and from the Tehachapi Mountains to the western edge of the Sacramento–San Joaquin Delta. The story of this aquifer and the history of its use has been told in several USGS publications. A groundwater aquifer is the equivalent of an underground reservoir. The “underground reservoir” under the San Joaquin Valley is a large porous body of loosely packed alluvium (clay, silt, sand, and gravel) that is saturated with water. This groundwater aquifer functions much more like a sponge than like a surface reservoir. The water storage occurs in all the pore spaces of this alluvium.

The San Joaquin Valley aquifer averages about 2,400 feet in thickness, but increases from north to south to a maximum thickness of more than 9,000 feet near Bakersfield. The aquifer has two zones: a mostly unconfined upper zone (nearer the surface), separated from a confined lower zone by the Corcoran Clay. The water below the Corcoran Clay generally has distinctly different water chemistry from the water above the clay. The main source of groundwater in the San Joaquin Valley is the upper 1,000 feet of basin-fill deposits. The water that fills this alluvium comes largely from mountain rivers and streams on the east side of the San Joaquin Valley. These are the streams we are all familiar with that flow down out of the Sierra.

However, some of the ephemeral streams on the west side of the valley that flow down out of the Coast Ranges can also have large flow events in certain years. Some of these events in the ancient past resulted in huge volumes of water. Flows from these west-side streams have contributed to groundwater recharge on the upper end of the alluvial fans and have created groundwater bodies of distinct chemical types related to the soils and geology of the upper watershed. Examples are Los Gatos Creek and Warthan Creeks (tributaries of Arroyo Pasajero) in western Fresno County. Coalinga and the lands east around Interstate 5 have used groundwater from that system. The quality was not great, but usable for agriculture and industry until the California Aqueduct came along to provide a better supply.

Some of the water now stored in the alluvium flowed down from the Sierra and the Coast Ranges quite sometime ago. In that sense, some of the water being pumped up now is being mined from the distant past.

Very little of the runoff soaks into the ground as the rivers flow through the mountains and foothills; those areas tend to be underlain with granitic, metamorphic, and other relatively impermeable rocks. However, once the rivers emerge out onto the valley floor, they historically flowed over the highly permeable sand and gravel of the alluvial fans that soaked up the water. An alluvial fan is a distributary system made of multiple channels that allow for large areas of shallow inundation. This water then became part of the groundwater aquifer and continued moving between the sand grains.

This situation changed dramatically with the coming of irrigated agriculture to the San Joaquin Valley, beginning in about the 1870s. Settlers dug canals to tap the streams and rivers, spreading that water onto fields and orchards. This spreading of water over the surface of the ground permitted a steady agricultural growth that required more and more water. As more water was spread, less reached the valley’s wetlands and less soaked into the groundwater aquifer. Then the four federal reservoirs began operation during the 1954–61 period, allowing large quantities of water to be stored for use later in the year.

Today, rivers are largely cut off from the natural process of the routine spring flooding of the alluvial fans with the most permeable materials. Groundwater recharge has become largely dependent on irrigation events during the summer irrigation months. And to a lesser but important extent when there is surplus surface water by percolation into man-made recharge basins or by supplying surface water to areas that traditionally pump groundwater (in-lieu recharge).

The amount that can be percolated is substantially less than the natural process because the crops use some of the water, and because the most permeable areas on the alluvial fans don't receive as much total water for the length of time that the wet year inundation (overland flow) once caused. There are many areas of the San Joaquin Valley that cannot contribute significantly to the groundwater aquifer. The heavy Corcoran Clay layer is distributed throughout the central and western San Joaquin Valley. It varies in thickness up to 160 feet under the Tulare Lakebed. Ken Schmidt has found places in the valley where clay is vertically continuous for 1,000 feet; no decent water-bearing zones to be found. Similarly, older hardpan soils in many areas of the valley make little contribution to the groundwater.
It takes a long time for precipitation that falls on the surface of the Earth to percolate down to the wellhead hundreds or thousands of feet below. The average horizontal hydraulic conductivity in the Central Valley is about six feet per day.\(^{465}\) Shallow wells tap groundwater reserves that have percolated into the soil over recent decades or centuries. However, as wells have gone deeper, they are reaching much further back in time.

Tom Knudson has found that the deeper wells in the San Joaquin Valley are tapping groundwater reserves that were generally deposited 10,000 to more than 30,000 years ago.\(^{466}\) Similar ages have been found for groundwater being extracted from many desert basins, including Coachella Valley and Owens Valley. That water fell to Earth during the last ice age, the period of the Wisconsin glaciation.

Tapping such water is more than a scientific curiosity. It is one more sign that we are living beyond our sustainable water supply. The implication is that the current climate does not provide sufficient precipitation to meet our needs (the amount of water we choose to apply); we have to reach back to the last ice age to supplement our current available level of precipitation.

**History of the Groundwater Aquifer**

A 1989 USGS report by Alex Williamson and others developed a groundwater budget for the entire Central Valley that accounted for the relation between the surface-water and groundwater systems.\(^{467}\) During predevelopment conditions, the Central Valley groundwater aquifer was recharged by an average of about 2.0 million acre-feet per year. In the language of the water plan updates, that was natural recharge, the percolation to groundwater basins from precipitation falling on the land and from flows in rivers and streams. That natural recharge had of two components:

- Infiltration from precipitation falling on the land averaged about 1.5 million acre-feet per year.
- Infiltration from rivers and streams (aka streamflow losses) averaged about 0.5 million acre-feet per year.

An average of about 2.0 million acre-feet per year of groundwater was discharged (withdrawn) to streams, springs, or seeps, evaporated to the atmosphere, and transpired by plants. That groundwater discharge had two components:

- An average 1.7 million acre-feet per year of groundwater was evapotranspired.
- An average 0.3 million acre-feet per year was lost to gaining reaches of streams.

Prior to development, the Central Valley’s groundwater system generally was in equilibrium. Except for fluctuations caused by climatic changes, discharge was approximately equal to recharge, and the volume of water in storage remained relatively constant. At the time of Euro-American settlement, the groundwater table came right up to the surface in places. Even in dry years when the rivers were low, the groundwater aquifer still contained so much water that it spilled out into numerous springs.

When irrigated agriculture started in the San Joaquin Valley, this relationship began to change. The first canals in the Tulare Lake Basin were constructed in the 1870s. The first well in Tulare County was constructed in the 1890s.\(^{468}\) The early wells in the southern San Joaquin Valley were hand-dug pits. Wells could only be dug or drilled 20 to 30 feet deep until they hit the Monterey Shale formation. Because of the dropping water table, those wells did not last beyond 1910. In the early days of well-digging, there were more wells in California than in all the rest of the U.S. There were 300 to 400 documented wells in the San Joaquin Valley.\(^{469}\)

As well-drilling technology improved, wells could go deeper. This was one of the keys to tapping the Artesian Belt (see the section of this document on Artesian Wells — Discovery of the Artesian Belt). The valley’s artesian wells continued to flow naturally until they all stopped for uncertain causes during the first decade of the 20th century. The probable cause seems to be that the water table dropped as the rivers were diverted and the wells went deeper.

The first version of the deep well turbine (or centrifugal) pump was introduced by Bryon Jackson in 1901.\(^{470}\) It got to the Westside Sub-basin of the San Joaquin Valley fairly soon after that. The town of Tranquility (30 miles west of Fresno) had such wells for drinking water in 1920.

There were 794 pumped wells in Tulare County in 1909, and in 1919 there were 4,515; or nearly six times as many as ten years before. In 1909 there were 38,999 acres irrigated by pumps; in 1912, three years later, this type of irrigated acreage nearly doubled to 75,300 acres, and in 1921 there were 159,200 acres irrigated by the use of groundwater. In the space of 12 years, the area irrigated by pumps in Tulare County had increased 420%.\(^{471}\)
Floods and Droughts in the Tulare Lake Basin

General Flood and Drought Notes

When groundwater first began to be used for irrigation, it was commonly believed that the supply was inexhaustible; but as development proceeded and the draft on the groundwater increased, a lowering of the water table was noted in many areas. Around 1930, the development of an improved deep-well turbine pump and rural electrification enabled additional groundwater development for irrigation. This was right when the valley was suffering through some of the worst of the 1918–34 drought. In the 1930s, the hand-dug pits in the Poplar area (northwest of Porterville) began to run dry as the groundwater table dropped. The post-WWII period brought a tremendous increase in the amount of groundwater withdrawals for irrigation, resulting in an ever-increasing decline in the level of the groundwater aquifer.

In 1955, about one-fourth of all groundwater extracted for irrigation in the U.S. was pumped from the San Joaquin Valley. California’s groundwater has been described as one of the least-regulated, least-monitored aquifers in the U.S. While possibly an exaggeration, that gives a sense of the challenges involved in trying to get a handle on the situation with the Tulare Lake Basin’s aquifer. Water managers typically don’t have current status and trend data on that aquifer; they have had to make decisions based on assumptions that sometimes turn out much later to be incorrect.

In 2009, the USGS released a report on a five-year study of groundwater availability of the Central Valley aquifer. After 1900, when large-scale farming began in the Central Valley, water tables dropped significantly as wells were drilled to feed crops. Groundwater levels in the deep aquifer system (below the Corcoran Clay) in parts of the western Tulare Lake Basin eventually dropped by more than 400 feet compared with pre-1900 levels.

Waterhead levels in the area of greatest overdraft had been above ground level prior to 1900. But by 1961, they were well below sea level. Until 1968, irrigation water in those areas was supplied almost entirely by groundwater. As of 1960, water levels in the deep aquifer system were declining at a rate of about 10 feet per year. The severity of the groundwater overdraft in the San Joaquin Valley was part of the impetus for building the state and federal canal systems in the 1960s that divert water from the water-rich northern half of the state to the arid southern half.

The problem was framed at the time as one in which there was insufficient available dedicated and developed water supplies to meet our needs (the amount of water we chose to apply). If society were to provide us with additional dedicated and developed water supplies, then we would stop overdrafting the groundwater aquifer. See the section of this document on the California State Water Project and the Central Valley Project for a description of those projects. Huge pumping stations located in the Sacramento–San Joaquin Delta are used to export a portion of the northern waters to the south via those state and federal systems.

Beginning in 1950, water was diverted through the Friant–Kern Canal from below Millerton Lake to the east side of the San Joaquin Valley. In 1951, surface water deliveries along the northwest side of the San Joaquin Valley began through the Delta–Mendota Canal. In 1967, surface water deliveries to farms along the west side and near the southern end of the San Joaquin Valley began through the California Aqueduct. The availability of imported surface water following the construction of these canals resulted in a decrease in groundwater withdrawals.

Surface water imports from north of the Tulare Lake Basin allowed groundwater levels to recover by more than 200 feet in some areas on the west side of the valley by 1974. Increased use of the groundwater aquifer system during droughts and periods of reduced imports is reflected in DWR’s statistics on new well construction. At the end of the 1976–77 drought, surface water availability increased and fewer wells were drilled until the 1987–1992 drought. New well construction peaked in 1991, when more that 1,100 new wells were drilled in the valley. After the 1987–1992 drought, fewer wells were constructed until the reductions in surface water imports prompted additional well construction beginning in 2007.

As described in the section of this document on Historic Areas of Land Subsidence, groundwater levels are no longer a good measure of the amount of water in the groundwater aquifer. When the water was pumped out in the 1960s, much of the pore space collapsed; the ability of the aquifer to store water was lost. When the water levels recovered, the aquifer was no longer able to hold as much water as before.

In subsequent droughts like 1976–77, water levels have dropped comparatively fast considering the amount of water that was withdrawn. In areas of current subsidence, we are continuing to destroy the ability of the aquifer to store water.
Sustainable Yield of the Groundwater Aquifer

A groundwater aquifer is like an underground reservoir, a storage place for water. It is not a source of water like rivers and streams or direct precipitation. Much of the water in our groundwater aquifer was placed there by rivers flowing out of the mountains long before EuroAmericans began to settle the Tulare Lake Basin.

In the long run, we can only afford to withdraw as much groundwater from the aquifer as flows back in on average. As shown in Table 15 on page 120, the groundwater aquifer is recharged in two ways: recharge from surface applications and recharge naturally. Based on the best available data, those two sources together average about 4.8 million acre-feet per year. That is the maximum amount we can withdraw on average without further lowering the groundwater table. That is the sustainable rate of withdrawals from the groundwater aquifer.

The use of surface water and groundwater must be considered together. The term for this is “conjunctive use.” Conjunctive use is the coordinated management of surface and groundwater supplies to increase the yield of both supplies and enhance water reliability. The phrase “conjunctive use” is particularly used to describe the practice of storing surface water in a groundwater aquifer in wet years and withdrawing it in dry years. Reliance on groundwater substantially mitigates drought impacts for many urban and agricultural water users, and local water agencies have widely practiced conjunctive management of groundwater and surface water either formally or informally for many decades.

Summary of Groundwater Overdraft

In the Tulare Lake Basin, we have been using more water than we have dedicated and developed water supplies since early in the 20th century. Our needs (the amount of water we choose to apply) usually exceed our available water supplies. By that definition, we have been in a long-term drought over that entire period. (This is a socioeconomic drought, not a meteorological or hydrological drought. See the section of this document on What Constitutes a Drought for a description of the different types of droughts.)

We make up the difference by routinely tapping our groundwater aquifer to satisfy our unmet needs. As shown in Table 14 on page 119, groundwater withdrawals make up 46% of the total water that we use, basin-wide (excluding direct precipitation).

Not all groundwater is the same. Much of the groundwater is recharged from surface waters; when we pump from that, we’re pumping from a sustainable supply. As explained above, our sustainable rate of groundwater withdrawal is about 4.8 million acre-feet per year, based on the best available data. Yet as shown in Table 15 on page 120, our average groundwater withdrawal is 6 million acre-feet per year. That difference (1.2 million acre-feet per year) is roughly 21% of the total average groundwater withdrawal.

As shown in Table 15, that 1.2 million acre-feet per year in average groundwater overdraft represents 10% of the total water that we use, basin-wide (excluding direct precipitation). This has resulted in a significant draining of our groundwater aquifer. There are two different ways to frame the issue of groundwater overdraft. One way is to view it as a large and sustained basin-wide overdraft. We choose to apply more water, or average, than we have available in dedicated and developed water supplies. For some, that is self-evident. For others, that is a pretty bold statement. An alternative way to frame the issue is one in which there are insufficient available dedicated and developed water supplies to meet our needs (the amount of water we choose to apply). If society were to provide us with additional dedicated and developed water supplies, then we would stop overdrafting the groundwater aquifer. That is how the groundwater overdraft issue was framed in the 1950s.

In water year 2011 the State Water Project reported record-high water exports from the Sacramento–San Joaquin Delta, 4.90 billion cubic meters of water, the highest export rate recorded since 1981. The federal Central Valley Project exported 3.13 billion cubic meters of water in 2011, an increase from exports in 2008–2011, but comparable to exports from 2002–07. Translated into acre-feet, the total exports via the state and federal Delta pumps was 6,520,000 acre-feet in 2011 – 217,000 acre-feet more than the previous record of 6,303,000 acre-feet set in 2005.481

For perspective, that 6.5 million acre-feet is nearly twice as great (1.9 times to be precise) as the combined average annual flow of the two largest rivers in the southern San Joaquin Valley. The San Joaquin River produces a long-term average annual flow (measured at Friant Dam) of about 1.8 million acre-feet per year. The Kings River has an average annual flow (measured at Pine Flat Dam) of about 1.7 million acre-feet. Together, these two rivers produce an average of 3.5 million acre-feet of water.
Switching farms to this new surface water supply allowed groundwater aquifers to recover somewhat. For a while, it was assumed that water tables had stabilized after about 1970 because the groundwater overdraft was thought to have largely stopped. Unfortunately, we now know that this was not a safe assumption; the groundwater table did not stabilize at that time.

A 2009 USGS study led by Claudia Faunt, a hydrologist with the agency’s California Water Science Center, developed a numeric model of the Central Valley’s hydrologic system, the Central Valley Hydrologic Model (CVHM). The CVHM can be considered a tool for identifying, organizing, and integrating the necessary monitoring data. This model can be used to assess the effects of supply and demand (see the section of this document that discusses Water Demand). The CVHM can also be used to address groundwater depletion issues such as critically low groundwater levels and other consequences of groundwater storage depletion including subsidence, the effect of groundwater pumpage on streamflow and groundwater levels supporting marshes and lakes, and reduced availability of water for evapotranspiration. Managers can use the CVHM to help address the implications of different management options for water use.

As shown in Figure 23, groundwater storage in the Tulare Lake Basin showed a steep and fairly steady overall rate of decrease between 1961 and 2003, the period covered by the USGS study. The study found that the San Joaquin Valley experienced a net loss of 59.7 million acre-feet of groundwater storage during the 41-year period (1961–2003) covered by the study. For comparison, that is over 37 times the combined current capacity of the four federal reservoirs in the Tulare Lake Basin.

Groundwater overdraft in the San Joaquin Valley during the 41-year period covered by the USGS study was estimated to average more than 1.5 million acre-feet a year (59.7 million acre-feet / 41 years). While the northern and western parts of the San Joaquin Valley saw water level recovery during this period, the report found that “overall, the Tulare Lake Basin part of the San Joaquin Valley still is showing dramatic declines in groundwater levels and accompanying increased depletion of groundwater storage.”

The groundwater aquifer levels recovered somewhat after 2003. However, during the three-year 2007–09 drought, farmers relied so heavily on groundwater that it brought the groundwater aquifer down to near the historic low. See the section of this document on the 2012–15+ drought for a discussion of how the aquifer has continued to decline and the number of wells that are going dry during that drought.

Some recovery occurs after each wet year, but overall the trend has been down. For example, the California Water Institute at CSU Fresno recently produced a map of the San Joaquin Valley showing the change in groundwater depth from 1983–2009. The only areas of first water that did not decline during this period were generally in the Westside Sub-basin where salty perched water resides on the first clay lens below the surface. ("First water” is the depth to the first fully sustained saturated zone in the subsurface, regardless of quality.)

A NASA/UC Irvine study published in 2009 concluded that for the period from October 2003 through March 2009, the groundwater supplies of the entire San Joaquin Valley were depleted by an average of over 2.8 million acre-feet per year. The data covered a period consisting of one very dry year, two moderately dry years, and two wet years, an average mix in our region. The lead author of the 2009 study was Jay Famiglietti, the director of Hydrologic Modeling at UC Irvine. That 2.8 million acre-feet overdraft is nearly as great (80%) as the combined average annual flow of the San Joaquin and Kings River (1.8 + 1.7 = 3.5 million acre-feet of water).

Bill Tweed described the situation in a column that he wrote for the Visalia Times-Delta. In the Tulare Lake Basin, water for agriculture, cities, rivers, wildlife refuges, etc. comes from three sources:
1. The most sustainable and local of these three is the Sierra, mainly in the form of water from the Tuolumne, Merced, San Joaquin, Kings, Kaweah, and Kern Rivers.
2. The second source of our water is exports from the Sacramento–San Joaquin Delta, which is the immediate source of water that we import from Northern California via the Delta-Mendota Canal and California Aqueduct. That water, moved south at considerable expense, is increasingly fought over and hard to get. See the section of this document on the Role of the Endangered Species Act in Reducing Delta Exports for a discussion of some of the issues being fought over that limit our ability to increase Delta water exports.
3. The third source is what we pump from the groundwater aquifer, much of which is never replaced. A century ago, much of the valley had groundwater almost to the surface; artesian wells were common. Now, many areas have been mined for water to a depth of several hundred feet. The groundwater situation in Tulare County is particularly well studied and understood.
When surface supplies are inadequate to meet our needs (the amount of water we choose to apply), we pump out of the groundwater aquifer. When surface supplies allow, water districts work to recharge the aquifer. Such efforts form an important part of water management in the San Joaquin Valley.

In recent decades, it has become increasingly apparent that we are withdrawing more from the aquifer than we are returning, we have a long-term groundwater overdraft. Thanks to the NASA/UC Irvine study, we now know how badly we are failing to replace the water that we pump for agricultural and urban use. We would have to divert most of the water from our two biggest rivers in order to cover the shortfall.

In 2012, Bridget Scanlon and her colleagues at the University of Texas produced the highest-resolution picture yet of how groundwater depletion varies across space and time in California’s Central Valley and the High Plains of the central U.S. The authors of that report used water-level records from thousands of wells, data from NASA’s GRACE satellites, and computer models to study groundwater depletion in the two regions. The 2012 study built on the 2009 NASA/UC Irvine study. Both studies identified the southern areas of the Central Valley – particularly the Tulare Lake Basin – as facing the most dire groundwater issues.

Scanlon’s study painted a stark picture of how much water has been removed from the area’s groundwater aquifer. In the mostly drought-defined years of 2006–09, water users in the Tulare Lake Basin used enough groundwater to fill Lake Mead, the nation’s largest man-made reservoir.

As shown in Figure 23, about 70 million acre-feet has been lost from the Central Valley’s groundwater aquifer during the last 52 years (1962-2014).

During 2014, a team of scientists led by Jay Famiglietti, a senior water scientist at the NASA Jet Propulsion Laboratory, used data from NASA’s GRACE satellites to measure the amount of groundwater lost during the 2012–15 drought. The team found that the Central Valley lost about 11 trillion gallons (10 cubic miles) in the three-year period 2011–14. That is about 1.5 times the capacity of Lake Mead, the nation’s largest man-made reservoir.

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Figure 23. Amount of groundwater lost in past 52 years: 1962-2014.

Sources: Claudia Faunt, USGS, 2009; Jay Famiglietti, Center for Hydrologic Modeling, UC Irvine, 2014
Floods and Droughts in the Tulare Lake Basin
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As shown in Figure 23, about 73 million acre-feet has been lost from the Central Valley’s groundwater aquifer during the last 52 years (1962-2014). This is an average, long-term overdraft of 1.4 million acre-feet per year.

Famiglietti has compared the long-term decline in the Central Valley’s groundwater to a tennis ball bouncing down a flight of stairs — there are temporary bounces when rain is plentiful, but high water demand is ensuring that the overall direction is downward.

Groundwater levels appear to be sinking faster in the Central Valley than anywhere else in the U.S., according to a 2015 USGS report by Leonard Konikow. It is dropping three times faster than the much larger High Plains aquifer that stretches from Nebraska to Texas.

As shown in Table 15 on page 120, roughly 79% of our groundwater withdrawals are recharged either from surface water that has been applied to crops or by natural recharge (precipitation soaking into the ground, one way or another). However, roughly 21% of the groundwater withdrawals that we use is not being recharged by any of those sources. For decades, we have routinely used more water than is available in dedicated and developed water supplies, withdrawing the difference from the groundwater aquifer. This overdraft of the groundwater aquifer amounts to roughly 10% of the total water that we use, basin-wide (excluding direct precipitation).

As shown in Table 14 on page 119, about 46% of our region's water supply deliveries (excluding direct precipitation) comes from groundwater withdrawals (6 million acre-feet) in an average year. In a serious drought year like 1977 (see Figure 22), we have turned to groundwater withdrawals for up to 82% of our water supply.

By some estimates, the groundwater aquifer has been dropping an average of nearly 2 feet per year, basin-wide, in recent decades. The water table under Visalia has dropped an average of 3 feet per year over the last 25 years through 2013.

DWR uses monitoring well data to estimate the change in groundwater elevation for most of the valley portion of the Tulare Lake Basin. Their estimates showed that the groundwater elevation in our basin dropped by an average of approximately 3.5 feet per year between spring 2005 and spring 2010 (17.5 feet in 5 years).

That is equivalent to roughly 1.2 million acre-feet per year. For perspective, the Kings River has an average annual flow (measured at Pine Flat Dam) of about 1.7 million acre-feet. So we are depleting our groundwater aquifer at a rate equivalent to 70% of the flow of the largest river in our basin. As shown in Table 15 on page 120, this overdraft of the groundwater aquifer amounts to roughly 10% of the total water that we use, basin-wide (excluding 9.8 million acre-feet of direct precipitation).

DWR released a report on the state’s groundwater on April 30, 2014. That report found that groundwater levels are experiencing record historical lows in many areas of state, especially in the Tulare Lake Basin. In many areas of the San Joaquin Valley, recent groundwater levels are more than 100 feet below previous historical lows. The report found that the Kaweah and Kings sub-basins have the greatest number of deepened wells in an alluvial groundwater basin.

The Tulare Lake Resource section of the California Water Plan Update 2013 had this to say about groundwater overdraft:

During years of normal or above normal precipitation, or during periods of low groundwater extraction, aquifer systems tend to recharge and respond with rising groundwater levels. As groundwater levels rise, they reconnect to surface water systems, contributing to surface water base flow or wetlands, seeps, and springs. However, for much of the Tulare Lake Hydrologic Region, due to extensive pumping over the years the groundwater table has been disconnected from the surface water system for decades and provides no contribution to base flow. In 1980, DWR Bulletin 118–80 identified five of the seven southern San Joaquin Valley groundwater sub-basins (Kings, Kaweah, Tulare Lake, Tule, and Kern County), as being subject to conditions of critical overdraft. Thirty years later, things do not appear to have changed much. Although efforts have been made by local groundwater management agencies to reduce overdraft conditions in the region, a number of the groundwater management plans and more recent studies for these five key groundwater sub-basins acknowledge that groundwater overdraft continues.
The overdraft is not uniform throughout the Tulare Lake Basin; some areas have a very large overdraft while others have relatively little. However, taken as a whole, the Tulare Lake Basin has a large and sustained overdraft. The Tulare Lake Basin has by far the largest groundwater overdraft of any region in the state.

As described in the section of this document on Land Subsidence, 38 cubic miles of water have been removed from the Central Valley's groundwater aquifer over the past 150 years (since the 1860s). Most of that volume came out of the Tulare Lake Basin. That is the total amount of our groundwater overdraft since EuroAmerican settlement began.

That is a huge volume of water, equivalent to 130 million acre-feet — more than the volume of Lake Tahoe. It would cover the entire state of California to a depth of 15 inches.

**How Water Leaves our Basin**

As shown in Table 16 on page 121, about 21% of all the water used in our basin (4.8 million acre-feet) is recycled or reused water. It is applied twice. If you exclude direct precipitation (and only consider water delivered by conveyance infrastructure), then about 37% of all the water used in our basin is recycled or reused water. This water is used twice, then it disappears from our basin.

The remaining 63% of the water used in the Tulare Lake Basin (18.1 million acre-feet) is not reused. After this water is used once, it disappears from our basin. The Tulare Lake Basin has functioned largely as a closed basin since 1878 without a regular outlet to the ocean, essentially an inland sink. Surface water comes in, but it never flows out. Therefore, virtually all of this water must be leaving largely through exports or evapotranspiration.

Groundwater is like a bank account. Over the long run, we cannot afford to withdraw any more water from the underground aquifer than flows back into the ground. In that regard, it’s helpful to have a general accounting of some of the ways in which surface water leaves the Tulare Lake Basin. Surface water that leaves the basin has no opportunity to soak back into the aquifer.

Another way of looking at this issue is to think of the Tulare Lake Basin water supplies and where they go. Table 17 shows where the 13.1 million acre-feet of water comes from that is applied in the Tulare Lake Basin in an average year. We supplement our locally available supply (the 13.7 million acre-feet of precipitation that falls in our basin) with an additional 9.2 million acre feet of water. But at the end of an average year, none of that water is left over. (In fact, we are depleting our basin’s groundwater aquifer at an average rate of at least 1.2 million acre-feet per year.) Our basin has essentially sprung a huge leak. Therefore, it’s worth getting a handle on the accounting; where is the leak?

There are only two significant ways for water to leave a closed basin like ours:

1. **Exports of water outside the basin** (aka inter-basin transfers). This is very variable but seems to average less than one million acre-feet per year. Although Tulare Lake no longer overflows its sill, humans have altered the natural hydrography of the basin so that it is not completely closed. As detailed below, there are ways in which water leaves our basin.

2. **Evapotranspiration.** This is the sum of evaporation and plant transpiration from the earth’s land surface to the atmosphere. Evaporation accounts for the movement of water to the air from sources such as the soil, canopy interception, and water bodies. Transpiration accounts for the movement of water within a plant, and the subsequent loss of water as vapor through stomata in its leaves. Evapotranspiration in the Tulare Lake Basin is estimated to have doubled when we began irrigating crops on the valley floor during the summer. Some of this water vapor condenses and falls back within the basin as rain or snow.

But most of this moisture has been leaving the basin unseen. However improbable it may seem, that turns out to be how most of the water is leaving our basin; that is how we have sprung our huge leak. That means our basin is losing upwards of 22.9 million acre-feet through evapotranspiration in an average year. Plants are thirsty in our hot, arid environment. Agricultural irrigation in the Central Valley doubles the amount of water vapor pumped into the atmosphere. Water vapor flowing west over the Sierra from July through mid-September joins with the North American Monsoon water cycle. In a study published in 2013, Min-Hui Lo and Jay Famiglietti at the Center for Hydrologic Modeling at UC Irvine used a global climate
They demonstrated how this water vapor export affects the regional hydrological cycle of the Southwest. The result is that summer precipitation in the states affected by the monsoon is increased by 15%, the Four Corners region experiences a 56% increase in runoff, and runoff in the Colorado River Basin increases by 28%. All percent differences are the differences between applying irrigation to the Central Valley and not applying it.

This supplement to Southwestern precipitation by Central Valley irrigation has only existed since EuroAmerican settlement of our area. Furthermore, the water vapor exports sent over the Sierra are so great in part because we have found ways to supplement the amount of water available for irrigated agriculture in the Central Valley; we are not limited to the direct precipitation that we receive. Only 29% of the water applied in the Tulare Lake Basin in an average year comes from the direct precipitation that we capture (deliveries from local rivers plus groundwater withdrawals from natural recharge). As shown in Table 17, the remaining 71% comes from other sources (imports from north of our basin, reused and recycled water, and groundwater overdraft).

Just as some water users on this side of the Sierra may see the total amount of water that we currently apply as what they are entitled to, water users in the Southwest may have grown accustomed to the flow of water vapor that we send their way as the norm or what they are entitled to. If water users in the Central Valley were to reduce our use (such as by reducing the groundwater overdraft), this would reduce water vapor exports which would reduce rainfall in the Southwest. We have effectively created a culture of dependency.

We have a relatively good handle on exports because they are visible and can be gaged:

- Diverting Kings River floodwaters to the San Joaquin River in order to minimize flooding in the Tulare Lakebed. This is done using the North Fork / Fresno Slough / James Bypass channel. The James Bypass began operation in 1872, and the capacity of the associated system has since been increased several times. Prior to about 1872, all of the Kings River water flowed into Tulare Lake. See the section of this document on Pine Flat Dam for a more detailed description of the James Bypass. Water that is sent through this system winds up in San Francisco Bay; it is essentially a loss from the point of view of Tulare Lake Basin water users. With the construction of Pine Flat Dam in 1954, the need to divert water through this system was greatly reduced. Even so, diversions through this system have occurred in 38% of the years since the dam was completed.

- Diverting Kern River floodwaters into the California Aqueduct rather than into Buena Vista and/or Tulare Lakes. This is done using the Kern River Intertie and Cross Valley Canal. Once the water enters the California Aqueduct, it is pumped over the Tehachapi Mountains and sent to the Los Angeles area. (Los Angeles is always happy to receive high-quality water from the Tulare Lake Basin, especially when they only have to pay shipping costs.) The Kern River Intertie was completed in 1977. Prior to that, a big flood on the Kern would first fill Buena Vista and/or Goose Lakes, and then spill into Tulare Lake. The Kern Intertie was used to make large diversions to the Los Angeles area during the 1983 and 1998 floods. Lesser diversions may have been made in other floods. The shipping costs to get water to Southern California are not insignificant. We all see the big pumping plant and pipes west of Interstate 5 when we drive the Grapevine. The State Water Project (SWP) is the largest single consumer of energy in California with a net usage of 5.1 million mWh. The SWP consumes so much energy because of where it sends its water. To convey water to Southern California from the Sacramento–San Joaquin Delta, the SWP must pump it 1,926 feet over the Tehachapi Mountains, the highest lift of any water system in the world. Pumping one acre-foot to the region requires approximately 3,000 kWh. Southern California’s other major source of imported water is also energy intensive: pumping one acre-foot of Colorado River Aqueduct water to Southern California requires about 2,000 kWh. It requires an average of more than 9,000 kWh to move a million gallons of water to Southern California. The Metropolitan Water District of Southern California estimates that the amount of electricity used to deliver water to residential customers in Southern California is equal to one-third of the total average household electric use in Southern California. Twenty percent of all the energy in the state of California is used to move water.

- Transferring water from the Kings, St Johns, and Tule Rivers into the Friant-Kern Canal in order to minimize flooding in the Tulare Lakebed. This is done by using pumps at the point where each of those rivers crosses the canal. Once the river water enters the canal, it flows by gravity to the canal’s terminus near Bakersfield. There it is emptied into the Kern River. The water is then routed to the Los Angeles area using the Kern River Intertie and Cross Valley Canal as described above. A combined total of over 472,000 acre-feet of floodwaters was pumped into the canal during the years 1978, 1980, 1982, 1983, 1986, 1995, 1997, 1998
and 2006. (The total amount may have been a good bit more than this; records are incomplete.) Transfers may have been made in later years as well. Including all sources (four rivers), exports to areas south of the Tehachapis occurred in 30% of the years since the Kern River Intertie began operation in 1977.\textsuperscript{505} As Peter Vorster documented in the 2007 EPA report on the Tulare Lake Basin prepared by ECORP Consulting, smaller amounts leave the basin via the Arvin-Edison intertie and return flow from the Fresno Irrigation District irrigated lands and stormwater from Fresno into the San Joaquin River.\textsuperscript{506} Some Delta water imports is recharged into the Kern County groundwater banks and later exported out of the basin. We don't have anything like a full grasp of just how large the total exports were for most years. However, Table 18 gives examples of what the numbers add up to for three of the larger years for water exports from the Tulare Lake Basin.\textsuperscript{507}

<table>
<thead>
<tr>
<th>Source:</th>
<th>Kings River</th>
<th>Kings/Kaweah/Tule</th>
<th>Kern River</th>
<th>Total Exports</th>
</tr>
</thead>
<tbody>
<tr>
<td>Route:</td>
<td>James Bypass</td>
<td>Friant-Kern Canal</td>
<td>Kern Intertie</td>
<td></td>
</tr>
<tr>
<td>To:</td>
<td>SF Bay</td>
<td>LA</td>
<td>LA</td>
<td></td>
</tr>
<tr>
<td>Year</td>
<td>(million acre-feet)</td>
<td>(million acre-feet)</td>
<td>(million acre-feet)</td>
<td>(million acre-feet)</td>
</tr>
<tr>
<td>1969</td>
<td>1.6</td>
<td>not an option</td>
<td>not an option</td>
<td>1.6</td>
</tr>
<tr>
<td>1983</td>
<td>2.3</td>
<td>.2</td>
<td>.8</td>
<td>3.1</td>
</tr>
<tr>
<td>1998</td>
<td>1.0</td>
<td>.2</td>
<td>.1</td>
<td>1.3</td>
</tr>
</tbody>
</table>

Total exports of at least 700,000 acre-feet also occurred in 1978, 1980, 1986, and 2006. There are many years for which no data are available.

That isn't to say that exports (or inter-basin transfers) are necessarily a bad thing; just that we need to recognize that such diversions have consequences to the groundwater aquifer. Every time that water is transferred out of the Tulare Lake Basin, it's that much less water available for applied use or for groundwater recharge. The pressure to make certain types of diversions may increase in the future.

Potential for a Sustainable Water Supply

We always want more water than we have. However, in concept most of us could agree that our goal is something similar to sustainable groundwater extraction. Our basin currently uses more water than we have. Our average needs (the amount of water we choose to apply) exceed our average supply. As shown in Table 15 on page 120, our basin has been overdrawing the groundwater aquifer by an average of at least 1.2 million acre-feet per year in recent decades. We are effectively in drought conditions most of the time, even when precipitation is above average.

We consistently, decade after decade, use more water than our available surface supplies. The implications of this are hard to escape. Even if our temperatures and snowpack were to remain stable, the groundwater table in the Tulare Lake Basin will continue to drop due to the overdraft. One of the consequences of this is that the wells will continue to go ever deeper, and the cost of pumping will continue to rise. This race to the bottom is not sustainable. Eventually groundwater withdrawals will be limited by supply and demand. Agriculture (the valley’s single biggest water user) will be forced to reduce its reliance on groundwater.

The San Joaquin Valley groundwater system is the largest storage reservoir south of the Sacramento–San Joaquin Delta and, therefore, has a large potential interest from a statewide perspective. The groundwater overdraft in the Tulare Lake Basin is arguably due to a combination of using too much groundwater, not recharging enough, and not conserving enough. Many groups are working on different aspects of these various problems.

Increasing surface storage

One proposed solution is to build more water storage. This can be useful in helping to get through short-term droughts. However, it does not help with longer droughts and it does relatively little to address the issue that our average water use is greater than our average water supply.

Temperance Flat Dam is a proposed dam project on the San Joaquin River west of Auberry. The dam would be located at the back end of Millerton Lake and would inundate the area known as the San Joaquin River Gorge. It would inundate several PGE hydro plants. The primary purpose of the project is to increase storage capacity in the upper San Joaquin River Basin. Under the current proposal, Temperance Flat would add 1.3 million acre-feet of additional storage over and above the existing storage capacity of Millerton Lake at 525 thousand acre-feet, bringing total storage capacity to 1.8 million acre-feet.
The project would provide an estimated 61–76 thousand acre-feet per year of additional water. That additional water represents a little over 2% of average surface runoff of the Tulare Lake Basin and about 1% of the yield of the CVP. It has the potential to reduce groundwater recharge because it would turn the inexpensive surplus surface water that is recharged in wetter years into more expensive regulated water.

The current estimated cost of construction is $2.5 billion with an estimated annual operating cost of $115 million. Temperance Flat is one of three major storage projects that are being considered as potential candidates for funding by the $7.5 billion Proposition 1 water bond that was passed by voters in November 2014. The other two projects under consideration are the proposed Sites Reservoir (located in the Sacramento Valley west of Colusa) and the raising of Shasta Dam.

Much of our current surface storage is in the form of the Sierra snowpack. Extensive material has been published about the expected impacts of global climate change in California. As described in the section of this document on Long-term Temperature Changes, more than half of the Sierra snowpack is projected to be lost by the end of the century, with notable impacts being observed by mid-century. Runoff will occur earlier in the year rather than being held back in the Sierra snowpack. That will be the equivalent of lost surface storage.

**Increasing Delta imports**

Another proposed solution is to increase our water supply by increasing Delta water imports. Some advocates of this approach envision that it is as simple as getting politicians to allocate a greater percentage of the available water to our basin; taking it away from some other use. However there are significant obstacles to increasing our share of that water supply. There is only so much water in California, and our basin has to compete with other interests for it.

Intrusion of brackish water into the Sacramento–San Joaquin Delta is a recurring natural phenomenon; however, it became a serious issue after the development of agriculture in the upper Sacramento and San Joaquin Valleys reduced inflows. Multiple droughts between 1910 and 1940 caused significant salinity intrusion in the Delta because of the reduction of freshwater inflows. The growing Delta water quality issue provided the initial impetus for building dams on Central Valley rivers to boost dry-season freshwater flows. This eventually became the CVP.

Present-day Delta flow and water quality requirements arise from the original water rights granted by the SWRCB to the SWP and CVP to divert water upstream of the Delta, thereby raising the salinity of water used by in-Delta users. See the section of this document on Role of the Endangered Species Act in Reducing Delta Exports for a discussion of some of the issues being fought over that limit the ability of south-of-Delta water users to increase Delta water exports.

On January 31, 2014, the SWRCB approved an emergency rule change (a Temporary Urgency Change Petition or TUCP) to modify the conditions of the water right permits and licenses for the SWP and the CVP. The approval temporarily modified Delta flow and water quality requirements for these projects. By temporarily modifying the conditions of the water rights, the SWP and CVP were able to export more water to south-of-Delta water users than their permits and licenses otherwise allowed. This additional Delta pumping in 2014 and 2015 from the relaxation of the standards was primarily to more fully satisfy the San Joaquin River Exchange Contractors senior water rights. It did not increase the allocation of water to Tulare Lake Basin users.

The SWRCB accepted over 12 hours of often emotional testimony on February 18, 2015 about how to manage the dwindling supply of water in the Delta in 2015. SWP and CVP had requested changes to water right requirements for their projects so that they could export upward of 80,000 extra acre-feet of water in spring 2015 much as they had in 2014. South-of-Delta water users strongly supported this request; they really needed more water. However, Delta farmers and environmentalists strongly supported keeping that water in the Delta for the benefit of the environment; they felt they had an equally strong need for the water.

Future Delta water imports to our basin will probably be one of opportunistic flows, rather than dependable steady-state flows. Making the best use of this resource will require unprecedented cooperation and coordination of surface water and groundwater users south of the Delta. That will be challenging. However, given the right mix of infrastructure and institutional arrangements, the San Joaquin Valley groundwater system conditions could improve significantly compared to present conditions.

It has become popular to say that liberal environmentalists have been responsible for preventing construction of water conveyance systems that would have brought a more reliable water supply to Southern California. The
issue is more complicated than that. There has been little construction of major dams or water conveyance systems since a voter referendum that failed in 1982. California voters rejected a proposal to build a 43-mile-long, 500-foot wide water conveyance system called the Peripheral Canal that would have improved the quality and quantity of water delivered to Southern California. Environmentalist groups opposed the proposal, but the major opposition — and the majority of money funding the opposition — came from California’s farmers. Governor Brown has placed a high priority on implementing the proposed Twin Tunnels Project, a component of the Bay Delta Conservation Plan. The tunnels are not likely to increase the average imports to the San Joaquin Valley. However, they would increase the reliability of imports, and that would be very beneficial.

Delta imports can be brought in during the dry season when they are needed, but water supplies in the Delta are least available at that time. An alternative approach is to import some water during the wet season and store it for use later when it is needed. Imported water could be stored either in surface reservoirs or in groundwater banks.

One proposal for increasing Delta water imports is the Citizen’s Water Plan, formulated by Rob Simpson, Steve Haze, and others. That plan proposes to increase imports during the wet months and store the water in the southern part of the Tulare Lakebed near the present South Wilbur Flood Area, south of where farming currently occurs, just north of Sand Ridge. It would then be used to recharge the groundwater aquifer and distributed to farms during the dry months.

This proposal has many advantages. However, water engineers Dick Schafer and Dennis Keller have observed that there are also serious problems with it. The water would be stored at the low point in the Tulare Lake Basin, losing the advantage of gravity. Pumping would be required to get the water up to farms. The proposed reservoir in the Tulare Lakebed would be wide and shallow, so evaporation would be relatively high. It could also encourage algae growth which would cause problems for pumping. Water quality in the lakebed has historically been poor.

Project proponents acknowledge these shortcomings, but believe that they can be addressed. Furthermore, they believe their proposal has fewer shortcomings than the proposed Temperance Flat Dam and the Bay Delta Conservation Plan with its Twin Tunnels.

Agriculture users in the San Joaquin Valley have seen a decrease in Delta imports over time, especially during droughts. This has happened for two reasons:

- Less water is being exported from the Delta. As shown in Table 10 and Table 11, less total water has been exported south of the Delta in dry years since 1990. See the section of this document on the Role of the Endangered Species Act in Reducing Delta Exports for a discussion of some of the issues being fought over that limit our ability to increase Delta water exports.
- Urban water interests, especially those in Southern California, have been taking a larger share of the water sent to south-of-Delta users in recent decades. In the first 20 years of the SWP, the majority of deliveries were for agriculture; in most of the last 20 years, the majority of the deliveries have been for urban use. Southern California has three sources of water: the Delta, Owes Valley, and the Colorado River Basin. During earlier droughts, they were able to rely on water from the Colorado River in excess of the state’s basic interstate apportionment — Lower Basin water that was either hydrologically surplus or unused apportionment of Nevada or Arizona. This additional supply helped protect the Metropolitan Water District of Southern California (MWD) service area against shortages and allowed MWD to participate in exchange agreements to assist other agencies in the Tulare Lake Basin and elsewhere that were experiencing critical shortages. Drought in the Colorado River Basin and increasing water usage by Nevada and Arizona has brought this era of additional supplies to a close. As a result of such decreasing water supplies elsewhere coupled with population increases, urban water users have been taking a greater percentage of the south-of-Delta exports.

Increasing water yield
As shown in Table 12, the Tulare Lake Basin has 9.8 million acre-feet in direct precipitation. There is potential to convert some of that precipitation into dedicated and developed water supplies.

Due to fire suppression, there is currently a greater density of shrubs and trees in the upper watersheds than there used to be; than there was before fire suppression. The amount of overstocking differs by species and by area, but perhaps it is something on the order of 20%–40%. In any case, it is thought that most forest types would be more resilient to drought, fire, insects, etc. if they were significantly less dense than they currently are.
Roger Bales (a hydrologist with UC Merced) and Scott Stephens (a professor of fire science at UC Berkeley) were interviewed on NPR’s Morning Edition on October 14, 2014, about use of water by conifers and the effect of fire suppression. Bales said their hypothesis was that if there were half as many trees, there would be 20%-40% more runoff. Stephens said that the water piece of fire suppression is really huge; it is underappreciated, but it is massive. A recent study from UC Irvine found California’s forests will be using even more water by the end of the century because warming temperatures will make the growing season longer.

The term “potential evapotranspiration” can be thought of as the amount of evapotranspiration that occurs in much of the forests now, with an unnaturally dense stocking rate. The term “actual evapotranspiration” represents the amount of evapotranspiration that would occur if forest density were reduced to a lower stocking rate. If this difference between potential evapotranspiration and actual evapotranspiration were maintained, then the increased runoff described by Bales and Stephens could be realized.

Reducing water use
During periods of severe drought and decreased water availability, agricultural water users tend to fallow land, decreasing the acreage that is under irrigation. When more water becomes available, more acreage is brought under irrigation. The land easiest to remove from irrigation is that land which has been planted in annual crops. There is a greater investment in permanent crops such as trees, grapes, or asparagus, so it is a greater financial cost to remove that land from irrigation.

As water has become more expensive, there has been greater motivation for farmers to convert their land from row crops such as tomatoes and cotton to more valuable permanent crops such as pistachios, almonds, and grapes. (Or the motivation may have been simpler: just to increase profits.) This conversion of agriculture land from row crops to permanent high-value crops can make financial sense and can be the best use of the water available to a farmer. (For example, almonds offer price stability and consistent profitability. They are second only to wine grapes for having the highest crop value per unit of water.) But the effect is to harden water demand so that these lands cannot readily be fallowed during periods of drought.

In a drought, farmers can choose not to plant row crops if they aren’t going to get enough water that year. However, they cannot fallow orchards. Water is needed year-round to keep trees alive, and those trees are a significant investment. For example, you have to provide water to new pistachio trees for 7 years before you see production; but then they live for upwards of 100 years, so water demand is hardened for that time. Water reliability is a primary concern when a farmer makes the decision to convert land from row crops or rangeland to permanent crops.

As shown in Figure 19, the total amount of water delivered by conveyance infrastructure in the Tulare Lake Basin typically varies less than 5% from year to year even though our precipitation varies dramatically. When surface supplies are inadequate to meet our needs (the amount of water we choose to apply), water users rarely cut back significantly on total water use. Instead, they pump more heavily from the groundwater aquifer to make up the difference, always applying about 10% more water than the available sustainable water supply.

If this approach could be changed, we could come much closer to the goal of a sustainable water supply. There are at least two different ways that this approach could be applied:

- The amount of applied water could vary each year to equal the available sustainable supply. If the available supply in a given year decreased by, say, 20%, then the amount of applied water would decrease by 20%.
- Alternatively, the amount of water applied could be roughly the same in each year, much as shown in Figure 19. However, it would be reduced to equal the average sustainable supply. Under this approach, the groundwater aquifer would be recharged in a wet year. In a dry year, there would be increased withdrawals from the aquifer. This would effectively be water banking.

If either of the above approaches were taken, this would eliminate the groundwater overdraft. That would achieve the goal of sustainable groundwater management. However, that would require a fairly significant reduction in the amount of irrigated acreage in the Tulare Lake Basin along with greater efficiencies in applied water by reducing irrecoverable losses.

Increasing groundwater recharge
As described in the section of this document on Groundwater Overdraft, rivers today are largely cut off from the natural process of the routine spring flooding of the alluvial fans with the most permeable materials. There are several projects currently underway to reconnect the more permeable alluvial soils with more water for groundwater recharge.
A 2011 state law (AB 359) promotes the management and protection of the state’s groundwater supplies by requiring local water agencies to map groundwater recharge areas and to submit those maps to local planning agencies.

Some Integrated Regional Water Management (IRWM) groups are believed to have mapped their recharged areas. Eric Osterling with the Kings River Conservation District said that a California Water Foundation study will soon be released on this subject.

Investments in utilizing such lands for recharge are underway in some areas. Some of the more promising projects seek to work with farmers to have them use their land in wet years with the promise to avoid or mitigate any damages to crops.

Many water districts have constructed shallow ponds (aka recharge basins) to put excess water into the groundwater aquifer. These recharge basins somewhat mimic the intermittent wetland habitat that once existed on the alluvial fans historically. As an example, the Kaweah Delta Water Conservation District has some 40 recharge basins that total approximately 5,000 acres.515

Recharge basins such as these are typically small and somewhat widely distributed. They are generally located over the alluvial fans because that is where the rivers are. The water that gets stored in the groundwater aquifer is retrieved during dry years by individual wells operated by agricultural users and cities. It is low tech, but it works.

Water banking takes this concept one step further. It combines wells with the recharge basins; all of which are concentrated in one place and operated by the water district. The concept is that the water placed in the groundwater aquifer won’t move significantly; it will be there to be retrieved when desired. It uses the aquifer as the equivalent of an underground reservoir. Kern County is the lead in water banking. Virtually all of the water districts in the San Joaquin Valley portion of Kern County are involved in water banking at some level.

The biggest water bank in the San Joaquin Valley is the Kern Water Bank. It covers about 20,000 acres of the Kern River’s sandy alluvial fan southwest of Bakersfield and has about 7,000 acres of recharge basins. Up to 72,000 acre-feet per month can be recharged at the beginning of a recharge program, a rate that gradually declines as the recharging progresses.516, 517

When Secretary of the Interior Bruce Babbitt visited the Kern Water Bank in January 2000, he hailed it as “the most effective groundwater storage program in the United States, probably the whole world”. Water recharged by the Kern Water Bank comes primarily from the State Water Project via the California Aqueduct, but some floodflows from the Kern River and the Central Valley Project have also been recharged.

Most of the water we recharge should be thought of as very short-term storage; it is usually extracted the same year. As shown in Figure 19, it is the rare year when precipitation is great enough to recharge the groundwater aquifer for as much as a year. Increasing groundwater recharge only provides a net increase in total dedicated and developed water supplies or other long-term benefits in the following situations:

- Captures water that would otherwise evaporate before plants can use it.
- Captures floodflows that would otherwise be exported outside of the basin.
- Allows for increased imports into the basin by providing storage space.
- Recharge aquifer to keep it from being occupied by brackish water. Many groundwater basins, including the Tulare Lake Basin, are less limited by their total volume than by their quality. Our basin is a mixture of pockets of high-quality water and brackish water. In some areas, when a pocket of high water is emptied, it is at relatively high risk of being filled by migration of an adjacent pocket of brackish water. Periodically recharging the aquifer with high quality surface water reduces the risk that the brackish groundwater will migrate into the emptied space. As long as we periodically refill the emptied space with high-quality water, we can keep using it. But once we allow brackish water to migrate into that space, we lose the ability to use it for storage. It is all about the axiom of nature abhorring a vacuum; nature will tend to fill it with something, but it may not be what we want.

Water conservation
One opportunity available to us for increasing our available water is through water conservation. All of us have a role to play in conservation (homeowners, agriculture, federal land managers). It is always worth stepping back and looking at the overall system. Where are the big potential savings?
Floods and Droughts in the Tulare Lake Basin
General Flood and Drought Notes

As described in the section on How Water Leaves our Basin, the Tulare Lake Basin loses upwards of 22.9 million acre-feet of water through evapotranspiration in an average year. If we found ways to reduce evaporative losses (or to make better use of that water before it evaporates), that would be the equivalent of increasing our water supplies. Some of the ways that this could be done (and has been done), include:

- Planting crops that make more efficient use of water. The concept of crop water use efficiency (WUE), determined by the ratio between the marketable yield and the seasonal values of actual evapotranspiration, has become a tool for analyzing the strategies that allow attaining the best use of water in agriculture. Water can be used more efficiently by selecting species and varieties that have higher WUE values.
- Likewise, crop WUE values can be increased by improving agro-techniques (irrigation techniques, water quality, altering soil structure and drainage patterns, mineral supply and water quality.
- Growing crops inside commercial greenhouses. This significantly reduces evapotranspiration loss.
- As shown in Table 16 on page 121, only 59% of net water use (the amount of water actually consumed) is applied a second time. The rest of that water is applied once, then leaves our basin via evapotranspiration. There may be opportunities to increase the amount of water that is applied a second time before it is leaves our basin. Some water managers are doing just that.
- Capturing excess floodflows which would otherwise be lost, and banking them for future use.
- Capturing water which would otherwise be exported from our basin, and banking it for future use.

Accomplishing any of those tasks would be the equivalent of building more water storage or securing additional water supplies from north of our basin. Although not flashy, those all offer practical ways of getting us closer to achieving the goal of providing a more sustainable water supply.

Water conservation is a worthy practice that is well worth pursuing. However, from a basin-wide perspective, water conservation doesn’t provide any extra water. That is, water conservation doesn’t result in our having any additional water left over, or in our using any less water as a basin. Whatever water we conserve will most likely go into expanding the amount of irrigated cropland.

Agricultural water users aren’t like urban users. Because of their organizational structure, urban water users can act as a group and truly conserve water. That is what happened in Los Angeles.

But we shouldn’t think of agricultural users as potential water conservers in that sense:
- First, that is because agricultural users act individually based on their own interests rather than as a group. It is an attitude of use it or lose it. If they don’t use the water, their neighbor will use it.
- Second, if agricultural users save water, there is often excess land available that is not under irrigation due to lack of water. Therefore, they apply the water that they save to irrigate the acreage that had been uncultivated due to lack of irrigation water.

It is not that agricultural users are incapable of reducing water use. They can and do reduce water use when they don’t have enough water. However, if they have a reliable supply of water, they will typically take savings gained from the application of efficiency measures and use that water to increase overall production. It is not in an individual agricultural user’s best interest to give up water to someone else.

For an agricultural user, it is a tough sell to take water rights away in exchange for saving water. Those water rights mean the opportunity to make more money. In domestic and industrial use, if you can continue your quality of life or a business’ net earnings while using less water, it is not a difficult sell to give up water to somebody else. 518

Requirement to Manage Groundwater Sustainably

Existing law authorizes local agencies to adopt and implement groundwater management plans which are required to contain specified components. Local agencies that seek state funds administered by DWR for groundwater projects or groundwater quality projects must do certain things, including preparing and implementing a groundwater management plan that includes basin management objectives for the groundwater basin. There are 27 groundwater management plans for the Tulare Lake Hydrologic Region. 519

While this creates a structure for managing groundwater, it does not achieve the goal of managing groundwater sustainably. California was the last Southwestern state to pass a law to begin managing groundwater. The current unregulated groundwater pumping situation in California is one in which people acting rationally for their own self-interest deplete a common resource. It is a classic tragedy of the commons. People have the right to pump water underneath their own land until their neighbor comes along with a deeper well or a bigger motor and suck the water from under their property. Then they have to drill a deeper well. This race to the bottom is not sustainable.

Of the more than 400 groundwater basins in California, the law is pointed at 127 as high- or medium-priority targets for regulation. Many have overdraft and contamination problems.

Without some regulation to share the underground water, water users have been drilling deeper and deeper for water as wells go dry. The new law is intended to help to bring that under control. However, it will take five years just to come up with plans to manage those groundwater basins.

The law requires regions to form their own groundwater sustainability agency or agencies (GSAs) by January 1, 2017. By 2020, each GSA will have to present the state with a groundwater sustainability plan that is adequate to manage, restore and protect the groundwater in their region. The GSAs aren’t required to achieve a groundwater/surface water balance until 2040.

“It’s really, we’re talking about after 2020,” he said. “Until then, it’s whoever’s got the biggest pump or the deepest well.”

Each plan will designate improvement targets every five years for 20 years. The law gives all the power to the local groundwater agencies to figure out the best way to achieve those targets. Water users at the local level will have to work together to figure out how to achieve the goal of sustainable groundwater extraction. In concept, almost everybody agrees that sustainability is a good idea. However, some water users don’t like the idea that there might be extraction limits or fees.

The legislation contains a provision allowing the state to eventually step in and limit withdrawals if a local groundwater agency fails to act. Some water users object to that provision. Congressman Devin Nunes said that he views the Sustainable Groundwater Management Act as part of a plan by extreme environmentalists to remove 1.3 million acres of San Joaquin farmland from production by limiting the availability of water.

Groundwater Management in Fractured Bedrock Aquifer

As explained in the section of this document that describes the Groundwater Overdraft, the valley portion of the San Joaquin Valley is underlain by an immense groundwater aquifer of loosely packed alluvium (clay, silt, sand, and gravel). But those deep sediments end at the edge of the valley floor. In the foothills and mountains to the west of the valley, the groundwater is found primarily in fractured bedrock.

In 2013 and 2014, DWR conducted a pilot water supply study: Improving Groundwater Management in the Southern Sierra Fractured Bedrock Aquifer. John Kirk, engineering geologist with DWR's Fresno office, led this study effort. This was a groundwater and watershed study focused on the community of Three Rivers. That community was chosen because of its central location and community interest in learning about the local and regional water supply.

Bobby Kamansky wrote up the interim results of that study in The Kaweah Commonwealth. The study seeks to understand the quantity and quality of the water in the Kaweah River Basin as it flows from the Great Western Divide and nine smaller watersheds and into the underground aquifer. The study examined several aspects of the area: local geography, geology, land use, precipitation, hydrographs, and water demand. Geology plays an important role in water quantity and quality because much of the area’s water comes from water stored in the fractures and fissures of the hard rock aquifer.

The study examined the logs for 500 wells (about half of the wells in the area), which offered valuable information about the geology where the wells are drilled and the water quality and quantity in the wells.

The study identified that 54% of the land in the watersheds was in public ownership (national parks and BLM), while 46% was privately owned. The study area has a total of 1,575 parcels, 81% of which are smaller than 10 acres and located primarily along Kaweah River tributaries.

The study examined water demand and compared it to water availability at different times of year. The water flow along the Kaweah River in Three Rivers is subject to dramatic increases in March through June when abundant water flows from snowmelt and swells rivers and fills underground rock fractures.
Floods and Droughts in the Tulare Lake Basin

General Flood and Drought Notes

The average precipitation in the Kaweah River Basin ranges from over 55 inches at the crest of the Great Western Divide to a low of 14 inches at the base of the local watershed near Lake Kaweah with an average of 22 inches for the whole watershed. The study estimated that the watershed recharges an average of approximately 4 inches of that 22 inches of average precipitation across the watershed.

Three Rivers residents collectively use over 300 acre-feet of groundwater each year. With a population of more than 2,000 residents and approximately 1,000 households, annual use in Three Rivers is estimated to be 110,000 gallons (0.34 acre-feet) per resident. Three Rivers residents use 200 gallons per day in the winter and nearly 500 gallons each day in the summer.

Unfortunately, the greatest water demand occurs when the quantity of water in the river and in the groundwater aquifer is lowest. In other words, demand is highest when supply is lowest.

In Three Rivers, 10% of wells surveyed were less than 50 feet, 22% had depths of less than 100 feet, but the majority of wells (68%) were between 100 and 500 feet. Half of the wells had yields greater than 15 gallons per minute, 42% of the wells had yields between 2 and 15 gpm, and 8% of the wells yield less than 2 gpm. Very few dry wells were reported from the watershed around Three Rivers, while over a thousand wells in the valley went dry in 2014.

Water quality in the Kaweah River Basin is best at the highest elevations where precipitation is greatest and recharge exceeds water demand. In and around Three Rivers, a number of wells have poor water quality, including high salt content (exceeding drinking water standards), sulfur, hydrogen sulfide, or bacteria.

The information in the study provides a pilot for other communities in the Southern Sierra who wish to better understand their water supply. The report is expected to be completed in early 2015.

Land Reforming

Changes to the Soil — Salinization of Farmland

Soil salinization is caused by two primary factors.

- Our valley heat causes water to evaporate from the surface of the ground and the surface of leaves (evapotranspiration). When the water evaporates, this leaves salts behind and these accumulate in the soil.
- Not all irrigation water is the same quality. When there isn’t enough high-quality fresh water to meet our needs (the amount of water we choose to apply), farmers often turn to saltier water to irrigate their crops. This practice directly adds salt to the soil where crops are being grown.

The accumulation of salts in some large agricultural areas south of the Delta has long been noted. For decades, approximately half a million tons of salt annually have accumulated in the San Joaquin River and Tulare Lake Basins. For the San Joaquin River Basin, more salt enters the basin through irrigation water than leaves via drainage into the San Joaquin River.

The problem of soil salinization is especially severe in the Tulare Lake Basin because we have functioned largely as a closed basin since 1878 without a regular outlet to the ocean. Surface water comes in, but it never flows out. Therefore, the Tulare Lake Basin retains almost all the salt that enters the basin. When water used to flow out of the basin during high water years, it removed some of those salts. This accumulation of salts has created saline soils on the west side of the San Joaquin Valley.

Irrigation of those saline soils has led to severe declines in productivity (crop yields) and large scale fallowing of impaired land with toxic loads of salt, boron, selenium, and arsenic. (This problem is compounded when irrigating with relatively salty groundwater instead of fresh water.) Between 2002–2014, over 100,000 acres of impaired land was bought by Westlands Water District and retired from irrigated production. Depletion and contamination of groundwater supplies with salt and boron has also severely limited the productivity of the remaining land in drought years. Further reductions in agricultural acreage can be expected as salts continue to accumulate. Roughly a million acres of irrigated farmland are susceptible to this problem.

Like most towns in western Fresno and Kings County, Mendota lost its municipal groundwater supply over a decade ago because the deep groundwater became too salty after decades of irrigation of saline soils in the area. The city now relies on a small supply of imported surface water from the CVP. There is little water available for expansion of businesses or residences in the town.
The productive life of much of this area has already been extended by improvements in agricultural water use efficiency (which results in not only less water, but less salt, being applied to the soils), set-asides of some local areas for salt disposal, improved leaching methods, and retirement of some lands with high natural soil salinity. Maintaining a sustainable salt balance in remaining agricultural areas would require techniques such as the development of drains from the basin, reductions in salt loads entering the basin, or further reductions in irrigated area.

In a portion of Kings and west Fresno County, farmers are pumping water up from a great depth. This water is 3–4 times saltier than the water in the California Aqueduct. That is the reason the ground in those orchards is turning white in places. The water contains a mixture of salts, but especially sodium sulfate. It is the sodium ions that particularly cause problems. It is so much easier to put salty water on the land than to get rid of it.

The sodium ions will eventually cause two problems:
- a toxic effect on almonds, stone fruit and other crops except pistachios and cotton which are salt-tolerant
- a change of the soil structure so that the soil is less permeable

There are a number of things that farmers can do to get rid of the sodium in the upper part of the soil profile. All of them depend on applying more water than is needed for evaporation plus what the plants need. This is termed the leaching requirement. Drip irrigation doesn’t provide enough water to flush the sodium down past the root zone.

Farmers can also address salinity to some extent by adding other chemicals. The soil on the west side is relatively high in lime (calcium carbonate). When you have soil that is relatively high in lime and has an organic component, then you can add dilute sulfuric acid to the water and get the sodium ions and total salinity to go further down in the soil profile. This depends in large part on having enough water to push the sodium and total salinity down past the root zone. Otherwise, the wetting front will build up relatively rapidly.

These techniques works best if the soil has been prepared in advance using a slip plow or equivalent to disrupt the soil layers so that there will be better drainage in the root zone.

This same situation applies when the soil has lots of sodium ions, but the water is relatively fresh, such as down below the 230-foot contour line. The same basic principles and techniques that are used for applying salty water to low-salt ground can be used to grow crops using fresh water in soil that has lots of sodium ions.

Much of Hanford is located between 245-250 feet elevation. Mussel Slough at Highway 198 and 13th Avenue in Hanford is located at 235 feet elevation. There the soil is still well-drained sand and farming is quite feasible. But below about 230 feet elevation, the salts start picking up on the (Arroyo Pasajero) alluvial fan and it becomes a tough environment for most plants. This salinization below 230 feet elevation is natural, not human-caused. Much of the area below this elevation has a perched water table due to poorly drained soils. The water doesn’t drain due to strata of textures different than the ones above them. Sometimes those strata are clay, but not always. Over time, the perched water has become salty due to evaporation. As a result, the water is only useful for salt-tolerant plants. With good farming practices, it is possible to farm down to about 220 feet elevation. Below that, it becomes too salty.

The Blakeley Canal is located at 190 feet elevation. The soils from there up to about 216 feet elevation (the high stand of Tulare Lake) mark the basin rim of Tulare Lake. This is the highest area of lacustrine clay. This is the formation you encounter when you’re going east from Kettleman City on Highway 198, immediately after you drop into the Tulare Lakebed.

These soils vary from sand to clay with virtually all of the in-between possibilities. Close to the 190-foot contour, you will generally encounter gray nasty clay under the surface soils at some point. As you move up in elevation, the profiles have a variety of textures, but they are generally more sandy. Some of these soils have remnants of burnt tules, and all of them are salty.

**Land Subsidence**

A NASA/UC Irvine study published in 2009 concluded that for the period from October 2003 through March 2009, the groundwater supplies of the entire San Joaquin Valley were depleted by an average of over 2.8 million acre-feet per year. The lead author of the 2009 study was Jay Famiglietti, the director of Hydrologic Modeling at UC Irvine.
Floods and Droughts in the Tulare Lake Basin
General Flood and Drought Notes

In 2012, Bridget Scanlon and her colleagues at the University of Texas published a study that built on the NASA/UC Irvine study. Both studies identified the southern areas of the Central Valley — particularly the Tulare Lake Basin — as facing the most dire groundwater issues. Among other things, Scanlon’s study calculated that 140 cubic kilometers (34 cubic miles) of groundwater had been depleted from the Central Valley between the 1860s and 2003 (60 cubic kilometers from the 1860s to 1961 and 80 cubic kilometers from 1962 to 2003). Water level changes were dynamic with declines focused during droughts (1976–1977, 1987–1992, 1998–2003) and recovery at other times.

In 2011, Famiglietti and his associates published the results of another study about groundwater withdrawals in the Central Valley between 2003 and 2010. Among other things, this study found that 20 cubic kilometers of groundwater (5 cubic miles) had been depleted from the Central Valley between the 2003 and 2010. These researchers found that most of this volume came out of the Tulare Lake Basin.

That 38 cubic miles is a huge volume of water. It is equivalent to 130 million acre-feet — more than the volume of Lake Tahoe. It could cover the entire state of California to a depth of 15 inches. It is as if there were an enormous, unregulated mining operation going on under our feet, leaving lots of voids.

Groundwater withdrawals at unsustainable rates and volumes have resulted in a long-term economic boom for California’s agriculture economy that enabled the San Joaquin Valley to become one of the world’s most productive agricultural regions. Just in the San Joaquin Valley alone, the Gross Agricultural Production for 2012 was $32.4 billion, 76% of the total agricultural production for the entire state of California ($42.6 billion).

However, the groundwater extraction far exceeds natural aquifer recharge, and the depleted system has not been replenished by actively recharging the aquifer via conjunctive management practices. Efforts have been made, but they have been inadequate. As a result, these economic benefits have not gone without a broader cost to the infrastructure affected by land subsidence, to the quantity and quality of groundwater resources, to the increased energy required to pump groundwater, and to the decline in ecosystem services provided by the interaction of groundwater-surface water systems.

The San Joaquin Valley has the largest vertical subsidence and the largest areal extent of subsidence in the world because of groundwater withdrawal. This land subsidence from groundwater extraction in the San Joaquin Valley has been called the greatest human alteration of the Earth’s surface. Although this was the case in 1970s, it is conceivable that some other part of the world may have surpassed us since.

Land subsidence in the San Joaquin Valley is inexorably linked to the development of agriculture and the availability of water. Because the valley is semi-arid, and streamflow into the east side of the valley varies substantially from year to year and is mostly not available on the west side, agriculture developed a reliance on the groundwater aquifer system. That is, we use substantially more water than our average surface supplies provide.

Unconsolidated sediments composing a groundwater aquifer system in an alluvial basin like ours are often sorted into layers of similarly sized particles, such as gravel, sand, silt, and clay. Water moves most easily through permeable coarse-grained deposits of sand and gravel and much more slowly through finer-grained deposits of silt and clay. The fine-grained silt, silty-clay, and clay units function as “aquitards” that confine and separate groundwater flowing through the coarser-grained aquifers that underlie or overlie them. Sediments with layers or strata with significant changes in soil texture also act as “aquitards” even when coarser grained sediments are underneath finer grained sediments.

More than half of the thickness of the San Joaquin Valley groundwater aquifer system is composed of fine-grained sediments, including clays, silts, and sandy or silty clays, that are susceptible to compaction if depressured by pumping wells. Think of these like cells in a sponge that smash flat when water is removed, but lose the ability to ever accept water again.

Throughout most of the valley, a thick compressible clay, the Corcoran Clay member of the Tulare Formation, confines and separates deep aquifer sediments from a shallow unconfined or partly confined aquifer. Most of the chronic groundwater level decline, and compaction due to decline of hydraulic head, has occurred in the deep
confined aquifer. Most subsidence has probably resulted from compaction of relatively thin aquitards within
the deep aquifer system rather than in the Corcoran Clay confining unit, because the Corcoran’s large thickness
and low permeability inhibited drainage of water from its interior.

Subsidence related to groundwater withdrawals began in the mid-1920s. Subsidence rates increased as
agricultural development intensified after World War II and more wells were drilled to supply water for
irrigation. Subsidence rates eventually exceeded one foot per year in some places. In 1955, about one-fourth
of the total groundwater extracted for agricultural uses in the U.S. was pumped from the San Joaquin Valley,
and regional aquifer compaction was occurring at a rate of about one foot per year.

Land subsidence due to groundwater withdrawals in the San Joaquin Valley locally exceeded 28 feet by 1970. By December 1977, subsidence had reached a maximum of 29.6 feet southwest of Mendota in western Fresno County. More than 5,200 square miles of irrigable land, one-half the entire valley floor (10,000 square miles), had been affected by subsidence by then.

Those measurements were as of 1977 when the last comprehensive surveys of land subsidence were made. For
lack of better data, modern-day references to land subsidence continue to use those measurements even
though much more subsidence has occurred in the meantime.

Studies of Land Subsidence

Land subsidence in the San Joaquin Valley began sometime shortly after 1926. It was first noted in the Delano
area in 1935 where it was associated with groundwater overdraft. Despite this early recognition of the
relationship between groundwater-level decline and subsidence, subsidence from groundwater overdraft was not
investigated regionally until the early 1950s, when government agencies became concerned about a 30% reduction in the design capacity of the San Joaquin River, about the effect of subsidence on the Delta-Mendota Canal (then under construction), and on the California Aqueduct (then in the planning stages).

Part of the reason for the delayed reaction to subsidence was that it generally occurred uniformly and over such
a broad area that few residents or agencies realized that it had happened. As late as the mid-1950s, subsidence
in the Arvin-Maricopa area was ascribed to tectonically uplifted survey control points in the Tehachapi Mountains
rather than to the actual causative agent: decline of groundwater levels.

In 1954, a federal-state interagency committee and the USGS "Mechanics of Aquifers Project," headed by
Joseph Poland, began studying land subsidence by conducting both field monitoring and research. The studies
identified the magnitude and extent of subsidence and its quantitative relationship to groundwater overdraft,
and they developed new monitoring methods and techniques for analysis of field data that allowed accurate
computer models of aquifer system compaction to be built. These studies also provided information that allowed
optimal siting of the California Aqueduct which, along with the Delta-Mendota Canal and the Friant-Kern Canal,
brought surface water into the valley to diminish reliance on groundwater.

Comprehensive leveling surveys of the valley ended in 1970 and, over time, funding for coordinated subsidence
investigations also ended, and field installations such as borehole extensometers and water-level monitoring
wells were decommissioned or fell into disrepair. DWR continued to collect compaction data from a few
extensometers and from deep monitoring wells where available, and state, federal, and local water agencies
continued to run surveys on canal alignments intermittently, but analysis of this information was not
centralized.

In recent years, the NASA Jet Propulsion Laboratory (JPL), Stanford University, and USGS have carried out
studies of land subsidence. Much of the new work has been informed by analysis of Interferometric Synthetic
Aperture Radar (InSAR) satellite data. Tom Farr at NASA-JPL has led that analysis effort. InSAR provides the
most cost efficient method to generate high-resolution land surface deformation information over large areas
with high spatial detail.

Despite these recent studies, there has been no coordinated monitoring of land subsidence since the 1970s.
Much of the following discussion was informed by a comprehensive 2014 overview report prepared by James
Borchers and Michael Carpenter. A 2014 DWR summary report on subsidence generally substantiated the
findings of the much more detailed Borchers and Carpenter report.

DWR contracted with NASA in 2014 for mapping of recent subsidence in parts of the Central Valley where
satellite-based InSAR imagery was available. The work, to be completed in 2015, is a drought-response action
and screening effort to identify areas of ongoing relative land surface displacement.
Historic Areas of Land Subsidence
There were three primary areas of subsidence in the San Joaquin Valley during the 1926–70 period:

1. Los Banos-Kettleman City area (including western Fresno and Kings Counties)
2. Tulare-Wasco area in the southern valley, east of the Tulare Lakebed
3. Arvin-Maricopa area in the extreme southern end of the valley

The biggest and best known of these was the Los Banos-Kettleman City area. The greatest subsidence during the 1926–70 period was located about 10 miles southwest of Mendota on Panoche Road, one mile northeast of the California Aqueduct. That was the location of the well-known photograph (Figure 24) of Dr. Joseph F. Poland standing next to a power pole signed to indicate the former elevation of the land surface in 1925, 1955, and 1977.

![Figure 24. Dr. Joseph F. Poland on Panoche Road southwest of Mendota.](image)

One of the best places to see an example of early subsidence is in the community of Three Rocks, along Highway 33 on the west side of Fresno County. Three Rocks is located in the Westside Sub-basin (aka Westlands Water District) which extends roughly from Firebaugh on the north to Kettleman City on the south. This type of subsidence caused by applications of water on loosely consolidated sediments is known as hydrocompaction, shallow, or near-surface subsidence. It often occurs within a much larger area of subsidence related to ground water withdrawal. This area of hydrocompaction is a contiguous area that is 43,550 acres in size. Subsidence has continued off and on in that community since the initial hydrocompaction.

For a while, it had been assumed that subsidence in the San Joaquin Valley had largely stopped after the big federal and state canals were completed in about 1970. That assumption turns out to have been based on wishful thinking and a lack of monitoring. It is now recognized that further significant subsidence has occurred along the western side of the Tulare Lake Basin since the 1970s. It has been speculated that some areas in western Fresno and Kings Counties have now subsided upwards of 50 or 60 feet.
The magnitude of this subsidence is just speculation; no survey has been made in the area since 1977. Making a repeat survey on Panoche Road now means overcoming several challenges. First, PG&E has replaced the power pole that Joseph Poland was standing next to when his picture was taken in 1977. More important, the bench mark was not found in 1988 when the NGS survey crew visited the area; it has likely been destroyed.

The main Arroyo Pasajero channel begins at the confluence of Los Gatos and Warthan Creeks in Coalinga (near the southern part of the Los Banos-Kettleman City area). Its other two tributaries, Jacalitos, and Zapato Chino Creeks join the arroyo further downstream. Coalinga and the upper reaches of Arroyo Pasajero lie in Pleasant Valley. Pleasant Valley is separated from the rest of the Tulare Lake Basin by Anticline Ridge, a feature that parallels I-5 on its west side. The arroyo passes out of the valley through a narrow gap in this ridge. The arroyo then flows northeast across a broad alluvial fan toward Lemoore at the head of the Tulare Lakebed.

Since the 1920s, the alluvial fan formed by the arroyo has experienced as much as 18 feet of land subsidence due to groundwater extraction for irrigation, the majority occurring prior to 1968. Groundwater extraction is a continuing practice in this area. In the absence of surveys, it is uncertain how much additional subsidence has occurred in this area since 1980.

Although increasing water levels slowed subsidence by the early 1970s, subsidence continued due to delayed drainage of water from compacting clayey aquitards, particularly in the three areas identified above. For example, although no long-term water-level decline occurred at an extensometer installation near Pixley (Tulare-Wasco area) between January 1959 and February 1971, almost 3 feet of aquifer-system compaction occurred during this period.

Three fissures formed in the large subsiding area near Pixley (Tulare-Wasco area), about 25 miles north of Wasco. The easternmost fissure which formed in 1969 was believed to be related to differential subsidence that probably was triggered by groundwater extraction.

Movement during 1977–1978 on the northern end of the Pond-Poso Creek fault, an active tectonic fault located about 7 miles north of Wasco, was attributed to stresses imparted by declining groundwater levels.

As discussed in the section of this document that describes the Groundwater Overdraft, groundwater levels are no longer a good measure of the amount of water in the groundwater aquifer. When the water was pumped out in the 1960s, much of the pore space collapsed; the ability of the aquifer to store water was lost. Therefore, when the water levels recovered, the aquifer was no longer able to hold as much water as before. In subsequent droughts, water levels have dropped comparatively fast considering the amount of water that was withdrawn. In areas of current subsidence, we are continuing to destroy the ability of the aquifer to store water.

Although only about one-third of the peak annual groundwater withdrawals of the 1960s was pumped during the 1976–1977 drought, water levels fell more than 150 feet over a large area on the west side of the valley, and subsidence rates increased. The droughts of 1987–1992 and droughts and reductions in surface water imports during 2007–10 had similar effects, despite the fact that water levels never approached the historically low levels of the 1960s.

The rapid decline of groundwater levels during post-1975 droughts in response to relatively small volumes of pumping (compared to those of the 1960s) resulted from a loss of storage space in the aquifer system — mostly from inelastic compaction of aquitards during the 1950s and 1960s — and from reduced hydraulic conductivity (permeability) of those compacted aquitards that restrict drainage of water to permeable parts of the aquifer system. Water levels were considerably higher than during the 1960s, yet there was renewed land subsidence during droughts. This illustrates the complex effects of unequal distribution of pre-consolidation stress within the aquitards, and between the aquitards and more permeable units of the aquifer system.

New Areas of Land Subsidence
A.A. Swanson summarized land subsidence that occurred between 1970 and 1995. He gleaned information from road, canal and levee surveys, documentation of canal and bridge infrastructure repairs, and reports of changing canal gradients from water agency managers. Swanson’s information indicated that subsidence was continuing in each of the three main areas identified in the earlier comprehensive subsidence studies (Los Banos-Kettleman City, Tulare-Wasco, and Arvin-Maricopa) and in an additional area near the San Joaquin River north of Mendota.
Floods and Droughts in the Tulare Lake Basin

General Flood and Drought Notes

The current drought and cropping patterns that have changed from row crops and rangeland to tree and other permanent crops have again forced reliance on groundwater aquifer systems in the San Joaquin Valley for agricultural irrigation supplies. By planting permanent crops, we have hardened water demand so that these lands cannot easily be fallowed during periods of drought. Recent satellite radar analyses show that two large areas in the San Joaquin Valley are currently subsiding substantially.557

Subsidence rates had reduced for a while in the three original subsiding areas after completion of the major federal and state canals. Recent studies indicate that land subsidence rates of one foot per year have returned to those San Joaquin Valley basins that are highly reliant on groundwater supplies.558 Rates are highest in two new subsiding areas, one in the north and one in the south.

The northern of these two areas is in Merced County near El Nido; locals often call this area Red Top. This 1,200-square-mile area stretches from the cities of Merced on the north, to Los Banos on the west, Madera on the east and Mendota on the south. This is the general area that Swanson identified in 1995; it is just larger than was previously realized. In recent years, this area has been sinking at a rate of nearly one foot per year.559

The other area is between Tulare and Kettleman City, centered near Corcoran. This 2,700-square-mile swath of subsidence appears to have begun sometime after the early 1990s.

The two currently subsiding areas are shifted substantially from the three locales of major subsidence during the 1926–70 period.

The northern subsidence area near El Nido subsided between 1–4 feet during 1926–70; subsidence in that area was centered much further to the southwest during that period. The greatest subsidence during 1926–70 was located one mile northeast of the California Aqueduct on Panoche Road southwest from Mendota. This was the location of the photograph of Joseph Poland standing next to a power pole signed to indicate the former height of land surface during 1925–77.

Currently, subsidence occurs in a broad swath of the mid-valley area further to the northeast. Results of satellite radar measurements indicate at least 1.8 feet of subsidence at rates of 11 inches/year near the San Joaquin River and the Eastside Bypass during 2008–10, including the southern part of the Delta-Mendota Canal.560

Chris White of the Central California Irrigation District (CCID) said that the current subsiding area is apparently correlated with increased groundwater extraction needed for changing land uses and cropping patterns in areas east of the San Joaquin River that lie outside water district boundaries and have no access to surface water for irrigation. As is the case in the area of recent subsidence further south in the San Joaquin Valley, open land and seasonal crops have been supplanted by perennial crops and orchards that require irrigation year-round. These farms are now relying on extraction of water from the deep confined part of the aquifer system for irrigation.

In 2002, the USACE predicted that 17 feet of subsidence would occur between 2000–2060 where Route 152 crosses the San Joaquin River and the East Side Bypass. Initial satellite radar measurements have shown substantial subsidence is occurring in that area.

Investigation of subsidence along the Delta-Mendota Canal has determined that the northern portion of the canal is relatively stable.561 Chris White of the CCID said that historical land surface deformation measurements indicate that the southern portion of the canal subsided substantially. More recently, slight subsidence has occurred, probably in response to the large subsidence feature south of the town of El Nido.

Sack Dam is a small diversion dam on the San Joaquin River near Mendota that captures water for 45,000 acres of farmland. This 60-mile stretch of the river along the valley's west side dried up in the early 1950s after Friant Dam was built. It is the most challenging stretch for the San Joaquin River Restoration Program.

Sack Dam presents an obstacle to salmon swimming upstream on the San Joaquin River. As part of the restoration program, a new Sack Dam is to be built with both a fish ladder and a system to raise and lower the dam so that salmon can pass through.562

Recent surveying work by DWR and USBR indicated that subsidence is occurring at about 0.5 feet per year near Sack Dam and 0.9 feet per year near the Eastside Bypass. Discovery of subsidence in this area halted redesign efforts for the dam as agencies consider how to adapt to the lowered land surface and prepare for likely
continued subsidence. The design process is complicated by the need to build a taller dam to contain a deeper and broader pool of water that would result from continued subsidence. Dams taller than six feet are regulated by more rigorous safety regulations and approvals.

West of El Nido, the reach of the Eastside Bypass downstream from the Sand Slough Connector has been affected by subsidence since before 1980. As the levees subside, they are more susceptible to being breached during floods. This could inundate up to 170 square miles. Subsidence has also caused huge quantities of sediment to accumulate in this reach, significantly reducing the channel's ability to carry water and reducing levee freeboard (the distance between the water surface and the top of the levees).

USACE has removed 650,000 cubic yards of sediment from this reach of the Eastside Bypass, and the Lower San Joaquin Levee District (LSJLD) and DWR have raised the levees, but flow capacity has not been fully restored to design capacity.

The reach of the Eastside Bypass upstream from the Sand Slough Connector has been even more severely impacted by subsidence. Modeling studies by DWR predict that if subsidence continues at current rates, flow capacity of this reach will be reduced to only 57% of design capacity by 2016.

Subsidence has also likely reduced the flow capacity of the San Joaquin River east of El Nido to less than half its design capacity.

CCID’s Poso Canal, which runs parallel to the San Joaquin River, has lost 10% of its flow capacity. Embankments were raised on CCID’s Main Canal near the Russell Avenue bridge, and flow capacity has been reduced near Los Banos. Russell Landon of CCID said that embankments were raised on CCID’s Outside Canal in 1971 and 1994.

Chris White of CCID said that a 2007–14 subsidence remediation program to raise 16 miles of embankments, construct two new weirs and 20 service turnouts will cost $5.4 million when completed.

When the Russell Avenue bridge over the Outside Canal was constructed in 1954, the flow capacity of the canal was 620 cfs. The bridge now restricts canal flow to 340 cfs, partly because subsidence has changed the structure from a free flowing conduit to a siphon. In 1960, 2-foot high-sidewalls were added to the Russell Avenue bridge to prevent water from the canal from flowing across road surfaces. Currently, water seeps through the roadbed onto the bridge, and on windy days waves splash over the sidewalls and wet the road surface.

Chris White said that CCID staff routinely conducted facility inspections by examining the canal and turnouts from a boat which passed easily under the bridge — something that is obviously impossible now that the bridge is partly submerged. The bridge has been judged structurally deficient and is scheduled for replacement by CCID, Caltrans, and the Fresno County Department of Public Works at a cost of $2.5 million.

The current southern subsidence area is northwest of the Tulare-Wasco area where maximum subsidence during 1926–70 was about 13 feet. More than 1.5 feet of subsidence occurred over a large part of the southern subsiding area during 2007–11.

The maximum rate of recent subsidence in the southern area is about twice the maximum rate that occurred historically in that area. Corcoran is near the center of the southern subsidence area. The area near Corcoran subsided 3.9 feet during a 3½ year period. The casings on some deep wells in Corcoran have been left exposed about 2 feet above ground in the last few years. That is, the ground has settled, leaving the well casings sticking up in the air.

Another example of subsidence in this area is the Angiola Water District canal in Corcoran. Historical photographs show this canal when it was full. Children are fishing in the canal with their legs dangling over the canal embankment; the water 12–15 feet below them. Matt Hurley of the Angiola Water District said that if a person sat there in the spring of 2013, they would get wet up to their knees.

Periodically Caltrans resurveys the vertical elevation of its highways. Their last resurvey of Highway 198 showed that the area between Hanford and Lemoore experienced up to 9 feet of subsidence between the 1960s and 2004.
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Lemoore Naval Air Station (LNAS) lies in the northwest edge of the southern subsiding area. Subsidence studies at LNAS indicate that total subsidence there between 1925–2010 exceeded 10 feet — considerably more than the 3.5–4.0 feet that Caltrans surveys indicated had occurred in that area between the 1960s and 2004. The difference is likely due to subsidence that occurred prior to the 1960s and after 2004.

Because the LNAS and Highway 198 are located on the northern fringe of the subsiding area, neither the LNAS studies nor the Caltrans survey capture the maximum subsidence in this region. No independent surface measurements of land subsidence or aquifer-system compaction have been acquired to confirm the magnitude of subsidence in the southern subsiding area.

According to John Kirk at DWR, the southern subsiding area is correlated with increased groundwater extraction needed for changing land use and cropping patterns; open land and seasonal crops have been supplanted by perennial crops and orchards that require year-round irrigation.

Cost of Land Subsidence
Subsidence has caused major impacts to infrastructure and physical features, including the San Joaquin River, Delta-Mendota Canal, Friant-Kern Canal, and San Luis Canal, as well as numerous privately owned canals and related infrastructure such as turnouts, bridges, pipelines, and storm sewers.

Historically, subsidence impacted about 30 miles of the Delta-Mendota Canal upstream from its terminus at Mendota Pool by submerging canal service turnouts, drain inlets, bridges, pipelines, and check structures used to control water surface elevation in the canal, and by overtopping the concrete lining of the canal.

In the 1960s, USBR raised bridges and relocated pipeline crossings. In 1977, USBR made $30 million (2013 dollars) in modifications to the canal as a result of subsidence. In 1992, operation of the Delta-Mendota Canal passed to the San Luis Delta-Mendota Water Authority (SLDMWA). In 2004, the SLDMWA installed 10,000 feet of concrete canal-lining extensions in areas of reduced freeboard. Since then, SLDMWA has identified many areas where canal freeboard has been reduced to less than ½ foot. Additional subsidence has been documented along the canal during the period 2003–10.565

The San Luis Canal, a shared asset of the federal CVP and the state SWP that was completed in 1968, has been affected by subsidence along 85 miles of its length between the Los Banos and Kettleman City areas. The canal, now considered the middle section of the California Aqueduct, passes through three major subsidence bowls: 1) southwest of Mendota, 2) near the town of Cantua Creek, and 3) near the town of Huron.566

Because subsidence had adversely affected the earlier-built Delta-Mendota Canal, designers incorporated up to 10 feet of extra freeboard into the San Luis Canal, adding $31 million (2013 dollars) to construction costs. Additional subsidence required raising of canal linings, bridges, and other canal structures and rehabilitation of roads in the early 1980s at a cost of $56 million (2013 dollars). Subsidence due to hydrocompaction caused two sags in the California Aqueduct, centered near Cantua Creek, USGS, in cooperation with DWR, studied subsidence during 2003-2010 along this middle reach of the California Aqueduct.

Subsidence caused by structure settlement, groundwater extraction, hydrocompaction, and hydrocarbon production has been described for the California Aqueduct in reaches south of the San Luis Canal. DWR is currently compiling historical information on subsidence magnitude, repairs, and operational impacts in order to assess the reliability of water deliveries from the entire California Aqueduct. The performance of all engineered structures, (check dams, turnouts, siphons, concrete lining, etc.) will be evaluated.

The total cost to the U.S. government to account for or repair subsidence damage from groundwater extraction to major canals and drains built by the federal government on the west side of the San Joaquin Valley has been $88 million (2013 dollars). These are contract construction costs and do not include costs for design, inspection, or studies. They also do not include the considerable cost of pre-compacting the San Luis Canal and major lateral canal alignments by diking and flooding to avoid subsidence caused by hydrocompaction of moisture deficient soils in the western San Joaquin Valley.

About 30 miles of the Friant-Kern Canal, from 95 to 125 miles downstream of the Friant Dam, was impacted by subsidence in the Tulare-Wasco area. Between the end of construction in 1951 and January 1975, parts of the affected reach of the canal subsided 5.5 feet, interfering with operations. Subsidence also affected parts of the canal farther south toward Bakersfield. A 17-mile reach of the canal was rehabilitated during 1976–80 at a cost of $15 million (2013 dollars).
Costs for repairs to other infrastructure in the San Joaquin Valley, including replacement and repair of wells damaged and destroyed by land subsidence, have not been compiled systematically. Realistically, costs probably cannot be determined because there is no centralized repository of this information. In the 618 square mile region of maximum subsidence in the valley, 275 wells reported failed casings due to subsidence-induced compressive rupture between 1950–61. The current costs to replace an 18-inch diameter agricultural well average $200–$250 per foot. Assuming those wells averaged 1640 feet deep, a conservative cost estimate to replace them would be $90 million dollars (2013 dollars).

One very large 120,000-acre farm on the west side of the San Joaquin Valley had the casings fail on 60 wells in the 1970s. Those wells had an average depth of 2,000 feet; the cost to replace them would be $24 million (in 2013 dollars).

The most comprehensive and systematic estimate of the economic costs of land subsidence in the San Joaquin Valley was done by Gilbert Bertoldi of the USGS. Bertoldi collected billing invoices from land surveyors and other contractors and repair estimates from county agencies for remediating subsidence damage during 1955–1972 in areas of Fresno and Kings Counties that had subsided more than 4 feet between 1925–1972. The invoices and repair estimates showed costs for periodic surveying and regrading of agricultural fields to enable proper flow of water during flood irrigation, replacing networks of broken 8 to 10-inch ceramic pipes that were buried trenches to transport irrigation water to fields, and broken sanitary sewers in urban areas. (Fresno and Kings County agencies had reported that sanitary sewers had broken for undetermined reasons; Bertoldi determined that broken sewers were in the subsiding areas.) Costs to repair or replace wells damaged by subsidence, the value of structures lost by condemnation, and decreased property values as a result of zoning changes where subsidence increased the extent and depth of flooding were also included in the damage estimates.

The costs due to the impact of subsidence on infrastructure totaled more than $1.3 billion (2013 dollars) during the 18–year period 1955–1972. This represents only a fraction of the total costs to date due to land subsidence in the San Joaquin Valley. It reflects only the costs for these two counties for this 18–year period. It excludes damages for areas that subsided less than four feet. Cost data for the period from 1973–present are mostly unavailable. This cost estimate also excluded the indirect costs of land subsidence such as flood damage to inundated farm equipment and long-term environmental effects.

**Gradient Change**

Land subsidence on the west side of the Tulare Lake Basin steepens the gradient of streams coming off the Coast Ranges. This increases the ability of those streams to erode their channels. This is most evident on the Arroyo Pasajero where recent floods have left huge sediment deposits adjacent to Lassen Ave. and the California Aqueduct.

Significant subsidence has occurred in this portion of the Tulare Lake Basin. This subsidence has resulted in increased gradient for the Arroyo Pasajero and similar stream courses on the west side of the Tulare Lake Basin. This increased gradient results in greater erosion. Many of the soils in this area have silty textures and are cut like butter under these conditions. As a result, the groundwater overdraft and consequent subsidence has increased sources of sediment in the Arroyo Pasajero.

**Land Uplift**

The consequences of overdraining the groundwater aquifer aren’t just a lack of water and increased water cost. The earth has responded in various ways. One of the most noticeable ways is when the earth settles, filling in the voids where the groundwater used to be. The removal of groundwater has caused subsidence over a large portion of the valley floor.

But research has recently shown that the earth has responded in another way as well. Researchers had noted that small earthquakes are more common along the Parkfield segment of the San Andreas Fault in the late summer and fall. This area marks the transition region between two segments of the San Andreas Fault: the one that is creeping near Parkfield and the locked segment responsible for the 1857 Fort Tejon Earthquake. Both segments experience an increase in small temblors in the fall. This seasonal variation was presumably related to precipitation in some way, but it wasn’t clear how.

At the same time, researchers had been trying to understand why the mountains around the San Joaquin Valley were slowly rising according to modern GPS measurements. Initially it was assumed that the current uplift of the Coast Ranges was related primarily to plate motion, the same force that drives horizontal movement along
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the San Andreas Fault. However, tectonic models didn’t fully explain the observed uplift in the Coast Ranges or in the Sierra.

In May 2014, Colin Amos, a geologist at Western Washington University in Bellingham, and his co-authors published a paper in the journal *Nature* on how the removal of groundwater appears to have affected the southern San Joaquin Valley. 571

The researchers analyzed years of data from hundreds of GPS stations at sites from the Pacific Ocean into Nevada that cross a southern portion of the San Joaquin Valley.

During most years, we extract more water from the groundwater aquifer than gets recharged. Over 150 years, we have removed 38 cubic miles of water from our groundwater aquifer. That is the total amount of groundwater overdraft since settlement began in the San Joaquin Valley.

The researchers discovered that the average uplift rates (1 to 3 millimeters per year) closely match what would be expected when that much weight was removed from the valley’s groundwater aquifers.

The authors of the Amos paper looked at the cumulative effect of removing that much water. All that groundwater did more than just fill voids in the earth. It also had weight. That much water weighed about 175 billion tons. It pressed down on the earth’s crust. The weight of all that water extracted from the earth and evaporated has caused the underlying crust, the entire North American Plate, to rise.

That is called elastic rebound. It is the same process that happened in the Arctic when the great ice caps melted and the ground rebounded. Now you can see the beach lines extending hundreds of feet above your head in that area. We have done something similar in our area with pumps and evaporation.

As the crust has lifted under the San Joaquin Valley, it has raised the Coast Ranges, the valley floor, and the Sierra Nevada. The mountains surrounding the San Joaquin Valley may have risen as much as 6 inches since the 1860s, far faster than they otherwise would have, given the plate boundary convergence rate and erosion rates.

There are two related processes at work here, both driven by the groundwater overdraft: long-term uplift and flexing.

In addition to the long-term average uplift described above, Amos and his co-authors used the data from the long-term GPS monitoring stations to show that our part of the North American Plate is flexing each year. The valley floor rises in spring when the runoff occurs, and the Coast Ranges and the Sierra rise in the late summer and early fall.

Annual uplift of 3–8 mm (roughly ⅛–¼ inch) occurs in the Coast Ranges and western Sierra during the end of the summer growing season, corresponding with peak groundwater withdrawals in the San Joaquin Valley.

Each year we go through this cycle. This means that when the runoff occurs, the valley floor rises slightly. And when we turn on the pumps, we actually cause the mountains, the big mountains to rise. Our water management activities have that kind of power.

When the mountains are at their maximum point of uplift in the late summer and fall, this flexing slightly reduces the stress where the North American Plate joins the Pacific Plate. This stress change corresponds with an increase in the number of small temblors on the San Andreas Fault near Parkfield. The Amos study suggested that this seasonal uplift of the Coast Ranges reduces the stress on the adjacent San Andreas Fault, which may explain some of the annual increase in small temblors observed in that area.

Although the Coast Ranges are slowly rising, it could be just enough to unsettle the fault. As the ground expands upward, it pulls away from the fault, slightly reducing the forces clamping the plates tightly together. That makes it a bit easier for the plates to slide, which might set off subtle shudders.

That means that when we turn on the pumps in the Tulare Lake Basin, we cause the North American Plate to flex, uplifting the Coast Ranges, and causing an increase in small earthquakes on the San Andreas Fault on the other side of the Coast Ranges. We move so much water that we cause the San Andreas to tremble.
Summary of Droughts

Summary of Past Megadroughts
Megadroughts are defined more by their duration than their severity. They are extreme dry spells that can last for a decade or longer. As shown in Table 19, the Southern Sierra has experienced five megadroughts since A.D. 800. The flows in this table are based on tree-ring reconstructed flows for the upper San Joaquin River.

Table 19. Summary of past megadroughts.

<table>
<thead>
<tr>
<th>Drought (A.D.)</th>
<th>Length (years)</th>
<th>Upper San Joaquin River % of average flow (900–2012)</th>
</tr>
</thead>
<tbody>
<tr>
<td>832–1074</td>
<td>243</td>
<td>97%</td>
</tr>
<tr>
<td>1122–1299</td>
<td>178</td>
<td>99%</td>
</tr>
<tr>
<td>1444–1483</td>
<td>40</td>
<td>79%</td>
</tr>
<tr>
<td>1566–1602</td>
<td>37</td>
<td>85% (based on the years 1569–1595)</td>
</tr>
<tr>
<td>1918–1934</td>
<td>17</td>
<td>68%</td>
</tr>
</tbody>
</table>

The two extremely long droughts of the 900s and 1100s clearly had severe and prolonged impacts to the north and east of the Tulare Lake Basin. Studies conflict as to whether those droughts impacted our basin; authors of those studies are trying to resolve the issue. In any case, the Little Ice Age (approximately 1450–1850) apparently affected the tropical Pacific, resulting in the storm tracks tending to come further south in future years. The three megadroughts that have occurred since then have been of much shorter duration. The megadroughts of 1444–1483 and 1566–1602 were much worse in the southern part of the Central Valley than further north. This suggests that the storm tracks did not come as far south in those years.

Summary of Droughts in the San Joaquin Valley: 1400–1900
The droughts in Table 20 were compiled from a variety of sources, including tree-ring reconstructions, Tulare Lake elevations, and settler accounts. This is not meant to be a complete listing of all the droughts that have occurred during these 500 years. It captures most of the larger and more important droughts, but omits some of the shorter droughts, especially in the earlier centuries.

Table 20. Selected droughts in the San Joaquin Valley: 1400–1900.

<table>
<thead>
<tr>
<th>Drought (A.D.)</th>
<th>Length (years)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1444–1483</td>
<td>40</td>
</tr>
<tr>
<td>1527–1533</td>
<td>7</td>
</tr>
<tr>
<td>1540–1548</td>
<td>9</td>
</tr>
<tr>
<td>1566–1602</td>
<td>36</td>
</tr>
<tr>
<td>1618–1619</td>
<td>2</td>
</tr>
<tr>
<td>1629–1632</td>
<td>4</td>
</tr>
<tr>
<td>1652–1659</td>
<td>8</td>
</tr>
<tr>
<td>1721–1722</td>
<td>2</td>
</tr>
<tr>
<td>1728–1729</td>
<td>2</td>
</tr>
<tr>
<td>1735–1737</td>
<td>5</td>
</tr>
<tr>
<td>1753–1757</td>
<td>5</td>
</tr>
<tr>
<td>1776–1778</td>
<td>3</td>
</tr>
<tr>
<td>1780–1783</td>
<td>4</td>
</tr>
<tr>
<td>1793–1796</td>
<td>4</td>
</tr>
<tr>
<td>1807–1809</td>
<td>3</td>
</tr>
<tr>
<td>1822–1824</td>
<td>3</td>
</tr>
<tr>
<td>1827–1829</td>
<td>3</td>
</tr>
<tr>
<td>1840–1846</td>
<td>7</td>
</tr>
<tr>
<td>1855–1861</td>
<td>7</td>
</tr>
<tr>
<td>1862–1864</td>
<td>3</td>
</tr>
<tr>
<td>1869–1871</td>
<td>3</td>
</tr>
<tr>
<td>1873–1879</td>
<td>7</td>
</tr>
<tr>
<td>1882–1883</td>
<td>2</td>
</tr>
<tr>
<td>1887–1888</td>
<td>3</td>
</tr>
<tr>
<td>1898–1900</td>
<td>3</td>
</tr>
</tbody>
</table>
Floods and Droughts in the Tulare Lake Basin

Summary of Droughts

Dave Meko and others used tree-rings to reconstruct the flow on the San Joaquin River and its major tributaries for 1113 years (900–2012). Table 21 gives the 20 driest years, plus the driest 3-year, 10-year, and 20-year periods.

Table 21. Driest 20 time periods on the upper San Joaquin River for 1113 years (900–2012).

<table>
<thead>
<tr>
<th>Rank</th>
<th>Driest Year</th>
<th>Driest 3-year period (ending year)</th>
<th>Driest 10-year period (ending year)</th>
<th>Driest 20-year period (ending year)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>1580</td>
<td>983</td>
<td>1933</td>
<td>1936</td>
</tr>
<tr>
<td>2</td>
<td>1795</td>
<td>981</td>
<td>1934</td>
<td>1935</td>
</tr>
<tr>
<td>3</td>
<td>1924*</td>
<td>982</td>
<td>1931</td>
<td>1465</td>
</tr>
<tr>
<td>4</td>
<td>1532</td>
<td>980</td>
<td>1932</td>
<td>1466</td>
</tr>
<tr>
<td>5</td>
<td>1829</td>
<td>981</td>
<td>1935</td>
<td>1468</td>
</tr>
<tr>
<td>6</td>
<td>1126</td>
<td>1480</td>
<td>1461</td>
<td>1469</td>
</tr>
<tr>
<td>7</td>
<td>1729</td>
<td>1846</td>
<td>1482</td>
<td>1937</td>
</tr>
<tr>
<td>8</td>
<td>1632</td>
<td>1934</td>
<td>1459</td>
<td>1934</td>
</tr>
<tr>
<td>9</td>
<td>1782</td>
<td>1481</td>
<td>1460</td>
<td>1158</td>
</tr>
<tr>
<td>10</td>
<td>1864</td>
<td>984</td>
<td>1483</td>
<td>1467</td>
</tr>
<tr>
<td>11</td>
<td>1783</td>
<td>1148</td>
<td>1481</td>
<td>1464</td>
</tr>
<tr>
<td>12</td>
<td>1841</td>
<td>1992</td>
<td>1480</td>
<td>1157</td>
</tr>
<tr>
<td>13</td>
<td>957</td>
<td>1461</td>
<td>984</td>
<td>1463</td>
</tr>
<tr>
<td>14</td>
<td>1931</td>
<td>1145</td>
<td>1783</td>
<td>1156</td>
</tr>
<tr>
<td>15</td>
<td>1655</td>
<td>1239</td>
<td>986</td>
<td>1462</td>
</tr>
<tr>
<td>16</td>
<td>1777</td>
<td>1929</td>
<td>1930</td>
<td>1159</td>
</tr>
<tr>
<td>17</td>
<td>1579</td>
<td>933</td>
<td>1465</td>
<td>1441</td>
</tr>
<tr>
<td>18</td>
<td>1843</td>
<td>1483</td>
<td>1784</td>
<td>1483</td>
</tr>
<tr>
<td>19</td>
<td>1059</td>
<td>1459</td>
<td>1929</td>
<td>1482</td>
</tr>
<tr>
<td>20</td>
<td>954</td>
<td>1845</td>
<td>1462</td>
<td>1162</td>
</tr>
</tbody>
</table>

*Table 21 is based on the location of the SJF gage, the San Joaquin River at the inflow to Millerton Lake. As shown in Table 23, total flows in the Tulare Lake Basin in 1977 were slightly less than in 1924. Therefore, the four driest years in our basin were 1580, 1977, 1924, and 1795. The flow for 1795 (reconstructed from tree-ring data) was only 2% less than the flow actually measured by stream gages for 1977 and 1924. So we can’t say with confidence where 1795 falls in this order, especially in the Tulare Lake Basin. Water year 2015 will almost certainly be drier than 1977, 1924, and 1795.

The San Joaquin and Sacramento River Basins share many of the major droughts and wet periods. Some differences, however, are evident in relative severity of droughts and wet periods. The drought of the 1920s–30s appears in both the basins, but is more severe in a long-term context in the Sacramento than in the San Joaquin. Both basins have dry conditions in the 1100s, but less so in the San Joaquin. Dry periods on the order of 30-40 years near the end of the 1400s were among the most severe periods of this length in the San Joaquin record.

The 1920s–30s and 1990s contained periods of drought notably severe in both the Sacramento and San Joaquin River Basins, even in a centuries- to millennium-context. The record-low flow is 1580 in both the Sacramento and San Joaquin River Basins.

The tree-ring reconstructions of flow contain no strong, regular, statistically significant cycles over their full lengths from 900–2012.

Summary of Droughts since 1901

The Tulare Lake Basin has experienced nine droughts since 1901. Table 22 shows how the San Joaquin Valley Water Year Index categorized the drought years in that basin. It also shows total runoff for the four major rivers in our basin (Kings, Kaweah, Tule, and Kern) combined during this drought. Conditions for 2015 are projected.
Table 22. Summary of droughts in the Tulare Lake Basin since 1901.

<table>
<thead>
<tr>
<th>Drought</th>
<th>Water Year</th>
<th>Water Year Classification</th>
<th>San Joaquin River Basin</th>
<th>Tulare Lake Basin Drought Average</th>
</tr>
</thead>
<tbody>
<tr>
<td>1912–13</td>
<td>1912</td>
<td>Below normal</td>
<td>1,608,840</td>
<td>55%</td>
</tr>
<tr>
<td></td>
<td>1913</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td>1918–34</td>
<td>1918</td>
<td>Below normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1919</td>
<td>Below normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1920</td>
<td>Below normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1921</td>
<td>Above normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1922</td>
<td>Wet</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1923</td>
<td>Above normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1924</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1925</td>
<td>Below normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1926</td>
<td>Dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1927</td>
<td>Above normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1928</td>
<td>Below normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1929</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1930</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1931</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1932</td>
<td>Above normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1933</td>
<td>Dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1934</td>
<td>Critically dry</td>
<td>2,047,511</td>
<td>70%</td>
</tr>
<tr>
<td>1947–50</td>
<td>1947</td>
<td>Dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1948</td>
<td>Below normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1949</td>
<td>Below normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1950</td>
<td>Below normal</td>
<td>1,781,568</td>
<td>61%</td>
</tr>
<tr>
<td>1959–61</td>
<td>1959</td>
<td>Dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1960</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1961</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td>1976–77</td>
<td>1976</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1977</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td>1987–92</td>
<td>1987</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1988</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1989</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1990</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1991</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>1992</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td>1999–04</td>
<td>1999</td>
<td>Above normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>2000</td>
<td>Above normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>2001</td>
<td>Dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>2002</td>
<td>Dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>2003</td>
<td>Below normal</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>2004</td>
<td>Dry</td>
<td>2,074,538</td>
<td>71%</td>
</tr>
<tr>
<td>2007–09</td>
<td>2007</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>2008</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>2009</td>
<td>Below normal</td>
<td>1,818,822</td>
<td>62%</td>
</tr>
<tr>
<td>2012–15+</td>
<td>2012</td>
<td>Dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>2013</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>2014</td>
<td>Critically dry</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>2015</td>
<td>Critically dry</td>
<td>1,003,578</td>
<td>34%</td>
</tr>
</tbody>
</table>

Table 22 summarizes the droughts in the Tulare Lake Basin since 1901. Research suggests that that the 20th century may well have been an outlier, an unusually wet century. Overall, California experienced less drought in the 20th century than most of the preceding 4 to 20 centuries. Table 23 compares runoff in the Tulare Lake Basin for the ten driest years since record-keeping began in 1894.
Floods and Droughts in the Tulare Lake Basin
Summary of Droughts

Table 23. Driest 10 years in the Tulare Lake Basin since 1894 — Total runoff.

<table>
<thead>
<tr>
<th>Water Year</th>
<th>Kings River at Pine Flat</th>
<th>Kaweah River at Terminus</th>
<th>Tule River at Success</th>
<th>Kern River near Bakersfield</th>
<th>Tulare Lake Basin Combined</th>
</tr>
</thead>
<tbody>
<tr>
<td>1. 2015*</td>
<td>341,000</td>
<td>83,700</td>
<td>11,300</td>
<td>110,000</td>
<td>546,000</td>
</tr>
<tr>
<td>2. 1977</td>
<td>386,007</td>
<td>93,641</td>
<td>15,884</td>
<td>201,040</td>
<td>696,572</td>
</tr>
<tr>
<td>3. 1924</td>
<td>391,920</td>
<td>101,650</td>
<td>24,700</td>
<td>190,810</td>
<td>709,080</td>
</tr>
<tr>
<td>4. 1931</td>
<td>465,640</td>
<td>114,270</td>
<td>24,730</td>
<td>184,130</td>
<td>788,770</td>
</tr>
<tr>
<td>5. 2014</td>
<td>538,359</td>
<td>116,760</td>
<td>14,550</td>
<td>178,159</td>
<td>830,549</td>
</tr>
<tr>
<td>6. 1961</td>
<td>555,392</td>
<td>146,916</td>
<td>42,419</td>
<td>252,360</td>
<td>977,066</td>
</tr>
<tr>
<td>7. 1976</td>
<td>646,620</td>
<td>130,760</td>
<td>20,300</td>
<td>232,570</td>
<td>1,030,250</td>
</tr>
<tr>
<td>8. 1934</td>
<td>684,030</td>
<td>141,203</td>
<td>29,805</td>
<td>209,729</td>
<td>1,064,767</td>
</tr>
<tr>
<td>9. 1990</td>
<td>689,580</td>
<td>152,761</td>
<td>32,627</td>
<td>216,403</td>
<td>1,091,371</td>
</tr>
</tbody>
</table>

*projected runoff from NOAA River Forecast Center

Table 24 presents the same information as in Table 23, but as a percentage of the long-term average runoff (1894–2014).

Table 24. Driest 10 years in the Tulare Lake Basin since 1894 — Compared to average.

<table>
<thead>
<tr>
<th>Water Year</th>
<th>Kings River at Pine Flat</th>
<th>Kaweah River at Terminus</th>
<th>Tule River at Success</th>
<th>Kern River near Bakersfield</th>
<th>Tulare Lake Basin Combined</th>
</tr>
</thead>
<tbody>
<tr>
<td>1. 2015</td>
<td>21%</td>
<td>20%</td>
<td>8%</td>
<td>15%</td>
<td>19%</td>
</tr>
<tr>
<td>2. 1977</td>
<td>23%</td>
<td>22%</td>
<td>12%</td>
<td>28%</td>
<td>24%</td>
</tr>
<tr>
<td>3. 1924</td>
<td>24%</td>
<td>24%</td>
<td>18%</td>
<td>27%</td>
<td>24%</td>
</tr>
<tr>
<td>4. 1931</td>
<td>28%</td>
<td>27%</td>
<td>18%</td>
<td>26%</td>
<td>27%</td>
</tr>
<tr>
<td>5. 2014</td>
<td>32%</td>
<td>23%</td>
<td>11%</td>
<td>25%</td>
<td>28%</td>
</tr>
<tr>
<td>6. 1961</td>
<td>33%</td>
<td>28%</td>
<td>18%</td>
<td>26%</td>
<td>30%</td>
</tr>
<tr>
<td>7. 1976</td>
<td>32%</td>
<td>35%</td>
<td>31%</td>
<td>35%</td>
<td>33%</td>
</tr>
<tr>
<td>8. 1934</td>
<td>39%</td>
<td>31%</td>
<td>15%</td>
<td>32%</td>
<td>35%</td>
</tr>
<tr>
<td>9. 1990</td>
<td>41%</td>
<td>33%</td>
<td>22%</td>
<td>29%</td>
<td>36%</td>
</tr>
<tr>
<td>10. 2013</td>
<td>42%</td>
<td>36%</td>
<td>24%</td>
<td>30%</td>
<td>37%</td>
</tr>
</tbody>
</table>

All of these critically dry years were part of multi-year droughts:
- Three were part of the 17-year-long megadrought of 1918–34.
- Two were part of the 1976–77 drought, the driest two years in California’s history prior to 2014–15.
- Three were part of the 2012–15+ drought.
- One was part of the 1959–61 drought.
- One was part of the 1987–92 drought.

The next four driest water years after these were, in order: 2007, 1992, 1960, and 1959. While these were all very dry years, only three of them, 2015, 1977, and 1924, would rank as among the top dozen most severe drought years in the 1115-year period 900–2014. That is a reminder of just how severe a drought can be in a given year.

Tree-ring reconstruction shows that 1580 is the drought year of record in the Central Valley and the Southern Sierra. Water year 2015 will almost certainly be the second-driest. As explained under the section of this document that describes the 1918–34 drought, there is virtually a three-way tie among 1795, 1924, and 1977 as to which is the third-driest year in the San Joaquin Valley in the 1115-year period 900–2014. Based on stream gage data, we know that 1977 was a slightly drier year than 1924. However, we can’t say with any confidence where 1795 falls in this order, especially in the Tulare Lake Basin.

Based on tree-ring reconstructions, we know that the reconstructed flow on the upper San Joaquin at the inflow to Millerton Lake in 1580 was only 36% of that of the reconstructed flow in 1795 and 1924.
Summary of Floods

Summary of Past Megafloods
Botanic and geomorphic evidence indicates that a very large flood occurred in Northern California in about 1600. As detailed in Table 25, geomorphic evidence indicates that megafloods occur in Southern California on approximately a 200-year cycle. This bicentennial flooding was skipped only three times since 212 and never twice in a row.

The last skip was in the early 1800s. Even that skip may have only been a skip from the perspective of Southern California. It’s possible that the huge flood that the Central Valley experienced in 1805 belongs in this series.

<table>
<thead>
<tr>
<th>Approximate Date of Flood</th>
</tr>
</thead>
<tbody>
<tr>
<td>212</td>
</tr>
<tr>
<td>440</td>
</tr>
<tr>
<td>603</td>
</tr>
<tr>
<td>1029</td>
</tr>
<tr>
<td>1418</td>
</tr>
<tr>
<td>1605</td>
</tr>
</tbody>
</table>

It is uncertain whether the Northern California megaflood of about 1600 was one and the same as the Southern California megaflood of about 1600. In any case, it’s tempting to think that there would also have been floods in the Tulare Lake Basin during the 1600–1610 time period since the regions to our north, east, and south were experiencing immense precipitation events at this time. The Tulare Lake Basin would have been under the same general storm tracks.

Summary of 19th-century Flood History
Table 26 presents what we know about those floods that occurred between 1800–1849. Information on floods from this period is very incomplete. The information about these floods typically came from American Indians and from the earliest settlers such as John Sutter.

<table>
<thead>
<tr>
<th>Date of Flood</th>
<th>Type of Flood</th>
</tr>
</thead>
<tbody>
<tr>
<td>1805</td>
<td>Rain</td>
</tr>
<tr>
<td>1826</td>
<td>Rain</td>
</tr>
<tr>
<td>1847</td>
<td>Rain</td>
</tr>
</tbody>
</table>

High-water marks observed in the San Joaquin Valley and attributed to the 1805 flood were some six feet higher than the huge 1861–62 flood reached.
Floods and Droughts in the Tulare Lake Basin

Summary of Floods

Extensive settlement in California began around 1850 following the discovery of gold. Settlement of the Tulare Lake Basin began about that time as well. Between 1850 and 1900, a number of great floods occurred in the Central Valley. For this time period, we have much better data for the San Joaquin River Basin than for the Tulare Lake Basin.

Table 27 lists 12 of the major floods in the San Joaquin River Basin during the period 1850–1900. It excludes floods that clearly don’t appear to have been major floods in the Tulare Lake Basin. We have fairly decent to very good descriptions of five of those 12 floods: 1850, 1852–53, 1861–62, 1867–68, and 1890. The others were included in this table because they were identified by the USGS and/or USACE as having been a major flood in the San Joaquin River Basin. The effect of these floods on the elevation of Tulare Lake is reflected in Figure 15.

Table 27. Selected major floods in the San Joaquin River Basin: 1850–1900.

<table>
<thead>
<tr>
<th>Date of Flood</th>
<th>Type of Flood</th>
</tr>
</thead>
<tbody>
<tr>
<td>May–June 1850</td>
<td>Snowmelt</td>
</tr>
<tr>
<td>December 1852–February 1853</td>
<td>Rain</td>
</tr>
<tr>
<td>1861</td>
<td>Unknown</td>
</tr>
<tr>
<td>January–February 1862</td>
<td>Rain</td>
</tr>
<tr>
<td>December 1867–January 68</td>
<td>Rain</td>
</tr>
<tr>
<td>1869</td>
<td>Unknown</td>
</tr>
<tr>
<td>1872</td>
<td>Unknown</td>
</tr>
<tr>
<td>1878</td>
<td>Unknown</td>
</tr>
<tr>
<td>1884</td>
<td>Unknown</td>
</tr>
<tr>
<td>1886</td>
<td>Unknown</td>
</tr>
<tr>
<td>1889–90</td>
<td>Rain</td>
</tr>
<tr>
<td>1893</td>
<td>Unknown</td>
</tr>
</tbody>
</table>

In the northern part of the Central Valley, the 1861–62 flood is the flood-of-record on most rivers. The weather conditions during 1861–62 resulted in above-average precipitation between the Columbia River and the Mexican border. Major flooding was widespread throughout this area. However, the 1867–68 flood was especially severe on Sierra Nevada streams tributary to the southern part of the Central Valley. During recorded history, the 1867–68 flood was one of the greatest in the Tulare Lake Basin. Peak stages in that region during December 24–25 were the highest of record.

The 1867–68 flood also resulted in the deepest flood depths ever on the streets of Visalia. The 1861–62 flood put a maximum of 24 inches on Main Street while the 1867–68 flood put 5–6 feet.
Selected Floods in the Tulare Lake Basin Since 1905

Table 28 lists floods that generally had very high floodflows on at least one of the major rivers in the Tulare Lake Basin since 1905. Selecting floods to include in this list was somewhat arbitrary, just as our concept of what constitutes a flood is arbitrary.


<table>
<thead>
<tr>
<th>Date of Flood</th>
<th>Type of Flood</th>
<th>Kings River at Pine Flat (cfs)</th>
<th>Kaweah River at Terminus (cfs)</th>
<th>Tule River at Success (cfs)</th>
<th>Kern River at Bakersfield (cfs)</th>
</tr>
</thead>
<tbody>
<tr>
<td>March 1906</td>
<td>Rain</td>
<td>8,861</td>
<td>7,260</td>
<td>9,500</td>
<td></td>
</tr>
<tr>
<td>May–June 1906</td>
<td>Snowmelt</td>
<td>24,900</td>
<td>5,170</td>
<td>337</td>
<td>3,868</td>
</tr>
<tr>
<td>January 1909</td>
<td>Rain</td>
<td>9,578</td>
<td>8,226</td>
<td>12,787</td>
<td></td>
</tr>
<tr>
<td>December 1909</td>
<td>Rain</td>
<td>10,275</td>
<td>9,714</td>
<td>9,890</td>
<td></td>
</tr>
<tr>
<td>January 1914</td>
<td>Rain</td>
<td>10,540</td>
<td>13,520</td>
<td>11,232</td>
<td></td>
</tr>
<tr>
<td>January 1916</td>
<td>Rain</td>
<td>10,450</td>
<td>7,260</td>
<td>5,500</td>
<td></td>
</tr>
<tr>
<td>February 1937</td>
<td>Rain</td>
<td>13,520</td>
<td>13,520</td>
<td>17,727</td>
<td></td>
</tr>
<tr>
<td>February 1938</td>
<td>Rain</td>
<td>11,232</td>
<td>8,226</td>
<td>9,890</td>
<td></td>
</tr>
<tr>
<td>March 1943</td>
<td>Rain</td>
<td>9,714</td>
<td>9,714</td>
<td>9,890</td>
<td></td>
</tr>
<tr>
<td>February 1945</td>
<td>Rain</td>
<td>9,890</td>
<td>9,890</td>
<td>10,275</td>
<td></td>
</tr>
<tr>
<td>November 1950</td>
<td>Rain</td>
<td>16,640</td>
<td>5,918</td>
<td>5,918</td>
<td></td>
</tr>
<tr>
<td>January, 1952</td>
<td>Rain</td>
<td>5,918</td>
<td>5,918</td>
<td>5,918</td>
<td></td>
</tr>
<tr>
<td>May–June 1952</td>
<td>Snowmelt</td>
<td>15,500</td>
<td>5,170</td>
<td>337</td>
<td>3,868</td>
</tr>
<tr>
<td>December 1955</td>
<td>Rain</td>
<td>72,589</td>
<td>44,512</td>
<td>6,100</td>
<td>15,612</td>
</tr>
<tr>
<td>February 1963</td>
<td>Rain</td>
<td>34,612</td>
<td>18,405</td>
<td>12,227</td>
<td>22,259</td>
</tr>
<tr>
<td>December 1966</td>
<td>Rain</td>
<td>64,564</td>
<td>53,280</td>
<td>40,085</td>
<td>72,787</td>
</tr>
<tr>
<td>January 1969</td>
<td>Rain</td>
<td>40,513</td>
<td>22,437</td>
<td>22,259</td>
<td></td>
</tr>
<tr>
<td>September 1978</td>
<td>Rain</td>
<td>19,205</td>
<td>3,890</td>
<td>337</td>
<td>3,868</td>
</tr>
<tr>
<td>January 1980</td>
<td>Rain</td>
<td>33,283</td>
<td>16,933</td>
<td>13,036</td>
<td></td>
</tr>
<tr>
<td>April 1982</td>
<td>Rain</td>
<td>48,909</td>
<td>18,514</td>
<td>6,638</td>
<td></td>
</tr>
<tr>
<td>September 1982</td>
<td>Rain</td>
<td>30,415</td>
<td>6,308</td>
<td>586</td>
<td>6,673</td>
</tr>
<tr>
<td>December 1982</td>
<td>Rain</td>
<td>24,682</td>
<td>8,325</td>
<td>6,638</td>
<td></td>
</tr>
<tr>
<td>May 1983</td>
<td>Snowmelt</td>
<td>24,218</td>
<td>6,671</td>
<td>2,036</td>
<td>13,812</td>
</tr>
<tr>
<td>February 1986</td>
<td>Rain</td>
<td>25,060</td>
<td>9,428</td>
<td>7,528</td>
<td></td>
</tr>
<tr>
<td>March 1995</td>
<td>Rain</td>
<td>26,970</td>
<td>8,369</td>
<td>7,347</td>
<td></td>
</tr>
<tr>
<td>January 1997</td>
<td>Rain</td>
<td>50,217</td>
<td>17,948</td>
<td>18,780</td>
<td></td>
</tr>
<tr>
<td>November 2002</td>
<td>Rain</td>
<td>10,969</td>
<td>9,436</td>
<td>4,906</td>
<td>10,306</td>
</tr>
</tbody>
</table>

The flow data shown above are the peak daily flows (the average hourly flow for the peak day of the flood). Flows are expressed in cubic feet per second (cfs). Where necessary, these flows have been adjusted to remove the effects of dams upstream of the gage. Source: USACE, Sacramento District, and USGS.
### Table 29. Flood exceedence frequencies for selected floods: 1905–2011.

<table>
<thead>
<tr>
<th>Type of Flood</th>
<th>Kings River at Pine Flat</th>
<th>Kaweah River at Terminus</th>
<th>Tule River at Success</th>
<th>Kern River near Bakersfield</th>
</tr>
</thead>
<tbody>
<tr>
<td>March 1906</td>
<td>Rain 16%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>May–June 1906</td>
<td>Snowmelt</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>January 1909</td>
<td>Rain 14%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>December 1909</td>
<td>Rain 18%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>January 1914</td>
<td>Rain 13%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>January 1916</td>
<td>Rain 13%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>February 1937</td>
<td>Rain 10%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>February 1938</td>
<td>Rain 12%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>March 1943</td>
<td>Rain 14%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>February 1945</td>
<td>Rain 13%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>November 1950</td>
<td>Rain 8%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>January, 1952</td>
<td>Rain 25%</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>May–June 1952</td>
<td>Snowmelt</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>December 1955</td>
<td>Rain 1%</td>
<td>1.2%</td>
<td></td>
<td>8%</td>
</tr>
<tr>
<td>February 1963</td>
<td>Rain 5.3%</td>
<td>6.4%</td>
<td>20%</td>
<td>5%</td>
</tr>
<tr>
<td>December 1966</td>
<td>Rain 1.4%</td>
<td>0.6%</td>
<td>0.5%</td>
<td>0.3%</td>
</tr>
<tr>
<td>January 1969</td>
<td>Rain 4%</td>
<td>4%</td>
<td>1.5%</td>
<td>1.8%</td>
</tr>
<tr>
<td>September 1978</td>
<td>Rain 12%</td>
<td>29%</td>
<td>93%</td>
<td>30%</td>
</tr>
<tr>
<td>January 1980</td>
<td>Rain 5.5%</td>
<td>8%</td>
<td>9%</td>
<td>8%</td>
</tr>
<tr>
<td>April 1982</td>
<td>Rain 2.9%</td>
<td>6%</td>
<td>13%</td>
<td>13%</td>
</tr>
<tr>
<td>September 1982</td>
<td>Rain 6.5%</td>
<td>25%</td>
<td>84%</td>
<td>16%</td>
</tr>
<tr>
<td>December 1982</td>
<td>Rain 10%</td>
<td>18%</td>
<td>30%</td>
<td>25%</td>
</tr>
<tr>
<td>May 1983</td>
<td>Snowmelt</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>February 1986</td>
<td>Rain 9.5%</td>
<td>15%</td>
<td>17%</td>
<td>15%</td>
</tr>
<tr>
<td>March 1995</td>
<td>Rain 9%</td>
<td>18%</td>
<td>40%</td>
<td>15%</td>
</tr>
<tr>
<td>January 1997</td>
<td>Rain 2.5%</td>
<td>7%</td>
<td>8%</td>
<td>4%</td>
</tr>
<tr>
<td>November 2002</td>
<td>Rain 30%</td>
<td>15%</td>
<td>20%</td>
<td>10%</td>
</tr>
</tbody>
</table>

The flood exceedence values shown in the above table reflect the chance of a flood of a certain size or larger occurring in any given year. The exceedence frequencies were calculated by combining the peak daily flow rates from Table 28 with the appropriate flood frequency curve.

The rain-flood frequency curves were calculated by the USACE based on observed rain-floods for the following periods of record:
- Kings River below Pine Flat Dam: water years 1955–1978
- Kaweah River below Terminus Dam: water years 1905–2004
- Tule River below Success Dam: water years 1953–1988
- Kern River below Isabella Dam: water years 1953–2008

Table 30 presents the flood recurrence intervals for selected floods. Although that is the term used in this document, this concept will be found referred to elsewhere under a wide variety of terms, including:
- exceedence interval
- occurrence rate
- 50-year return period, etc.
- 50-year flood event, 50-year-flood, etc.
Table 30. Flood recurrence intervals for selected floods: 1905–2011.

<table>
<thead>
<tr>
<th>Flood Recurrence Intervals</th>
</tr>
</thead>
<tbody>
<tr>
<td>Type of Flood</td>
</tr>
<tr>
<td>----------------</td>
</tr>
<tr>
<td>March 1906</td>
</tr>
<tr>
<td>May–June 1906</td>
</tr>
<tr>
<td>January 1909</td>
</tr>
<tr>
<td>December 1909</td>
</tr>
<tr>
<td>January 1914</td>
</tr>
<tr>
<td>January 1916</td>
</tr>
<tr>
<td>February 1937</td>
</tr>
<tr>
<td>February 1938</td>
</tr>
<tr>
<td>March 1943</td>
</tr>
<tr>
<td>February 1945</td>
</tr>
<tr>
<td>November 1950</td>
</tr>
<tr>
<td>January, 1952</td>
</tr>
<tr>
<td>May–June 1952</td>
</tr>
<tr>
<td>December 1955</td>
</tr>
<tr>
<td>February 1963</td>
</tr>
<tr>
<td>December 1966</td>
</tr>
<tr>
<td>January 1969</td>
</tr>
<tr>
<td>September 1978</td>
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<tr>
<td>January 1980</td>
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<td>April 1982</td>
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<td>September 1982</td>
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<td>February 1986</td>
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<tr>
<td>March 1995</td>
</tr>
<tr>
<td>January 1997</td>
</tr>
<tr>
<td>November 2002</td>
</tr>
</tbody>
</table>

The recurrence intervals in the above table were calculated by combining the peak daily flow rates from Table 28 with the appropriate flood frequency curve. An alternative — and equally valid — approach would have been to combine the peak hourly flow rates with the corresponding flood frequency curve. This can result in noticeably different results. (See the January 1997 flood on the Kaweah for a comparatively extreme example. Using the daily flow rate, that flood had a recurrence interval of 14 years; but using the peak hourly flow rate, it had a recurrence interval of 25 years.)

One of the big lessons learned from preparing this document is that our rivers have been relatively quiet of late. Table 30 shows that the Tulare Lake Basin hasn’t really experienced any big floods in over 40 years. There haven’t been any 50-year floods or 100-year floods. The Kaweah and Tule Rivers haven’t even seen any 20-year floods during this time period.

That probably doesn’t mean anything; it’s just a statistical coincidence. The problem is more psychological. When we don’t experience a big flood for a while, we tend to forget just how big our floods can be. We have come to think of the federal reservoirs and our levees as protecting us from the effects of big floods, and that isn’t necessarily realistic when we consider our flood history.

Those of us who live and work above the federal reservoirs have a greater responsibility to become involved in planning for floods than those who live below the dams. The last really big flood in the Tulare Lake Basin was the December 1966 flood. It’s sobering to reflect back on the experience of that flood. Fifteen-foot waves were reported to have been common on the mainstem of the Kaweah in Three Rivers. Today we think of Dry Creek below Terminus Dam as not much more than a quiet foothills stream. However, in the 1966 flood, Dry Creek carried 44% more water than the Merced River in Yosemite Valley during the much more famous January 1997 flood.

The take-away message is that it would be prudent to prepare for big floods, much bigger than we have been experiencing during the last 40 years. This is particularly important for those of us who live and work in areas that aren’t protected by a federal reservoir.
Summary of Floods and Droughts since 1849

This document describes what we know about approximately 188 floods and 36 multi-year droughts that have occurred during the last 2,000 years. The floods of 1849–50 are the first for which there are fairly accurate historic descriptions. Figure 25 illustrates the 178 floods and 17 multi-year droughts that have occurred since 1849.

This document sometimes groups floods from adjacent years together. This can occur when floods wrap over the winter months as in 1955–56. This can also occur when two years of floods run together as in 1982–83. In Figure 25, the floods are shown by the first year of the two-year flood grouping.

Droughts are generally illustrated for the number of years that the drought was active somewhere in the state. The drought may not have been active each of those years in the Tulare Lake Basin. See the section of this document that describes each drought to learn what we know about local conditions during those years.

This graph illustrates how often floods occur during multi-year droughts. That was one of the lessons learned from preparing this document. Floods occur at all manner of times. When they occur varies so widely because there are such a variety of causes for floods.

Figure 25 illustrates far more individual flood events in recent years than in earlier years. Certainly 1982–83 was a period of exceptional flooding events. However, in general this document reflects more flood events from recent years because of data availability. In the last few decades, there has been an explosion in information that society records and puts on the Internet. In many ways, this graph of flood frequency reflects that explosion in information availability.

This document describes 10 floods that occurred before 1849; the earliest of those was dated about A.D. 212. See the section of this document that describes the California megafloods for more about that event. Slightly over half of the droughts described in this document, 20, occurred before 1849; the earliest began in about A.D. 900. See the section of this document that describes Megadroughts before the Little Ice Age for more about that drought. In general, the farther back that you look, the bigger the flood or drought has to be in order to be detectable.
Figure 25. Known floods and multi-year droughts for past 167 years: 1849–2015.
Specific Floods and Droughts

Megadroughts before the Little Ice Age

Parts of California experienced two extremely long megadroughts during the Medieval Warm Period aka Medieval Climatic Anomaly (MCA). The MCA lasted from about A.D. 950–1250 (or 800–1300, depending on the source). The MCA was the warmest period of the last 2,000 years prior to the 20th century. After a transition period, the MCA was followed by the Little Ice Age (approximately 1450–1850).

These two megadroughts occurred during the MCA, a period of above-average warmth in the Western U.S. However, there is no evidence that the MCA was a time of global warming of strength comparable to the present day or of model projections of the current century. Climatologists think that the drought conditions of the MCA were caused by more than just elevated temperature.\(^{582}\)

Climatologists think that something about the climate during that period affected the tropical Pacific Ocean in such a way as to push the storm tracks further north, and this brought drought to much of the Western U.S. Parts of California, including the Sierra, were subject to prolonged, severe droughts from about A.D. 800–1400.\(^{583}\) The driest two periods in western North America were centered on the mid–1100s and the mid–1200s. Both of these periods are reflected in the Sierra.\(^{584, 585}\) These prolonged droughts caused large lakes to shrink or dry out completely, and more frequent wildfires.\(^{586}\)

A 1994 study by Scott Stine showed that these were epic drought periods in California.\(^{587, 588}\) Stine’s research was based on drowned tree stumps rooted in Mono Lake, Tenaya Lake, West Walker River, and Osgood Swamp in the Central Sierra. He concluded that runoff from the Sierra during those periods was significantly lower than during any of the persistent droughts that have occurred in the region over the past 140 years. Stine found that the first of these droughts lasted more than two centuries before the year 1112; the second drought lasted more than 140 years before 1350. Both of these droughts are reflected in parts of the Central Sierra. The first of these droughts also impacted Patagonia in South America.

At Fallen Leaf Lake (located near the southern end of Lake Tahoe), trees up to 82 feet in height have been found rooted upright 118 feet below the current lake surface. Those trees became established and grew during this megadrought.

By the early 1980s, the Los Angeles Department of Water and Power’s (LADWP) diversions had caused Mono Lake to drop to elevation 6,372 feet, roughly 50 feet below its natural elevation. This drawdown exposed thousands of acres of lake bottom along with many cottonwood and Jeffrey pine stumps. Scott Stine sampled those stumps and had them radiocarbon dated. They fell into two groups: one that had died in about 1100 and the other in about 1350. He concluded that there had been two megadroughts in the area that brought Mono Lake down far enough to allow trees to become established and grow to considerable size.

Scott found a similar situation at Tenaya Lake in Yosemite National Park, where a dozen trees protrude from this glacial lake. Some individual trees are standing in nearly 70-foot-deep water. When Scott had those trees radiocarbon dated, he found that they fell into similar groups as at Mono Lake. One group had died in about 1100 and the other in about 1350. A count of the annual rings showed that the first drought (the one ending in about 1100) had lasted at least 140 years, and the second drought (the one ending in about 1350) had lasted at least a century. Phil Catarino and Eric Henningsen, the divers who worked at Fallen Leaf Lake, have confirmed that the bases of the Tenaya trees are rooted in growth position.

Scott found relic stumps of Jeffrey pines in West Walker River along U.S. Highway 395.\(^{589}\) Since the root systems of Jeffrey pines can tolerate complete inundation for only a couple weeks at most, he reasoned that there must have been an extended drought in the area while these trees were growing. Sure enough, they dated to the same periods as the two megadroughts at Mono and Tenaya Lakes.

During much of the 20th century, Owens Lake has been reduced to a playa due to water diversions by LADWP. In the 1990s, Scott went out into the Owens Playa as part of an archeological assessment. The search turned up not only archeological materials (some of them at very low elevations on the playa), but a rooted, tufa-encrusted shrub stump as well. Radiometric analysis of the shrub stump indicated that it comported in age with the circa 1100 megadrought sites from farther north.
Susan Lindstrom, while diving in Independence Lake north of Tahoe, discovered stumps that may have been comparable with the second MCA megadrought that ended around 1350.\textsuperscript{590} The dates given by Stine for these two megadroughts were based on many calibrated radiocarbon dates with approximate 50-year standard deviations.

In a study published in 2007, Nick Graham and Malcolm Hughes used long tree-ring-derived streamflow reconstruction for the Merced River to much more accurately date the two droughts that caused the low stands at Mono Lake during the MCA.\textsuperscript{591} The Merced River and Mono Lake Basins share a common boundary along the Sierra. The Graham study showed that there was a very close correlation between fluctuations in runoff for that of the Merced River catchment on the western slope of the Sierra and the Mono Lake immediately to the east.

This allowed the researchers to use the reconstructed flows for the Merced to estimate the flows feeding Mono Lake. The results showed two deep low stands of the lake during MCA times that agreed closely in magnitude and timing with those that Stine inferred from relic vegetation:
- 110-year drought lasting from AD 900–1009
- 99-year drought lasting from AD 1176–1274

The Merced River annual streamflow deficits during these MCA droughts were far more severe than anything experienced during historic times, and have not been closely approached in the past 550 years, with (reconstructed) centennial average flow reaching 75% of modern annual averages and decadal averages reaching 60% of that value. In comparison, during the most severe drought in the observed record, decadal average discharge reached 75% of the 1901–1994 mean.

The Graham study also showed that the MCA droughts (as they appear in the flow reconstruction) were marked more by the paucity of years with near-normal flow than by the magnitude of flow deficits in a relatively few dry years. There were occasional periods of average or above-average precipitation. However, despite occasional relief, it was generally dry for decade after decade. There were also periods of particularly dry years that occurred during these very long-duration megadroughts. These can be thought of as mini-droughts within the megadroughts. This emphasized the fact that centuries of MCA drought were apparently characterized by a marked and persistent change (relative to today) in winter circulation patterns over the North Pacific and North America.

After determining the date of the Mono low stands, the researchers used a 2004 PDSI dataset from Edward Cook, Dave Meko, and others to determine the area impacted by these two MCA droughts.\textsuperscript{592}
- The first MCA drought (AD 900–1009) extended from a core region in Central and Southern California into Arizona and northeast through the Great Basin and into Wyoming and Montana. PDSI averages were near their long-term mean in Oregon and Washington; that area was not affected by the drought.
- The second MCA drought (AD 1176–1274) was less severe than the first, was again focused most extensively in Central and Southern California and extended northward through western Nevada and into Idaho. This second drought was thus confined more to the far Western U.S. than the first. As in the first drought, conditions in the Pacific Northwest were near normal.

The study showed the Tulare Lake Basin and the San Joaquin River Basin to be located in the core region of both droughts, the area that had the greatest PDSI anomalies. Based on those results, the average runoff from AD 900–1009 and 1176–1274 on the upper San Joaquin River at the inflow to Millerton Lake should have been about 75% of the long-term average. But Dave Meko ‘s recent tree-ring reconstructions for this basin found that flows were actually 97% and 103% for these two drought periods.\textsuperscript{593}

In a study published in 2009, Edward Cook and his team used a new national drought atlas (created from tree-ring reconstructions) to better estimate the dates of the two MCA megadroughts that Stine had described:\textsuperscript{594}
- 243-year drought lasting from AD 832–1074
- 178-year drought lasting from AD 1122–1299

Cook’s study used the national drought atlas to map those droughts and indicated that the San Joaquin River and Tulare Lake Basins were located within the most intense portion of both.

As described in the section of this document that discusses Runoff Reconstructions, a team led by Dave Meko later used tree-rings to reconstruct the flow on the San Joaquin River and its major tributaries for 1113 years (900–2012).\textsuperscript{595} This provided significant clarity about what the climate was in the San Joaquin basin. When doing 30-year filtering, Dave’s results showed a dry period from the beginning of the reconstruction in 900 to
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

about 1000, and another dry period in the mid-1000s. But the severity was only about 80% of average flow even with that narrow smoothing. When averaged over longer periods (e.g., 100 years), wet periods cancelled the dry, giving near-average reconstructed flows.

Dave said that there is a strong correlation between flows in the Merced and San Joaquin Rivers. He thinks it is unlikely that a massive, multi-century drought would affect Mono Lake levels and Merced River flows but have no footprint in the San Joaquin. Drought might be more or less severe from one basin to the next, but it should show up to some extent. He can only say that the tree-ring chronologies his team used that had data covering the MCA drought periods do not show those droughts with similar severity and duration as implied by the Graham and Hughes Mono Lake study. Dave speculated that because the two earlier studies had few tree-ring chronologies to represent the San Joaquin River Basin, they had to reach further out to find one that they could use. Depending on how far the search radius went out, it could have pulled in chronologies from far outside the San Joaquin River Basin. If so, that could be an explanation for the discrepancy in reconstructions.

The research by Scott Stine, Nick Graham, Malcolm Hughes, and others found very good evidence for the two epic MCA droughts at Mono Lake, Merced, Tuolumne, and Owens Lake. Stine’s work on Mono and Owens Lake fluctuations showed that there were some very wet periods between those droughts. The lake levels in both of those lakes rose rapidly in relatively short time periods. Owens Lake rebounded between the droughts to one of its highest Holocene levels which would require very wet conditions.

Clearly the climate in those areas was going through major changes. It would seem that for a century plus, the storm tracks didn’t come as far south as the Tuolumne; then they resumed coming as far south as Owens Valley for a century plus; then they again failed to come as far south as the Tuolumne for a century plus; then they resumed coming as far south as the Owens Valley.

Given our experience with the modern climate and storm tracks, we would expect that the climate patterns that affected those basins (Tuolumne south to the Owens Valley) in the MCA would have also impacted the Tulare Lake Basin. From what we know about how the storm tracks work, it seems like our basin should have been impacted when the other basins were.

But that leaves us with a conundrum. We have the work by Stine, Graham, Hughes and others that represent solid evidence for the two period of epic drought to the north and east of our basin. It seems very likely that the San Joaquin River and Tulare Lake Basins would have been within both of those droughts. However, we have Dave Meko’s study clearly showing that the drought was not present at all in the San Joaquin River Basin, not even in the upper basin. The San Joaquin River Basin immediately to our north was not experiencing any significant change in precipitation during the MCA; the storm tracks were a relative constant.

The scientists who deal with these issues don’t yet know how to explain that unexpected finding, or what it means for our basin.

• One possibility is that both the San Joaquin River and Tulare Lake Basins really were in severe drought despite the findings of the Meko study.
• Another possibility is that the drought affected the Merced River and Tulare Lake Basins, but not the San Joaquin.
• A third possibility is that neither the San Joaquin nor the Tulare Lake Basin experienced drought during the time that those huge droughts were active to our north and east.

We just have to wait until the scientists do more work to resolve this issue.

From reviewing studies such as Scott Stine’s, DWR concluded that California is subject to droughts far more severe and prolonged than anything witnessed in the historical record.596

In a 1996 study, Lynn Ingram and others concluded from their analyses of variations in the isotopic composition of fossil shells in the San Pablo Bay (the northern extension of San Francisco Bay), that “alternate wet and dry (drought) intervals typically have lasted 40 to 160 years” in the San Francisco Bay drainage area (Sacramento River and San Joaquin River Basins) over the past 750 years (since about 1250).597

George Durkee (national park wilderness ranger) collected a core from a stump above the Crabtree Ranger Station and sent this to Scott Stine who dated it. Radiocarbon dating showed that the tree died in the 1100s, so it may have been contemporaneous with the megadrought of the 1100s. That is, that stump apparently became rooted during that drought.
Tony Caprio and Linda Mutch studied how the Mountain Home Grove of giant sequoias responded to a fire that occurred in 1297. Their research suggested that this fire event was of unusually high severity, not equaled over the last 2,000 years. It appeared to result in the mortality of most non-sequoias and a considerable number of giant sequoias.

Time series of precipitation, reconstructed from tree-rings, showed that a long drought occurred between 1292–96 throughout the Southern Sierra. Tony and Linda speculated that this drought was a contributing factor in creating the conditions for the Mountain Home fire that occurred in 1297.

As shown in Table 21:
- In the first of the two MCA megadroughts, the 243-year drought that lasted from AD 832–1074, conditions were particularly dry on the upper San Joaquin during the 10-year period 975–984.
- In the second megadrought, the 178-year drought that lasted from AD 1122–1299, conditions were particularly dry on the upper San Joaquin in the 20-year period 1149–1158. This was also an exceptionally dry period in the Sacramento River Basin. A reconstruction of the Sacramento River streamflow developed from tree-rings found that the most intense 20-year drought of about the last 1,000 years was between 1140–1160.

The megadrought in the mid-1100s was simultaneously in effect on the drainages of the Colorado, Sacramento, and San Joaquin River Basins. However, the drought was less severe on the San Joaquin than on the Sacramento River Basin. Or at least that is how it has generally been perceived.

As shown in Table 19, average reconstructed runoff for these two drought (832–1074 and 1122–1299) on the upper San Joaquin was only 1%–3% below the 1113-year average (900–2012). These two megadroughts were severe events to our north and east, but Dave Meko's research shows that they really didn't happen in the San Joaquin Basin, not even in the upper San Joaquin River Basin. Based on that research, the San Joaquin Valley was apparently right on the edge of a record-setting megadrought.

The last severe megadrought that impacted the Great Basin and adjacent parts of California ended in about 1350. One source said that just as that drought was ending, the Sacramento Valley experienced an extended drought from 1350–1400. But that does not appear to be the case.

The San Joaquin River Basin would experience a megadrought that began in about 1444. The driest 50-year period in the San Joaquin River Basin as a whole was between 1450–1500. The years 1450–1500 were also the driest 50-year period ever on the upper San Joaquin at the inflow to Millerton Lake during the 1115-year period 900–2014.

"Prolonged droughts — some of which lasted more than a century — brought thriving civilizations, such as the ancestral Pueblo (American Indians) of the Four Corners region, to starvation, migration and finally collapse, “Lynn Ingram, a geologist at the UC-Berkeley, wrote in her recent book The West Without Water.

The California megadroughts of the 900s and 1100s coincided roughly with a warmer climate in Europe, which allowed the Vikings to colonize Greenland and vineyards to grow in England, and with a severe dry period in South America, which caused the collapse of that continent’s most advanced pre-Inca empire, the rich and powerful state of Tiwanaku.

A study published in 2011 by David Stahle and others provided a 1,238-year-long tree-ring chronology for ancient Mexico and Central America (or Mesoamerica). The study was the first to reconstruct the yearly climate of pre-colonial Mexico over more than a millennium, pinning down four ancient megadroughts to their exact years.

The study team looked at 30 specimens of millennium-old Montezuma bald cypress trees growing near Tenochtitlan, capital of the Aztec empire, and Tula, the Toltec state’s main city. This tall tree species, related to North American giant sequoias, is the only plant in Central America that frequently lives up to 1,000 years or more.

One large ancient drought previously confirmed for the U.S. Southwest was shown to have extended into central Mexico from 1149–1167. That drought may have devastated the local maize crops, a potentially fatal blow to the declining Toltec culture.
The new record also pinned down more precisely than ever before the timing of two other severe dry periods, possibly leading to new insights into the Aztecs’ rise to power and into the spread of exotic diseases that Spanish conquistadores brought to America. The paleoclimate reconstruction study also confirmed the so-called Terminal Classic drought that some anthropologists relate to the collapse of the Mayan civilization.

The 10th to the 14th centuries, encompassing the two prolonged MCA droughts reported by Scott Stine, saw dramatic changes in Western Hemisphere civilizations, and some have been attributed by archeologists to changes in rainfall. One example is the sudden decline of the Anasazi cliff-dwellers in the American Southwest at about the year 1300. An even more striking example is the collapse of Tiwanaku.

Tiwanaku was a flourishing empire that lasted seven centuries and reached its zenith near the end of the first millennium A.D. It commanded an area about the size of California, extending from the Andean plateau around Lake Titicaca to the Pacific Coast and covering parts of present-day Bolivia, Chile, and Peru.

The empire’s economy was based on intensive agriculture carried out on raised fields: acres of end-to-end rectangular platforms created by digging the dirt from areas between them. The dug-out areas became canals from which silt was taken to provide fertilizer. This highly productive and environmentally sound system dominated Latin American agriculture for centuries.

But an extended drought afflicted that region starting between A.D. 950–1000 and continued, with fluctuations, until 1410. That period mostly overlaps the one in which California’s two epic megadroughts occurred.

The South American drought was of horrendous proportions, and it destroyed Tiwanaku’s agricultural base. The empire’s cities were abandoned by about 1000. The raised fields could no longer support the cities, and archeological evidence shows that the fields were abandoned between 1000–1100. The political state collapsed, the population dispersed, and with agriculture no longer possible, the people relied on raising alpacas and llamas.

Tiwanaku’s agricultural system had been able to adjust to the less drastic cycles of drought and inundation that were thought to be normal, but Tiwanaku agro-engineers were incapable of responding to a drought of unprecedented duration and severity.608

The downfall of the Mayan civilization has been linked with two prolonged droughts that occurred at roughly the same time as the California and Tiwanaku droughts.609 Following the first Mayan drought, probably between A.D. 800–1000, the Mayans emigrated north, only to be hit by another prolonged drought between A.D. 1000–1100. That proved to be the terminal blow for their civilization, coinciding with the fall of the city of Chichen Itza.

During the MCA, parts of California experienced two epic megadroughts: roughly 243 and 178 years long. Such extremely long megadroughts ended with the beginning of the Little Ice Age (approximately 1450–1850). Climatologists think that this cooling of the climate affected the tropical Pacific Ocean which resulted in the storm tracks coming further south, and this brought increased moisture to much of the Western U.S. including the Tulare Lake Basin.

Subsequent megadroughts (droughts of much longer than average length) have been significantly shorter in duration. They have also been more severe in the southern San Joaquin Valley then further north, suggesting that the storm tracks have not extended as far south during these episodes.

What would happen if a megadrought were to occur in today’s California? Jay Lund (director of the UC Davis Center for Watershed Sciences), Scott Stine (professor emeritus of geography and environmental science at Cal State East Bay), and other researchers decided to answer that question. They used computer modeling to simulate what would happen if the California of today experienced a 72-year-long megadrought, one in which annual runoff amounted to only about half the historical average.610

The results were somewhat surprising. Based on their simulation, the California economy would not collapse. Traumatic changes would occur as developed parts of the state shed an unsustainable gloss of green and adjust to the new reality. However, the state would not shrivel into a giant, abandoned dust bowl.
Under the model scenario,

- Urban water rates would climb. The iconic suburban lawn would all but disappear. Cities largely would do okay aside from higher water costs, since they have the most financial ability to pay for water. They would do more water conservation and wastewater reuse, a little ocean desalination, and would purchase some additional water from farms.
- In Southern California, withering decades of drought would speed up the region’s current move to expand local water sources and reduce dependence on increasingly erratic supplies from Northern California, the Eastern Sierra, and the Colorado River.
- As the sector with the greatest water use in California by far, agriculture would be most heavily impacted by the drought. Agriculture would shrink, but by no means disappear.
- The state’s 8 million acres of irrigated cropland would decrease by as much as half. Farmers would largely abandon relatively low-value crops such as cotton and alfalfa and use their reduced water supplies to keep growing the most profitable fruits, nuts and vegetables. That is what happens now in the most severe droughts.
- Farmers would either let abandoned farmland revert to scrub, or dry-farm them with wheat and other crops that once predominated in the Central Valley.
- Some farm communities would turn to ghost towns. Farm workers would suffer.
- There would be significant economic impacts within the agricultural sector. However, agricultural production and associated industries such as food processing make up only about 4% of California’s overall economy. Overall, California has a remarkable ability to weather extreme and prolonged droughts from an economic perspective. (We need to remember that much of the country and the world has come to depend on California for a significant portion of their food production. There would be significant impacts on food supply and price if that food source were greatly reduced due to a severe and prolonged drought.)
- Aquatic ecosystems would suffer, with some struggling salmon runs fading out of existence.

**Summary of megadroughts since the Little Ice Age**

Megadroughts are defined more by their duration than their severity. They are extreme dry spells that can last for a decade or longer, according to research meteorologist Martin Hoerling of NOAA and climate scientist Toby Ault at Cornell University. There is no universally accepted definition for what qualifies as a megadrought. Some climatologists define a megadrought as a drought that is of much longer than average duration. Others define a megadrought as a drought that is longer than 30 years. This document generally uses the former definition: a drought that is of much longer than average duration.

As shown in Table 20 and Table 22, most of the droughts in the Tulare Lake and San Joaquin River Basins last 2–4 years. (A single dry year isn’t generally considered a drought.) We are only aware of 10 droughts in the last 11 centuries that have lasted 6 or more years. As shown in Table 19, only five of those have been megadroughts, lasting longer than a decade.

As described in the section of this document on the Megadroughts before the Little Ice Age, the duration of megadroughts has become shorter with time. The first two lasted several centuries each. The next two lasted about four decades each. The only one that has occurred in the last four centuries lasted just under two decades. Our standard of what constitutes an unusually long drought has changed as the climate has changed.

The 40-year-long megadrought from 1444–1483 occurred just as the Little Ice Age was beginning. It was more severe in the San Joaquin River Basin than in the Sacramento River Basin. This suggests that the storm tracks did not come as far south during those years.

The 37-year-long megadrought from 1566–1602 was more severe in the Tulare Lake Basin than further north, again suggesting that the storm tracks did not extend as far south in those years.

The 17-year-long megadrought from 1918–34 has been the only megadrought in the last four centuries so far. It is generally considered the most severe decade-long drought experienced in the Central Valley during the 1115-year period from 900–2014.

At decadal and longer time scales (e.g., driest 20 years ever), pre-20th century low-flow extremes are dominated by periods in the mid- to late-1100s in the Sacramento River Basin, while the late 1400s were more severe in the San Joaquin River Basin. However, prolonged drought occurred for both of these periods in both basins.
Floods and Droughts in the Tulare Lake Basin
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That comparison was made treating the San Joaquin River Basin as a whole. That is, the equivalent of the modern SJQ4 gage, a summary series defined by CDEC as the total San Joaquin River runoff. However, the various sub-basins within the San Joaquin River Basin were not equally affected during past droughts.

For example, the years 1451–1465 were the driest 15-year period ever on the total San Joaquin River Basin. However, the years 1920–34 were the driest 15-year period on the upper San Joaquin at the inflow to Millerton Lake.

Likewise, the years 1446–1465 were driest-ever 20-year period on the total San Joaquin River Basin. However, the years 1917–1936 were the driest ever on the upper San Joaquin.

Those two comparisons reflect that the 1918–34 drought was more severe in the southern part of the San Joaquin River Basin than even the driest part of any of the preceding megadroughts.

Table 21 illustrates how severe different droughts were on the upper San Joaquin as measured over different time periods.

Park Williams, a bioclimatologist at the Lamont-Doherty Earth Observatory of Columbia University, said more area in the Western U.S. (defined as west of 95 degrees West) has persistently been in drought during the 15-year period from 2000–14 than in any other 15-year period in more than 850 years, since the 1150s and 1160s. The year 2015 will almost certainly be a drought year, extending this to a 16-year period.

This finding is consistent with the hydroclimatic reconstruction from tree-rings in the Colorado River Basin. It may well apply to many other river basins in the West as well. As described in the section of this document on the Megadroughts before the Little Ice Age, a severe drought in the mid-1100s was simultaneously in effect on the drainages of the Colorado and Sacramento Rivers.

The Western U.S., taken as a whole, has been in drought for the 16-year period 2000–15. However, the specific area affected by the drought has moved around each year. The San Joaquin River Basin has been on the edge of the drought and has only been affected by it for 12 of those 16 years: 2000–04, 2007–09, and 2012–15 (based on the San Joaquin Valley Water Year Index and/or total runoff for our four major rivers).

From our perspective, this means we haven’t been in a megadrought. Instead, we have experienced three droughts of relatively average duration: the 2000–04, 2007–09, and the 2012–15+ droughts. However, if we were to step back and look at the bigger picture, we are really on the edge of a record-setting megadrought. This is similar to the situation when we may have been on the edge of the megadroughts of 832–1074 and 1122–1299.

Potential for Future Megadroughts

According to a 2007 study by Richard Seager and others, the Southwest’s aridity is about to get worse. The study predicted that climate change will permanently alter the landscape of the Southwest so severely that conditions reminiscent of the Dust Bowl days of the 1930s could become the norm within a few decades.

The study suggested a perpetual arid condition over the Southwest. Of the 19 different computer models that the research team used for the study, all but one showed a drying trend in the swath of North America between Kansas, California, and northern Mexico. The models predicted an average 15% decline in runoff for the Southwest between 2021–2040, compared to the average surface moisture between 1950–2000.

The Southwest’s future droughts are expected to be of a different nature than those that have afflicted the region in the past. Scientists attribute past droughts to variations in sea surface temperatures caused by El Niño and La Niña events in the Pacific Ocean. La Niña is especially influential as it tends to shift precipitation belts north, leaving the Southwest thirsty.

As the climate warms, however, the basic dynamics of the atmosphere change, particularly in regard to the Hadley cell, a powerful circulation pattern that drives weather in the tropics and subtropics. The model projections in the 2007 study were based on an understanding of the fundamental dynamics of the Hadley cell.

Warm, moist air from near the equator normally rises into the atmosphere until it reaches the stratosphere, the second layer of Earth’s atmosphere. The air then spreads north and south toward the poles, descends over the subtropics, and flows back toward the Equator in the form of trade winds, completing the cell. The Hadley cell is
this pattern of atmospheric circulation in which warm air rises near the equator, cools as it travels poleward at high altitude, sinks as cold air, and warms as it travels equatorward near the surface.

Because the descending air over the subtropics suppresses rain by drying the lower atmosphere, the Saharan and Arabian deserts and the deserts of northern Mexico are located in these regions.

As the atmosphere warms from climate change, scientists expect the Hadley cell to expand its reach, bringing hot, dry air to a larger swath of the Middle East, Mediterranean, and North America, including the Southwest. The 2007 study found that in the future warmed climate, the Hadley cell and the subtropical high should expand poleward, which tends to block rain coming through from the Pacific.

The various model projections were not in agreement as to when the drought effects would begin or how broad an area would be affected. The Southwest is already experiencing changes that scientists link to climate change, including more severe wildfires, earlier winter snowmelt, the destruction of heat-weakened trees by beetles, and a loss of biodiversity in southern Arizona’s high-elevation sky islands.

A study published in 2014 by Toby Ault, a climate scientist at Cornell University, and others looked at the risk of megadroughts in the near future. The authors found that state-of-the-art global climate models appear to underestimate the risk of persistent droughts in semiarid regions.

The Ault study developed methods to more accurately assess the risk of such events in the coming century using climate model projections as well as observational (and reconstructed paleoclimate) information. The authors found that the chance of the Southwest experiencing a decade-long drought is at least 50%, and the chance of a drought that lasts over 30 years ranges from 20% to 50% over the next century.

The authors said that as we continue to add greenhouse gases into the atmosphere, we are weighting the dice for megadrought conditions for the Southwest. With ongoing climate change, long-drought scenarios become increasingly likely.

The Tulare Lake Basin is right on the edge of the highest-risk Southwest drought area identified in the Ault study. There is a possibility that we are already seeing a replay of a megadrought similar to those of 832–1074 and 1122–1299. As described in the section on Megadroughts since the Little Ice Age, California’s last three droughts (2000–04, 2007–09, and the 2012–15+ droughts) have been part of a longer-term megadrought across most of the Western U.S. since 2000.

A study published in 2015 by Ben Cook, a climate scientist at NASA’s Goddard Institute for Space Studies, and others looked at the risk of megadroughts in the Southwest and Central Plains during this century. Cook’s team analyzed climate models that include historical records and looked at drought trends revealed in tree-rings over the last 1,000 years. Based on this, they predicted an 85% chance of a drought lasting 35 years or more between 2050–2100.

The future of drought in western North America is likely to be worse than anybody has experienced in the history of the U.S. These are droughts that are so far beyond our contemporary experience that they are almost impossible to think about.

Climate models indicate that megadroughts threaten the Southwest this century. These models predict the Southwest will become drier in the future as climate change shifts global weather patterns. In the Central Plains, the northern areas may become wetter and the southern areas may become drier.

However, Cook’s study found that rising temperatures are a bigger threat than a lack of rain and snow. Compared with the great dry spells that hit 1,000 years ago, this century will see higher temperatures in the Southwest because of greenhouse gas emissions. The warmth will increase drought severity by increasing evapotranspiration.

Calendar year 2014 was remarkably warm in California, the West, the contiguous U.S., and for the Earth as a whole. This fits within a context of a long-term warming trend that has been going on for several centuries and has been accelerating in recent decades (see the section of this document that describes Long-term Temperature Changes).

The rising temperatures will favor longer, more severe droughts. The Cook study found that this will happen even in those areas of the West and Central Plains that may actually see more, not less, winter rain and snow in
the future. The lack of precipitation is not the biggest concern. It is the rise in temperature that is driving the model simulations toward dramatic aridity.

The study results are more conclusive for the Southwest than for the Central Plains. One reason for the uncertainty is that tree-ring records in the West are exquisitely detailed compared with the history available for the Plains states.

The Cook study findings reveal future droughts will arise from different causes than previous droughts. The megadroughts during the 1100s and the 1200s were triggered by natural ocean-atmosphere cycles, such as cooler temperatures in the Pacific Ocean that shifted weather patterns toward drier conditions in the Southwest. But the drying effect from warming will eventually overwhelm those cycles in the Southwest and Central Plains.

Those historical droughts eventually ended, and we moved on to wetter periods. This future drying represents a fundamental change and shift toward drier average conditions in western North America. Because of greenhouse gases, there is little evidence we're going to bounce back to the conditions we have experienced in historic times.

The main distinction between Ault's 2014 paper and Cook's 2015 paper is that Cook focused on soil moisture, which factors in temperature, whereas Ault only considered the precipitation "supply side" of megadrought. The two studies therefore bracket a pretty realistic range of risks, with Cook's being more appropriate for most applications. The chief source of uncertainty comes from the frequency of El Niño events, but even if these do become more common under rising temperatures, it is unlikely that they will offset future evaporative demands of the atmosphere. Cook's paper also shows that the megadroughts of the past are likely to be much smaller in amplitude than the ones we could see this century if climate change goes unmitigated.

The 2012–15+ is the most severe drought in at least the last century and probably much longer. Given the increased temperatures that we have been experiencing during the last several decades, it is possible that this could be a precursor of extended, severe droughts similar to those that have occurred in the past millennia.

**California's Six Megafloods (A.D. 212–1605)**

Arndt Schimmelmann did research based on sediment cores taken in the Santa Barbara Basin and published his results in two scientific papers. Schimmelmann found evidence for a huge flood striking Southern California approximately every 200 years, centered on the years shown in Table 31. This cyclic megaflood is generally described as a Southern California event because that is where the research was done. However, the flood presumably strikes Northern and Central California as well.

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<th>Approximate Date of Flood</th>
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The bicentennial flooding in Southern California was skipped only three times since 212 and never twice in a row. The quasi-periodicity of approximately 200 years for Southern California floods recorded in the Santa Barbara Basin matches the approximate 200 year periodicities found in a variety of high-resolution paleoclimatic archives. More importantly, it matches the roughly 208-year cycle of solar activity (the Suess Cycle) and inferred associated changes in atmospheric circulation. The last skip was in the early 1800s, leading the authors of one of the Schimmelmann papers to conclude that "we foresee the possibility for a historically unprecedented flooding in Southern California during the first half of the 21st century."

The skip in the early 1800s may have only been a skip from the local perspective of the Santa Barbara Basin. The Central Valley experienced a huge flood in 1805, one that was even bigger than the huge 1861–62 flood. Perhaps the storm track that year just didn’t extend far enough south to be recorded in the Santa Barbara Basin.

The 1605 event was measured with a precision of ±5 years. The dating of that flood event is consistent with tree-ring evidence for a wet and cold paleoclimate elsewhere in the region. The depth of the silt layer deposited...
in the Santa Barbara Basin in the 1605 event implied an intensity of precipitation and flooding of the Ventura and Santa Clara Rivers unmatched in the last 1,000 years.

We tend to think of floods as relatively short-term events, lasting for no more than a few months. However, judging from what we know of the 1605 flood, megafloods appear to be a much longer-term type of event. Once such an event begins, it can last for up to 10 years. During that period, multiple episodes of flooding and extreme runoff can occur as well as other unusual climatic events. The 1605 flood was associated with a large-scale change in climate that affected the Northern Hemisphere from roughly 1600–1610.

The decade 1600–1609 stands out as the coldest in a 570-year (A.D. 1400–1970) comprehensive record of summer temperatures across the Northern Hemisphere, based on tree-ring and ice-core data. The years 1601 and 1605 produced unusually narrow tree-rings in the Sierra, suggesting very cold growing seasons.

The 16-year period from 1597–1613 in the Sacramento River Basin had the maximum reconstructed riverflow for the 420-year (1560–1980) time period. There was a major flood on several Northern California rivers, including the Salmon and Klamath, in about 1600. Those rivers wouldn’t see another flood that big until the December 1964 flood.

The year 1602 was the end of a 37-year drought (1566–1602) drought in the Southern Sierra. Mono Lake rose to elevation 6456, the highest level of the past millennium, around the year 1650.

The authors of one of the Schimmelmann papers speculated that the 1605 flood might have represented a double whammy. The low in the 208-year cycle of solar activity created a period of preexisting cooling conditions across the Northern Hemisphere. By coincidence, a cluster of volcanic events may have then combined with the preexisting cooling to intensify the flooding conditions.

In any case, it is presumed that the cooling in the climate was probably accompanied by a shift of prevailing wind patterns and associated storm tracks toward the Equator during the time period 1600–1610. That appears to be the unifying link in the paleoclimatic evidence listed above and from other sources.

In addition to the above research, USGS has apparently done sediment research in San Francisco Bay that also showed evidence of past megafloods. This may be a reference to the work done by Lynn Ingram at UC Berkeley. The details of that research are unknown.
Norm Miller said that the California megafloods are known to be due to atmospheric rivers.

### 1444–83 Drought

This 40-year megadrought affected at least the Sacramento and San Joaquin River Basins and is known from tree-ring reconstruction. At decadal and longer time scales (e.g., driest 20 years ever), pre-20th century low-flow extremes are dominated by periods in the mid- to late-1100s in the Sacramento River Basin, while the late 1400s were more severe in the San Joaquin River Basin (see Table 21). However, prolonged drought occurred for both of these periods in both basins.

Since the drought of the late 1400s was more severe in the San Joaquin River Basin than in the Sacramento, this suggests that the storm tracks didn’t come as far south during the late 1400s. The years 1451–1500 were the driest 50-years ever on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012. The was also the driest 50 on the entire San Joaquin River Basin. As shown in Table 19, average reconstructed runoff for the 40 years that this drought was active on the upper San Joaquin (1444–83) was 79% of the 1113-year average (900–2012). There were 7 non-drought years during this 40-year period.

### 1527–33 Drought

This seven-year drought affected at least the San Joaquin River Basin and is known from tree-ring reconstruction. The year 1532 was an extreme drought year. As shown in Table 21, it was the fourth driest years on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012.

### 1540–48 Drought

This nine-year drought affected at least the San Joaquin River Basin and is known from tree-ring reconstruction. 1546 was a non-drought year.

This and the next two droughts were associated with a megadrought that extended across much of North America from roughly 1540-1590. That drought has been linked with the disappearance of the English Lost Colony of Roanoke Island (1587), the abandonment of the Spanish colony of Santa Elena at Parris Island, South Carolina (1587), the abandonment of Tewa, Keres, and other Puebloan villages in New Mexico (mid-16th century), and with two of the greatest human mortality events in New World history. These were the cocoliztli epidemics of 1545 and 1576 in Mexico when an estimated 5 to 15 million people died from hemorrhagic fevers, possibly associated with a rodent vector leveraged by extreme drought conditions.

### 1566–1602 Drought

Lisa Graumlich reported a 37-year megadrought from 1566–1602 that affected the Tulare Lake Basin and is known from tree-ring reconstruction. Graumlich’s research showed that this was a continuous 37-year megadrought in Sequoia National Park. We’re not entirely certain how much reliance to put on Graumlich’s reconstruction. It may be correct, but it doesn’t seem to produce results that are very consistent with the tree-ring reconstructions done by Meko, Graybill, or Hughes and Brown.

In any case, this drought was less continuous to the north. Hydroclimatic reconstruction from tree-rings showed that in the Sacramento and San Joaquin River Basins, this period consisted of a number of discontinuous drought years. In the upper San Joaquin at the inflow to Millerton Lake, drought conditions existed for only 20 of these 37 years.

The upper San Joaquin arguably experienced a 26-year megadrought lasting from 1569–1595. As shown in Table 19, average reconstructed runoff for that period was 85% of the 1113-year average (900–2012) on the upper San Joaquin at the inflow to Millerton Lake. That gives an indication of how dry conditions may have been during the 37-year megadrought from 1566–1602 that affected the Tulare Lake Basin. Since the drought was more severe in the Tulare Lake Basin and on the upper San Joaquin than further north, that suggests that the storm tracks didn’t come as far south during these years.

Tree-ring reconstruction shows that 1580 is the drought year of record in the Central Valley and the Southern Sierra. The tree-ring for that year is barely detectable, demonstrating the severity of the drought. As shown in Table 21, 1579 and 1580 were both extreme drought years. 1580 was the driest year ever, and 1579 was the seventh driest year on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012. 1580 is the low flow-of-record on every river in the Southern Sierra for which we have data.
Reconstructions for the Sacramento and San Joaquin River Basins flag 1580 an exceptionally dry single year — far drier than any experienced since record-keeping began.

- On the Sacramento River, the reconstructed flow for 1580 was only 45% of that of the reconstructed flow in 1924, the second-driest year of the reconstruction.
- The relative severity of low flow in 1580 is almost as great on the total San Joaquin River Basin, where flow in 1580 is reconstructed at 54% of the flow in 1924.
- On the upper San Joaquin at the inflow to Millerton Lake, the reconstructed flow for 1580 was only 36% of that of the reconstructed flow in 1924.  

The portion of the drought that included 1580 was of shorter duration in the San Joaquin River Basin than in the Sacramento River Basin.  

1618–19 Drought

This drought is known from tree-ring reconstruction. It was a three-year drought on the Sacramento River Basin, active from 1618–20. But it was somewhat different in the San Joaquin River Basin; there it was a two-year drought lasting from 1618–19. The year 1620 was a non-drought year on the San Joaquin including on the upper San Joaquin at the inflow to Millerton Lake.

1631–32 Drought

This two-year drought affected at least the Sacramento River and San Joaquin River Basins and is known from tree-ring reconstruction. Dave Meko recommended treating this as a four-year drought in the San Joaquin River Basin. The year 1632 was an extreme drought year. As shown in Table 21, it was the eight driest year on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012.

1652–59 Drought

This was an eight-year drought on the San Joaquin River Basin, and is known from tree-ring reconstruction. It contained one non-drought year (1656). In the Sacramento River Basin, this drought began in 1651. The year 1655 was an extreme drought year. As shown in Table 21, it was the fifteenth driest year on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012. The pair of years 1654–55 was the third driest year on the upper San Joaquin at the inflow to Millerton Lake.

1721–22 Drought

This drought is known from tree-ring reconstruction. This was a six-year drought on the Sacramento River Basin, active from 1719–24. But it was different in the San Joaquin River Basin. In that basin, the years 1718–20 had average or a little below-average precipitation. Then the drought hit with a vengeance. Combined precipitation for water years 1721–22 on the upper San Joaquin at the inflow to Millerton Lake was just 55% of the average for the 1113-year period: 900–2012. The year 1723 was a non-drought year.

1728–29 Drought

This two-year drought affected at least the upper San Joaquin River Basin and is known from tree-ring reconstruction. The year 1729 was an extreme drought year. As shown in Table 21, it was the seventh driest year on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012.

1735–37 Drought

This could be viewed as either a three-year drought (1735–37) or a seven-year drought that contained two non-drought years (1934 and 1938). It affected at least the Sacramento River and San Joaquin River Basins and is known from tree-ring reconstruction. Flow reconstruction based on tree-rings show that the drought was primarily in effect on the upper San Joaquin at the inflow to Millerton Lake during the years 1735–37.

1753–57 Drought

This was a five-year drought on the San Joaquin River Basin, and is known from tree-ring reconstruction. On the Sacramento River Basin, this drought lasted seven years, from 1755–61.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

1776–78 Drought
This three-year drought affected the Sacramento River and San Joaquin River Basins and is known from tree-ring reconstruction. The year 1777 was an extreme drought year. As shown in Table 21, it was the sixteenth driest year on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012.

1780–83 Drought
This four-year drought affected the Sacramento River and San Joaquin River Basins and is known from tree-ring reconstruction. Dave Meko recommended treating this as a four-year drought in both basins. The years 1782 and 1783 were both extreme drought years. As shown in Table 21, they were the ninth and eleventh driest years on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012. The pair of years 1782–83 was the second-driest ever on the upper San Joaquin at the inflow to Millerton Lake. Only the pair of years 1579–80 was drier.

1793–96 Drought
This drought is known from tree-ring reconstruction. In the Sacramento River Basin, this drought lasted from 1793–95. Dave Meko recommended treating this as a four-year drought in the San Joaquin River Basin.

Tree-ring reconstruction shows that 1580 is the drought year of record in the Central Valley and the Southern Sierra. Water year 2015 will almost certainly be the second driest year. The year 1795 was an extreme drought year. As explained under the section of this document that describes the 1918–34 drought, there is virtually a three-way tie among 1795, 1924, and 1977 as to which is the third-driest year in the San Joaquin Valley in the 1115-year period 900–2014. Based on stream gage data, we know that 1977 was a slightly drier year than 1924. However, we can’t say with any confidence where 1795 falls in this order, especially in the Tulare Lake Basin.

Based on tree-ring reconstructions, we know that the reconstructed flow on the upper San Joaquin at the inflow to Millerton Lake in 1580 was only 36% of that of the reconstructed flow in 1795 and 1924.

1805 Flood
Based on research done in the Santa Barbara Basin, California experiences a megaflood approximately every 200 years. (See the section of this document that describes the California megafloods.) The last such flood to occur in Southern California was in about 1600. The 1805 flood (a Central Valley flood) may have been the next one in that series. In any case, 1804–05 was an unusually wet winter throughout the Central Valley. There was heavy runoff the following year.

Histories of early settlements state that California Indians spoke of a great flood, which was supposed to have occurred about the beginning of the 19th century and to have drowned thousands of them. This reference may have been to the flood of 1805, which is said to have covered the entire Sacramento River Valley except the Sutter Buttes. The Sutter Buttes is a volcanic plug that rises about 2,000 feet above the valley floor near the center of the Sacramento Valley. High-water marks observed in the San Joaquin Valley and attributed to the 1805 flood were some six feet higher than the huge 1861–62 flood reached.

1826 Flood
1825–26 was reported to have been an unusually wet winter throughout the Central Valley. (One source said that the wet winter was 1824–25, but that was almost certainly an error.) There was heavy runoff and flooding the following year.

According to the Yuba County history, the American Indians said that the Sacramento Valley had a large flood in the winter of 1825–26. One trapping party was compelled to camp in the Marysville Buttes because of high water. Those hills were full of grizzlies, elk, antelope, and smaller game that had taken refuge there. The American Indians recalled the flood of 1826 as a devastating one.

We have no records of what was occurring in the Tulare Lake Basin at the time of this flood. The flood may have extended this far south, but we really don’t know that.

1807–09 Drought
This three-year drought affected at least the San Joaquin River Basin and is known from tree-ring reconstruction.
1822–24 Drought

This three-year drought affected at least the San Joaquin River Basin and is known from tree-ring reconstruction.674

1827–29 Drought

This three-year drought affected at least the San Joaquin River and Tulare Lake Basins. The year 1829 was an extreme drought year. As shown in Table 21, it was the fifth driest year on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012. Annie Mitchell wrote that American Indians said that Tulare Lake went dry about 1825.675 Possibly that was associated with the 1827–29 drought. Or she may have gotten her date wrong.

1839–40 Flood

County histories and journals of pioneers mention floods in the lower Sacramento River Basin during the 1839–40 season.676

1840–46 Drought

This seven-year drought was active in both the Sacramento and San Joaquin Valleys. This drought is not well understood because the state was so lightly populated at the time. What is known comes from scattered pioneer accounts and tree-ring reconstruction:677, 678

- The year 1842 was a non-drought year.
- John Bidwell came to California via the California Trail in 1841. He recalled that 1841 was an extremely dry year in the Sacramento area.
- The years 1841 and 1843 were both extreme drought years. As shown in Table 21, they were the twelfth and eighteenth driest year on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012.
- John Bidwell recalled that 1843 and 1844 were extremely dry years in the Sacramento area.
- Drought ruined the crops at Sutter’s Fort in 1844.
- The dry season of 1844 lasted so long that John C. Fremont was able to take his mounted expedition (complete with cattle) to elevation 11,000 feet in the vicinity of Hell for Sure Pass at the west boundary of present-day Kings Canyon National Park at the end of December 1844. See the section of this document on the Chronology of Tulare Lake for a description of that expedition.
- The years 1841–46 were the second-driest six-year period on the upper San Joaquin River at the input to Millerton Lake during the 1113-year period 900–2012.679 Only 1926–31 was drier.

1847 Flood

1846–47 was reported to have been an unusually wet winter. (One source said that the wet winter was 1845–46, but that was almost certainly an error.) There was heavy runoff the following year.680

The winter of 1846–47 was the one that trapped the Donner Party in the Sierra. A severe blizzard during the last week of October 1846 buried the upper elevations of the Sierra and blocked the trail into Northern California. During the winter of 1846–47, the snowline east of Sutter’s Fort (located in present-day Sacramento) was typically about 3,000 feet, indicating a severe winter.

According to the Yuba County history, the early settlers spoke of floods in the winter of 1846–47, which did but little damage, simply because there was not much to be injured. John Sutter described the area near present-day Sacramento as a vast expanse of water.681

One account reported that the Stanislaus River, at a point about 1½ miles upstream from its mouth, overflowed the country for miles beyond its channel, and that the San Joaquin River was about three miles wide at the crest of the flood. This is the earliest flood mentioned in historical accounts of the settlement of the San Joaquin River Basin.682 As in the Sacramento River Basin, the extent of the overflow in the lower reaches of these rivers in the early days provides little indication of the discharge, as the minor floods would have spread beyond the normal channels in many places almost as far as the major floods.

This flood is known from its effect on rivers farther north in the Central Valley. Rivers in the Tulare Lake Basin may have flooded as well, but no settlers were living here to record the event.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

1849–50 Floods (3)

There were three floods in 1849–50:
1. December 1849 – January 1850
2. May–June, 1850
3. December 1850

The floods of 1849–50 are the first for which there are fairly accurate historic descriptions, and the one of January 1850 undoubtedly was of major proportions. However, it would be exceeded at Sacramento in 1852.683 1849–50 was reported to have been an unusually wet winter. In December 1849, John Benjamin Hockett and a party of immigrants from Arkansas camped on the Tule River when the whole valley appeared a vast sheet of water.584

According to the Yuba County history, the winter of 1849–50 was a wet one, and the streets of Marysville were for a time so muddy that they were almost impassable. The miners along the river were compelled to work in the creeks and ravines in the hills until the water subsided.585 There was heavy runoff throughout the Sierra in 1850, including in the Tulare Lake Basin.

John Sutter had placed his fort on high ground several miles from the river to avoid flooding problems. Sam Brannan decided to locate the new city of Sacramento next to the river for financial reasons; that is where commerce was. Many new residents believed Brannan’s claim that the area was not subject to flooding. They were caught by surprise when the city was hit by a major flood on January 7–8, 1850.686

An excerpt from Exceptional Years: A History of California Floods and Drought by J.M. Guinn, 1890:687

In January 1, 1850, the "Argonauts of '49" had their first experience of a California flood. The valley of the Sacramento was like an inland sea, and the city of Sacramento became a second Venice. But, instead of gondolas, the honest miners navigated the submerged streets in wagon-boxes, bakers’ troughs, crockery crates, and on rafts made of whisky-kegs. Whisky in hogsheads, whisky in barrels and whisky in kegs floated on the angry waters and the gay gondolier as he paddled through the streets drew inspiration for his song from the bung-hole of his gondola.

This was the first major flood to inundate Sacramento’s waterfront since Euro-American settlement. Fundraising began for building levees on the Sacramento and American Rivers. Some very low-standard levees would be built in March 1850.

It seems likely that there was also flooding on the rivers in the Tulare Lake Basin, but we have no records of that. There were no missions in the San Joaquin Valley, nor were there any American settlements of note south of Sacramento. Fort Miller wouldn’t be established on the San Joaquin River until May 1851. Fort Visalia wouldn’t be established on the Kaweah Delta until November 1852.

In May 1850, Lieutenant George H. Derby of the U.S. Army's Topographical Engineers explored the Tulare Lake Basin. His assignment was to find a practical location for a wagon road to the Kings River and to find a location for a military post to control the American Indians.688

On May 7, Derby observed that the Tule River was 100 yards wide, 12–20 feet deep, and very rapid. Two days later, he came to the Kern River, which he described as very broad and deep, and with a 6 mph current. It was running so full that it couldn't be crossed by his mules. It was discharging into Buena Vista Lake which was 10 miles long and 4–6 miles wide.

Returning north, Derby’s party reached the Kaweah Delta on May 14. Including the main river, there were five distinct channels. In his report, Derby referred to the area as "the five creeks of the River Frances".689 Derby described four of those channels as being much wider than the Tule River. All five appeared to be at their height, and all were deep and rapid. Derby would later conclude that this was still several weeks before the peak of the runoff.

(When Lt. R.S. Williamson of the U.S. Army’s Topographical Engineers mapped the Tulare Lake area three years later in 1853, the Kaweah Delta was already known as the Four Creeks Country. That name stayed in popular currency for several decades, although there was never general agreement as to which of the various creeks in the area were the four creeks.)690
Derby crossed the Kings River by boat on May 18. It was about 300 yards wide, rapid, and as cold as ice. While exploring farther downstream, the American Indians told him that the Kings was higher than they had ever seen it. Derby then turned west, cutting across the swampy portion of the San Joaquin Valley. He discovered that all of the water of the Kings was flowing toward Tulare Lake. In addition, a large amount of overflow from the flooding San Joaquin was also flowing toward Tulare Lake with a strong current.

When Derby finally reached the outlet for Tulare Lake (what we now call the Fresno Slough), he discovered that it had only an extremely slow current flowing toward the San Joaquin River. Derby’s party became entrapped in the Fresno Slough area by the rising waters and barely escaped with their lives.

Clearly Derby found extremely high water when he encountered the Kern, Tule, Kaweah, and Kings Rivers during the runoff of May–June, 1850. The American Indians said that the Kings was higher than they had ever seen it. We have encountered no reports of subsequent spring snowmelt floods that resulted in such high water. Spring snowmelt floods take place at the onset of hot weather after a wet winter has built up a larger-than-average snowpack in the mountains. The May–June 1850 flood may have been one of the largest such floods to occur in historic times.

S.T. Harding researched the total runoff for water year 1850. He estimated the total runoff for that year to be 1,420,000 acre-feet. A glance at Table 7 will show that since Derby’s visit, there have been 19 years with more than three times that much total runoff. However, that was still a huge snowpack, and could have generated a very large flood given the right temperatures.

Since the late 1800s, spring runoff waters have largely been diverted onto irrigated lands. For an explanation of how this came to be, see the section of this document: Why is there no lake in the Tulare Lakebed today? The May–June 1850 flood was one of the last great snowmelt floods before these irrigation diversions largely captured the runoff.

Harding found that Tulare Lake’s lowest level in water year 1850 was about elevation 208.0. The runoff that year was sufficient to raise the lake to a maximum elevation of 211.5 feet. Derby’s experience suggests that Tulare Lake was at a relatively low level prior to the flood and was now being refilled when he was crossing the distributaries of the Kings and San Joaquin Rivers. Although the lake was one foot above the delta sill at the beginning of the water year (elevation 208 – 207 feet), C.H. Lee found that significant outflow didn’t really start until the lake reached an elevation of 210 feet. That was because the delta sill was so heavily vegetated with tules. That could help to explain why Derby first observed rivers flowing toward Tulare Lake, then flowing back toward the Sacramento–San Joaquin Delta.

Panoche/Silver Creek west of Mendota flooded at sometime in 1850.

In December 1850, a major flood struck Sacramento; it was bigger than the flood which had struck that city in January. The levees that had been built in March of that year failed. The flood destroyed most of Sacramento. It seems likely that there was also flooding on the rivers in the Tulare Lake Basin, but we have no records of that.

1852–53 Floods (2)
There were at least two floods in the Central Valley in 1852–53:
1. December 1852
2. March–April 1853

Flooding was widespread in the Sacramento and San Joaquin Valleys in the winter of 1852–53. Based on fragmentary accounts, it is possible that the December to April period might best be viewed as a more or less continuous series of flood events rather than as individual floods. The December and March–April floods cited above were just the best documented events.

The December 1852 flood in Sacramento was bigger than the January 1850 flood had been. An article in a Red Bluff newspaper in 1861 described the 1852 floods as being the highest known to the oldest residents prior to December 1861.

That may have been true on the Upper Sacramento. However, on the lower river at Sacramento, the March–April flood 1853 flood was larger than the December 1852 flood. On March 29, Sacramento residents watched the river rise 12 feet in 24 hours. This time the city stayed submerged for six weeks.
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

The floods were closely followed on the East Coast. Newspapers in Oregon, San Francisco, Sacramento, and Nevada would write up the latest news. Those newspapers would be carried by steamer to the Pacific side of the Isthmus of Panama. From there, they would go overland to the Caribbean side, then by steamer to New York. The articles would then be reprinted in the New York Times, about two weeks after the original event had occurred. News from the Sandwich Islands (Hawaii) and Australia was relayed in a similar manner.

The USACE identified the 1852–53 flood as a major flood in both the Sacramento and the San Joaquin River Basins. Heavy snows, flooding, property damage, and hardship were also reported from Oregon and Nevada. The Willamette River had a major flood in January 1853.

1852–53 was an unusually wet winter, the third such winter in eight years (1846–47, 1849–50, and 1852–53). The San Francisco Alta California newspaper said that this was the most severe winter since California had been inhabited by the Americans. The surgeon at the newly established Fort Miller on the San Joaquin River reported that 40 inches of water fell during the months of January and February, 1852.

At Gilroy’s Ranch (site of the present-day town of Gilroy), the waters were higher than they had been known for 25 years. (Trivia: John Gilroy, a Scotsman, was the first non-Spanish settler in Alta California to be legally recognized by the Spanish crown. He arrived in Monterey in 1814, 25 years before John Sutter established his fort in what would become Sacramento.)

There were heavy snows in the higher portions of the mining districts. The New York Times reported that snows were over 10 feet deep by the end of December, and numerous roofs collapsed. Snow was within six miles of Fort Miller (now covered by Millerton Lake).

The mud, floods, and heavy snows made travel very difficult; many roads were virtually impassable even within Sacramento. Provisions were scarce to nonexistent in the mining communities, and prices quickly soared to exorbitant levels. Some miners retreated from the worst-hit areas so that the remaining provisions could be shared by those who couldn’t get out. Immense rains combined with the melting snows to cause widespread flooding, loss of life, and property damage in the mining districts. Many bridges were washed out and ferries swept away.

At Marysville (north of Sacramento at the confluence of the Yuba and Feather Rivers), the waters were reported to be two feet higher in February 1853 than they had been in March 1852. The Feather River was carrying such an immense quantity of logs and driftwood that it looked like it was full of porpoises.

Sacramento experienced severe flooding in both the December 1852 and March–April 1853 floods. This prompted the city to raise the main levee. After four years of regular flooding (1850, 1851, 1852, and 1853), property owners agreed to pay to raise the streets in the business district (now partly Old Town Sacramento) by up to 5 feet. Buildings in the raised district had to also raise their street-side entrances to meet the level of the new sidewalks. The streets would be raised again by a much larger amount between 1862–1869.

Four floods hit Marysville during the winter of 1852–53, and the surrounding country was more or less under water the whole season. The rains commenced early in November 1852, and toward the latter part of the month the water was as high as it had reached the season before. Again, a week or two later, the water rose 6½ inches higher than at first. The waters then subsided, but the last week in December was one of continued rain, and on December 31 the water from the Feather and Yuba Rivers began to come into the city. The next day the water was 20½ inches higher than at the last flood, and was from 6–10 inches deep on the floors of the buildings about the plaza. There had been a grand ball planned at the Merchants’ Hotel on New Year’s Eve, 1853; but when the hour arrived, the hotel was surrounded by water. Several young gentlemen, loath to lose the anticipated pleasure, proceeded to the hotel in boats, and with a number of ladies residing there, danced merrily until morning.

All the low and bottom lands were completely submerged by this flood. As it was the first experience that the new ranchers had of this kind, they lost very heavily in stock, crops, etc. Communication from the city with the outside world, and among the farmers, had to be maintained by boats. People were compelled to come to the city in boats in order to obtain supplies, and trading with the mines was effectively blockaded for some time. The continuous rains and almost impassible muddy roads were such an impediment to freighting that a great shortage of supplies was caused in the mines.
At the earliest possible moment, a number of energetic and enterprising men started out with trains of supplies, hoping to reach the destitute regions before the markets were supplied, and thus reap a bountiful harvest of gold to reward them for their labor. Those who reached the mines first were amply rewarded for their exertions, and were able to secure any price that their conscience would permit them to ask, such as one dollar per pound for flour, and twenty cents per pound for hay.

The fourth and last flood of the season in Marysville commenced on Saturday, March 25, 1853; and on Tuesday the water reached a point eight inches higher than in January. The country on all sides of Marysville and Yuba City was under water. By Saturday, the waters had subsided sufficiently to permit the pack trains to leave the city.

One source said that Sacramento was virtually wiped out by the 1852–53 flood, just two years after the devastating flood of December 1850. In any case, it was heavily damaged. The warm rain that struck the city on December 29–30 was described as being of unprecedented severity. Not even the new brick houses provided a place of refuge. Most of the foundations settled due to the saturated ground. Those buildings were roofed with tin, and the storm rolled the tin up like parchment. In many houses, the occupants were obliged to go out into the streets to seek shelter from the rain. There were only two brick buildings in the entire town that didn’t leak during the storm.705, 706

(Sacramento would rebuild after the flood, and in 1854 it would become the fourth and final capital of the state. In 1852–53, the twin towns of Vallejo and Benicia were serving as the capital of the newly formed state. The initial capital had been San Jose.)

There was extensive flooding and property damage in the region of Colusa in the lower Sacramento Valley.

The Alta California newspaper said that floods were widespread in both the northern and southern mining districts. Ferries were destroyed on a number of rivers. Bridges were washed away on the Stanislaus, Calaveras, and other rivers. Flood levels were higher than in the memorable winter of 1849–50.707

The Sacramento and San Joaquin Valleys formed a “world of water.” The bottoms on the San Joaquin River were under one vast sheet of water, estimated to be some 20 miles wide. At the mouth of the Merced, the owners of the ferry took up residence on their boats. Food supplies were low for the 200 settlers on the San Joaquin.

A Belgian gold miner, Jean-Nicolaus Perlot kept a diary of events in the Mariposa area:708

Never in my life have I seen it rain more heavily or for a longer time. From the sixth of December (1852) to the first of March (1853), the rain didn’t stop for as much as three hours, unless it was during my sleep, which is hardly probable; how many times, during those three mortal months, how many times I awakened at some hour of the night! And always I heard the monotonous sound of the rain falling on the roof of our house.

The flood of 1852–53 raised Tulare Lake by 11.5 feet. At that point, the lake had an elevation of almost 216 feet and a depth of about 37 feet at its deepest point (216–179 feet). There was 9 feet of water in the outlet channel flowing over the delta sill (216–207 feet). From there, the water connected through the Fresno Slough to the San Joaquin River and flowed on to San Francisco Bay.

The White River is the next river south of the Tule. 1852 was also a major flood in the mining district on the White.711 (The village that supported the mines is located 10 miles east of Delano. It was then known as Tailholt but was later renamed White River.)
Gordon’s Ferry was established on the Kern River just north of present-day Bakersfield College in the spring of 1852. Eight months later, rain fell for three weeks across California. An observer wrote: “The rivers have been swelled to such an extent as to inundate all the low lands, causing immense damage, destroying stock and agricultural products.” According to José Jesús López, early pioneers said that the Kern River swept Gordon’s “perfectly bare of all signs of improvements.”

Below Gordon’s Ferry, the Kern River flowed through Kern and Buena Vista Lakes on its way to Tulare Lake. Tejón Creek flowed into the southeast end of Kern Lake in a channel two feet deep and ten feet wide. In the 1852–53 flood, Tejón Creek overflowed its channel for more than two months.

During the height of the 1852–53 flood, some sailors jumped ship in San Francisco. They stole a whaleboat, hoisted the sail, and headed inland. Taking advantage of the prevailing winds, they sailed south up the San Joaquin River, through the Fresno Slough, and entered Tulare Lake. This was the first of six documented trips between that lake and San Francisco Bay to occur in historic times. (The other five trips were in 1868, 1938, 1966, 1969, and 1983.)

But the sailors didn’t stop in Tulare Lake: they continued south up the Kern River to Buena Vista and Kern Lakes. And since Tejón Creek was in flood, they kept going up that creek (east) another 15 miles or so until they were about two miles north of an American Indian village. That village was located where the old Tejon Ranch headquarters would later be built (at the end of present-day Sebastian Road, due east of where Interstate 5 and Highway 99 diverge). There the sailors beached their boat and walked over to the village.

After staying with the villagers for about two weeks, the sailors returned to their boat to go back downstream. However, by then, Tejón Creek had gone down so much that the water was back in the channel, and they almost didn’t make it. About 10 Tejón Indians helped them. Santiago Montez was one of them, and he later recalled the event:

> Them sailor pretty smart. When water not deep, they put boat across creek and sit on boat. That make dam. That back water up. They all jump out of boat, grab boat by sides and run ahead of water fast they can. Then boat go maybe hundred yards and stick again. They do that lots time before they get back in Kern Lake.

The sailors stayed on in Kern and Buena Vista Lakes where they trapped beavers and otters. At least some of them married American Indians and raised their children in the local schools. The whaleboat remained in use until it was apparently swept away in the 1867–68 flood. Perhaps it washed down to Tulare Lake or even back to San Francisco Bay. If that boat could talk, what a tale it could tell.

### 1855–61 Drought

This seven-year drought affected the Central Valley and maybe other areas of the state. It is sometimes described as having multiple components, 1855–57, 1860–61, etc. depending on area. Tulare Lake continued its steady decline during The 1855–61 drought.

Based on tree-ring reconstructions, we know that the reconstructed flow on the upper San Joaquin at the inflow to Millerton Lake in water years 1856 and 1858 was less than half of the 1113-year average (900–2012).

Floyd Otter said that Robert Glass Cleland documented the effects of the 1855–59 drought. During the drought, cattle grazed everything that they could reach and then died by the tens of thousands.

Hale Tharp was a cattleman. Tharp said that 1858 was a dry year. Because of that (and because he knew that other herdsmen would be coming into the Kaweah canyons), Tharp decided to investigate the stories the American Indians told him about the perennial meadows and big trees in the high mountains. Guided by American Indians, Tharp explored the Crescent Meadow portion of Giant Forest in 1858, becoming the first white man to see giant sequoias.

Scott Mensing studied the relationship between blue oak regeneration and fire pulse on Tejon Ranch at the time of settlement. He reported that the droughts of the 1850s were so severe as to suppress regeneration from acorns.

Floyd Otter said that Robert Glass Cleland documented the effects of the 1860–61 component of the drought. This was to be a virtual replay of the 1855–59 component of the drought which had resulted in the deaths of so many cattle. During the 1860–61 drought, cattle once again grazed everything that they could reach and then...
died by the tens of thousands. Large numbers of cattle starved to death during the droughts of 1855–61 and 1863–65. A Mussel Slough woman vividly recalled life on the shore of Tulare Lake during this period.  

The country was nothing but a dry, barren desert with bands of wild roving cattle that would come out of the timber along the river in the morning and go out to the lake to feed. Where the water of the lake had receded a little grass would spring up and they would get a little feed. The poor things were almost starved…so we could not blame them for eating the hay we had stacked for our horses. The settlers all dug ditches for fences to keep them out but without much effect.

This drought affected much of the state. J.M. Guinn wrote that the year 1856 was a particularly exceptional year, including a severed drought, intense heat, multiple intense earthquake shocks, and severe sand storms. It was considered the driest season that the country had known for 20 years. During the summer of 1856 and the ensuing winter, the loss of cattle by starvation in Los Angeles County alone was estimated at 100,000,000.

Guinn reported that Los Angeles County experienced extreme weather in late 1859. The temperature reached 110 degrees in October of that year. The most remarkable rainstorm ever known in the county occurred in December. An estimated one foot of water fell within 24 hours. The rivers overflowed the lowlands, doing considerable damage. The starving cattle and sheep, unsheltered from the pitiless rain, chille d through and died by thousands during the storm. Large tracts of the bottom lands were covered with sand and sediment.

An article in the Red Bluff Independent reported that summer of 1861 was the hottest, driest season since California became a state in 1850. The fall rains were late in coming, and cattle starved to death in large numbers in early November.

The drought would end decisively when a series of epic storms moved into the state at the beginning of December, unleashing the great 1861–62 flood.

1861 Flood
This flood occurred during the 1860–61 drought. Relatively little is known about this flood. It is different from the much more famous 1861–62 flood. The USACE identified it as a major flood in both the Sacramento River and the San Joaquin River Basins. How it affected the rivers within the Tulare Lake Basin is unclear. Runoff during water year 1861 caused Tulare Lake to rise 2.3 feet in elevation.

1861–62 Flood
Flooding occurred from December 1861 through January 1862. For reference, 1861 was the first year of the Civil War. General Beauregard fired on Fort Sumter on April 12, 1861, igniting that war. California managed to participate in the Civil War in various ways despite conflicted loyalties and the effects of the disastrous flood of 1861–62.

1861–62 was an incredibly wet winter. Extremely wet winters in California are often associated with an El Niño weather pattern. However, research done at Oregon State University indicates that was not the case with the 1861–62 storms. The polar jet stream apparently swept up and down the West Coast during that winter, causing the temperatures to vary wildly, from very warm to very cold. The warm storms brought the typical rain-on-snow events. However, the cold storms brought snow down nearly to sea level in the Sacramento Valley. San Francisco recorded nine days with temperatures below freezing in January alone. On January 28, San Francisco registered 22 degrees, a full 5 degrees colder than any temperature ever measured in the modern era in that city.

The atmospheric mechanisms behind the storms of 1861–62 are unknown; however, the storms were likely the result of an intense atmospheric river, or a series of atmospheric rivers. Immense quantities of water were delivered during this storm event.

The mining community of Sonora received 102 inches of rain (8½ feet) in a 74–day period (November 10, 1861 – January 23, 1862). San Francisco recorded 28.25 inches in 30 days. This was 6.48 standard deviations above the average rainfall for 30 consecutive days. The associated recurrence interval is 37,000 years.

Most of the states of Oregon and California were affected by the flooding. Record stages resulted on the major rivers throughout those two states. It almost certainly had a recurrence interval greater than 100 years.
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The 1861–62 flood was a huge event. One source said that it stretched from Canada to Mexico, but that overstates the case. The storm tracks responsible for the flood were generally aimed at the southern part of Oregon. As a result, the northern part of Washington received less than average precipitation during the period of this flood. Overall, the 1861–62 flood had minimal impact on Washington.

So it is more accurate to think of this as a flood — or a series of floods — that stretched from the Columbia River to Mexico; that is a region over 1,000 miles wide.

November brought one storm after the other to Oregon, resulting in a marked excess in precipitation over most of the state. In the Cascades, temperatures were cold enough to result in well above-average accumulations of snow. December turned warm, and the rains melted much of that snow. There were a series of major floods throughout the state. The meteorological conditions of the Pacific Northwest during the winter of 1861–62 were summarized by Edward Lansing Wells, the chief meteorologist for many years at the Portland office of the U.S. Weather Bureau.728

The storms that struck Oregon in November moved in rather far south. However, the one that hit at about the beginning of December seems to have passed just far enough north to produce strong, warm southerly winds with extremely heavy rains that reached into eastern Oregon.

Of all the floods to hit Oregon that winter, the most impressive occurred on the Willamette River. (Based on fragmentary evidence, that was also the one that was most closely followed in the new town of Visalia, far to the south.) A Belgian gold miner, Jean-Nicolaus Perlot, left the California gold fields to settle in Portland, Oregon in time to witness the flood on the Willamette:729

The peaceful Willamette became, by the fifth of December, an impetuous torrent; leaving its bed, it upset and carried away the establishments which bordered its banks. It was, for two days, a curious and heart-rending spectacle: the river was covered with strays of all kinds, trees, animals, fences, provisions, houses, sawmills, flour mills all that was floating pell-mell, and passed before Portland with a speed of three leagues an hour.

Some 353,000 acres were inundated; “the whole Willamette valley was a sheet of water.” It was the largest flood on that river in recorded history. Many towns were damaged or destroyed.

The Willamette peaked at Oregon City on December 4, 1861: 635,000 cfs, 35% greater than the average flow of the Mississippi River. Oregon City sits at the base of Willamette Falls, the largest waterfall in the Pacific Northwest (based on volume). One night during the flood, the residents of that town watched a number of houses come over the falls, with lights still burning inside. Then, on December 5, they watched as the side-wheel steamer St. Clair ran the 40-foot-high falls "with great rapidity."730

December 1861 to January 1862 constitutes one of the greatest flood periods in the history of California. The 1861–62 flood period was remarkable for the exceptionally high stages reached on nearly every stream, for repeated large floods, and for the prolonged and widespread inundation in the Sacramento River and San Joaquin River Basins. Rainstorms were heavy in the lower elevations and snowfall continuous in the upper elevations throughout the two basins.731

The summer of 1861 was the hottest, driest season that the northern Central Valley had experienced in over a decade. The fall rains were late in coming and cattle deaths were high in November. However, by December 10, the drought was over in the Red Bluff area and flood damage from the Sacramento River was extensive.732, 733 The settlers wanted an end to the drought, and they got it.

Northern California experienced record-setting precipitation and flooding. The initial flood began late in November 1861 when storms brought rains to the lowlands and covered the mountains with up to 20 feet of snow. This was followed by warm rains in the mountains which melted the accumulated snow. This pattern was repeated several times through the winter. The storms kept coming right through January.

Flooding was severe in the North Coast of California. Flooding on the Klamath River was particularly impressive. A 500-foot-long wire suspension bridge spanned that river in a canyon below Weitchpec (east of Trinidad). That bridge was 99 feet above low-water and thought to be safe from any possible flood. However, the Klamath overtopped that bridge and swept it away. Most of the American Indian ranches on the Klamath Indian
Reservation were located along the Klamath River. The Klamath destroyed all of those ranches that were within 25 miles of the river mouth.

Sacramento was built where the American and Sacramento Rivers meet. It had experienced severe flooding in the 1852–53 flood, causing the city to raise its streets and strengthen the main levee. Despite these precautions, Sacramento was one of the hardest hit cities in the 1861–62 flood. It was flooded about five times during that winter. The first inundation occurred at 6 a.m. on the morning of December 9, 1861, when the American River breached the east levee. By 10 a.m., many houses were floating or overturned.

The flooding was made worse because the railroad’s R Street levee prevented the floodwaters from draining. To drain the city, engineers directed a prison chain gang to cut through that levee between 5th and 6th Streets. When they did, the water rushed through the opening to the Sacramento River, sucking about 25 floating houses through the gap.

Leland Stanford was elected governor in November 1861 and was just taking office when the flood hit. The Sacramento River was flooding so badly that Stanford had to crawl out the window of his home and row himself to his inauguration.

But this was just the beginning of the flooding. The Sacramento area got another 25 inches of rain during the following two months, almost four times the average rainfall. The Sacramento River surged at three times its average seasonal flow of 285,000 cfs, inundating the Sacramento–San Joaquin Delta region. Sacramento was under water for three months. After much debate about appearances and propriety, the State Legislature voted on January 23 to abandon the state capital and move to San Francisco. The California Supreme Court also moved its operations to San Francisco, but it never moved back.

The flooding prompted Sacramento to raise the streets of its business district (now partly Old Town Sacramento) by up to 15 feet between 1862–1869. The tunnels under present-day Old Town Sacramento are reminders of the original downtown buildings and streets.

In the Sacramento River Basin there was a succession of floods starting on December 8, 1861 and continuing into March 1862. Many reports published during the period described the lower Sacramento and San Joaquin valleys as one vast sea of water. Probably as much as 5,000 to 6,000 square miles of the valley floor were submerged.

In 1860, the State of California had hired Josiah Whitney and William H. Brewer, both Yale graduates, to conduct a long-term, in-depth investigation of the state’s resources. They were just two years into their studies when the great flood of 1861–62 bankrupted the state and soon thereafter terminated their project. Brewer was a botanist and an agriculturist. He was also a compulsive diarist — keeping detailed notes of his experiences from 1860–64. They were mostly letters to his brother, which were assembled into a book: Up and Down California. This book was printed by the Yale University Press in 1930. It provides detailed accounts of both the 1861–62 flood and the ensuing drought of 1863–65.

A defining feature of the flood in the Central Valley was that the rains came down far faster than the Sacramento and San Joaquin Rivers could drain the floodwaters to San Francisco Bay. Brewer was there to describe the resulting enormous lake that swelled up in the Central Valley. Nothing remotely like it has ever been seen since. The prolonged period of flooding in the lower Sacramento Valley lasted from December 13, 1861 to about February 1, 1862.

Brewer wrote from San Francisco on January 19, 1862:

The amount of rain that has fallen is unprecedented in the history of the state... The great central valley of the state is under water — the Sacramento and San Joaquin valleys — a region 250 to 300 miles long and an average of at least twenty miles wide, a district of five thousand or six thousand square miles, or probably an area of three to three and a half millions of acres!

Brewer wrote of the Central Valley on February 9, 1862:

Nearly every house and farm over this immense region is gone. There was such a body of water — 250 to 300 miles long and 20 to 60 miles wide, the water ice cold and muddy — that the winds made high waves which beat the farm homes in pieces. America has never before seen such desolation by flood as this has been, and seldom has the Old World seen the like.
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That lake that formed in the Central Valley in 1862 was roughly three times larger than the Great Salt Lake (5,500 square miles versus 1,700). The low-elevation lakes and wetlands in the Tulare Lake Basin were part of that big lake. The Sacramento Valley was so inundated that steamers ran back over the ranches fourteen miles from the river, carrying stock, etc., to the hills. Approximately 100,000 cattle died in the valley.

Brewer reported that "It is supposed that over one-fourth of all taxable property of the state has been destroyed." Brewer kept in touch with the State Treasurer and news of the dwindling state government income because he was having long delays in being paid for his work.

The magnitude of flooding in Sacramento was due in part to the sediment coming from the hydraulic mining. Brewer reported that the river was choked with sediments before the flood and that the riverbed was raised by at least six feet during the flood. Sacramento responded by once again strengthening the levees and raising the elevation of the streets by as much as 10 feet.

According to the Yuba County history, long and incessant rains ushered in the rainy season, and the Feather River started to rise rapidly on Saturday, December 7, 1861. All day Sunday the rain poured down, and that night the city was nearly under water. Early Monday morning several buildings, undermined by the water, fell crumbling to the ground, creating great consternation. The floors of the Merchants' Hotel fell through to the basement, carrying with them the sleeping occupants, several of whom were severely injured by the fall, although no one was killed. A great many frame houses were floated from their positions, and some of them were carried down the stream. The steamer Defiance, playing tunes on her calliope, made its way through the streets giving assistance to those who were rescuing the unfortunate.

The condition of the country was described in the Marysville Appeal:

Westward one vast water level stretched to Yuba City, where a kindred inundation was raging; the entire town site being under water. Beyond this to the foothills of the Coast Range there appeared to be no dry land. Northward the plains were cut up into broad streams of running water, which were swiftly coursing toward the great sheet of water stretching between the Yuba and Feather rivers, up as far as the residence of Judge Bliss, unbroken except by the upper stories of houses, trees and floating debris. Southward the whole plain toward Eliza was one sheet of water, dotted with trees, roofs of houses, floating animals and wrecks of property of every description. Where Feather River sweeps past Eliza, stock of every kind could be seen constantly passing downstream, some alive and struggling and bellowing or squealing for life. Hare and rabbits were destroyed by thousands.

The people in the country had to leave everything and flee to high ground for safety; many who were too late for this, climbed trees and remained perched among their branches until rescued by friends. Nearly all the bridges on the Yuba and Bear Rivers were carried away, and drift timber and saw-logs came down the streams in great quantities, some of which were left in gorges thirty feet high when the water fell. A deposit of sand up to six feet thick was left on the bottomlands when the waters retreated.

On January 11, 1862, the water raised six inches higher than before, but the warning of the previous flood had caused the merchants and farmers to move everything perishable beyond the reach of danger. The loss of stock that winter and the next summer was very great, and in Sutter County was estimated to be three-fourths the entire number. Few animals escaped except those able to reach the Marysville Buttes, and the cold weather nipped the grass, causing large numbers of the cattle to die from starvation.

For a week, the tides at the Golden Gate did not flood; rather, there was continuous and forceful ebb of brown, fresh water 18–20 feet deep pouring out above the salt water. A sea captain reported that his heavily laden ship foundered in the Gulf of the Farallones off of San Francisco due to the layer of fresh water. Tule islands floating across the bay and out to sea were crawling with rattlesnakes. Some of these islands came to rest under the San Francisco wharves where the snakes were a menace for months. Freshwater fish were caught in San Francisco Bay for several months after the peak of the flood. Such events have not happened since.

The 1861–62 flood is known as a flood of Northern California because that is where the population of the state largely lived at the time. However, the flooding from that event was also quite severe in Central and Southern California.
In the San Joaquin River Basin, there were extreme, successive floods during December 1861 and January 1862. By early January, snow had accumulated to unusual depths in the Sierra. Much of this snow deposit was melted by warm rains and helped to swell the flood volume.

Out of 100,000 head of cattle in San Joaquin County, only 10,000 survived. William Knight, a fur-trader, came to California in 1841. Caught up in the gold fever of 1849, he was heading south when he was stopped by the wide Stanislaus River. Seeing a business opportunity, he began a ferry operation using an old whaling vessel. That ferry was later replaced with an open-truss bridge. Legend has it that the plans for that bridge were drawn up by Ulysses S. Grant. In any case, the 1861–62 flood washed away that first bridge and all but one of the other bridges on the Stanislaus. (A new covered bridge was constructed at the same site in 1863. That bridge has withstood many subsequent floods and is in remarkably good condition. However, it was closed to vehicular traffic in June 1981.)

At daylight on Friday, January 10, the crest of the Stanislaus hit the town of Knights Ferry like an avalanche, rising rapidly until it covered the business section, which was built on a flat above the river. At dusk, the river fell about four feet, and the residents thought that the worst was over. Then at 2 a.m. Saturday, the Stanislaus rose again and carried off all the remaining buildings on the flat. Only the buildings on the hill remained. Every bridge on the Tuolumne River was washed away except the one at Steven's Bar.

The Mariposa area was hit by a heavy storm in late December 1861, resulting in dramatic flooding on the Merced River and at the mining community of Coulterville. The tale was told in the January 7, 1862 issue of the Mariposa Gazette as reprinted in the February 6, 1862 issue of the Visalia Delta:

It was the hardest storm, particularly that part of it occurring Thursday night, Dec. 26th, that has ever swept these mountains within the recollection of that very respectable individual "the oldest inhabitant." The Merced River rose fearfully high, sweeping off every bridge upon it, tearing out dams, etc. In Coulterville, the gale of Thursday night was terrible, accompanied by a heavy rain. That night was the most hideous we have ever known. It was worse there than further south. The town, it might be said, was afloat. The rear portion of all the establishments bordering on (Maxwell) Creek, went along with its turbulent waters. Bell's Saloon was flooded, and Cashman & Co.'s barn, a large building, raised anchor and took a notion to sail. R. McKee, Esq., however, with characteristic intrepidity, hitched it to the liberty pole.

The Merced River, downstream from the mouth of its canyon, flooded the town of Snelling. The flood widened and changed the course of the Merced River channel. Reports stated that the whole country surrounding lower Mariposa Creek and the Fresno and Chowchilla rivers, as seen from the foothills, was one vast sheet of water.

During December 1861 and January 1862, the San Joaquin River rose from 24,000 cfs to approximately 133,000 cfs, a fivefold increase. The city of Stockton and the surrounding country were inundated for many miles. Floating farmhouses broke the telegraph wires on the outskirts of Stockton. A steamboat ran through the back wall of the Russ Hotel in the town of Hill’s Ferry (northwest of Los Banos). The flood destroyed nearly all the bridges, mills, and other structures along the channels of the San Joaquin River and its major tributaries.

Panoche/Silver Creek west of Mendota flooded in December 1861 and January 1862.

In the Tulare Lake Basin, there was an exceptionally great flood on January 11, 1862. The Kings, Kaweah, and Tule Rivers brought down tremendous quantities of timber from the Sierra and deposited them on the plains.

The 1861–62 flood on the Kings River began the formation of Cole Slough, cutting the head of that slough. (See the section of this document on Pine Flat Dam for a more detailed description of the formation of Cole Slough and associated waterworks.) The entire town of Scottsburg was washed away by the Kings during this flood and was subsequently rebuilt at a safer location.

In the winter of 1861–62, Joseph Hardin Thomas had just completed construction of a sawmill on Mill Flat Creek, downstream from present-day Sequoia Lake. It was a double-circular sawmill with a 40 horsepower steam engine, one of the two primary sources of lumber for Visalia. Thomas’s mill was destroyed by 30-foot-deep floodwaters resulting from a debris slide in January 1862. The flood also destroyed Feggan’s Mill, an older sawmill located six miles farther downstream nearer the Kings River.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

From the number of large trees washed down from the mountains by the floods on the Kings, Kaweah, Tule, and White Rivers, the settlers inferred that this was the greatest flood for many years. From the number of large trees washed down from the mountains by the floods on the Kings, Kaweah, Tule, and White Rivers, the settlers inferred that this was the greatest flood for many years.  The 1861–62 flood brought plenty of destruction, but it also brought opportunity. Thomas Flaxman, the owner of the 80-foot-long sternwheeler Alta, decided to use the flood to bring the Alta to Tulare Lake. The crew was composed of men familiar with freighting from Firebaugh to Stockton on the San Joaquin River. The route that Captain Giddings chose was through the San Jose Slough to Summit Lake and from there to Tulare Lake. Upon entering the San Jose Slough, Giddings took aboard several vaqueros to act as pilots. The first day went fine, but that evening the Alta got into the wrong channel.

The captain resorted to kedging: carrying the anchor ahead, dropping it, and then pulling the larger vessel ahead by reeling in the anchor cable. It was arduous work. At midnight, the weary crew lay down to sleep until morning, when they were to resume their task. However, when they woke, the floodwaters had gone down, leaving the vessel high and dry. The boat was stranded about four miles southeast of present-day Burrel. The following years were ones of drought. There the Alta sat, just two miles from Elkhorn Station. A strange sight indeed to stage passengers who passed on the dusty road nearby. The Alta was gradually picked apart for lumber and firewood. In 1875, her engine, boiler, and pilot house were salvaged for a side-wheel steamboat — the Mose Andross — that A.J. Atwell was having built at Buzzards Roost on the northeastern shore of Tulare Lake.

In the Visalia area, rain started early in October 1861. By the end of that month, the ground was wet down to a depth of eight inches. By the end of November, sufficient rain had fallen to wet the ground down to a depth of 2½ feet. The rain started again in mid-December and continued to fall for several weeks. January brought a week of warm gentle rain which filled the creeks to the banks. Heavy rains continued until March 1862. The wind remained southerly from mid-December throughout most of January.

Byron Allen’s mother, Marjorie Houston Allen Oakes, described “The Big Flood of 1862” at a 1929 meeting of Three Rivers residents who were worried about lack of rain as of mid-December, 1929. Several Three Rivers pioneers tried to encourage hope by telling about other years when rain did not come until late, but then very wet seasons followed. Marjorie said she was a young girl in 1862. She said there was no rain until January 2, but then it rained continuously for 35 days and nights. The water was the highest the white man had ever seen in this region. She said looking west from the foothills the whole valley seemed to be one sheet of water. This account of the dry spell prior to January 2 doesn’t seem to agree with other accounts of precipitation in the area. Anecdotal accounts such as this should generally be taken with a grain of salt. Memories can change with the passage of time.

The first heavy storm of the season in the Tulare Lake Basin occurred on December 23-25, 1861. It was reported that there was a flood on the White River and a damaging flood on the Tule River which overflowed farms to a depth of several feet. The Kings and Kaweah Rivers apparently did not reach exceptionally high stages during this event. Then on January 11, 1862, there was an exceptionally great flood, which probably has since been equaled or exceeded only in December 1867. The flood followed a general storm, which resulted in record-breaking stages on tributaries of the lower San Joaquin and Sacramento Rivers.

Prior to the 1861–62 flood, the Kaweah River waters spread at high stages over what was known as the Kaweah River Swamp (aka Visalia Swamp). The swamp commenced near present-day Terminus Dam and extended southwesterly about 9 miles with a width of 1–3 miles. The spreading waters were reunited in various channels in and below the swamp. Channel capacity was inadequate to pass floodwaters, not only within the swamp, but also below it.

The Shipp Cut had been made in 1854; it was a small drain ditch from the Kaweah River Swamp near Rocky Ford (north of present-day Kaweah Oaks Preserve) west to Canoe Creek. The 1861–62 flood cut a new channel along the northern border of the swamp. Shipp Cut and a section of Canoe Creek were enlarged by the floodwaters and became a part of this new channel, and finally a connection was established with the Cross Creek channel downstream of Visalia, creating what is now known as the St. Johns River.

This rerouting of the floodwaters to the north of the Kaweah Delta may have reduced the flooding in Visalia. In any case, surprisingly little water came down Mill Creek.
Mill Creek flooded downtown Visalia three times during the 1861–62 flood:
- evening of January 11–12 (22 inches deep on Main Street)
- January 17–18 (20 inches deep on Main Street)
- night of January 20 – January 23+ (24 inches deep on Main Street)

This information serves chiefly to establish the dates of the three larger floods. However, because of the probable variable influence of extensive overflow below the foothills, it indicates only very roughly their magnitude.

The lowlands along the tributaries of Tulare Lake were probably flooded continuously from January 11 until about the end of the month. It is not clear from the contemporary accounts when the maximum stages occurred in the foothills. It is probable, however, that the Kings River reached its greatest stage on January 11, the day when the highest stage was observed on the adjacent San Joaquin River. The White River and Poso Creek, in the southeastern part of Tulare Lake Basin, were reported to have been at their highest on January 18.

Most of the wells in Visalia were contaminated by the floodwaters, so drinking water was hard to obtain. Since rain continued to pour down during the flood, some people caught rainwater for their drinking water. There was one pump in town (at the corner of Encina Avenue and Oak Street), and it was the duty of the young boys in many families to fetch the unpolluted water from that pump, carrying their loads in boats.

The floodwaters caused significant property damage in the Visalia area as well. The flood destroyed many irrigation ditches, a lot of fencing, and four bridges in and around town. The flood in the Visalia area was described in some detail in the January 23, 1862 issue of the Visalia Delta.

Some 42–46 homes as well as some businesses were destroyed in Visalia during the flood. Many homes of this period were made of adobe. As the water came up about these, they began melting and sinking, necessitating immediate departure of their inhabitants. At the time of the flood, there were only a few brick structures in Visalia; most mercantile buildings were made of wood or adobe. The floodwaters melted away the foundations of the adobe buildings and toppled them over, so to speak, on their heads. However, not a single wood or brick building came down in Visalia during the flood. It was a hard-learned lesson for the town, rather like the moral of the Three Little Pigs and the Big Bad Wolf. After the 1861–62 flood, most of the homes and businesses in Visalia were constructed of either brick or wood. Adobe was a building material reserved for high ground.

This was the first major flood to come into Visalia since settlement was begun, and there was significant property damage. However, the floodwaters were shallow and the flood damage was trivial compared to what was happening farther north and south. The residents of the town realized their good fortune. The Visalia Delta reflected on this in late January 1862:

> The more we learn of the late terrific storm, both at home and abroad, the more do we find matter for congratulation that Tulare (County) has escaped so cheaply. The loss to the mass of citizens is absolutely nothing, as compared with less fortunate localities.

People who lived around Tulare Lake were also keenly aware of the benefits that the floodwaters brought. The Visalia Delta highlighted one example:

> The recent high water has had the good effect of stocking all the creeks and small streams with an abundance of lake trout. In such vast numbers did they ascend that we are informed that in Antelope Valley (vicinity of present-day Elderwood) and at the Cottonwoods, the receding flood has left wagon loads of them standing high and dry. It has been so long since the creeks have been full enough to allow them to ascend, that they had nearly disappeared from the smaller streams in this vicinity.

By February 1862, the flood was over. In an effort to allay the fears of new immigrants to Visalia, the Visalia Delta wrote:

> Many boats are still observable lying high and dry at the gates and steps in front of the residences of some of our citizens. Put them away, gentlemen, there is no chance of their being wanted very soon again and their presence might give strangers the impression that floods were a regular institution in Tulare (County).
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Visalia was built on the alluvial fan of the Kaweah River: the Kaweah Delta. An alluvial fan is a distributary system made of multiple channels that allow for large areas of shallow inundation. Because the flooding in the 1861–62 flood was shallow, no horse, cow, or even full grown hog was known to have drowned in the country around Visalia. In fact, the Tulare area had thousands of fat hogs ready to be driven north to where prices for all manner of foodstuffs had just skyrocketed. The cry went out: Drive up the hogs.\textsuperscript{772}

The 1861–62 flood deposited considerable quantities of drift, silt, and sand on the Kaweah Delta. A portion of the channel of the Lower Kaweah River was obstructed by these deposits, significantly reducing its ability to carry flows. During low-flow periods, that section of the Kaweah would now be dry; it would only serve to carry water during flood periods. The St. Johns would become the principal channel of the Kaweah River across the delta to Tulare Lake for the next 15 years.

Prior to the 1861–62 flood, the Kaweah divided into its distributaries (the Four Creeks, if you will) at a point more or less in the middle of the delta. However, after the flood, the primary distributary point moved up near the top of the delta, well up into the swamp, about a mile below present-day McKay's Point. The four channels continued to come back together to reform into a single river channel on the lower side of the delta.

There was a fifth channel that came directly out of the foothills south of the Kaweah. Presumably that followed the course of Yokohl Creek at least as far as present-day Highway 198. In any case, that fifth channel flowed along the south side of the Kaweah Delta (apparently following a course similar to present-day Outside Creek) and then joined the Kaweah in the marshy ground near where that river flowed into Tulare Lake.

W.B. Cartmill recalled what the 1861–62 flood was like. The Cartmill Ranch was located north of Tulare on present-day Cartmill Road (Ave 248), about four miles west of Highway 99. It was located on Packwood Creek, well back from Tulare Lake. The Cartmill family arrived at the ranch on October 26, 1861. It started raining shortly after their arrival and continued raining almost every day until about Christmas. The flood came on Christmas Eve.

The Cartmill parents struggled all that night to keep the floodwater out of their cabin, but to no avail. When W.B. and his sister woke on Christmas morning, they were surprised to find that the water had entered their cabin and was nearly up to the bed. For a 4½ year old boy, this was an excuse for fun, jumping from the bed into the water. However, the family was compelled to abandon their cabin that morning and move in with a neighbor. When they left the cabin, they had to wade several hundred yards to reach dry land. W.B. recalled that the water was up to his armpits and running so strong that he had to hold onto his mother's dress. It was two weeks before they could return to their cabin.\textsuperscript{773}

During the January 1862 floods, the Tule River changed its channel for a considerable distance downstream from the foothills.\textsuperscript{774} Prior to the flood, the Tule ran northward between what are now Second and Third Streets in Porterville, turned west past the foot of Scenic Hill and thence in a northwesterly direction across the plain.\textsuperscript{775} After the flood, the river continued directly west as it left the foothills in a new channel some distance south of the settlement. The settlement that came to be called Porterville was relocated as a result of the flood.\textsuperscript{776}

The Butterfield Overland Mail route crossed the Tule River at the foot of Scenic Hill (at the junction of present-day Main St. and Henderson Ave). The Tule River stage stop was located at that point, and was operated by Porter Putnam for a time. This portion of the stage route was discontinued in March 1861 due to Civil War fighting in the South. However, Porter stayed to found the town named for him.\textsuperscript{777}

In the Tulare Lake Basin, reports stated that a damaging flood on the Tule River overflowed farms to a depth of several feet. The lowlands along the tributaries of Tulare Lake were probably flooded continuously from the middle to the end of January 1862.\textsuperscript{778}

After the Tule changed course, the area between the old and new channels was declared swampland under terms of the Swamp Land Act of 1850, which provided for the reclamation of such land. To satisfy the terms of the law, applicants for such land were said to have loaded a rowboat into a wagon and, sitting in the boat, driven over the land which they were interested in claiming. They could then swear to an affidavit that they had gone over the land in a boat and that it now had been reclaimed. It is likely that the federal authorities in Visalia knew of the little joke, but settlers were wanted, and the land was valued at only $1.25/acre anyway.\textsuperscript{779}
The White River had a major flood in the gold mining district. The river rose 5 feet higher than it had in the 1852 flood. There was great property damage.\textsuperscript{780} Poso Creek flooded on January 18 with a rush of logs and water 60 feet high.\textsuperscript{781}

The 1861–62 flood was a major flood on the Kern, causing huge property damage in the mining country. Virtually all the bridges, dams, and mills were destroyed. The river was 45 feet higher than ever previously recorded.\textsuperscript{782}

The rain and flooding on the Kern lasted for two months. There were only a few settlers living in the lower portions of Kern River Island (present-day city of Bakersfield), and for the few years that they had lived there, the rising winter runoff had spared their tule and adobe homes. Christmas Day, 1861, though, was not like the past light floods that had occurred regularly. The floodwaters rose during the night. Within a few hours, every home in the low-lying areas was washed away. The plight of the settlers was described in the \textit{Visalia Delta}:\textsuperscript{783}

\begin{quote}
The settlement known as Alkali City or Kern River Island is also ruined. They have lost all — stock, grain, and everything else — scarcely escaping with their lives. Several of the inhabitants were forced to remain on a very small island ten or twelve days, with nothing to eat except half rations of roasted corn.
\end{quote}

The flood of December 25 changed the river channel at the site of the present-day city of Bakersfield and inundated all except the higher knolls in the vicinity. It seems certain however, that the flood of 1861–62 flood was not as great as the 1867–68 flood would be.\textsuperscript{784}

Gordon’s Ferry (aka Gale’s Ferry) was located just north of present-day Bakersfield College. The Sinks of the Tejón was the first Butterfield Overland Mail stop north of Fort Tejon. It was located at the intersection of present-day David and Wheeler Ridge Roads, roughly 10 miles northeast of where Interstate 5 and Highway 99 diverge. When the Kern River came out of the canyon, it created one vast sea of water from Gordon’s Ferry to the Sinks of the Tejón. Somewhere within that huge sheet of water was Kern Lake. Buena Vista Lake backed up to within 12 miles of Fort Tejon.\textsuperscript{785} A major debris slide formed on the South Fork Kern in January 1862. This was described in the \textit{Visalia Delta}:\textsuperscript{788}

\begin{quote}
The cause of this disaster is owing to a slide from the mountain, filling up the bed of the stream, the water forming in the immense reservoir above, and after forcing its way through the obstruction, forbid all opposition...The crumbling of the mountain is described by those who saw it as a grand and terrific sight. Huge masses of rock were hurled from their base, trees uprooted, were sent whirling through the air, and this mass of matter gathering force, as it went, came down the steep mountain declivity, with wild and terrific confusion, indescribable. Mr. Jacob Macomb and family, residing on the South Fork of Kern, were awaked at the midnight hour, by the water and had barely time to leave their house before it fell.
\end{quote}

Reed Tollefson thinks that the above landslide may have occurred fairly far upstream on the South Fork Kern, perhaps on land that is now within Sequoia National Forest. The South Fork Valley has very flat topography. However, just above these private lands on the national forest, the river enters a deep gorge for many miles. Presumably that is where the slide occurred.

Tulare Lake had been at a very high stage after the 1852–53 flood, the second highest ever recorded. After 1853, there was a gradual shrinkage of the lake until the fall of 1861. Over those eight years, the lake dropped about 13 feet in elevation.
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

There were multiple causes of this:
- The maximum elevation in 1853 (215.5 feet) was higher than the elevation of the delta sill (207 feet). The water above this elevation simply flowed out of the lake and connected through the Fresno Slough to the San Joaquin River, and from there it flowed on to San Francisco Bay.
- Normal evaporation in our hot valley summers.
- Eight years with only low or average runoff and no floods.
- Two years of drought (1856 and 1857).
- Diversion for irrigation was just getting underway.

The 1861–62 flood raised the lake by 15.7 feet to elevation 216, the highest that the lake has been during historic times. At elevations above 207 feet, the lake over-topped the lowest point on the delta sill. At the lake’s highest stage, about 9 feet of water was flowing in a broad expanse northerly over this ridge (216–207 feet). From there, the water flowed into the Fresno Slough and the San Joaquin River. At the height of this flood, the lake was about 37 feet deep at the deepest point (216–179 feet). The surface area increased from about 350 square miles in 1861 to about 790 square miles in July 1862.

C.E. Grunsky estimated that more than 5,000,000 acre-feet of water flowed into Tulare Lake in the single season 1861–62. For comparison, that is 3.1 times greater than the combined current capacity of all four of the federal reservoirs in the Tulare Lake Basin. Using more precise data, S.T. Harding later recalculated the total inflow as being 6,290,000 acre-feet of water, the equivalent of 3.9 times the combined current capacity of our present-day reservoirs.

The 1861–62 flood was a record flood in the Tulare Lake Basin not just because of its volume. The force of the flood was such that all four of the major rivers (Kings, Kaweah, Tule, and Kern) cut new channels. Given that, one can only wonder that Visalia received only the most minor of flooding, and that essentially no livestock drowned in the vicinity of the Kaweah Delta.

One source said that thousands of cattle were drowned in the Tulare Lake Basin during the 1861–62 flood. No details were provided to support that claim. Livestock deaths on the Kaweah, Tule, and White Rivers appear to have been minimal. The Kern did experience serious flooding, so perhaps some cattle drowned in that area.

Rain and snow alternated on the floor of the Owens Valley with some precipitation every day from late December 1861 through mid-February 1862. Owens River was ¼–½ mile wide and Owens Lake rose 10–13 feet during that winter.

The winter storms engulfed all of Southern California. Beginning on Christmas Day, 1861, the Los Angeles area had 15 days of gentle rain followed by 28 continuous days of heavy rain. During the course of the 1861–62 season, Los Angeles received over 66 inches (5½ feet) of rain. The Mojave River rose 20 feet above average in present-day Oro Grande. Lakes formed in the Mojave Desert. Planes were cut by gulches and arroyos from Ventura to San Luis Rey. (Mission San Luis Rey de Francia is located just south of Marine Corps Base Camp Pendleton in San Diego County.)

The Santa Ana River flooded catastrophically on the night of January 22, 1862, sweeping away the village of Agua Mansa (literally "Gentle Water"), located on the Santa Fe Trail just south of present-day Colton. Hearing the roar of the river that night, Father Borgotta frantically rang the church bell, sounding the alarm. The inhabitants of the village ran or swam to high ground. "The gentle Santa Ana River became a raging torrent which washing, swirling and seething, swept everything from its path." One writer reported that there were "billows fifty feet high." Peter C. Peters of Colton recalled that "when morning came — (there was) a scene of desolation." Only the church and a house near it survived.

USGS reconstructed the cross-section of the flood and determined that the normally placid Santa Ana had been flowing at approximately 320,000 cfs that night. That was three times greater than any subsequent flow in the area, even the 1938 flood.

The Santa Ana flood formed two large lakes south of that river — one in the Inland Empire and another in the floodplain of Orange County. The lake in Orange County lasted about three weeks with water standing four feet deep up to four miles from the river. Sediment cores suggest that the last time the area saw a flood that big was in about 1600.
In San Diego, over seven inches of rain fell in January alone. All of Mission Valley was underwater, and Old Town was evacuated.\textsuperscript{792} The 1861–62 flood was the flood-of-record for much of Southern California.\textsuperscript{793}

Figure 26 shows approximately how much of the state was under water during the 1861–62 flood.

\begin{center}
\includegraphics[width=\textwidth]{figure26.png}
\end{center}

Figure 26. Map of areas under water during the 1861–62 flood.

\textbf{1863–65 Drought}

This three-year drought followed on the heels of the huge 1861–62 flood. It affected at least the Central Valley and possibly a larger area.

This drought is often referred to as the Great Drought of 1863–64. Tulare Lake continued its steady decline during The 1855–61 drought.\textsuperscript{794}

During this drought, Sacramento received less than half its average rainfall and the San Joaquin Valley was even drier. In the first year of this severe drought, William Brewer described the San Joaquin Valley as "a plain of absolute desolation." On February 27, 1864, the \textit{Stockton Independent} reported that there had been no rain of consequence for 11 months. By the end of March, the wheat and barley fields around Stockton were dried out completely.

An excerpt from \textit{Exceptional Years: A History of California Floods and Drought} by J.M. Guinn, 1890:\textsuperscript{795}

\begin{quote}
1862-63 did not exceed four inches, and that of 1863-64 was even less. In the fall of 1863 a few showers fell, but not enough to start the grass. No more fell until March. The cattle were dying of starvation.... The loss of cattle was fearful. The plains were strewn with their carcasses. In marshy places and around the cienegas, where there was a vestige of green, the ground was covered with their skeletons, and the traveler for years afterward was often startled by coming suddenly on a veritable Golgotha – a place of skulls – the long horns standing out in defiant attitude, as if protecting the fleshless bones.
\end{quote}

The year 1864 was an extreme drought year. As shown in Table 21, it was the tenth driest year on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

In the Tulare Lake Basin, 1864 is remembered as the most severe year of the 1863–65 drought by far. Most of the streams dried up, the feed either did not mature or withered, and there was not even sufficient water for drinking. Kathleen Small said that the drought was remembered as the Drought of 1864.796

Cattle died in large numbers during 1864, bringing their numbers down dramatically. Sheep came through the drought in better shape, their numbers increasing. The double whammy of the 1861–62 flood followed by the 1863–65 drought is one of the main reasons that the Sierra grazing business changed from primarily cattle to primarily sheep.

Table 32 uses the livestock censuses to illustrate the dramatic change that occurred between 1860 and 1870.

Table 32. Livestock censuses of the San Joaquin Valley.

<table>
<thead>
<tr>
<th>Year</th>
<th>Cattle</th>
<th>Sheep</th>
</tr>
</thead>
<tbody>
<tr>
<td>1860</td>
<td>226,248</td>
<td>78,568</td>
</tr>
<tr>
<td>1870</td>
<td>288,483</td>
<td>901,892</td>
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</tbody>
</table>

This drought played a major role in shaping the state’s historical development by contributing to the demise of the cattle rancho system, especially in Southern California. The widespread economic damage that this drought caused to California agriculture reflects the dominance of non-irrigated agriculture at the time, the limited extent of water infrastructure, and the absence of groundwater pumping technology.797

William Brewer wrote about the drought and the effect of sheep grazing in 1864. Brewer’s party arrived in Visalia on June 6 of that year. A few days later, they undertook what has become one of the great steps in American mountaineering, entering the Sierra at Big Meadow. They finally emerged at Mariposa at the end of the summer. When Brewer eventually got around to counting up his miles of travel, he found that he had ranged over 15,000 miles within the state — but that summer of 1864 spent exploring the crest of the Sierra had to be among the best.798

W.B. Cartmill recalled how severe the drought was on the lower part of the Kaweah Delta. Tulare Lake served as a gigantic watering hole. Thanks to the huge 1861–62 flood, the lake was brimful when the drought set in. Vast herds of cattle would spread over the country for miles, traveling as far back from the lake as they could go without water in search of the scant grasses. Then they would rush back to the shore each day to quench their thirst. By 1864, grass had disappeared completely on the plains, with oak leaves, acorns, and salt grass serving as fodder of last resort. Cattle died in large numbers, their stench filling the air.799 Their hides were taken, but the meat was left to rot. It’s easy to visualize the scene with huge flocks of turkey vultures and California condors, reminiscent of the Serengeti or the Pleistocene.

David Campbell, an early Tulare County pioneer, recalled that 1864 was so dry that the grass did not even get started that year.800

After the Kaweah River reaches the area of the present-day Terminus Dam, it flows onto its delta and divides into distributaries. The 1861–62 flood deposited considerable quantities of drift, silt, and sand on the Kaweah Delta. A portion of the channel of the Lower Kaweah River was obstructed by these deposits, significantly reducing its ability to carry flows. During low-flow periods, this section of the Kaweah was now dry; it only served to carry water during flood periods.

In addition, the 1861–62 flood had created a new distributary, which came to be called the St. Johns River, along the north side of the delta. The St. Johns would remain the principal channel of the Kaweah until 1877 when the Fowler Cut would reopen the old Kaweah River channel. This worked out great for north-side farmers, but not well at all for farmers on the south side of the delta.

As the low-water period of 1862 approached, irrigation water was scarce for those who depended upon Deep Creek, Packwood Creek, and Visalia Creek for their supply, and many projects were proposed for relief. (Visalia Creek was originally known as Tiber Creek. We now know it as Mill Creek.) During 1862 and the years immediately following, a number of ditches were constructed from the St. Johns River near Rocky Ford, southwesterly to the original main channel of the Lower Kaweah River. All of those ditches were intended to increase the flow in the delta streams in the Visalia area and on the south side of the delta.

After much contention between settlers on these several streams as to the apportionment of their scant supply of water, they finally reached agreement. In 1867, a gate was constructed in the head of Visalia Creek
(presumably where the Lower Kaweah River ends, just north of the Ivanhoe turnoff (Road 156/158) on Highway 198). The timing was not good; that gate would last less than a year before being swept away by the 1867–68 flood.

The drought, especially with its near total failure of winter pasture grasses, forced ranchers to look elsewhere for previously unused rangelands. What resulted was the first major utilization of the Sierra for large-scale livestock feeding. During the drought years, hungry cattle from the lowlands swarmed over the Sierra foothills and forests while the high country suddenly found itself assaulted by huge herds of domestic sheep. Within a few years, much of the herbaceous vegetation of the Sierra had either been destroyed or replaced. In the foothill grasslands, annual Eurasian grasses replaced the grazing-sensitive native perennial species. In the high country, entire basins were so thoroughly denuded that parties traveling on horseback lamented the almost total lack of feed for their animals.\(^{801}\)

In 1859, Paschal Bequette, Sr. brought his family to Visalia and became a cattle and horse breeder. He recalled that they saved their horses during the 1863–65 drought by taking them up the South Fork trail to Hockett Meadow where there was good feed and water.\(^{802}\)

Floyd Otter said that valley ranchers also drove their hogs into the high country during drought years.\(^{803}\) In July 1864, at the height of the worst drought year, Clarence King took a trip from Visalia to the Mt. Whitney country. In a letter to Josiah Whitney (Chief of the California Geological Survey), he wrote:

> I rode until nine in the evening, when we came to the “Hog Ranch,” two acres of tranquil pork, near a meadow in the most magnificent forest in the Sierras.

That “ranch” was probably a temporary drover’s camp on the South Fork of the Kaweah. Floyd Otter thought that it might have been in or near Hockett Meadow. King later described this pig-herd in the words of its owners as “The poottiest hogs in Tulare County — nigh three thousand.” One can only imagine what 3,000 “half-wild boars, sows, and pigs” could do in a summer on the Hockett Plateau.

The route that Paschal Bequette, Clarence King, and the hog drovers took was presumably along the newly constructed Hockett Trail. This was a toll trail from Visalia up the South Fork of the Kaweah to the Cerro Gordo silver mines in Inyo County. The trail was built by John Benjamin Hockett and his partners under a charter granted by the Tulare County Board of Supervisors in December 1862.\(^{804}\) It was built in less than two years and was open for use in August 1864.

According to Samuel Thomas Porter’s history of the Mineral King mining rush, Pleasant Work (what a nice name, he was one of the sons of Hop Work) was also known to run hogs in Hockett Meadow before 1867.\(^{805}\)

Stockmen learned at least three important lessons from the great droughts of 1855–61 and 1863–65:
- Severe droughts were common.
- Sheep came through the droughts in better shape than cattle.
- The Sierra provided grazing for sheep, cattle, horses, and pigs during droughts.

The 1863–65 drought was viewed as something of a blessing by the folks trying to drain the swampland around Bakersfield.

The drought finally ended in November 1864, and three years of average precipitation followed.

**1867–68 Flood**

Flooding in 1867–68 occurred primarily in December 1867. On the lower parts of the Sacramento — San Joaquin River systems, the floods carried over into January 1868.\(^{806}\)

The floods resulted from exceptionally heavy rain during the period December 21–25, 1867, which extended throughout Northern California.\(^{807}\) One feature of this storm event is that the intense part of the storm was apparently preceded by an extended period of soaking rain, at least in the Southern Sierra. The uncommonly heavy precipitation occurred in December 1867 and January 1868, and it fell on both the west side of the Sierra and the east side. Camp Independence received a record-setting 19.39 inches of precipitation in water year 1868 compared with an average annual of about 5.47 inches. Of that total, 12.19 inches fell in December 1867, and another 5.46 Inches fell in January 1868.\(^{808, 809}\)
Hopkins (Hop) Work and his family settled in the Three Rivers area in 1858 or 1859. They built an adobe cabin on the east side of the South Fork, close to where it joined the mainstem of the Kaweah. That adobe would later gain a reputation of being notorious for some unknown reason. One of their sons, (presumably Enoch), built a home for his family on the west side of the South Fork. Enoch’s family planted an apple orchard in the vicinity of present-day Cherokee Oaks in 1865. Sophie Britten thought that was the first orchard in Three Rivers. However, Earl McKee thought that Ira Blossom planted a pear orchard on the opposite side of the river a year or so before this. 810, 811

Joseph C. (Joe) Palmer arrived on the South Fork Kaweah at roughly the same time as the Work family, but settled 13 miles upstream near the present-day South Fork Campground. He called his homestead Rose Bud. Work and Palmer were involved in a fight with the American Indians shortly after the Work family settled in the area. (Palmer would later acquire property in the Three Rivers Area next to Ira and Julia Blossom.) 812, 813

The second family to settle in the Three Rivers area was that of Ira and Julia Blossom. The Blossoms arrived after the summer of 1860, but were there by 1861. Their homestead was on the lower slopes of Blossom Peak. Their first house, an A-frame, was in the floodplain of the South Fork. But after the December 1867 flood destroyed their home, they rebuilt on higher ground nearby. 814, 815

Bob Barton (one of James Barton’s sons) and Muriel Kenwood (one of Julia Blossom’s great-granddaughters) gave an account of the December 22, 1867 landslide dam flood on the South Fork Kaweah to Frankie Welch. According to their account, there had been an abundance of rain in November and again in December; thus accumulating a heavier than normal snowpack in the high mountains. Then the weather turned warm in mid-December and a heavy rain started falling; it rained steadily for four days and four nights. The warm rain, falling on all that snow, melted it and raised the levels of the rivers to flood stage. Barton and Kenwood said that the warm rain fell on December 23, but they may have gotten that date off. 816 Given the flood conditions in Visalia, it seems like the rain should have occurred at least one day earlier than this.

Joseph Palmer gave an account of the landslide dam flood to Judge Walter Fry on October 9, 1890; this was nearly 23 years after the flood occurred. He told Fry that it had been raining in the Three Rivers district almost steadily for 41 days and nights (presumably from November 9 through December 20), with heavy snows above the 5,000 foot level. All the rivers were very high. The weather turned warm on December 21, and a hard rain fell all day, even at high elevation. 817

Palmer’s account of the flood conflicts in significant ways with the account of Bob Barton and Muriel Kenwood. See the section of this document that describes the Landslide Dam Failure #1: South Fork of the Kaweah. Palmer was almost certainly an eyewitness to the flood, but the dramatic account that he gave Walter Fry was apparently fabricated to fit the facts. However, his account of the precipitation seems to be reasonably consistent with the accounts of others.

The flood was apparently caused by the exceptionally heavy warm rains falling on a big snowpack and saturated soil. That combination resulted in major floods on all the main tributaries of Tulare Lake. 818

The winter of 1866–67 had been rough on the Chinese laborers constructing the Transcontinental Railroad over Donner Summit. Avalanches had wiped out two of their work camps. The winter of 1867 proved equally challenging. Sub-tropical storms deluged the region with more than 40 inches of rain in December 1867, causing extensive flood damage.

In the northern part of the Central Valley, the 1861–62 flood is generally the flood-of-record. In the Sacramento River Basin, the main river and its lower tributaries were at extreme flood stages between December 22, 1867 and January 2, 1868. However, the floods of 1867–68 are believed to have been generally lower in discharge and volume than those of 1861–62. At some points the American and lower Sacramento Rivers were reported at higher stages in 1867–68, but it is probable that these high stages were caused by aggradation of stream beds or by channel contraction due to levee building. In the Sacramento River Basin upstream from the Feather River, the floods of December 1867 were definitely secondary to those of 1861–62. 819

The weather conditions during 1861–62 resulted in above-average precipitation between the Columbia River and the Mexican border. Major flooding was widespread throughout this area. However, the 1867–68 flood was especially severe on Sierra Nevada streams tributary to the southern part of the Central Valley. In the foothills, the flood on the San Joaquin River exceeded considerably any other known flood and was probably higher than
any known flood at all points upstream from the mouth of the Merced River. However, the San Joaquin River stages downstream from the mouth of the Stanislaus River were not as high in 1867 as in 1862.

During recorded history, the 1867–68 flood was one of the greatest in the Tulare Lake Basin. Peak stages in that region during December 24–25 were the highest of record. Major floods occurred on all the main tributaries in the Tulare Lake Basin. The Kings, Kaweah, Tule, and Kern Rivers carried floodflows in 1867–68 that are believed to be the greatest known, exceeding those of the 1861–62 flood.

Flooding occurred throughout the San Joaquin Valley in December 1867, extending barely into January 1868. The San Joaquin Valley was described as looking like an ocean. Unlike the 1861–62 flood, this flood lasted only weeks rather than months.

The preceding multi-year drought had ended in November 1864. One account said that the high country then experienced two consecutive years of heavy snows with virtually no summer between. This supposedly resulted in a huge accumulation of snow in the Sierra.

Whether or not that was true, there are multiple accounts that rain and snow began in mid-November 1867 in the Kaweah River Basin and came down almost continuously through December. One account said that the snowline was at about 5,000 feet until December 20, at which point the weather turned warmer. Presumably similar weather conditions were happening throughout the Central and Southern Sierra.

As on the other rivers in the Tulare Lake Basin, flooding in 1867 on the Kings was greater than the 1861–62 flood. It is considered to be the greatest flood on the Kings since at least the flood of 1805. The 1867–68 flood completed the formation of Cole Slough. (See the section of this document on Pine Flat Dam for a more detailed description of the formation of Cole Slough and associated waterworks.) An outstanding characteristic of the flood of 1867 on the Kings River, as well as on the Kaweah, Kern, and upper San Joaquin Rivers, was the tremendous quantity of timber brought down from the Sierra and deposited on the plains.

Descriptions of the flood of 1867 on the Kings River do not given a definite comparison with the flood of 1862. The Kings engulfed and destroyed the newly rebuilt town of Scottsburg. (Even though that townsite had been selected because it was thought to be safe from flooding.) From this fact it appears that the 1867 flood was at least as severe as that of 1862, and probably reached a greater height. The community was then rebuilt at an even safer location and renamed Centerville.

From reliable accounts by an eye-witness of the flood of 1867, the Kings River reached a stage about 3 feet greater than in 1937 at a point one mile downstream from Piedra. At the Pine Flat dam site on the Kings, about three miles upstream from Piedra, it was later determined from the position of cedar and other drift logs deposited along the channel, that a previous flood had exceeded that of 1937 by at least 7 feet. The rise in December 1937 was about 18 feet above low-water. From statements of ranchers who settled in this vicinity about 1875 it is believed that those logs were deposited either in 1862 or 1867.

As on the other rivers in the Tulare Lake Basin, flooding in 1867 on the Kaweah was greater than the 1861–62 flood. It is considered to be the greatest flood on the Kaweah since at least the flood of 1805, 1827, 1828, 1829. Again, as with the upper San Joaquin, Kings, and Kern Rivers, the Kaweah brought tremendous quantities of timber down from the Sierra and deposited them on the plains. Smith Mountain is about a mile east of Dinuba. By some accounts, flooding was so extensive in the 1867–68 flood that one could have ridden in a boat from Smith Mountain to the Tule River.

The town of Visalia was partly flooded by water from the Kaweah River on December 23, and by December 24 the flood stage in the town had exceeded the record of 1862 by 4 inches. After receding about 2 feet, the water again rose to about the same stage on December 26. The stage of the Kaweah River in its channel downstream from the foothills was reported to have exceeded the stage in 1862 by 2 feet. As determined by the position of the sequoia and cedar logs deposited by this flood, the maximum stage on the Kaweah near Three Rivers referred to the datum of the old Horse Creek gage (now submerged under Lake Kaweah) was about 20.0 feet.

According to the book Floods of the Kaweah, the Kaweah rose 17.5 feet above the average low-watermark in the foothills. Despite considerable searching, Valerie McKay was unable to find the source of that figure or where the watermark was measured. The earliest gage on the Kaweah was a USGS staff gage (USGS gage #11-2105) that was located near Three Rivers just above the junction with Horse Creek; that gage was established.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

in October 1903 and was read twice daily. Harry H. Holley established the measuring station at McKay's Point in October 1916 for the River Association. Perhaps the watermark reported above was measured at the Horse Creek gage.

The flood deposited considerable quantities of drift, silt, and sand on the Kaweah Delta, raising its elevation. The head of the Lower Kaweah River channel had been partially closed by the 1861–62 flood. The 1867–68 flood further obstructed that channel by depositing still more drift and silt in it.

The flood refilled and otherwise destroyed some of the ditches that had been constructed to bring water from the St. Johns River back into the old channel of the Lower Kaweah River. It washed out the new gate in the head of Visalia Creek (what we now know as Mill Creek). It partially closed the head of Packwood Creek (just north of the Ivanhoe turnoff (Road 156/158) on Highway 198). It also partially closed the head of Deep Creek (northeast of the present-day Kaweah Oaks Preserve).

The 1867–68 flood further enlarged the St. Johns River that had been created in the 1861–62 flood. It eroded a new head of the St. Johns River about a mile farther upstream, farther into the swamp. This new point of separation of the two channels became known as McKay's Point. That was the original spelling, but it is now sometimes written as Mckays Point (as on the Woodlake USGS quad) and McKay Point (the preferred spelling used by the Kaweah Delta Water Conservation District and at least sometimes by the U.S. Army Corps of Engineers). This document uses the original spelling: McKay's Point.

McKay's Point is located about a mile northwest of present-day Lemon Cove and three miles below Terminus Dam. A variety of structures would later be built at this location in an attempt to control the Kaweah and split the flow between the Lower Kaweah River and the St. Johns River. The river's natural tendency to relocate this point has been actively resisted, similar to the much bigger struggle to keep the Mississippi from following its preferred course down the Atchafalaya.

In 1870, a brush and rock diversion weir was built at McKay's Point. Presumably this weir served to allocate the water between the St. Johns and the various minor distributaries (aka creeks and ditches) that still connected at that point. It would remain that way until 1877 when the Fowler Cut (built by Samuel Fowler under contract to the Kaweah Canal and Irrigation Company) would reopen the Lower Kaweah River channel. The brush and rock diversion and weir at McKay's Point was also rebuilt in 1877. Once the Lower Kaweah River channel was rewatered, then the irrigation ditches attached to that channel got reliable water for the first time in 15 years.

No doubt the brush and rock diversion weir at McKay's Point had to be frequently repaired and it apparently had to be completely reconstructed in 1884 and 1897. The first concrete weir was built at McKay's Point in about 1909. It had to be replaced after the 1937 flood. That diversion weir was destroyed or at least bypassed in the 1955 flood. The current concrete weir diversion at McKay's Point is maintained and operated by the Kaweah Delta Water Conservation District.

Stringtown was a settlement of five families living in a line south of present-day Woodlake along the Kaweah River, east of Bravo Lake. (This reference to Stringtown's location is somewhat unclear. Bravo Lake was actually located adjacent to the St. Johns River.) The 1867–68 flood came into several of their houses; only one of which was on sufficiently high ground to survive completely intact. After the flood, the other four families (J.W.C. Pogue and his relatives) relocated to the Dry Creek area. This gives the 1867–68 flood the often cited reputation of having destroyed Stringtown.\(^{834, 835, 836}\)

In 1864, the newly created People's Ditch Company had built 12 miles of ditch to what is now Farmersville. That system failed during the 1867–68 flood. Logs from the Sierra reached the Farmersville area. Some of those logs were apparently giant sequoias from the South Fork of the Kaweah. In the years after the flood, ranchers would use this wood to build fences.

During the 1867–68 flood, all the streams in Tulare County were reported to be on a rampage with great loss of property. The Kaweah and St. Johns Rivers made a vast expanse of waters. The Visalia area was awash with water; boats were widely used for transportation, and there was significant loss of property.

There are multiple accounts of the flooding which clearly relate to the general flooding that was occurring throughout the area. David Campbell, an early Tulare County pioneer, recalled that the floodwaters formed almost a solid sheet of water from Porterville to Visalia.\(^{837}\)
Another account was about three Visalia families who lived near where Packwood Creek crosses the present-day Mineral King Road (about two miles east of Lovers Lane on Highway 198). All of those families gathered at A.H. Broder’s place during the day because his was situated on higher ground. They then built a three-foot-high embankment, enclosing about half an acre of ground. The siding from the barn was removed, and a raft built. If the river continued to rise, they planned to move to a still higher sand knoll which lay to the southwest. By 9:00 the following morning, Broder, who had been keeping tabs on the water level by means of sticks, reported that it had receded half an inch and that it would not be necessary to move.

About 200 American Indians took refuge on the same high mound as Broder and his neighbors. They made a gala festival of the predicament. Squirrels and rabbits in great numbers were caught and hung on lines to dry; the flood affording both amusement and provender.

Another account that has survived was about the residence of the Evans family, which was located on high ground near present-day Tulare Avenue in the general vicinity of Ben Maddox or Santa Fe. (Later this location was known as the Evansdale Orchard.) The water had risen previously, and then it rose again suddenly during the night. It surrounded their home and almost engulfed some of their neighbors’ homes. The Prothero family lived on the Bentley place and there the water ran through the windows. They moved to the Evans home for shelter.

Then came a call for help from the home of Mrs. Williams, who lived adjoining. This was about 1:00 in the morning, pitch dark and the swirling waters icy cold. Mrs. Williams had a baby but four or five days old and was unable to walk. Samuel and James Evans waded over, and placing her in a rocking chair, carried her to safety. Tom Robinson, with his wife and family, also took refuge with the Evenses, making a total of 25 gathered there. The barn, several hundred yards away, half full of hay, provided the only place for sleeping quarters for so many people.

Between it and the house, the water ran two or three feet deep. Luckily, a boat had previously been constructed in which to go to Visalia, and so the half-dried refugees cuddled around the stove in the Evans’s kitchen were enabled to get to bed without again getting wet. Jim Evans, acting as gondolier, conducted his guests to their hay mow lodgings. This nighttime flooding event may have occurred on the night of December 23–24.

As on the other rivers in the Tulare Lake Basin, flooding in 1867 on the Kern was greater than the 1861–62 flood. It is considered to be the greatest flood on the Kern since at least the flood of 1805. In the Kern River Basin, the flood was at high stage from December 25, 1867, to January 1, 1868. A remarkable feature of the flood was the large quantities of logs from the Sierra, including cedar and giant sequoia, that were deposited on the overflowed lands of Kernville and Bakersfield.

All of the streams in the southern part of the Central Valley reached peak stages during December 24–25. Three witnesses told Walter Fry that the flood on the Kaweah arrived in Visalia late on the evening of December 23. Ira Blossom was one of those three witnesses. He used to work at the Visalia grist mill in the early days after his family moved to Three Rivers. That is why he was not home to help his wife when the flood swept through Three Rivers.

The 1867–68 flood resulted in the deepest flood depths ever on the streets of Visalia. The 1861–62 flood put a maximum of 24 inches on Main Street. When the peak of the 1867–68 flood arrived in Visalia, it flooded the development along Mill Creek 5–6 feet deep as measured at the grist mill. That mill was on Mill Creek, on the southeast corner of Main and Santa Fe Streets.

Walter Fry recorded the experiences of several families who lived in the countryside around Visalia. One of those, Betty Townsend, lived near Cutler Park on the St. Johns. On the evening of December 23, 10 people sought refuge in her house and stayed there for a week. Her account included:

> Our Christmas dinner, in part, consisted of a turkey feast. The turkey was captured by one of the party from a bale of hay which was being swept down the torrent. A pig was similarly rescued and consumed.

Bob Barton (one of James Barton’s sons) and Muriel Kenwood (one of Julia Blossom’s great-granddaughters) said that the streets of Visalia were under water for six weeks.

From the meager reports available, it is probable that the Tule and White Rivers and Deer and Poso Creeks reached exceptionally high stages in December 1867.
The Tule River spread all over the Poplar and Woodville sections.

Deer Creek and the White River merged their waters in their lower courses.

Newspaper accounts stated that the Tule River was higher in 1867 than in 1862. Downstream from the foothills it flooded farmlands and, as in 1862, cut a new channel for a portion of its course. The lowlands between the Tule and Kern Rivers were described as having been almost completely flooded. As with the upper San Joaquin, Kings, and Kaweah Rivers, the Kern brought tremendous quantities of timber down from the Sierra and deposited them on the plains.

The Kern River had rerouted to the west in the 1861–62 flood. The 1867–68 flood moved the river channel even farther north to its present location, ending in the Buena Vista Slough a few miles north of Buena Vista Lake and entered that lake from the northwest as shown in Figure 14.

Tulare Lake gradually declined in elevation after the 1861–62 flood. In the summer of 1867, the lake level was 200.7 feet. However, the 1867–68 flood raised it by 14.7 feet, bringing it back to a maximum elevation of 215.4 feet. At the height of this flood, Tulare Lake was almost 37 feet deep at the deepest point. The lake has not been that deep since.

The onset of the flooding in the lakebed swept in at night, catching people by surprise. Jack Phillips (son-in-law of early pioneer Dan Rhoades) and four other guys had their hogs together and were camped near each other on the lakeshore below present-day Stratford. When they went to sleep at night, there was no water in their camp. By morning, the hog camp was going under water and a hurried retreat was made toward high ground. The water was at their heels all the way to the Dan Rhoades adobe in spite of the best time they could make. The same morning, Doc Vaness came out of the lakebed with his family just in time to avoid drowning.

Based on limited information, the flood described above most likely occurred on Christmas Eve, 1867. Peak stages on all the main tributaries in the Tulare Lake Basin occurred during December 24–25. The flood on the Kaweah arrived in Visalia on the evening of December 23. Those floodwaters would probably have reached the Tulare Lakebed on the morning of December 24 (having traveled roughly 30 miles at about 2.5 m.p.h. = 12 hours). The much larger volume carried by the Kings River would have had a longer route to travel, so those floodwaters presumably took somewhat longer to reach the lakebed. That would fit with a possible arrival time of the evening of December 24.

In any case, flooding was eventually so extensive that boats bearing supplies were reported to have passed freely from Visalia to places in Kings and Fresno counties. One account said that the flooding made the San Joaquin Valley a continuous lake of water from Buena Vista Lake to the San Joaquin River; that seems quite plausible given the height of Tulare Lake.

In 1868, Richard Smith loaded a 16-foot scow with a one-ton cargo of honey and made the 170-mile journey from Tulare Lake to San Francisco Bay. That remains the only recorded commercial trip ever made between Tulare Lake and San Francisco Bay to occur in historic times. (There were five non-commercial trips: 1852, 1938, 1966, 1969, and 1983.) Apparently Smith was able to make the return trip back through the tules to Tulare Lake.

How different people viewed the 1867–68 flood was a matter of perspective. And their perspective was determined in part by the elevation where they were living when they experienced the flood.

- Joe Palmer lived at an elevation of 3,600 feet near present-day South Fork Campground. Therefore, he experienced the flood from a canyon perspective; a brief event that passed by causing him no harm (see description below).
- Visalia was about 331 feet in elevation. The new settlers there experienced the flood as the Kaweah spread across its delta, causing destruction.
- Tulare Lake was about 215 feet in elevation at its peak. For people who lived in the vicinity, it was lake flooding. The lake was a major resource that was sustained by such periodic flooding.

The storm extended to the east side of the Sierra. Bishop received more than 16 inches of rain from December 1867 through January 1868.
In addition to the widespread flooding described above, a very rare landscape-wide event occurred. After some six weeks of steady rain, the ground had become saturated to a considerable depth. Slopes across the Southern Sierra became unstable, resulting in an apparently large number of landslides during a nine-day period in December 1867. A small portion of those slides dammed flooding rivers, and then failed when those dams were overtopped.

The Southern Sierra was very sparsely populated in 1867, so the chance of any one of those landslide dam failures being noted and recorded, and that record surviving into the 21st century is slim. Through good fortune, we have been able to reconstruct four of those landslide dam failures to varying degrees.

Such events don’t add any water to the flood that is already occurring down below. But they withhold the water for a while, and then release it all at once, resulting in a wall of water, a pulse.

**Landslide Dam Failure #1: South Fork of the Kaweah**

One of those landslide dam events occurred on the north side of Dennison Ridge Peak on the night of December 20, 1867.

This is the largest landslide to have occurred in the national parks in historic time. This event is included in a USGS report of documented historical landslide dams from around the world. The event was analyzed by Walter Fry.

Joseph Palmer had a homestead near the present-day South Fork Campground. He told Walter Fry that it had been raining in the Three Rivers district almost steadily for 41 days and nights, with heavy snows above the 5,000 foot level. All the rivers were very high. The weather turned warm on December 21, and a hard rain fell all day, even at high elevation.

The soil involved in this landslide was described as a sandy loam. If so, it would have had lots of voids that could hold water. Based on Palmer’s account of 41 days of steady rain, the ground would have been saturated to depth. The storm on December 20 would then have been the triggering mechanism. That was the way that the somewhat similar Mill Creek Landslide began on the South Fork of the American River on January 24, 1997.

When the slope became unstable, a mass of dirt and vegetation broke loose from near the crest of Dennison Ridge. The head of that landslide began on a 45 degree slope. It swept 2½ miles down into the canyon of the South Fork of the Kaweah. The landslide had a total estimated mass of about 580,000 cubic yards (445,000 cubic meters).

The landslide stripped the steep hillside of a thick forest of giant sequoia, pine, and fir in a path of devastation that ranged from 1,500–4,000 feet in width. This included the westernmost 300–400 acres of Garfield Grove. Walter Fry calculated that 350 million board feet of timber came down in that slide.

Most of the big landslides in the Southern Sierra have contained a large component of rock, including huge boulders. However, this landslide was markedly different. It consisted largely of trees, a thick layer of sandy loam, and relatively small rocks. Many of the pines, firs, and sequoias were quite large, including sequoias up to 30 feet in diameter. There were relatively few large rocks in this debris to provide structural stability. Therefore, this landslide dam failed much more catastrophically than other large landslide dams, and it left relatively little evidence of its presence immediately adjacent to the river. A large chunk of the dam washed out with the first failure, but it apparently took several days to completely wash out the dam.

Walter Fry analyzed the site in 1931, 64 years after the event occurred. Presumably there were still large logs and other debris present to allow him to determine that the top of the dam had been over 400 feet high at its highest point. Some of that material may still be there today.

Fry described the top of the dam and its volume, but he did not describe its shape. We can conjecture that it had a relatively steep slope and was much higher on the side adjacent to Garfield Grove. It may have flattened out near the bottom of the slide, typical of other landslides. If that were the case, then the amount of water impounded might have been only on the order of 200 feet or so; we have no way of knowing the low point on the dam. However, there is no apparent evidence of the dam on the opposite side of the river (the Ladybug Trail side); presumably all of those logs were removed by the flood.

Although new growth has disguised most signs of the landslide, its effects are still dramatically apparent in the vicinity of Snowslide Canyon, where dense sequoia forest ends abruptly at a landslide boulder field. The absence...
of large, old-growth trees is apparent throughout the disturbed area. There appears to be an even-age stand of giant sequoias covering some of the lower part of the landslide’s path, presumably having germinated or taken root in the freshly disturbed mineral soil.

The landslide would have exposed a lot of mineral soil in and immediately adjacent to a giant sequoia grove. Significant sequoia germination would have been expected following this event. Walter Fry conducted a detailed survey of the area some 60 years later, but did not mention seeing sapling sequoias. Nevertheless, those even-age stands of relatively young sequoias are there today. Dating these trees could quantify how the grove responded to this large-scale disturbance.

The landslide occurred just before midnight on December 20. When it came to rest, it formed a landslide dam that was ½ mile wide and over 400 feet high at its highest point. The South Fork Kaweah was presumably running at flood or near-flood stage because of all the previous rain and snow. It didn’t take the river long to fill the temporary reservoir. The dam failed about 1:00 a.m. on December 22, just 25 hours after the slide occurred.

The collapse of the landslide dam produced a flood surge about 40 feet deep that rushed down the South Fork Canyon. Joe Palmer’s homestead near the present-day South Fork Campground was several miles below the slide. (He also had some property 13 miles downstream in Three Rivers near that of Ira and Julia Blossom.) Nearly 23 years later (on October 5, 1890) he gave a dramatic account of the flood to Judge Walter Fry. According to that account, just before midnight on December 20, Palmer:

"was aroused by a heavy rumbling sound such as I had never heard before, and which lasted for an hour or more. Then a great calm set in, and even the roaring of the river ceased. On leaving my cabin in the morning, I found that despite the heavy rain the river was low. From this I knew that a great slide had blocked the canyon above and that later the dam would give way and cause a flood...About 1:30 a.m. I was aroused by a tremendous thundering and rumbling sound which made my hair stand on end. I jumped out of bed, grabbed my clothing, and ran for safety up the mountain side some 200 yards from the river. In a few minutes the flood came along with a breast of water some 40 feet in depth that extended across the canon, carrying with it broken-up trees which were crashing end over end in every direction with terrific force and sound. The river remained high for several days, and all the while timber was going down and being swept clear out to the Valley."

The bursting of the landslide dam at 1:30 a.m. on the morning of December 22 let loose a great flood, and the impounded water spilled and smashed its way down the South Fork Canyon, carrying everything before it, including giant sequoia logs. Some of those sequoia logs can still be found along the South Fork. A particularly good place to see them is the peninsula of land at the confluence of Grouse Creek and the South Fork (the old Hat Maxon Ranch; Hat was Kirk Stiltz’ great uncle). That peninsula was right in the path of the flood and must have gotten hammered.

The land has since revegetated to some degree, but evidence of the flood is still quite visible. James and Kathleen Seligman, the present-day owners, have set the flood area aside as the Shangrila Nature Preserve. The preserve consists of several acres of land that were swept over and carved by the flood. Large rocks are strewn all about, some quite far from the river. Most impressive are several tall rows of rocks, each a couple hundred feet long, that may have been formed as levees on the side of high-flow channels. Some of the rocks on these levees are extraordinarily large, approaching the size of small cars.

Four large ponderosa pines are growing on the preserve; they appear to have germinated from seeds swept in by the flood. There aren’t any other ponderosas growing in this part of the canyon. Chunks of logs of various species lie scattered about the preserve. The sequoia logs are the largest and best preserved of these. Some of the giant sequoia logs are roughly 40 feet above and several hundred feet back from the river. Even when you are standing next to these logs, it is hard to imagine a flood big enough to have put them where they are. This is truly an awe-inspiring site to visit.

From here, the flood swept on down the canyon and through the small community of Three Rivers, 15 miles below the landslide.
The dramatic account that Joe Palmer gave to Walter Fry in 1890 placed him at his homestead 13 miles up the canyon near the present-day South Fork Campground. He also reported near-Biblical amounts of rain: 41 days of steady rain followed by a day of hard rain. Those aspects of Palmer’s account conflict with the account that Bob Barton (one of James Barton’s sons) and Muriel Kenwood (one of Julia Blossom’s great-granddaughters) gave to Frankie Welch.  

According to their account, the weather turned warm and a heavy rain started falling; it rained steadily four days and four nights. There had been an abundance of rain in November and again in December; thus accumulating a heavier than normal snowpack in the high mountains. The warm rain, falling on all this snow on December 23, melted it and raised the levels of the rivers to flood stage.

Sometime early that day, the settlers on the South Fork noted that although it was still raining hard, the river was dropping rapidly and had receded to a low level. They knew that something drastic had happened up the canyon and that disaster was imminent. They immediately started moving their belongings to higher ground.

Sometime during that night an unbelievable torrent of water came rushing down the South Fork, demolishing everything in its path and leaving nothing but devastation in its wake.

(Joe Palmer’s account said that the landslide dam occurred just before midnight on December 20 and broke at 1:30 a.m. on the morning of December 22. The Bob Barton / Muriel Kenwood account said that the landslide occurred early on the morning of December 23 and broke that night. Therefore, their account has everything happening about two days later than Palmer’s account. December 23 is the night that the big flood on the Kaweah peaked in Visalia, so it is conceivable that this date got attached to the Three Rivers event.)

Although the settlers in Three Rivers all escaped with their lives, they lost almost all of their buildings, fruit trees, fences, livestock and land. Joe Palmer, a large man, carried Julia Blossom and several of her children to safety through waist-deep floodwaters.

That account makes it pretty clear that Palmer was present in Three Rivers at the time of the flood instead of at his homestead 13 miles up the canyon. Joe may well have been a hero in Three Rivers, but the dramatic account he gave Walter Fry was apparently fabricated to fit the facts. Joe couldn’t have been two places at once. And if that account was fabricated, then we shouldn’t necessarily take Joe’s report of 41 days of steady rain followed by a day of hard rain literally. It may or may not be factual.

George Cahoon was an early settler on the South Fork. Garry Kenwood was one of the grandsons of Pansy Blossom. Pansy told her grandchildren she had heard that Cahoon rode up, saw the landslide dam, and realized that it was close to failure. He then rode downstream, warning the settlers about the impending flood. That story may or may not be factual.

The floodwaters swept away the Enoch Work home on the west side of the South Fork. Apparently Hop Work’s adobe cabin on the east side of the river survived the flood since it appears in later tales. The Hop Work family lost all of their out-buildings and much of their stock. The flood also destroyed the apple orchard that Enoch had planted in present-day Cherokee Oaks in 1865; one of the first two orchards in Three Rivers. (Earl McKee said that Ira Blossom planted a pear orchard on the east side of the river a year or so before that.)

Before the flood, the flat floodplain of the South Fork was very fertile. After the flood, it was covered with sand, boulders, and wood, including huge logs of mangled sequoia trees. This is presumably the flood that deposited the big rocks along Cherokee Oaks Drive.

Afterwards, Hop Work and his wife left in discouragement. They moved down to the valley, eventually settling near Dunlap. Their son Enoch continued on with the ranching pursuits in Three Rivers.

Ira Blossom and his family stayed on, rebuilding their home on higher ground (500 feet east of Earl McKee’s present-day house), planted new trees and started again. They constructed their new home of poles, shakes, and clapboard, all made from the huge sequoia logs brought down by the flood.

Earl McKee said that in all the old photographs taken around Three Rivers, you see lines of tall sequoia picket fences, miles of them, bordering yards, pastures, roads, etc. Folks retrieved hundreds of sequoia trees from where the flood had deposited them. They cut the logs into rings and made shakes, or split 4- or 5-foot long sections for pickets, got a roll of wire, and made a picket fence. There are still some portions of sequoia picket fence over at the Barton’s.
From Three Rivers, the flood turned down-canyon toward Horse Creek. Hale Tharp and his family lived at the confluence of Horse Creek and the Kaweah River on the west side of Horse Creek a few hundred yards upstream from the river. At the time of the flood, Tharp's daughter, Fanny Ann Tharp, was 6 years old. In 1924, Mrs. Bernard Mehrten (the former Fanny Ann Tharp) gave an account of the flood to Frankie Welch:

It had rained steadily for nearly two weeks. Suddenly the swollen waters of the river began to rise and soon a widespread roaring torrent was sweeping past. Great logs and trees tossed like chaff upon its surface.

Higher and higher the mad waters climbed toward the Tharp's little cabin. Their Indian helper began carrying provisions up from the cellar and piling them on the table. Presently the roof of a house belonging to a neighbor four miles up the river floated by with an Indian clinging to it. Then Mrs. Tharp, staunch pioneer woman though she was, fainted. Mrs. Mehrten remembers climbing on to the clock shelf for the camphor which her father then used to revive her mother.

The floodwaters reached the doorstep of the cabin and then began to recede.

The Tharp's first thought was for the neighbors up the river. Loading a pack animal with food and bedding, they took a trail across the mountain and found the family in dire need of their aid.

The family's name was [presumably Enoch] Work, and their house had stood where the state highway now crosses the South Fork just below Three Rivers. Mrs. Work and the six children were saved by an Indian called Cherokee Nelson (or Nels), who waded waist deep through the swirling waters and carried them out. Mitch Works (the husband and father) [seems like this should have been Enoch Work] was in Visalia at the time and did not get back to Three Rivers for two weeks.

The children's grandparents [the Hop Work family] lived on the other side of the river, across the tree-lined roaring flood. There seemed to be no way to communicate to them that the family had been saved. But Cherokee Nelson, with Indian ingenuity, found a way. Taking the mother and children to a nearby treeless hilltop, he stood them in a row across its bald crest. Seven figures silhouetted against the skyline told the anxious watchers on the other side of the raging waters that all had been saved. [The nearest hill meeting this description is the one above present-day Crystal Drive.]

To this day, evidence of the big flood can be seen all the way from the high canyon in the South Fork, to where a landslide held back the gathering waters until they broke loose with a rush, to the fields about Visalia where great sequoia logs can be found buried in the sand. Fences made of these logs are still to be seen. Above Three Rivers are great sequoia logs lying on top of boulders twenty feet high. In other places the reverse is seen — huge boulders resting on fallen sequoias.

The streets of Visalia ran like rivers during the flood; the citizens went about in boats. Much livestock was lost; Mrs. Mehrten's father lost 80 head of goats — all the band but one little fellow who managed to ride a log to safety.

The flood moved rapidly through the steep canyons, but slowed dramatically when it emerged from the canyon and spread out onto the gentle slopes of the Kaweah Delta.

Harry H. Holley, the Kaweah Watermaster for many years, said that water takes about 6 hours to travel the 15 or so miles from McKay's Point to Visalia. Presumably the flood was traveling at roughly the same speed, about 2.5 mph, as it moved across the delta.

Traveling at that speed, the flood would have arrived in the small town of Visalia (42 or so miles downstream from the landslide) toward the end of the day on December 22, roughly 12–18 hours after the landslide dam had failed.

Many sources attribute the widespread flooding in the Tulare Lake Basin to the failure of the South Fork of the Kaweah landslide dam. That represents a misunderstanding of this event. The landslide dam failure, while dramatic, did not cause the flooding. All that it did was to impound the flow on one tributary for 25 hours and then release that water in a surge. It didn’t contribute any additional water to the flooding that was already occurring throughout the basin.
There is no credible record that anybody in Visalia even took note of the increased volume of water; it probably wasn’t all that much. (It would have been spread out many miles wide by then, so its height would have been greatly reduced.) But they definitely noticed the huge increase that occurred when the peak of the flood occurred on the following evening, December 23.

No doubt the mainstem of the Kaweah at peak flood (the biggest flood in that river’s recorded history) had the power to pick up and carry many sequoia logs out onto the delta, far more than had been moved on December 22. It’s easy to see how it became folklore that the flood of December 23–25 was caused by the spectacular landslide dam failure on the South Fork of the Kaweah. However, that flood was really caused by the same events that caused the flooding occurring on the other rivers in the Tulare Lake Basin. It was just another of our rain-on-snow events.

When the floodwaters subsided, a huge number of logs were left scattered widely about the Kaweah Delta. One big sequoia log came to rest right beside the grist mill at Main and Santa Fe. The trees lasted for years and appeared in numerous pioneer tales. Some accounts indicate that some of the trees were sawed for lumber. One portable sawmill was ordered specifically to see if it were feasible to mill the logs that had been left in the upper part of the Packwood Creek swamp. However, others recalled that so much sand and rock was imbedded into the trunks that the trees could not be sawed for lumber. When the 1874 No Fence Law made stockmen liable for the damage caused by their trespass cattle, many of the sequoia logs were split and used to fence the open range.

Landslide Dam Failure #2: San Joaquin River

The story of this dramatic landslide dam failure is known from two sources:
- Lilbourne Alisp Winchell’s 1920 History of Fresno County. This document has proved quite a challenge to track down. Fortunately, its account of the event was reprinted in Gene Rose’s San Joaquin: A River Betrayed.
- Floyd Otter’s The Men of Mammoth Forest.

The town of Millerton was created near Fort Miller, on the banks of the San Joaquin River. It was the original county seat of Fresno County, formed in 1856.

The fall rain came early and became more frequent by November, when the rains turned to snow in the Sierra. By December, record amounts of snow had been observed in the mountains. Then the temperature moderated and warm rains began sending the San Joaquin surging. In breathless prose, Winchell described the flood from the viewpoint of the residents of Millerton.

The San Joaquin at Millerton steadily grew in volume and height. Day after day the rains came. Anxiously the people awaited abatement of the storms. Each hour the angry stream reached higher and higher. The occupants of the buildings along the lower street began moving their most valued possessions, yet hoping for relief from the merciless encroachment.

Nightfall came — black under the overcast skies. It was Christmas Eve; but there were no devotional offerings. The harassed people were beyond joyous expression; though, from the women, there may have been silent prayers for mercy. There was universal vigilance and excited effort, and concern for community safety. Despairingly, as the black night measured the hours, they watched the unceasing advance of the surging torrent. Lanterns gleamed through the street; lights shone in all the upper houses; and the rain fell, and splashed in sheets in the frowning earth!

At eleven o’clock that night — Christmas Eve — the river was higher than the white men had ever seen it. Suddenly, crashing, roaring, sounds came to the ears of the wakeful villagers. Rushing with appalling speed and force a high wall of water, bearing on its surface an overwhelming tangle of broken and twisted trees from the forests of the high mountains. The whole blossom of this avalanche flood was thickly covered with the smashing, grinding, tearing logs — trunks, tops, roots were whirled along with the destroying speed of a tornado. Greater than the combined blows of all the batter rams and catapults of old, the massed projectile struck the town. Nothing in its tracks resisted it. In a few moments the awful work was done. Millerton was wrecked.

Floyd Miller’s account said that there were multiple slides on December 24, temporarily damming the San Joaquin. It was the failure of that landslide dam or dams that made the event so catastrophic, resulting in the destruction of Millerton.
The San Joaquin River passes through the narrow, granite San Joaquin River Gorge above Millerton. The San Joaquin is a huge river, and it was at flood stage. However, the large quantity of debris that slid off the mountains combined with the narrowness of the gorges allowed the landslide dam(s) to block the river. Eventually the river overtopped and breached the dam, sending a tidal wave of water and debris down the canyon.

At Millerton, Jones’s trans-river ferry was swept all the way to Sycamore Point. The same thing was repeated downstream at Hill’s Ferry, where the ferry had been destroyed during the 1861–62 flood. Debris from the 1867 siege damaged paddlewheel steamers plying the river. The steamers that were not damaged chugged around the inland sea plucking those residents lucky enough to have a second story home to which they could escape. Much of the port city of Stockton was inundated; floods had long been part of the Stockton scene. Boats were torn from their moorings and left as derelicts below. Stumps and debris from the Christmas catastrophe were seen as far away as Suisun Bay.871

Some Millerton residents rebuilt, some moved. However, as a result of the flood, the county seat was soon relocated to Fresno. The townsite of Millerton was inundated after Friant Dam was completed in 1942, forming Millerton Lake.

**Landslide Dam Failure #3: Mill Flat Creek**

There are several “Redwood Mountains.” The particular Redwood Mountain referred to in this story appears to have been the mountain that the Big Stump Grove is located on within present-day Giant Sequoia National Monument.

Forest Mill was a sawmill constructed on Mill Flat Creek, perhaps 5–10 miles downstream from present-day Sequoia Lake. This is the same location where Feggan’s Mill had been destroyed after a flood resulting from a January 1862 landslide dam failure. (It’s easy to confuse Mill Flat Creek with the similar-sounding Mill Creek. Mill Creek intersects the Kings River about two miles below present-day Pine Flat Dam. Mill Flat Creek intersects the Kings about two miles above the reservoir.)

In 1867, Forest Mill was owned by Jasper (Barley) Harrell and D.V. Robinson. It was powered by a 26-foot overshot water wheel. The following story of the debris slide and resulting flood was told in an 1881 issue of the *Visalia Weekly Delta*.872

’Twas on Christmas night 1867. It had rained all day, as it had done for about a week. The clouds were low; the day was dreary and lonesome; the night was one of those intensely dark, stormy nights that occasionally come in the pine forest, that one has to see in order to realize. Sometime in the fore part of the night, quite a tract of land with heavy timber, on the side of Redwood Mountain, slid into the creek, forming a dam which collected a large head of water, then giving way started down the creek crashing the timber before it. The first habitation it came to was an old log house inhabited by S.B. Corderoy. He heard it coming and caught his clothes and ran for life. The water just caught him as he reached high ground. The next place reached was (Michael) Hart’s. He heard the noise, thought it was a tornado, tore up the floor, and put his family under it and took his gun and stood in the door to await his doom, but the water did not reach his house. Next it came to the house of a man by the name of Root. He heard a great noise as of many waters, and jumped out of bed into it knee deep, where he stood fishing after his clothes and trying to convince his wife that it was better to lie a bed than to get out. Next it struck the Forest Mill, leaving it a complete wreck, and rose to the doorstep of the house where D.V. Robinson dwelt with his family, then passed on (to the Kings River) doing no more damage.

Floyd Otter said that this flood destroyed another sawmill in addition to the Forest Mill.873 This flood is apparently documented in detail in a 1906 *Lumbering in Tulare County* report prepared by H. Barton for the Fresno office of California’s Department of Forestry. However, CalFire has been unable to locate that file.

**Landslide Dam Failure #4: North Fork of the Kern**

While the South Fork of the Kaweah landslide dam flood was happening in Three Rivers, a similar drama was about to unfold on the North Fork of the Kern River. The location was about three miles downstream from the present-day Kern Ranger Station, within what is now Sequoia National Forest. The walls of the Kern Canyon at this point are about 3,000 feet high.
Three huge mass wasting events have occurred at this location:

1. An enormous landslide on the west side of the canyon that swept across the canyon. This occurred sometime well before 1800. This slide apparently created a landslide dam which caused a large lake to form, which was eventually breached by the Kern River. The narrow gorge of the Kern in this reach is a remnant of that landslide; the slide has forced the river against the east wall of the canyon. The pass where the trail goes in this reach might also be a remnant of the flood that resulted when the landslide dam was overtopped.

2. An enormous landslide on the east side of the canyon that swept across the canyon. This was immediately downstream from the western slide described above. This appears to have occurred between about 1800 and 1830. This slide apparently created a landslide dam, probably opposite Devil’s Staircase. This caused a large lake to form, which was eventually breached by the Kern River. The levee that forms Little Kern Lake is a remnant of this slide. The odd depositional formations in Grasshopper Flat below Devil’s Staircase are apparently a remnant of the flood that occurred when the landslide dam failed.

3. Two large debris flows that came out of small stream channels on the east side of the canyon and blocked the flooding Kern. These debris flows occurred opposite the toe of the western slide described above. They occurred on about December 28, 1867. They created one or two parallel dams which caused a large lake to form, which was breached by the Kern River after about a day. The dam that forms Kern Lake is a remnant of one or both of these debris flows.

Andrew C. Lawson visited the site of these three mass wasting events in 1903 as part of a geological study of the Upper Kern. He worked out most of the basic geology and the relative dating of these three events. Some additional information has come to light since.

In addition to Lawson, we know of three people who have visited the area of these mass wasting events and floods and recorded their observations:

- Steve Moffit, Sequoia National Park’s former trails supervisor, traveled this section of the Kern Canyon for over 20 years. During that time, he tried to puzzle out the strange depositional rock formations that he saw there.
- Josh Courter is the Western District Divide hydrologist on the Sequoia National Forest. He made a field trip through this portion of Kern Canyon in August 2012 and has studied aerial photography of it.
- Fletcher Linton is the forest botanist for the Sequoia National Forest. In addition to being a talented field botanist, Fletcher has skills as a geologist. He made a field trip to this area in the fall of 2013.

Lawson concluded from trees and sediment that the eastern slide occurred between about 1800 and 1830. It was younger than the western slide. Josh said that this slide still looks visibly newer, less consolidated, less weathered, and more sparsely vegetated than the slide on the west side where the trail goes.

Lawson determined that the levee that created Little Kern Lake (the lower of the two lakes) resulted from the eastern slide. When Josh and Fletcher looked at it, they came to the same conclusion.

Josh analyzed the two big slides. He concluded that the major landslide dam from the eastern slide likely formed opposite Devil’s Staircase, right in the center of that slide and just downstream from present-day Little Kern Lake.

The western and eastern slides were both massive. We have no record of just how high the landslide dams were that were formed by either of those slides. The height of those dams would have determined the maximum size of the lakes behind them.

If one of the dams had been, say, 1,000 feet high, the resulting lake would have extended 12–16 miles upstream to the vicinity of Kern Hot Spring. However, there doesn’t appear to have been enough material in either slide to make a dam that high.

The river at the point of the western slide is about 6240 feet elevation; that is the elevation of Kern Lake and the point where Rough Creek enters the river. The toe of the western slide (its height) at this point is currently 6620 feet elevation, 380 feet higher than the river. (For comparison, the highest point on the South Fork Kaweah landslide dam was 400 feet.) If the western slide had formed a 380-foot-high dam at this point, that would have created a lake that could have backed up 7–8 miles to about Rattlesnake Creek (6,600 feet elevation).
A 200-foot dam would have created a lake that could have reached to about Lower Funston Meadow (6,469 feet elevation). These calculations about lake size all assume that the dam maintained its structural integrity until the lake filled to the top of the dam.

We don't have any direct evidence about how the dams associated with the western and eastern slides failed. The dams may have failed slowly or all at once, but there was probably a lot of water that came through those dams when they failed. This could have been a highly erosive event.

When the dam associated with the western slide failed, the Kern River eroded its present channel along the eastern side of the canyon. There is a pass (6560 feet elevation) where the trail goes from Kern Lake to Little Kern Lake. That pass is about 60 feet lower than the high point on the toe of the slide (6620–6560 feet elevation). It is possible that this pass was eroded during the episode when the landslide dam was overtopped. It is easy to imagine that this pass served as a spillway into the small basin that forms present-day Little Kern Lake. That is based largely on speculation from aerial photography and walking through the pass; we have no direct evidence of this.

Just as we don't know how the dams associated with the western and eastern slides failed, we don't know how long they withheld their integrity. These were massive slides with a lot of rocky material. It is conceivable that the lakes behind these dams might have lasted for multiple years. If that were the case, this might have significantly affected conditions upstream. That would have resulted in an increase in sediments and tree mortality.

Some of the Jeffrey pines on the floor of Kern Canyon between Kern Lake and Lower Funston Meadow appear to be on the order of 300–400 years old. The root systems of those pines can tolerate complete inundation for only a couple weeks at most. If the lake formed by the western or eastern slides lasted longer than that, this would be reflected by the age of the trees that are present in the area of inundation.

Josh recalled thinking how much the Kern River appeared to change south of Lower Funston Meadow. Although it seems unlikely, it is worth thinking about whether the lake created by the western or eastern slide might have lasted long enough to have that kind of effect on the landscape.

Lawson talked to W.T. Grant who had visited the area in 1867 before Kern Lake existed, and then returned in 1868 after the lake appeared. Lawson concluded that Kern Lake was formed by two debris flows that came from two small streams on the east side of the canyon, an unnamed creek at the outlet of the lake and Rough Creek, 200 yards downstream of the outlet.

Josh analyzed Kern Lake. He said that the debris would likely have come down largely from Rough Creek, the lower of the two creeks identified by Lawson. This is a high-energy creek, and you can see areas in its upper watershed that could have contributed lots of material toward a big debris flow. Toward the lowest part of its channel, it has been scoured to bedrock. Nobody has investigated the upper portions of this stream’s watershed; those areas may have been well-scoured as well.

Today there is a small debris fan on the eastern side of the canyon at the outlet of Rough Creek. The debris cones from both of these creeks were much larger when Lawson visited this area in 1903. There has been a lot of erosion since then.

These two debris flows were sufficient to block the flooding Kern River for a day or so. When the dam broke, the river flowed over the two debris dams, then over the rock levee that formed Little Kern Lake. When Lawson visited in 1903, he said that the trees on the Little Kern Lake levee seemed to be at least 70 years old.

Dating of those trees and the ones farther down the canyon might provide some interesting information. All the trees in areas that were exposed to the full force of the floods (typically the valley floor and the lower canyon walls) should date from the early 1800s or later. However, there should be a point on the side of the canyon wall — corresponding to the high-water mark of the floods — where those young trees would abruptly meet the older trees that survived the various floods.

Our hypothesis is that the flood associated with failure of the eastern slide landslide dam in the early 1800s was much bigger than the flood associated with the failure of the debris flows in 1867. If that is true, it should be reflected in the age of the trees.
Fletcher observed that the Kern River has been excavated down to bedrock east of Little Kern Lake's levee. Larger angular boulders are in the river in this area as well. This is evidence of some huge flood, but whether it was the flood of 1867 or one of the earlier floods isn't clear. This area is above Devil's Staircase, where the landslide dam was probably formed by the eastern slide. Therefore, when that dam failed, there would have been a rush of water through this area. Additional scouring might have been caused by failure of the 1867 debris flow dam or the west side landslide dam.

Steve Moffit recalled that evidence of flooding in the Kern Canyon below Kern Lake looks fresh and near biblical in size. One of the most impressive sites is an area just north of the Grasshopper Flat campsite in Sequoia National Forest, a couple of miles below Little Kern Lake.

Grasshopper Flat is the area of the Kern Canyon immediately below Devil's Staircase, where the landslide dam was probably formed by the eastern slide. When that dam failed, the flood would have hit this area first.

Some guidebooks refer to Grasshopper Flat as being a floodplain. It is a stretch of relatively treeless land with long windrows of large, very uniform-size river boulders. In some places, these windrows are 12 or more feet tall and a hundred or more yards long. They are mostly parallel to the canyon and, therefore, to the trail. As you ride through this maze of river rock, it is very impressive as being big, very, very different, and covering a large area.

When Steve first encountered these strange windrows, he thought that perhaps they were some type of odd parallel moraine. However, after he learned of the Kern Lake flood story, it all fell together for him. If you stick a garden hose in a pile of dirt and gravel, it will tail out in windrows of uniform-size gravel according to the rate and volume of water flowing at that point and time. That appears to be the Grasshopper Flat floodplain on an enormous scale. Perhaps geologists have developed a different explanation, but that was Steve's deduction.

Based on what Josh observed when he came through Grasshopper Flat in 2012, he thinks the scenario presented above is likely what occurred.

Josh and Fletcher have analyzed the Grasshopper Flat, both in the field and using aerial photography. The flat has "stripes" or layers of deposition. You'll have a layer of fines then a layer of coarser and larger material, then back to finer materials. This process repeats itself, giving a "tiger stripe" look. Their hypothesis is that the gray deposition was remnants from the eastern landslide. It looks similar to when a culvert blows out after being blocked. You get these stripes of deposition or even "steps" as the water recedes from the area.

Josh and Fletcher are convinced that the striping through Grasshopper Flat originated prior to the 1867 flood. The material is very similar to the eastern slide. Even though we think of the 1867 as being very large, it apparently wasn't large enough to significantly rearrange the after-effects of the flood that occurred when the eastern landslide dam failed. That gives an indication of just how big that flood may have been.

We know more about the third flood because it occurred in historic times.

The winter of 1867 had exceptionally high precipitation that helped to mobilize the soil material on the steep watershed that the two small streams drained. When the soil became unstable, a mass of dirt, rock, and other debris broke loose and swept down into the river canyon. These debris flows occurred on about December 28, 1867. There were no witnesses to the event. The nearest settlers were some 30 miles down-canyon in the mining community that we now call Old Kernville.

The temporary lake that formed behind this huge landslide was estimated by one source to have had a depth of about 1,000 feet. The reliability of that estimate is highly suspect; it seems improbably large, like the stuff of urban legend. Failure of such a huge landslide dam would have presumably done even more damage to Kernville and Bakersfield than was experienced.

The debris flows plugged the river in the gorge formed by the toe of the western slide. Therefore, the maximum possible height of the dam would have been the height of the toe of that western slide above the river at this point: 380 feet (6620-6240 feet elevation). As described above, a 200-foot dam would have reached to about Lower Funston Meadow (elevation 6,469 feet). This was the flood-of-record on the Kern. Nobody has calculated how much water the Kern River, during its biggest flood in historic times, could store behind a dam in 24 or even 48 hours. We don't really know how long it would take to overtop a 200-foot or a 400-foot dam.
The Kern River would have been running at flood stage after the previous 40 or so days of rain and snow. One source suggested that it took the river two days to fill the temporary reservoir. Harvey Malone’s father-in-law told him that the riverbed downstream in Kernville was only dry for about 24 hours. In any case, the Kern River appears to have filled its new reservoir in less than two days and then overtopped the landslide dam. Because it was winter, there were no witnesses to this event. But evidently the dam failed catastrophically. Judging from when the flood reached Bakersfield, this failure occurred on about December 30.

One source said that the failure of the Kern Lake dam caused the Little Kern Lake levee to fail as well. Based on the vegetation he observed, Lawson apparently thought that the Little Kern Lake levee sustained little damage from the 1867 flood.

In any case, the floodwaters poured into Little Kern Lake and over its levee. The flood continued down-canyon about 30 miles to Kernville, where it destroyed many homes. According to Harvey Malone (whose father-in-law's family was living in the Kernville area at the time of the flood), the floodwaters backed up when they reached the head of the narrow section of the canyon (where Isabella Dam would later be built).

Water was said to have been about 50 feet deep in present-day Weldon near the Kern River Preserve. None of the accounts that were passed down speak of any fatalities in Kernville. Presumably that was because the residents realized that when the river went dry, a flood would soon be following. (That is how the story was passed down in Harvey’s family. Nobody has calculated how much water the Kern River, during its biggest flood in historic times, could store behind a dam in 24 or even 48 hours. If that dam failed catastrophically, would that have been sufficient to flood the town of Weldon to a depth of 50 feet?)

From Kernville, the floodwaters turned west and poured into Kern Canyon, down toward the village that would become Bakersfield. The massive wall of water, ice, uprooted trees, soil, and rocks scoured the canyon walls. By the time it reached the mouth of the canyon, the floodwaters were 200 feet high.

It had rained continuously in Bakersfield for many weeks, having stopped only a few days earlier. One source said that some of the residents were living in thatched tule houses. The better homes like that of Colonel Thomas Baker were constructed from adobe, but with roofs made from thatched tules. The tule thatch was coated on the outside with mud. All the houses had dirt floors. Six weeks or so of rain trickling through the dirt on the roofs had created a miserable, muddy mess.

By Christmas Day, 1867, the Kern River had risen near the top of the levees, but no flooding had occurred. Volunteers constantly patrolled the river banks protecting against any signs of a breech in the levee system. The river then gradually receded until near the year’s end, when the river flow supposedly stopped totally for two days. Apparently this was on or about December 30.

The river presumably didn’t stop entirely, just decreased greatly in volume. Even with the landslide dam on the North Fork of the Kern, there would still have been significant water coming down from the smaller South Fork of the Kern. Presumably some of the residents took alarm over this unexpected reduction in the river’s volume and realized its significance, just as they had in Kernville.

On New Year’s Day, 1868, residents were awakened to loud roaring sounds that were accompanied by the earth trembling. When the 200-foot-high flood came out of the canyon, it spread out and dropped in height; but it was still quite impressive. The trees, boulders, ice and brush in the forefront of the flood created a 50-foot-high logjam near where the Chester Avenue Bridge now exists. This towering logjam dammed the channel and diverted the river north around the village and onward to Buena Vista Lake. Although Bakersfield was flooded by a foot of water on what is present-day Chester Avenue, the village escaped the majority of the flooding. By some measures, this was the most spectacular disaster to occur in Bakersfield prior to the 1952 earthquake.

Immediately east of Bakersfield, the flood left an uprooted incense-cedar wedged into the boulders high upon the north wall at the canyon’s mouth, about 200 feet above the riverbed (photograph on file in the national parks). That tree was about 40 feet long and four feet in diameter and was badly scoured and scraped from its long trip. It was deposited about 20 feet above the end of the powerhouse tunnel that would be built 30 years later. Local fishermen and hikers climbed the cliff to inspect that cedar log for years until the powerhouse’s wooden flume was constructed in 1894. At that time, the cedar was cut up for use in that waterway.
Thanks to the logjam formed when the flood came out of the canyon, the river cut a new channel north of town. As a result, the town that would become Bakersfield then had to be relocated to be near that new channel. The present Kern River channel is the one that was cut during the 1867–68 flood.

After the flood, hundreds of thousands of logs were scattered over the south end of Kern County as far as Buena Vista and Tulare Lakes. The mesa where the Kern River Golf Course is now situated was covered with hundreds of large uprooted trees. Small logs covered an area half a mile square on Tom Barnes’s ranch (later known as the Canfield Ranch) east of Elk Hills. Barnes had been born in North Carolina in 1827 and moved to what is now Kern County in 1859.

Two sources said that Colonel Baker built a sawmill to mill the large number of trees that had been washed down by the flood. The availability of lumber apparently greatly improved the quality of housing in the area. Another source questioned that because the sand in those logs would have caused problems with the saw blades.

In any case, Tom Barnes is known to have built a log cabin on his ranch in 1868. The walls were made from logs left by the flood. The lumber for the roof was milled wood from the mountains south of the San Emidio Ranch (located at the base of the mountains south of Bakersfield). The milled wood was hauled through Fort Tejon to his ranch by ox team. Apparently Bakersfield did not yet have a sawmill. Barnes’s house is now an attraction at Kern County Museum’s Pioneer Village.

The 1867–68 flood is fairly well known; the floods from the failure of the western and eastern slides are not. There is a slight possibility that some of the above accounts of the 1867–68 flood are really referring to the failure of the eastern slide in the early 1800s.

1869 Flood

Relatively little is known about this flood. It was not a major flood in the Sacramento River Basin. The USACE identified it as a major flood in the San Joaquin River Basin. We have not found any record of how the flood affected the rivers within the Tulare Lake Basin in detail.

John Xantus was stationed at Fort Tejon. He told of rain for weeks in torrents in March of 1869.

Runoff during water year 1869 caused Tulare Lake to rise a very impressive 9.7 feet in elevation.

1869–71 Drought

This drought is not recognized at the state level. This is sometimes referred to as the 1870 drought, suggesting that was the worst year of the drought. John Vankat identified this drought as the 1869–71 drought, suggesting that it began in the summer of 1869 and continued through at least the fall of 1871. Tree-ring reconstruction shows that this drought was most severe on the upper San Joaquin at the inflow to Millerton Lake during 1871. Water year 1969 was not a drought year on the upper San Joaquin, but it is possible that the drought set in that summer in the Tulare Lake Basin.

In any case, this was a severe drought in the Tulare Lake Basin. Many cattle died. The losses to local cattlemen were so great that Tulare County supervisors asked the State Board of Equalization to make general reductions in the taxes of Tulare County. In August 1870, the county supervisors made dramatic reductions in the assessed valuations of many farmers as a result of the drought.

The very high precipitation of the winter of 1861–1862 was followed by drought of 1863–65. While hundreds of thousands of livestock died in the Central Valley of California, the search for new forage eventually led to the Coast Range and the Sierra, where the herders found sufficient forage for their stock. This first drought was somewhat alleviated by the winter of 1867–1868, but the drought of 1869–71 served to establish the Sierra as summer grazing range for sheep and cattle. However, since sheep gave a better financial return, they soon came to predominate over cattle.

1872 Flood

Relatively little is known about this flood. The USACE identified it as a major flood in the San Joaquin River Basin. Moderate flooding was reported near Sacramento in 1871–72, but it was not considered a major flood in the Sacramento River Basin.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Water year 1872 was a very large runoff year, delivering an estimated 2.6 million acre-feet of water to Tulare Lake. Tulare Lake was at essentially full pool (elevation 207 feet) when the flood started. The runoff raised the lake 5.3 feet to elevation 212.3 feet.

Despite the huge volume, we have found no record of flooding along any of the rivers in the Tulare Lake Basin. Perhaps the record of this old flood has just been lost. Or perhaps it has to do with how the flood is viewed. Maybe the high runoff did little damage to Visalia and the other settlements on the delta areas. But down in the lakebed, floods brought the volume of water necessary to sustain the lake through periods of drought. From that perspective, the flood wasn’t damaging at all, it was a boon.

1873–79 Drought

It isn’t quite clear how to view this seven-year drought. It was followed just three years later by the 1881–83 drought. It’s tempting to view this as an 11-year megadrought lasting from 1873–83. However, there are at least two reasons not to do this:
- There were two non-drought years separating the two droughts: 1880 and 1881.
- As shown in Table 33, average reconstructed runoff for the 11-year period (1873–83) on the upper San Joaquin was only 1% less than the 1113-year average (900–2012).

The drought was active somewhere in the state from 1873–79. In the Tulare Lake Basin, this drought is often referred to as beginning in 1877. Several sources said that conditions didn’t really start to improve until 1883.

Based on tree-ring reconstruction, the drought seems to have waxed and waned over the 1873–83 time period. In the upper San Joaquin River Basin at the inflow to Millerton Lake, the years 1874, 1876, 1878, 1880, and 1881 were non-drought years.

There are indications that the drought may have been more severe further south. In particular, 1874 and 1876 may have been drought years, at least in some areas. As shown in Table 34, runoff was relatively high in those years. Flooding occurred in every year of this drought except 1873. The high runoff in 1874 and 1876 might reflect flooding that occurred in a drought year. There just isn’t enough data to be sure.

Following is the fragmentary information we have about the status of the drought in the Tulare Lake Basin:
- A severe drought was reported for the Visalia area in 1873–75.
- A severe drought was reported for the Kern County area in 1876–77.
- David Campbell, an early Tulare County pioneer, recalled that 1877 was so dry that the grass did not even get started that year.
- Kenzie Whitten “Blackhorse” Jones lived on the west side of Fresno County. His valley farm failed during the terrible drought of 1877.
- Kathleen Small said that the drought of 1877 was devastating for Tulare County farmers.
- Morgan Blasingame of Millerton recalled the conditions on the San Joaquin: The drought of 1877 was the reason the cattlemen started going to the mountains. A few sheep and cattlemen had been going earlier, but in 1877 conditions were so desperate that you either moved your cattle to the mountains — or they died.
- Vanishing Landscapes refers to the protracted drought of 1877–79.
- David Campbell recalled that 1879 was so dry that the grass did not even get started that year.
The relative amount of precipitation in individual drought years is available from two sources: Visalia weather records and reconstructed runoff on the upper San Joaquin River Basin at the inflow to Millerton Lake. Table 33 provides that information.

Table 33. Relative precipitation during 1873–83 period.

<table>
<thead>
<tr>
<th>Water Year</th>
<th>Visalia % of average precipitation (1878–1972)</th>
<th>Upper San Joaquin River % of average flow (900–2012)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1873</td>
<td>95%</td>
<td></td>
</tr>
<tr>
<td>1874</td>
<td>105%</td>
<td></td>
</tr>
<tr>
<td>1875</td>
<td>90%</td>
<td></td>
</tr>
<tr>
<td>1876</td>
<td>159%</td>
<td></td>
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<tr>
<td>1877</td>
<td>61%</td>
<td></td>
</tr>
<tr>
<td>1878</td>
<td>104%</td>
<td>190%</td>
</tr>
<tr>
<td>1879</td>
<td>39%</td>
<td>69%</td>
</tr>
<tr>
<td>1880</td>
<td>127%</td>
<td>121%</td>
</tr>
<tr>
<td>1881</td>
<td>116%</td>
<td>140%</td>
</tr>
<tr>
<td>1882</td>
<td>67%</td>
<td>72%</td>
</tr>
<tr>
<td>1883</td>
<td>83%</td>
<td>56%</td>
</tr>
</tbody>
</table>

The last half of the 19th century had a number of very dry years. The 11-year period 1873–83 is an example of that. However, these dry years were interspersed with wet years. So overall, the period 1873–83, and the late 1800s in general, were periods of average runoff.

The Tulare Lakebed went through a period of steady decline from full-pool in 1878 to bone-dry in 1898, drying up completely for the first time in August 1898. The 1873–79 and 1882–83 droughts occurred during this period. However, the drying of the lake was not due to below-average runoff. Runoff on the upper San Joaquin during this 21-year period (1878–98) was 6% above the 1113-year average (900–2012). The drying of Tulare Lake was caused by the myriad canals that were constructed during the late 1800s, tapping the rivers in the foothills. This greatly diminished the water supply to the lake.

One source said that there were 18 diversions from the Kaweah River in 1875.

In the fall of 1875, John Muir traveled south from Yosemite, this time in search of giant sequoia groves. (Thanks to Bill Tweed’s research, we know this was the third of eight trips that Muir would take into what are now Sequoia and Kings Canyon National Parks.) Traveling alone except for a reluctant mule, Muir explored southward following the sequoia belt. It was during this trip that he stayed at Tharp’s Log (probably with James Wolverton) and later claimed to have named “the Giant Forest.” His trip took him south into the Tule River country, still following the sequoias. Referring to the Tule River sequoia forests after his hike through that area, Muir later wrote:

> All the basin was swept by swarms of hoofed locusts, the southern part over and over again, until not a leaf within reach was left on the wettest bogs, the outer edges of the thorniest chaparral beds, or even on the young conifers, which unless under the stress of dire famine, sheep never touch.

A Sierra Club Bulletin from this era recommended carrying lemons on wilderness trips to hide the taste of sheep piss in mountain streams and lakes, specifically in the East Lake/Reflection area of what is now Kings Canyon National Park.

Forest fires were widespread in 1875. In the summer of 1875, the Visalia Weekly Delta carried the news that “Heavy fires are raging in the mountains east of here, (giving the appearance at night) of immense lanterns suspended from the heavens.” Tony Caprio, the national parks' fire ecologist, commonly finds burn scars for 1875. The fires extended at least as far south as the Green Horns, presenting what was described as “a grand sight at night.” Floyd Otter attributed the majority of those fires to sheepherders.

During the period of this extended drought, there were three major floods (1874, 1876, and 1878) and up to five small to moderate size floods (1875, 1877, 1879, 1880, and 1881).
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Table 34 presents inflow to the Tulare Lakebed for each year of the 1873–83 period.

Table 34. Inflow to the Tulare Lakebed during the 1873–83 period.

<table>
<thead>
<tr>
<th>Water Year</th>
<th>Inflow to Tulare Lake (thousand acre-feet)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1873</td>
<td>400</td>
</tr>
<tr>
<td>1874</td>
<td>3,130</td>
</tr>
<tr>
<td>1875</td>
<td>115</td>
</tr>
<tr>
<td>1876</td>
<td>2,595</td>
</tr>
<tr>
<td>1877</td>
<td>315</td>
</tr>
<tr>
<td>1878</td>
<td>1,032</td>
</tr>
<tr>
<td>1879</td>
<td>58</td>
</tr>
<tr>
<td>1880</td>
<td>410</td>
</tr>
<tr>
<td>1881</td>
<td>120</td>
</tr>
<tr>
<td>1882</td>
<td>80</td>
</tr>
<tr>
<td>1883</td>
<td>110</td>
</tr>
</tbody>
</table>

Artesian Wells — Early Attempts
The first intentional attempt to dig an artesian well took place in the center of the intersection of Main and Court Streets in Visalia in 1859. That effort didn’t hit artesian water, but it produced very good water and became the town well. Since the water table in much of the Tulare Lake Basin was high and ordinary wells were inexpensive, no further attempt was made to find artesian water for many years.

In 1877 the Southern Pacific Railroad brought in a successful artesian well at a depth of 310 feet, two miles south of Tipton on the arid west side of Tulare County. The railroad created a 40-acre oasis that came to be known as the Tree Ranch. An artificial lake was created, stocked with carp, and a boat was provided for the enjoyment of travelers along the dusty west side. Thousands of trees were grown there as was nursery stock that was then planted elsewhere along the railroad right-of-way. It was quite a novelty.

Artesian Wells — Discovery of the Artesian Belt
Finally, in 1881, the intentional search for wells resumed. That attempt took place on the Paige & Morton Ranch, three miles west of Tulare. The Water Professor, A.P. Cromley, used a water witch to locate the well. Water “in grand abundance” was struck at 330 feet, just as the professor had prophesied. It produced a flow of about 800,000 gallons per day.

The Board of Supervisors visited the sparkling and wonderful fountain, and many people from the country round assembled, toasts were drunk and speeches made in its honor.

The success of that well generated much excitement, inducing many others to search for artesian water. Some were successful, some were not. By trial and error, an area that came to be called the Artesian Belt was identified within a few years. It covered over one million acres.

The water from those wells was quite warm, warm enough to ward off frost. Mrs. Henry McGee, a pioneer of Buzzards Roost (on the northeast shore of Tulare Lake), recalled that the warm water made the top of the pipe a favorite play area for children, especially for little girls. And it served as a grand steam laundry.

The wells continued to flow until they all stopped for uncertain causes during the first decade of the 20th century. The reasons proposed for that stoppage included:
- Some said that the water table suddenly dropped. This does seem to be the probable cause. See the section of this document that describes the Groundwater Overdraft for a discussion of why the water table was dropping in this area at this time.
- Draining of Tulare Lake.
- Underground rivers that suddenly ceased to be, or other less fanciful subterranean hydrologic changes. The 1906 San Francisco Earthquake was proposed as a possible causative factor of these unseen changes.

The 1892 Thompson Historic Atlas Map and a Los Tulares bulletin provide a much more detailed account of the fascinating history of artesian wells in the Tulare Lake Basin.\(^{899, 900}\)
1874 Flood
This flood occurred during the early stages of the 1873–79 drought. It is relatively common in the Tulare Lake Basin for floods to occur during multi-year droughts.

Water year 1874 was a very large runoff year, delivering an estimated 3.1 million acre-feet of water to Tulare Lake. The lake was just below full pool (elevation 206.5 feet) when the flood started. The runoff raised the lake 6.0 feet to elevation 212.5 feet. Tulare Lake has not been this high since.

As with the 1872 flood, we have found no record of flooding along any of the rivers in the Tulare Lake Basin. It could be that little damage was done to Visalia and the other settlements on the delta areas, or it could be that we just haven’t found the record of this old flood.

A map drawn in 1874 shows Tulare Lake as being nearly 700 square miles in size. This apparently reflected the condition when the lake was roughly 212 feet in elevation.

The 1874 flood in the Tulare Lake Basin is different from the huge storm that struck the San Francisco area on November 22–23, 1874. That storm dropped over 18 inches of rain on Fort Ross in 24 hours, with a recurrence interval of 32,000 years.

1875 Flood
This flood occurred during the 1873–79 drought.

Hydraulic mining contributed to the flooding and shoaling of rivers in the Sacramento Valley. Marysville, located at the confluence of the Yuba and Feather Rivers, experienced a number of devastating floods from 1852 through 1875. Debris from the Malakoff mine choked the Yuba until the river bottom was higher than the town.

According to the Yuba County history, the 1875 flood was the greatest and most destructive flood to hit Marysville, even bigger than the 1861–62 flood. As a result of the previous floods, the city had spent an immense sum of money to surround itself with a seven-mile-long ring levee. Trusting to the levee, the residents did not take the precautions that they had in the previous floods. This proved their undoing. When the flood came, it swept everything before it. Even goods that were placed upon platforms above the traditional high-water mark were lost as the river continued to rise.

For a week, heavy and incessant rain and snow storms prevailed, accompanied in some instances by thunder and lightning, an unusual phenomenon in the valley. Warm rain on a heavy snowpack is the typical recipe for valley floods. On the morning of January 19, the flood was threatening the city’s levee. Despite frantic efforts to raise the low spots, the levee was overtopped before the end of the day. Many houses were abandoned that evening as the residents sought safety in large houses, churches, the courthouse, etc.

By noon on January 20, the water was three to ten feet deep in the streets. A strong current ran down the F Street slough to the Yuba River. The whole valley, including the city, was one vast sheet of water on a level with the rivers.

Enormous damage was done to the residences and to even the most substantial buildings of the city. The railroads were badly damaged, and in the country there was a great deal of destruction of stock and farm property. John Muir provided an eloquent description of the storm in Flood-Storm in the Sierra in the June 1875 issue of the Overland Monthly.

The levee was later strengthened and Marysville has only been flooded three times since 1875. Another major rebuilding of the levees in the area was conducted from 2004–2008.

The 1875 flood was undoubtedly of minor proportions in the Tulare Lake Basin. Rain began on January 15. By January 20, Visalia had received three inches of rain and the foothills east of town had received about 10 inches. The streams in the area were running higher than they had at any time since the 1867–68 flood, but no real damage had occurred as of that date.

The USACE said that 1875 was a rain-flood in northwestern Tulare County. Possibly this was a rain-on-snow event at higher elevations. Visalia was partly flooded. The town incurred only minor to moderate flood damage.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

One source said that there was high water and minor flooding on the Kaweah during the spring runoff of 1875.908

Inflow to Tulare Lake in water year 1875 was relatively low, only 115 thousand acre-feet.

1876 Flood
This flood occurred during the 1873–79 drought.

The winter of 1876–77 was a strong El Niño event. That was presumably the cause of the 1876 flood since strong El Niño conditions typically result in an abundance of moisture in the Tulare Lake Basin. However, the 1876–77 El Niño affected the climate far beyond our area: it was a global event of major significance.

The extreme weather produced by this El Niño gave rise to the most deadly famines of the 19th century. Australia, New Zealand, and the South Pacific were in drought conditions during 1876–77. The drought was so severe in the Northern Hemisphere that up to 10 million died in India and 20 million died in China from malnutrition and drought-related diseases.909

Water year 1876 was a very large runoff year in the Tulare Lake Basin, delivering an estimated 2.6 million acre-feet of water to Tulare Lake. The lake was just below full pool (elevation 206.3 feet) when the flood started. The runoff raised the lake 5.4 feet to elevation 211.7 feet.

As with the 1872 and 1874 floods, we have found no record of flooding along any of the rivers in the Tulare Lake Basin. It could be that little damage was done, or it could be that we just haven’t found the record of this old flood.

1877 Flood
This flood occurred during the 1873–79 drought.

During the spring runoff, there was high water on the Kaweah and probably on other rivers within the Tulare Lake Basin. The levee built to protect Visalia gave way, causing flooding.

That levee was presumably on the south bank of the St. Johns River. This is the earliest record we have of that levee failing. It would not be the last. See the section of this document that describes the St. Johns Levee — Condition in Recent Years, for an account of some of its continuing problems.

1878 Flood
This flood occurred during the 1873–79 drought.

The USACE identified it as a major flood in both the Sacramento River and the San Joaquin River Basins.910

The 1877–78 winter brought heavy snows to the Mineral King Valley. The New England Tunnel and Smelter Company operated several mines in the valley including the White Chief Mine. The company was a very poor investment for its stockholders; it was always short on cash and on the verge of bankruptcy. By the end of 1876, the company had been locally renamed the "New England Thieving and Swindling Company.”

The company finally filed for bankruptcy in 1877. Court-appointed management, under the control of the creditors, attempted to continue operations through the winter of 1877–78. On February 18, 1878, a powerful avalanche destroyed the company’s main facilities at the White Chief Mine. The 30-by-60-foot bunkhouse, already straining under 20 feet of snow, was shattered. Miraculously, no one was killed, but operations were suspended until spring; and when spring came, the New England Tunnel and Smelting Company was dead.911

The Zalda Canal was constructed on the north side of the Kings River Delta in 1872. The 1878 and 1884 floods substantially enlarged that canal for about four miles, enabling connections to be made with other channels.912

One source said that it was the 1879 flood (not the 1878 flood) that enlarged the Zalda Canal, but that was apparently an error.

The 1878 flood filled Tulare Lake, bringing it to elevation 207.5 feet, causing it to spill through Summit Lake and into Fresno Slough and the San Joaquin River. “Eating” Smith chose that opportunity to bring the 32-foot-long
schooner *Water Witch* (formerly the *Alcatraz*) from San Francisco to Fresno Slough. From there he had it loaded onto wagon beds and hauled overland to Tulare Lake.\(^913,914\)

Tulare Lake has never filled again since the 1878 flood. The *Water Witch* appears to have been the last boat of any significant size to have made it to the lake. Since 1878, the Tulare Lake Basin has functioned largely as a closed basin, an inland sink without a regular outlet to the ocean.

**1879 Flood**
This flood occurred during the 1873–79 drought.

The 1879 flood was undoubtedly of minor proportions in the Tulare Lake Basin.

In 1886, the state engineer listed the high-water of 1879 on the Kings River near Kingsburg as only 6.7 feet above the low-water of 1878. This compared with a height of 17.3 feet in 1867–68 at the same place.\(^915\)

The USACE said that there was a rain-flood in northwestern Tulare County in 1879.\(^916\) Another source said that there was minor flooding on the Kaweah River in 1879.\(^917\) Visalia was partly flooded that year.\(^918,919\)

Inflow to Tulare Lake in water year 1879 was very low, only 58,000 acre-feet. That also suggests that the 1879 flood was of a relatively minor nature.

**1880 Flood**
Flooding in 1880 occurred in April.

The storm or storms and resulting floods appear to have affected much of the Central Valley.

Donner Summit set the U.S. record for the snowiest April ever with 298 inches (24.8 feet) of snow falling in one month.

A low-pressure area came ashore west of Red Bluff. The heaviest rainfall was located in a west-to-east band extending from Mt. St. Helena (near Napa) to Nevada City. Sacramento received 8.37 inches of rain during the two-day storm event of April 20–21. That was 5.79 standard deviations above the average with a recurrence interval of about 3,500 years.\(^920\) Flooding was reported near Sacramento and likely occurred elsewhere in the Central Valley.

April 1880 also brought heavy snows to the Mineral King Valley. Thomas Fowler had purchased the Empire Mine near the south end of the Mineral King Valley in September 1878. During the next twelve months, Fowler poured most of his personal fortune (and a lot of borrowed capital) into the Empire Mine. Tunnel work began on the mine, high above timberline on Empire Mountain, and continued throughout that winter. Soon construction began on a fifteen-stamp ore processing mill and a mile-long bucket tramway connecting the mine with the mill in the valley below.

The Mineral King Road, built with his encouragement, opened on August 20, 1879. Fowler’s expensive mining equipment rolled into Mineral King Valley soon thereafter. By September 1879, all the materials for the mill and tram were on site, and two months later both facilities were operable. Fowler had built more in one summer in Mineral King than everyone else together had done in the previous five years. In December 1879 Fowler closed the mill for the winter but attempted to continue work at the mine. The risk involved was not fully apparent until a huge avalanche destroyed much of the Empire works late on the evening of April 16, 1880.\(^921\)

Mary Trauger was the first white woman to stay in Mineral King through the winter. Harry and Mary Trauger were in the mining company’s long bunkhouse and office building on Empire Mountain when the avalanche wiped out that building. They both survived and were involved in the search for those who were buried under the snow and debris. This was the event in which Mary earned the title of “Mineral King’s Angel.” She reportedly located one buried man by hearing his watch ticking underneath the snow.\(^922\)

W.G. Pennebaker recalled that the Kaweah peaked on April 12, flooding Visalia and part of the surrounding countryside. It seems probable that this flood occurred at the same time as the big flood in Sacramento. So perhaps there was an editing typo at sometime, and the flood really peaked on April 21. This date would also be more in sync with when the big snows were occurring in Mineral King.
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

1881 Floods (2)

In the Sacramento River Basin, there were two floods:
1. one in mid-January
2. one that lasted from late January through early February

Outstanding flood peaks occurred in the upper part of the Sacramento River valley between January 14 and February 4. In the San Joaquin River and Tulare Lake Basins, the floods of 1881 were not of major proportions, although flood conditions were reported at some points in the San Joaquin River basin. There is evidence that the Sacramento River at Red Bluff and the Pit River reached peak discharges greater than in the 1861–62 flood. It is believed that the record-high stages on the lower Feather, Yuba, American, and Sacramento River reflected changed channel conditions, and that these stages were not the highest discharges of record.\(^{923}\)

Based in part on research by the USGS, the USACE identified the 1881 flood as being a major flood in both the Sacramento River and the San Joaquin River Basins.\(^{924}\) However, it was not a major flood in either the upper part of the San Joaquin River Basin or in the Tulare Lake Basin.

We have found no historical accounts of flooding on the rivers within the Tulare Lake Basin during 1881.

Inflow into Tulare Lake in water year 1881 was 120 thousand acre-feet, similar to water year 1875. That suggests that the flooding in 1881 might have been on the order of the relatively minor flooding that occurred in 1875.

1882–83 Drought

This two-year drought occurred just three years after the 1873–79 drought.

See Table 33 for a description of the relative amount of precipitation during the drought. See Table 34 for the inflow to the Tulare Lakebed during this drought.

Following is the fragmentary information we have about the status of the drought in the Tulare Lake Basin:
- On the legal side, lawsuits asserting riparian rights were fought out in Fresno courts in 1883 because of a dispute over Kings River irrigation water.
- On the not-so-legal side, a canal brush dam was blown up in 1883 because of a dispute over Kings River irrigation water.\(^{925}\)

1884 Floods (4)

There were at least four floods in 1884:
1. February
2. May (a cloudburst)
3. Snowmelt flood
4. December

1883–84 was reported to have been an unusually wet winter. There was heavy runoff the following year throughout the San Joaquin River Basin including the Tulare Lake Basin. The USACE identified it as a major flood in the San Joaquin River Basin.\(^{926}\) It was not a major flood in the Sacramento River Basin.

Floods threatened Fresno every winter because it was located in the sink of four creeks:
- Dry Creek ran just to the north of town.
- Red Banks and Dog Creek merged in the flat lands to the east.
- Fancher Creek ran nearby.

The center of Fresno was the confluent point for these four creeks.

On February 16, floodwaters covered every street in Fresno. All basements and ground floors were flooded. The only means of transportation within the city was by boat. The national parks’ files contain three photographs of this flood in Fresno:
1. This photograph depicts the situation on J street looking northwest from Fresno Street.
2. This photograph depicts the situation on M Street, south of Merced Street. The Fresno County Courthouse is seen in the background to the left. The church seen to the right is the Southern Methodist Church at the intersection of L and Fresno Streets.
3. This photograph depicts the situation on K Street looking north from Fresno Street. The Mill Ditch is visible in the foreground on Fresno Street. The steeple of the Methodist Church can be seen in the distance.

In 1884 there was a prolonged season of high water that caused much damage to the farms in the lowlands. There is no information available as to peak stages of the rivers during this period. The state records of streamflow for this season, listed in Water-Supply Paper 299, show very high snow runoff on the Kings, Kaweah, and Tule Rivers.

The Zalda Canal was constructed on the north side of the Kings River Delta in 1872. The 1878 and 1884 floods substantially enlarged that canal for about four miles, enabling connections to be made with other channels. After those floods, the Kings River began sending a portion of its floodwaters along the north side of its delta via this channel.

In the spring, the Kaweah flooded lowland farms on the delta. A brush and rock diversion weir had been constructed at McKay’s Point in 1877. The flood damaged that structure so extensively that it had to be reconstructed in 1884.

The USACE said that there was a snowmelt flood in northwestern Tulare County in 1884. A prolonged season of high water resulting from snowmelt caused much damage to farms in the lowlands.

A severe cloudburst occurred in in the hills above Yokohl Valley in May 1884. The valley was relatively well populated at the time. The Peter Stewart family lived about 10 miles up the Yokohl Valley from Merriman Station. In a few minutes, the storm turned the usually placid Yokohl Creek into the proverbial raging torrent. The flood swept away their small house, drowning Peter and his wife, his mother, their two children, and seriously injuring "Rat" Weisner. Some of the bodies were found as far away as Merriman.

Parts of the Central Valley experienced a major storm in December 1884. The highest rainfall measurements of this storm were for Bowman Dam in the Tahoe National Forest where 33.8 inches was reported in six consecutive days. This is 4.94 standard deviations above the average with a recurrence interval of 1,900 years. This storm delivered over half of the average annual rainfall at that station during those six days. This was a big and powerful storm, so it might have extended as far south as the Tulare Lake Basin. Panoche/Silver Creek west of Mendota flooded in December 1884. Other than this record, we haven’t found any account of flooding on area rivers.

Inflow to Tulare Lake in water year 1884 totaled 1.5 million acre-feet. This raised the level of the lake by 7.6 feet (from elevation 188.0 to 195.6).

**1885 Flood**

There were high floodwaters during the winter of 1885 along Cottonwood Creek on the route between Visalia and Badger. This was described by a Mrs. Lizzie in a talk that she gave at a picnic at Big Stump on August 27, 1950. Presumably there were high floodwaters in other streams and rivers in the Tulare Lake Basin that winter as well.

Inflow to Tulare Lake in water year 1885 totaled 483 thousand acre-feet. This raised the level of the lake by 5.6 feet (from elevation 188.0 to 193.6).

**1886 Flood**

Flooding occurred on January 25-27.

Relatively little is known about this flood. The USACE identified it as a major flood in the San Joaquin River Basin. It was not a major flood in the Sacramento River Basin. Based on fragmentary accounts, flooding seems to have occurred primarily in the northern part of the Tulare Lake Basin. Judging from later comparisons, the Kings River did not reach an extremely high stage. There was an extensive inundation in the city of Fresno from streams in that vicinity.

Runoff during water year 1886 caused Tulare Lake to rise 4.5 feet in elevation.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

1887–89 Drought
This three-year drought is recognized at the state level. Tree-ring reconstruction shows that the drought affected the upper San Joaquin at the inflow to Millerton Lake during the years 1887–89.937

In 1878, Tulare Lake had been at least 30 feet deep. The New York Times reported in its August 11, 1889, issue that the lake was less than three feet deep at the deepest part. (S.T. Harding would later estimate the minimum depth in 1889 to be 4.5 feet (183.5–179 feet).)938

The Kings River dried up in 1889 below the mouth of the Fresno irrigation system. Those two events were presumably due in part to the 1887–88 drought.

1888 Flood
This flood occurred either during or right at the end of the 1887–88 drought.

We know nothing about this flood except for two photographs from Fresno (on file in the national parks):
1. The flooded Union Ice Company warehouse located just south of the Southern Pacific Railroad tracks. The photographer was on G Street looking northeast toward Kern. The sign in the background says “Planing Mill,” probably Madary’s Planing Mill located on the corner of H Street and Kern.
2. Shows the damage along Inyo Street. The Arlington Hotel is seen on the right at the intersection of Inyo and J streets. In the distance is probably a freight warehouse of the Southern Pacific Railroad.

1889 Flood
Relatively little is known about this flood. Based on research by the USGS, the USACE identified it as a major flood in both the Sacramento River and the San Joaquin River Basins.939 While this may have affected the San Joaquin River Basin, it does not appear to have had much of a presence in the Tulare Lake Basin. It had no net effect on the elevation of Tulare Lake.

1889–90 Floods (2)
There were two flood events in 1890:
1. a big rain-flood from December 1889 through February 1890
2. a moderate snowmelt flood during May and June

The USACE identified the January event as a major flood in both the Sacramento River and the San Joaquin River Basins. Accounts of the May/June flooding episode are somewhat vague.

Statewide, the two wettest water years during historic times were 1890 and 1983. The heavy rainfalls of water year 1890 (the winter season of 1889–90) were confined to the northern half of the state.941 It was reported to have been an unusually wet winter throughout the Sierra.

The winter season of 1889–90 was remarkable for exceptionally heavy precipitation in the Central Valley Basin which produced floods of considerable magnitude from December 1889 – February 1890. Heavy snowfall in the Sierra resulted in unusually high runoff from melting snow during May and June, 1890. The floods were relatively greater in the San Joaquin River Basin, and they are considered to rank as the largest in that area for the period between the floods of 1867–68 and those of 1907.942

The Transcontinental Railroad had been completed in 1869. In January 1890, a relentless barrage of blizzards and a derailed train shut down Donner Pass for 15 days. In addition to an armada of snowplows and railroad crews, nearly 5,000 snow shovellers were hired to help clear the tracks, but the 66 feet of snow that fell on the pass that winter overwhelmed their efforts. This nearly stymied the attempt by journalist Nellie Bly to circumnavigate the globe in less time than novelist Jules Verne’s fictional voyage Around the World in 80 Days. However, Nellie was a very determined woman. With the help of the Central Pacific, she made it back to New York City on January 25, 1890, having traveled 72 days, 6 hours, and 11 minutes in her epic, planet-circling journey.

The winter of the 1890 water year was remarkable for the exceptionally heavy and widespread precipitation that produced floods of considerable magnitude throughout Northern California in January and February 1890 and moderate floods at other times from December 1889 through May 1890. The winter season of 1889 to 1890 featured an exceptionally heavy snowfall in the mountains, and the snow runoff period was one of the heaviest and longest of record. Lowlands in the lower Sacramento River Basin were flooded for many weeks. In
December 1889, the Sacramento River reached flood stages from Tehama to Sacramento. The peak stages on the river at Colusa and Sacramento were the highest yet observed. However, these high stages were primarily due to reclamation work along the river.

There were many breaks in the levees from Colusa downstream, and considerable damage was done to grain lands. A large break on the right bank levee of the Sacramento River below Sacramento helped to reduce subsequent flood stages. In January 1890, the tributaries of the Sacramento River were again at high stages. Stony and Putah creeks were reported to have been at the highest stages known to local residents. Considerable overflow from Cache Creek near Yolo flooded farms and caused washouts along the railroad. In February 1890, a flood occurred on the upper Sacramento River. The Sacramento River at Redding washed out part of a bridge.\textsuperscript{943}

The January flood resulted from a warm rain falling on this heavy snowpack. The San Joaquin River flooded, enveloping Stockton and other valley communities.\textsuperscript{944} Large floods occurred throughout the San Joaquin River Basin during the latter part of January. The upper San Joaquin River possibly reached an extremely high stage. The Merced, Stanislaus, Tuolumne, and Mokelumne rivers were at dangerously high stages, and some of the foothill tributaries of these rivers were reported to have been at the highest known stages to date. Several towns were flooded and railroad and highway structures washed out. The maximum stage of the season, however, was reached, at least on the lower San Joaquin River, during the snow runoff period in May 1890.\textsuperscript{945}

The levee protecting Visalia had not received regular maintenance in a number of years. (This was presumably the south-bank levee on the St. Johns River.) As a result, that levee failed near the end of January, resulting in widespread flooding in Visalia that lasted for part of one day.\textsuperscript{946} The flood wiped out railroad tracks near the town. Photographs of Main and Court Streets indicate that floodwaters were running about a foot deep through town. Travel on Main Street was said to have been by boat.\textsuperscript{947}

There were floods of considerable proportions in the Kaweah and Kings River Basins in January 1890. Overflow from the Kaweah River caused damage in Visalia where it was reported that boats were used on Main Street. Railroad tracks were washed out in the vicinity of Visalia. About January 25, 1890, the Kings River reached a stage reported to have been the highest since 1867–68, although it may have been exceeded in 1914 and 1937 at foothill points.

The crest at the railroad bridge near Kingsburg was reported in a Fresno newspaper as 16 feet above low-water, exceeding by 2.5 feet any since construction of that bridge. This stage possibly is comparable with the peak stage of 1867–68 which was 17.3 feet above low-water of 1878 as mentioned above. The stage of 1890 was probably referring to the same gage datum used by the State Engineer at this railroad bridge from 1878–84, as described in Water-Supply Paper 299. However, the relative heights of the crests at this point do not necessarily indicate, even roughly, the relative heights of the crests in the foothill channel.\textsuperscript{948}

It seems probable that the Kings and Kern Rivers were flooding at the same time, but no accounts of that have yet been found.

The January 1890 storm extended to the east side of the Sierra. Bridges and railroad tracks were washed out, and Main Street in Bishop was described as a lake.\textsuperscript{949}

There was heavy runoff in 1890 on all the rivers within the Tulare Lake Basin. Presumably this happened during the April-July snowmelt period, but we haven’t seen any documentation as to the actual timing. Inflow to Tulare Lake in water year 1890 totaled 2.0 million acre-feet. This raised the level of the lake by 12.0 feet (from elevation 183.5 to 195.5).

Tulare Lake reached a particularly large extent on multiple occasions between 1850 and about 1877. It then entered a fairly steady period of decline, punctuated by floods. (see Figure 15). The 1890 flood would remain the largest lakebed inundation for at least 50 years. Because of the system of levees that were built in the lakebed, there may never have been another lakebed inundation greater than this. That is, there were much larger inundations in the lakebed before this date, just not after.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

1893 Flood
Flooding in 1893 occurred in February.

For reference, 1893 was the year that President Benjamin Harrison — responding to continuing lobbying by a group of Visalians — withdrew the area that would become Kings Canyon National Park from sale. The issue of how to manage that land was left unresolved, and it would not formally become a national park until 1940. However, the initial step at preservation had been taken.

The 1893 flood was not a major flood in the Sacramento River Basin. The USACE identified it as a major flood in the San Joaquin River Basin. The flood of 1893 was of brief duration and in general does not rank as a major one in the Tulare Lake Basin. The gage height on the Kings River at Kingsburg on February 10 was 11.5 feet, not an exceptionally high stage.

The Kaweah flooded on February 9. Bridges and roads were washed out or damaged in the Kaweah River Basin and it was reported that the Kaweah River was higher than it had been for 20 years. The USACE said that this was a great flood in Visalia and northwestern Tulare County. Only two rain-floods on the Kaweah in the 19th century — the 1861–62 and the December 1867 floods — are known to have been larger.

The (Visalia) Daily Morning Delta described the flood in Three Rivers:

The inhabitants of Three Rivers have been rained in, except those who during the late storm were flooded out. Beginning in early evening of February 7, the rainfall continued for 48 hours. The South Fork rose rapidly, and on the 9th at 4 p.m. had attained its greatest height.

The muddy water tumbled along with a constant roar, accompanied by the booming of the boulders as they were hurled upon each other; frequently being heard was the heavy thud of some large log, which had been caught up and carried along by the furious water and then, as if tired of its burden, would throw it headlong anyway and anywhere to be rid of it. Occasionally would be heard a quick succession of rumbling, cracking and crashing sounds, like the roll and scattering discharge of artillery, as a tall, heavy tree became uprooted and fell through branches of surrounding trees into the fast accumulating driftwood and debris. The scent of phosphorous ground out by the tumbling rocks was disagreeably noticeable.

By the morning of the 10th, the storm had ceased and the level of the river fell, revealing much damage. Foot bridges had been carried away and there were many impediments to travel over former trails. The new bridge at "Kirkpatrick's Crossing" was slightly damaged but was easily repaired.

Charlie Blossom, Alfred Curtis and Henry Alles were among those who were away from home during this time and were subsequently flooded out and unable to return home. They drove from Visalia in a buckboard about 10 a.m., not reaching Ira Blossom's barn until nightfall. They had to camp on the opposite side of the river until morning, when they found the foot log put there by Mr. Blossom. After tending to their horses, they arrived at the house in time for breakfast.

Fred Clough walked from Ira Blossom's to the Three Rivers post office starting on the morning of the 10th and did not get home until near night, having to climb over the mountainside on a new trail as there was no place to cross the river safely.

The Tule also flooded on February 9. Evidently the rainfall was especially heavy in the southern part of the Tulare Lake Basin, where the Tule River was said to have been as high as in 1867. Highway and railroad bridges were washed out, and parts of Porterville were flooded. The height of this flood on the Tule River is not known at a point that would be comparable with recent recorded peak stages.

The Kern River overran its banks on February 10 due to melting snow and heavy rains. It didn't quite get into downtown Bakersfield, but it came a block from 19th and Chester Ave. The Kern Valley Nursery was one of the areas damaged.

Troop B, Fourth Cavalry arrived in Three Rivers on June 20. The snow had been so heavy that the Mineral King Road was still blocked with snow. Many of the streams were running so high as to be impassable.
Despite the flood in February, 1893 was not a particularly heavy runoff year. The level of Tulare Lake increased only 2.9 feet that year.

**1894 Flood**
Flooding in 1894 occurred in February. It was a major flood in some portions of California. It may have affected the Tulare Lake Basin, but we haven’t seen any reference to it.

The cavalry reported that the winter season ended exceptionally early. As a result, sheep had entered General Grant National Park and the northern part of Sequoia National Park in considerable numbers. At least 300 of those sheep perished in a late snowstorm.

**1895 Flood**
Flooding in 1895 occurred in January. It was a major flood in some portions of California. It may have affected the Tulare Lake Basin, but we haven’t seen any reference to it.

Troop I, Fourth Cavalry arrived in Three Rivers on May 23. Many of the streams were running so high as to be impassable. On June 4, the runoff was still so high as to prevent the cavalry from going up the South Fork of the Kaweah to Hockett Meadow. It was July 10 before the Mineral King Road became passable.

A brush and rock diversion weir had been reconstructed at McKay’s Point on the Kaweah River after the 1884 flood. It had to be reconstructed again in 1897. Some flood, perhaps the 1895 flood, had evidently damaged it so extensively that repairs alone were insufficient.

Tulare Lake rose four feet in 1895, indicating that there was higher than average runoff.

**1898–1900 Drought**
This drought is recognized as affecting at least part of the state, and it certainly affected the Tulare Lake Basin. This drought is generally described as the 1897–1900 drought. But it didn’t set in in the Tulare Lake Basin until water year 1898. Total flow in water year 1897 for the four rivers in the Tulare Lake Basin (Kings, Kaweah, Tule, and Kern) was 117% of the 1894–2014 average. Based on runoff records, it makes more sense to think of this as a three-year drought, lasting from 1898–1900.

Calendar year 1898 is considered to have set the record for the driest year in the history of the state. Well, sort of. The period of record for California (and all the states) begins in 1895, when standardized national weather records began. So 1898 is the driest year since record-keeping began at the national level. California may well have experienced a drier year prior to 1895. For example, there were several very dry years in the period from 1829–1864.

In any case, the 1898 record would last until 2013. According to the National Climatic Data Center (NCDC), California had its driest calendar year on record in 2013 with 7.38 inches of precipitation. This was 2.42 inches below the previous record dry year of 1898 and 15.13 inches below average.

Fresno received 0.30 inch of rain on April 1, 1897, but then the rains stopped. It was a cruel April Fools’ joke. That would set a record as the earliest occurrence of the last measurable rain for the water year. In 1924, the *Visalia Morning Delta* referred to 1897–1898 as being a dry year in Tulare County. Floyd Otter said that valley ranchers drove their hogs to the high mountains in the drought years of 1897–99.

In 1878, Tulare Lake had filled (elevation 207.5 feet) and spilled through Summit Lake and into the Fresno Slough and the San Joaquin River. However, in 1883, just five years later, the lake had reached its lowest level in memory. With many minor fluctuations (and one big flood, the flood of 1890), the lake gradually dwindled in size over the next two decades.

The *New York Times* reported in August 1898 when Tulare Lake dried up completely. This was the first time that had happened in historic times. It had gone from full-pool to bone-dry in just 20 years (1878–98). See the section of this document on the Chronology of Tulare Lake for the various causes that contributed to this drying.

The lakebed would remain dry through 1900.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

1898 Flood
Flooding in 1898 occurred in September. This flood occurred during the 1898–1900 drought.

The storm occurred on September 26 and was centered on the town of Tulare. It was apparently an isolated low-elevation event. The atmospheric mechanism behind the storm is unknown. It could have been a thunderstorm embedded in a tropical storm remnant.

Tulare received 3.89 inches of rain during the storm event. That was 5.66 standard deviations above the average with a recurrence interval of 2,700 years. Several other stations in the vicinity had over 3 inches of rain on September 26. It was the wettest day ever at Dinuba.\footnote{961}

This storm almost surely resulted in localized flooding, but we haven’t seen any accounts to that effect.

1901 Flood
Flooding in 1901 occurred in January.

We only know about this flood from the gaging station on the Kings. We haven’t found any anecdotal reports of it.

The Kings River at Piedra peaked on January 14: 33,200 cfs. This was the largest flow on that river since record-keeping began in 1895. This would remain the flood-of-record until the 1909 flood and the even larger 1914 flood.\footnote{962}

Flow for water year 1901 was 184\% of the 1894–2014 average for the Kings, 172\% for the Kaweah, 127\% for the Tule, and 121\% for the Kern. This suggests that the storm event that caused the flood was located to the north of the Tulare Lake Basin.

The Tulare Lakebed had dried completely for the first time in August 1898 and remained dry through 1900. The 1901 flood brought the lake back to life, if only modestly. Inflow to Tulare Lake in water year 1901 totaled 408 thousand acre-feet. That left the lake six feet deep (elevation 185.5 feet) at its deepest point.

1906 Floods (5)
Depending on how you count them, there were at least five floods in 1906:
1. January (rain-flood)
2. March (2 rain-floods)
3. May (snowmelt flood)
4. June (snowmelt flood)

The January to March period might best be viewed as a more or less continuous series of flood events rather than as individual floods. See Table 28 for a comparison of the March flood with later peak stages in the Kaweah River Basin.\footnote{963} Likewise, the May to June period might also best be viewed as an extended snowmelt runoff flood with multiple peaks.

The winter of 1905–06 was a strong El Niño event. As shown in Table 35, precipitation during water year 1906 was heavier than average throughout the valley.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Total Precipitation (inches of rain)</th>
<th>Percent of Mean</th>
</tr>
</thead>
<tbody>
<tr>
<td>Fresno</td>
<td>13.54</td>
<td>140%</td>
</tr>
<tr>
<td>Visalia</td>
<td>13.85</td>
<td>143%</td>
</tr>
<tr>
<td>Hanford</td>
<td>11.65</td>
<td>142%</td>
</tr>
<tr>
<td>Tulare</td>
<td>14.78</td>
<td>178%</td>
</tr>
<tr>
<td>Porterville</td>
<td>17.82</td>
<td>170%</td>
</tr>
<tr>
<td>Bakersfield</td>
<td>8.72</td>
<td>148%</td>
</tr>
<tr>
<td>Kernville</td>
<td>16.26</td>
<td>158%</td>
</tr>
</tbody>
</table>

A massive snowpack accumulated in both the San Joaquin River and Tulare Lake Basins during the winter of 1905–06. There were only a limited number of weather stations in the snow country at that time to measure the
snow directly. The California Cooperative Snow Surveys wouldn’t begin until 1930. The winter of 1905–06 appears to be the snowpack of record in the Southern Sierra. After the heavy winter of 1951–52, the USGS reexamined available records of snowfall at stations in the San Joaquin River and Tulare Lake Basins from the winter of 1905–06.\textsuperscript{964} See the section of this document that describes the \textit{1952 flood} for more detail about the results of that study.

Based on the limited data available, the USGS study concluded that the 1952 snowpack appeared to equal or exceed the snowpack that caused the great snowmelt floods of 1906. However, the only way to be certain was to wait until the snowpack melted and ran off. The results turned out to be quite clear, and rather surprising. In no case where the period of record included the year 1906, did the 1952 snowmelt runoff exceed that of 1906 on any river. That confirmed that all those watersheds had a bigger snowpack in 1906 than in 1952.

The difference was particularly remarkable in the Kaweah River Basin. The winter of 1951–52 set a modern-day snowfall record at Lodgepole, one that would last until the winter of 2010–11. However, as shown in Table 52 on page 283, the snowmelt runoff in the Kaweah River Basin in 1906 was 38% greater than in 1952.

Table 36 gives the total runoff in 1906 for each of the major rivers in the Tulare Lake Basin.

<table>
<thead>
<tr>
<th>River</th>
<th>Total Runoff April 1 – July 31 (thousand acre-feet)</th>
<th>Maximum daily flow</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings River</td>
<td>2,980</td>
<td>June 20, 1906</td>
</tr>
<tr>
<td></td>
<td></td>
<td>24,900</td>
</tr>
<tr>
<td>Kaweah River</td>
<td>814</td>
<td>May 28, 1906</td>
</tr>
<tr>
<td></td>
<td></td>
<td>7,260</td>
</tr>
<tr>
<td>Tule River</td>
<td>192</td>
<td></td>
</tr>
<tr>
<td>Kern River</td>
<td>1,390</td>
<td>June 21, 1906</td>
</tr>
<tr>
<td></td>
<td></td>
<td>9,500</td>
</tr>
</tbody>
</table>

The first two big winter storms of 1906 that we know of occurred on January 12–13 and February 9. Heavy rain fell on Friday and Saturday, January 12 and 13. The rain was reported to be exceedingly heavy in the foothills and mountains east of Visalia.\textsuperscript{965} Table 37 provides the total precipitation for the January 12–13 storm event for selected reporting stations.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Visalia</td>
<td>2</td>
</tr>
<tr>
<td>Eshom Valley</td>
<td>7</td>
</tr>
<tr>
<td>Kaweah #1 Powerhouse in Three Rivers</td>
<td>8</td>
</tr>
</tbody>
</table>

According to official NWS records, it initially looked like the February 9 storm was a storm of near-epic proportions in Bakersfield, dropping a total of 4.1 inches on the city that day. However, the National Climatic Data Center has since questioned the 1906 records for that city. Several transcription errors have been noted, including the use of non-standard annotations. For example, the entry for February 9 was written as 4-10 instead of 4/10. Therefore, it appears that a monster storm didn’t occur in Bakersfield on February 9. That explains why there wasn’t a major flood reported in that month.

Bakersfield still experienced a storm on February 9, and it was the biggest storm of the month. However, total precipitation in the city that day was 0.4 inches. It’s reasonable to believe that the storm dropped a lot of snow up in the conifer zone, but we have little data for making an estimate of the amount.

Likewise, it originally looked like Bakersfield set a record with an outstanding total of 8.70 inches of precipitation during the month of February. However, that record turned out to be due to the non-standard annotations. The corrected total was a much more modest 0.70 inch, 57% of which had fallen in the big February 9 storm.

There was no weather station in the national parks in 1906. The Fourth Cavalry detachment wouldn’t arrive in the park until June. However, there were four civilian rangers employed year-round. Walter Fry was the chief of those rangers. One of the other rangers, Charlie Blossom, kept a dairy, and his diary sheds some light on weather conditions in the parks during 1906.

Charlie’s patrol district included the South Fork of the Kaweah. He noted many days of heavy rain from January into May. The El Niño conditions resulted in an unprecedented snowpack in the Southern Sierra. On June 26,
Charlie traveled from Hockett Meadow to where the South Fork Campground is today. That was a 24-mile day, 6 miles of which were over packed snow still 6 to 8 feet deep.

Those were far from normal trail conditions for the end of June. Something truly aberrant had happened. It was almost as if the Southern Sierra had been thrown back into the Little Ice Age.

That was just a hint of the phenomenal snowpack that had accumulated in the Southern Sierra that winter. There was no formal weather gaging station to make direct measurement of the national parks’ snowfall, but the parks do have measurements of the compacted snowpack.

Walter Fry reported that the winter of 1905–06 brought the heaviest snowfall in the sequoia groves of both Sequoia and General Grant National Parks ever recorded to that date. The snow was 29 feet (equivalent to roughly 348 inches) on the level in Giant Forest on February 25, 1906. Even by June 25, the snowpack in Giant Forest had only melted down to about 12 feet on the level.966

Conditions like that have not been seen since. Based on Fry’s report, the winter of 1905–06 far exceeded the record-setting winter of 2010–2011 or the winters of 1951–52, 1968–69 and 1982–83.

There is some precedent for this much snow falling in the Sierra. Donner Summit received 298 inches (24.8 feet) in April 1880.

Walter Fry is generally regarded as a very reliable source, but his report of 29 feet of snow in February is astounding. (Not to mention the 12 feet remaining on June 25.) It does seem generally consistent with the USGS finding that the Kaweah River Basin’s snowpack in 1905–06 was 38% bigger than the huge snowpack of 1951–52. And it fits with the big snowmelt floods that occurred down in the valley in 1906.

But it is still difficult for us to comprehend the sheer magnitude of the snowpack that accumulated in the winter of 1905–06. The depth of snow in the national parks was double any that has occurred since. The deepest the snow has ever been measured at Giant Forest (including snow surveys) was 151 inches on March 16, 1952.

From Fry’s report, we know that there was roughly 348 inches of snow in Giant Forest in the Southern Sierra. Usually there would be a greater — or at least a comparable — depth of snow in the Northern Sierra. As a comparison, the snow survey site at Lower Lassen Peak (which is easily, year after year, the deepest snow in California) has only exceeded 300 inches once (in March 1983). However, that was not the case in 1906; there was a complete disconnect between the Southern Sierra and the Northern Sierra that winter.

Thanks to the USGS study, we know that the record-setting snowpack that formed in the winter of 1905–06 affected most or all of the watersheds of the San Joaquin River and Tulare Lake Basins. However, it apparently did not affect the Northern Sierra, and it did not affect the area immediately west of Tahoe. Bob Meadows researched the available data at some of the other sites in the Sierra in 1906, but none of them showed highly elevated snow depths. At the two sites west of Tahoe (Blue Canyon and Emigrant Gap), the snowpack was 30 inches or less at the end of February 1906.

A phenomenal amount of snow had been delivered to the Southern Sierra by the end of February 1906. How to account for the 29 feet of snow in Giant Forest? Based on what we know, it appears that a substantial portion of it might be accounted for by just one huge storm. The January 12–13 storm brought “exceedingly heavy rain” to the foothills east of Visalia, including 7 inches of rain at Eshom Valley and nearly 8 inches in Three Rivers (see Table 37).

That storm would have delivered significantly greater levels of moisture to the conifer belt of General Grant and Sequoia National Parks. Based on a study conducted by USGS, Giant Forest averages about 2.4 times as much moisture from winter storms as Three Rivers, at least in November (see Table 48 on page 273).

If that ratio held true for this storm, then the 8 inches of rain that Three Rivers received might have resulted in 19 inches or so of moisture at Giant Forest, which would have been upwards of 19 feet of snowfall. (Snowfall to rain ratios vary from about 6:1 to 15:1 in the Southern Sierra.) Gary Sanger said that if the core of the storm moved over Giant Forest, snow production could have been greater than this. Otherwise, it would likely have been less, especially in a convective storm.
The snowpack would have peaked in April. Based on available information, the Sierra has never seen a greater snowpack in historical times. An interesting related detail is that the San Francisco Earthquake occurred on April 18. At what was very close to the moment of peak snowpack during this very exceptional year, the Sierra was shaken by a very powerful earthquake (magnitude 7.8). That event was centered north of San Francisco, but was strongly felt in the Tulare Lake Basin. The earthquake apparently triggered a very strong cycle of avalanches in our local mountains. Historical records document that many of the buildings remaining in the Mineral King Valley from the 1870s silver rush, including the Smith Hotel, were destroyed by those avalanches.

As incredible as the winter of 1905–06 was in the national parks, an even bigger surprise would be coming the following winter, some 200 miles to the north. The year after the national parks got its phenomenal snow dump, the winter of 1906–07 would bring 884 inches (74 feet) of snowfall to Tamarack, California (that is snowfall, not snowpack). That station is located at 8,000 feet elevation, about 20 miles northeast of Calaveras Big Trees State Park. The Tamarack station was only established at the beginning of the 1906–07 season. Think what it might have recorded if it had been established a year earlier.

The January 1906 flood was mild enough, just a nice way to break the drought. It resulted from the very heavy rain that fell in the foothills on January 12–13. The resulting floodwaters arrived suddenly, but didn't last very long. (The short duration of the flood suggests that precipitation at the mid- to upper elevations probably came as snow and didn't contribute to the flood.)

The Kings River at Piedra peaked on January 19: 24,000 cfs. That is a fairly modest flood by the standards of the Kings.

Virtually all the streams and rivers in the Kaweah and Tule River Basins flooded. Hundreds of acres were inundated. The heaviest damage occurred on the St. Johns River north of Visalia. At 8:00 a.m. on Sunday, following the two days of rain, the St. Johns was still a dry sandy streak. However, one hour later, the river was running bank-full. The rapidly growing flood soon proved too big to pass through the culvert under the Santa Fe track. The bed of the track gave way, creating a break of thirty feet. (Trains had to be detoured via Hanford until repairs could be made.) All of the residences east of the Santa Fe were surrounded by a small sea of water.

Visalia’s China Town was also flooded. China Town generally encompassed an area bounded by East Main on the south and Oak Street on the north, Santa Fe on the east and Bridge Street on the west. Newspaper accounts of the day said that the Celestials moved their possessions from their basements up above flood level.

("Celestial” was a term used to describe Chinese emigrants to the United States, Canada and Australia during the 19th century. The term was widely used in the popular mass media of the day. It was not a disparaging term. It was an attempt to translate a classical term in Chinese by which the emigrants referred to themselves in dealing with the non-Chinese. China was often called the “Celestial Kingdom.”)

The Tulare Irrigation District’s large flume across the St. Johns River near Venice Hill was swept away.

The Santa Fe track two miles south of Visalia was under water for a considerable length of time. The land near the power company's substation east of Visalia was under water. (This is presumably a reference to the Venida Substation located at what is today the intersection of Highway 65 and Highway 198. Todd Dofflemyer speculated that Yokohl Creek flowed in this area before it was later rerouted further north to its present location. Or possibly it was Outside Creek.)

Considerable damage was done to the Southern Pacific’s bridge across the Tule River south of Porterville. In general, this heavy rain was considered a great blessing for farmers and stockmen; it had been dry for way too long. However, if they had been praying for an end to the dry season, their wish was about to be granted.

Charles Tollerton, a Dinuba pioneer, recalled that it rained every day in February. (Anecdotal accounts such as this should generally be taken with a grain of salt. Memories can change with the passage of time. Bob Meadows checked the weather logs for the nearby town of Reedley and discovered that it actually only rained on eight days in February for a total of 2.28 inches.)

Sand Creek was a mile wide in the Orosi district for weeks. People in the northern section of the country couldn’t get to Visalia. A railroad engine bogged down in Monson (southwest of Orosi) for two or three days, and everything along Sand Creek and the Cottonwood Creek channel was a lake.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

There was a pair of rain-floods in March. The first flood was the larger of the two. The Kings River at Piedra peaked on March 15: 25,400 cfs. The Kaweah River near Three Rivers peaked on March 15: 8,520 cfs. The Tule River near Porterville peaked on March 16: 4,760 cfs. A less severe rain-flood occurred during the March 25–28 period.

A pair of snowmelt floods occurred in May and June. The combination of all these floods was considered a major flood event on the Kings, Kaweah, and Kern Rivers. In the Tulare Lake Basin, the flood of 1906 would be recalled as the big one for the next three decades (until the December 1937 flood).

Troop B, Fourth Cavalry, arrived in the village of Kaweah on June 3. The preceding winter had been exceptional in the amount of snow and rainfall, resulting in considerable damage to roads and trails. The Colony Mill Road was impassable due to landslides. Most of the national parks’ trails were impassable due to snow and fallen trees.

The channel of Mill Creek through Visalia had been deepened, straightened, and covered with planks in 1891. That had been more or less adequate to carry runoff until the 1906 flood. However, the Mill Creek channel overflowed four times that year:
1. March 16–20. Floodwaters reached a maximum depth of four feet in parts of the city.
3. May 29. Caused by snowmelt. Lasted one day and covered only a small portion of the city.
4. June 13. This flood followed unusually high flows from melting snow, resulting from an unusually hot spell. Pronounced cooling saved the city from an even greater flood.

The Kaweah’s peak natural flow occurred at McKay’s Point on March 15: 12,749 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 8,861 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 6 years for the Kaweah.

The June flood was apparently the most impressive of the four. Water flowed down both Center and Oak Streets in what the Visalia Times-Delta described as a river running past the Tulare County Courthouse (located on Court Street between Oak and Center). At the height of the June flood, the water was about one foot deep on Main Street.

Earl McKee, Jr. said that the 1906 flood inundated a large portion of Visalia including the area around the Tulare County Courthouse (located on Court Street between Oak and Center). He said that the floodwater came up to the knees of horses.

Eastman had just popularized the camera, so this was the best documented flood yet in Visalia (multiple photographs on file in the national parks). A steam engine of the Southern Pacific Railroad was photographed sitting in about a foot of water on Oak Street near Church Street. Boats became a commonplace means of transportation in the downtown area. One account of this flooding episode read:

In 1906, Visalia had a bad flood. My father and mother lived out at Venice Cove here and he and his brother built a boat. They went to put in the river and then right where the break was, up about where Ben Maddox is, they got out and went down Main and Court Street. And the only place they got stuck was about a block from Center Street. Somebody threw a horse line to them and pulled them over this low spot and they went on down Main Street in a boat.

As a result of the 1906 flooding, Visalia explored two options for diverting a portion of the floodwaters from Mill Creek into Packwood Creek. It isn’t apparent that either of those options was pursued. The main response to the flood was the construction of a new and bigger, half-mile-long, concrete aqueduct/conduit in Visalia in 1910.

South of Mill Creek are the drainages of Packwood and Cameron Creeks. Waukena was a small farming community on Cameron Creek (a distributary of Deep Creek) along the northeast shore of Tulare Lake. This area was inundated in the 1906 flood, killing some of the orchards.

According to Sophie Britten’s book Pioneers in Paradise, the original bridge across the mainstem of the Kaweah in Three Rivers was a trestle bridge that was finished September 10, 1897. For a while, it was known as the River Inn Bridge because the River Inn was located right where the bridge crossed the river. (The hotel was built in May 1910 and burned to the ground in September 1911.) According to Sophie’s book, that bridge survived until the December 1937 flood.
According to Jim Barton, the North Fork Bridge (aka the Three Rivers Bridge) was constructed across the mainstem of the Kaweah in Three Rivers after the 1906 flood. Presumably one of the 1906 floods (probably the June flood) destroyed whatever bridge had previously crossed that river. Jim said that the new bridge was a post and timber bridge anchored by four steel cylinders filled with concrete. It remained in use until washed away in the December 1937 flood.974

The Tule River is the next drainage south of Cameron Creek. The Tule caused serious flooding in Porterville during the 1906 flood. One woman recalled going to the Opera House to fetch the tables and chairs, only to find them floating near the ceiling.975

Orlando Barton was the superintendent of the Devil’s Den Oil Company and traveled by bicycle between Devil’s Den and Visalia. He recorded a vivid and detailed account of the flooding on the lower section of the Kern River. The floodwaters reached Sand Ridge on June 7, 1906. However, instead of passing through the natural gap in the ridge as expected, the floodwaters began pooling into Ton Taché, the southern part of Tulare Lake. Three weeks later, that lake was 7 miles wide and 15 miles long. Finally, on June 20, the floodwaters began to break through the gap and flow on toward what we now think of as Tulare Lake proper.976

As Tulare Lake filled up, Orlando continued his observations like a field naturalist:977

*Hay will float on water. Two stacks arrived at George Scherin’s ranch on the south shore of the lake last week. One of them has about 20 tons of barley hay in it. Neither of the stacks are much out of shape after their cruise from the north shore.*

Total flow for water year 1906 was 228% of the 1894–2014 average for the Kings, 260% for the Kaweah, 342% for the Tule, and 256% for the Kern. It was the third wettest year ever recorded for the Kings, Kaweah, and Tule, and the fourth ever for the Kern.

Tulare Lake was virtually dry when Hobart Whitley visited it in 1905. However, the high runoff of the 1906 flood brought the lake back. The flood left the lake about 12 feet deep at the deepest point (elevation 191 - 179 feet), submerging 300 square miles. (This compares to 790 square miles at its maximum in 1862 and 1868.) As a relative measure of the volume of the runoff, that was the biggest increase in the lake’s depth since the 1890 flood.

One authoritative source gives the combined runoff of the four rivers in the Tulare Lake Basin during water year 1906 as 7,360,000 acre-feet, the second-largest runoff of record. However, that is based on outdated information. The total runoff for the four rivers in water year 1906 was 7,195,240 acre-feet. (The runoff in water years 1969 and 1983 would be substantially larger: 8,379,585 and 8,746,222 acre-feet, respectively.) The total floodwater entering the Tulare Lakebed in water year 1906 was about 1,530,000 acre-feet. As detailed in Table 7, this inflow exceeded that of any year since that time.978

Many levees were constructed in the Tulare Lakebed between 1903 and 1905 when lake levels were low. Unfortunately, those levees were light and poorly constructed. As a result, they failed when the high flows of 1906 entered the lake. The failure of those levees resulted in large financial losses, as almost 175,000 acres of wheat and barley had been planted that year. Most of that land was flooded before the crops could be harvested.

Although the 1906 flood was a disaster for lakebed farmers, others saw a silver lining. A number of recreation pleasure craft boats were quickly built or purchased for use on Tulare Lake that summer. The boats had gasoline engines with one owner bragging that his “lightweight” engine would weigh only 110 pounds.979

The series of flooding events that occurred in the Tulare Lake Basin in 1906 was unrelated to the intense storm that occurred just to our north on December 11, 1906. That storm covered a narrow band extending in a northeasterly direction from Monterey to Ione in the Sierra foothills north of Sonora. On December 11, Forest Lake on the 17-Mile Drive in Pacific Grove recorded 6.07 inches during that storm event. This was 6.38 standard deviations above the average with a recurrence interval of 9,000 years.980
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

1907 Floods (2)
There were two floods in 1907:
1. March
2. July

This was the second wet winter in a row. On January 17, 1907, Yosemite Valley reported a 60-inch snowpack, the greatest ever measured on the ground.981

Tamarack, California (about 20 miles northeast of Calaveras Big Trees State Park), set a seasonal snowfall record for the Sierra, recording 884 inches (74 feet) of snow falling during the 1906–07 season.

The March 1907 event was a very destructive flood in the Sacramento River Basin. It was caused by a severe rain from March 16–20 followed by a period of comparatively high runoff. Stages were exceptionally high throughout that basin. On the Feather River at Oroville, the flood height was the greatest ever observed, although it was believed that the riverbed at that location had been raised since 1862 by deposition of mining debris. Flooding on the Sacramento River system was so extensive that this event has been cited as the inspiration for construction of a 1,100-mile system of levees and dams for flood control.

This flood was also significant in the San Joaquin River Basin. Only a moderate rise on the upper San Joaquin River was observed during this flood, but there were exceptionally high stages on the large tributaries in the lower part of the basin. From the Merced River to the Mokelumne River, stages peaked on March 19, and were followed by high stages for several days. The San Joaquin River downstream from Mendota was at or above flood stage from the middle to the end of March.

The California Water Plan calls out the 1907 flood as being one of the major floods in the San Joaquin River Basin. Presumably that was in reference to the March flood.

It is not clear to what extent the 1907 floods affected the Tulare Lake Basin. Most of what we know about the flood impacts comes from the valley lakebeds. Based on Tulare Lake elevation data, the March flood appears to have delivered far more water to the lake than the July flood did.

Total flow for water year 1907 was 162% of the 1894–2014 average for the Kings, 141% for the Kaweah, 151% for the Tule, and 147% for the Kern.

On July 3, the levee that constrained Buena Vista Lake failed. This event was written up in the New York Times. The resulting flood inundated 25,000–30,000 acres south and west of Bakersfield including the old bed of Kern Lake. It damaged twelve miles of the Sunset Railroad. The levee that failed was supposedly built in 1866–67. If so, it would presumably have been built by Colonel Baker.

Tulare Lake had come back to life in 1906 and it expanded even more in the 1907 floods. Total inflow to the lake in water year 1907 was 977,000 acre-feet, raising the elevation 6.7 feet. The lake probably reached its maximum elevation (193.1 feet) on June 21, 1907.982 It would be 31 years before the lake would reach an elevation higher than this.

In 1907, a massive levee was built around the four sides of Tulare Lake, containing it to a fraction of its full natural size. Ripley’s Believe It or Not featured the “Square Lake” in its syndicated cartoon. The lake was now harnessed, and the former lakebed was declared safe for growing orchards.

1909 Flood (3)
There were at least three floods in 1909:
1. January–February
2. October
3. December

The heavy rainfalls associated with the storm sequence of January 1–20, 1909 extended in a southwest to northeast direction from Fort Ross near San Francisco to Greenville in the Feather River Basin. Nine stations reported their highest-ever rainfall totals for 20 consecutive days. La Porte in the Feather River Basin had 57.41 inches during the 20-day storm event which was 5.38 standard deviations above the average. The associated recurrence interval is 12,000 years. The Sacramento River at Red Bluff responded to the heavy rainfall with a
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

Flood crest of 30.5 feet on February 3, 1909, one foot higher than the previous record of 29.5 feet on February 4, 1881.\textsuperscript{983}

The January flood was written up in the \textit{New York Times}. It was a major flood in both the Sacramento River and San Joaquin River Basins. The flood of 1909 is believed to have been as great as that of March 1907, at least in the Sacramento River Basin.\textsuperscript{984}

During January 1909, flooding occurred at several places in the Sacramento River Valley from Red Bluff to the mouth of the Sacramento River. The Sacramento River reached high stages at Red Bluff in January and continued to rise into the beginning of February. The Sacramento River at Red Bluff reached a peak stage that was the highest yet observed. The lower river at Sacramento reached the maximum stage of record in the middle of January, and exceptionally high stages were recorded on nearly all the main tributaries to the river. Flood conditions prevailed in the lower basin through the end of the month. However, damaging floods occurred again in the beginning of February. The floods of 1909 were the most disastrous of any for which there is an authentic account, although it is believed that the flood discharge from the Sacramento River Basin in 1862 was probably far greater than the discharge from the floods of 1907 or 1909.\textsuperscript{985}

It was a fairly large flood along the Kings, Kaweah and Tule Rivers and may have been a flood on the Kern as well. Levees failed at both Visalia and Porterville.

The Kings River at Piedra peaked on January 14: 32,800 cfs. This was about the same size as the January 1901 flow had been.\textsuperscript{986}

The Kaweah’s peak natural flow occurred at McKay’s Point on January 21: 12,227 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 9,578 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 7 years for the Kaweah.

The brush and rock diversion weir at McKay’s Point on the Kaweah River had been reconstructed in 1897. It was apparently severely damaged in the 1909 flood and required reconstruction yet again. This would be the fourth time in less than forty years that the structure had to be replaced. This time it was rebuilt using concrete instead of brush and rock. The replacement structure would last until the 1937 flood.

The levee protecting Visalia (presumably the one on the south bank of the St. Johns) broke on the afternoon of January 14, and the floodwaters swept into town. The northwestern part of the town was quickly submerged.

Porterville also flooded on January 14. Twenty-five families living in the lower part of town were rescued by citizens with rafts.

Ernest Clayton Northrop recalled the extensive flooding that occurred during the winter of 1908–09. At the time, he was living on Bear Creek, a tributary of the North Fork of the Tule River, downstream from present-day Mountain Home State Forest in the general vicinity of the SCICON school. He said that it rained for many days and nights, followed by extensive flooding. The flood washed down sequoia logs which his family later made into fenceposts. Looking out from a point near their farm, they saw Tulare Lake spreading over most of the valley; there was water as far as they could see.\textsuperscript{987}

Troop G, Fourteenth Cavalry, arrived in Three Rivers on May 7. They reported that there had been a great fall of snow during the preceding winter. This prevented them from reaching their outpost camps until June 15. Buck Canyon had snow for so much of the summer that they were never able to establish their usual outpost camp there. Because of the heavy snow, tourists didn’t begin arriving until about July 1.

There was a major flood on Garza Creek on the west side of Fresno County in 1909. That flood probably occurred on October 5. There was a cloudburst on the west side of Kings County on October 5.\textsuperscript{988} That cloudburst apparently contributed to the death of Kenzie Whitten “Blackhorse” Jones four days later on October 9 when he was trying to persuade a horse to cross Garza Creek.\textsuperscript{989}

We only know about the December flood from the gaging stations. We haven’t found any anecdotal reports of it.

The Kings River at Piedra peaked on December 9: 44,800 cfs. This was the largest flow on that river since record-keeping began in 1895. However, the 1914 flood would be larger.\textsuperscript{990}
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

The Kaweah’s peak natural flow occurred at McKay’s Point on December 9: 14,108 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 8,226 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 5 years for the Kaweah.

Total flow for water year 1909 was 165% of the 1894–2014 average for the Kings, 189% for the Kaweah, 283% for the Tule, and 245% for the Kern.

Tulare Lake had high water, bringing it even higher than after the 1906 and 1907 floods. Total inflows to the lake in water year 1909 were 1,175,000 acre-feet, raising the elevation 4.7 feet.

1911 Flood
Flooding in 1911 occurred in January.

The winter of 1910–11 was a moderate to strong La Niña event.

The extreme rainfalls of the storm of January 9–11, 1911 extended in a southwest to northeast line between Los Gatos and Galt. Los Gatos recorded 17.34 inches, resulting in a recurrence interval of 800 years. Thirteen stations reported their highest-ever six-day rainfalls during this storm. A total of 390 inches of snow (32.5 feet) fell in Tamarack, California during January 1911. That remains the greatest monthly snowfall record in North America.

This was one of the greatest floods of the 20th century in the lower San Joaquin Valley. During this flood, the upper San Joaquin River near Friant reached high stages at the end of January. The flood was higher downstream near Newman at the mouth of the Merced River. The peak stage of 1911 set a record. High stages were also reached on the Calaveras, Mokelumne, Stanislaus, Tuolumne, and Merced rivers. The floods on these tributaries combined to raise the San Joaquin River to a record-breaking stage. Reports estimated that 75,000 acres of land were flooded from the overflow of the San Joaquin, Mokelumne, and Calaveras rivers.

The Kaweah also flooded. The other rivers within the Tulare Lake Basin may have flooded as well.

At the least, it was a year with high runoff. Total inflows to Tulare Lake in water year 1911 were 724,000 acre-feet.

1912–13 Drought
Table 38 shows how the San Joaquin Valley Water Year Index categorized the drought years in that basin. It also shows total runoff for the four major rivers in our basin (Kings, Kaweah, Tule, and Kern) combined during this drought.

<table>
<thead>
<tr>
<th>Water Year</th>
<th>Water Year Classification</th>
<th>San Joaquin River Basin</th>
<th>Tulare Lake Basin</th>
</tr>
</thead>
<tbody>
<tr>
<td>1912</td>
<td>Below normal</td>
<td>1,660,670</td>
<td>56%</td>
</tr>
<tr>
<td>1913</td>
<td>Critically dry</td>
<td>1,557,010</td>
<td>53%</td>
</tr>
<tr>
<td>Drought average (1912–13)</td>
<td></td>
<td>1,608,840</td>
<td>55%</td>
</tr>
</tbody>
</table>

1901–11 had generally seen average to well above-average water flows for the rivers within the Tulare Lake Basin. In contrast, 1912–13 saw those rivers deliver only about 50% of average flows. While this may not have been a severe drought, it represented a significant change from previous years.

1913 Flood
The “Black Flood” happened in Coalinga on July 22–23 (multiple photographs on file in the national parks). This was caused by an intense rainstorm. Until fairly recently, it was standard practice to have an oil sump (a large hole in the ground for waste oil) associated with every oil pump in the San Joaquin Valley. As the 1913 flood swept through the oilfield adjacent to Coalinga, it accumulated the oil from these sumps. When the flood entered the town, it was black with oil and left a terrible mess in its wake. Afterwards, a levee was built along Monterey Avenue on the west side of town to divert future floods coming from the oilfield into town. Today, environmental regulations prevent open sumps in the oilfields.
This storm apparently covered most or all of the Tulare Lake Basin. Panoche/Silver Creek west of Mendota flooded sometime in 1913, probably in July.  

Fresno and Bakersfield each received 0.33 inches of rain on July 22, making that the wettest July day on record for both of those cities.  

**1914 Flood**

Flooding in 1914 occurred in January.

An exceptional feature of this flood was the high precipitation in the Sierra. A total precipitation of 22 inches was reported for the storm period at Hume Lake in the upper Kings River Basin. The floods were not of major proportions in the southern part of the Tulare Lake Basin.  

An intense rainstorm struck Fresno and Coalinga on January 25. It apparently covered much or most of the Tulare Lake Basin. Portions of Coalinga were flooded (photograph on file in the national parks).  

Panoche/Silver Creek west of Mendota flooded sometime in 1914, probably in January.  

The Kings River peaked near Piedra on January 25 at 59,700 cfs. By then, the bridge over the Kings near Reedley was awash (photograph on file in the national parks).  

This was the largest flow on that river since record-keeping began in 1895. A resident on the Kings River at Trimmer reported in 1914 that the flood of that year had exceeded by 3 feet any others that had occurred since 1867–68, when the Kings River at that point was 6 feet higher than in 1914.

The 1914 flood would remain the flood-of-record until the December 1937 flood. There was slightly more precipitation in the storm of January 1914 than in December 1937, and it fell on ground previously moistened. However, there was heavy snow at an elevation of about 6,000 feet at the beginning of the 1914 storm, and the snowline lowered about 2,000 feet in elevation during the storm, making the amount of precipitation available in the form of water approximately equal to that in 1937.

The 1914 flood caused major damage to county roads and bridges along the Kings River. The flood also did considerable damage to buildings and killed many animals.

The national park’s superintendent reported that the rainfall and snowfall of the preceding winter was greater than usual, resulting in heavy damage to roads and trails. The January 25 storm was presumably a major contributor to that damage.

The North Fork Kaweah peaked on January 25: 7,400 cfs. This was the largest flow on that river since record-keeping began in 1910. This would remain the flood-of-record until the 1937 flood.

The Kaweah River near Three Rivers peaked on January 25: 13,300 cfs. This was the largest flow on that river since record-keeping began in 1903.

The mainstem Kaweah’s peak natural flow occurred on January 25: 13,899 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 10,275 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 8 years for the Kaweah.

The USACE said that the 1914 flood was a significant rain-flood in northwestern Tulare County. It was larger than the 1906 flood or any other flood since the turn of the century.

Flow on the Kaweah for January was almost 10 times greater than for the previous month. The town of Lemon Cove suffered major damage, washing away a small resort and hotel. Thousands of acres of valley farms were flooded. There was widespread flooding between Visalia and Exeter, halting highway travel. (The first Model T began production in 1908. Some of this highway traffic presumably would still have been non-motorized in 1914.) Water was neck-deep in some parts of Exeter.

Visalia had constructed a new, half-mile-long concrete aqueduct/conduit in 1910 to carry Mill Creek underground through town. The 1914 flood provided the first major test of that structure. It worked as designed; Visalia reported no significant damage.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

The Tule River near Porterville peaked on January 25: 6,600 cfs. This was the largest flow on that river since record-keeping began in 1901.1005

The Kern River near Bakersfield peaked on January 26: 18,300 cfs.1006 This was the largest flow on that river since record-keeping began in 1893.1007 It would remain the flood-of-record until the 1937 flood.

Corcoran was incorporated in 1914, the year that the 1912–13 drought ended. It seemed to end rather dramatically; floodwaters from the Kings, Kaweah (via Cross Creek), and Tule all reached the lakebed in January. An article in the Hanford Journal said that farmers in the lakebed were “doomed.” However, an article in the Corcoran Journal said that assessment was premature; it was too soon to say how great the damage would be.1008 As it turned out, the flood was short-lived. A record barley and wheat harvest was brought in that summer, apparently having suffered relatively little from the flood.1009

The Owens Valley experienced two floods in January 1914.1010 The first event left water in Bishop streets several feet deep; people went about in boats. One week later, a more intense storm hit the area, but it caused the worst damage in Big Pine where roads were deeply eroded, water mains broke, and bridges were washed away.

1916 Flood
Flooding in 1916 occurred in January.

The winter of 1915–16 was a La Niña event.

1916 was a year of vigorous weather systems throughout California. There was a very rare snowstorm on the west side of the southern San Joaquin Valley at the end of December 1915. A man-made forest of 2,300 oil derricks occupied the west side of the San Joaquin Valley in 1916. Half of those derricks, which ranged in height from 70–130 feet, were destroyed in two big windstorms that occurred on January 17 and 27, 1916.1011

A huge snowstorm hit Bishop and Big Pine from January 14–18.1012 Newspapers in Inyo County reported it as the worst snowstorm in the valley since the 1873 storm. It dropped over four feet of very wet, heavy snow. When the storm ended, temperatures dropped to ten degrees below zero, the lowest since settlement began. Many buildings collapsed under the weight of the snow, and travel throughout the region came to a stop.

A flood similar to that of 1914 occurred on January 17, 1916, in the Tulare Lake Basin.1013 The January river flooding continued in the Tulare Lakebed for about four months thereafter.

The flooding resulted from two Pacific storms. The first storm lasted from January 14–20 and covered an area that extended at least from San Diego north to the Kern River Basin. The next storm struck on January 24. Although the second storm extended to the Canadian border, it may not have produced as much precipitation in the Tulare Lake Basin as the earlier storm did.

The first storm was unusual in covering such a large area and extending so far north. The entire water year was an anomaly in that the Kern River Basin received twice as much precipitation that year as did any of the other watersheds in the Tulare Lake Basin. The Kern River Basin presumably received the brunt of this mega-storm because it was south-facing and was the southern-most watershed in the Tulare Lake Basin.

The Kings River at Piedra peaked on January 17, 1916: 45,400 cfs.1014, 1015 Although this peak was lower than in 1914 (59,700 cfs), it was among the highest recorded at this station between the start of record-keeping in 1895 and the 1937 flood.1016

The 1916 flood significantly opened the Zalda Canal, and thereafter it became the main channel for the Kings River.1017 This reach is now known as the North Fork of the Kings or the Kings River North Channel.

In the 1916 flood, this channel is said to have discharged 60% of the Kings runoff into the San Joaquin River and thence to San Francisco Bay. These changes in the flow of the Kings River soon left the farmers in the Tulare Lakebed without sufficient water to irrigate the reclaimed grain land, forcing them to sink deep wells for their irrigation water.1018

The USACE said that the 1916 flood was a significant rain-flood in northwestern Tulare County. It was roughly equivalent to the 1914 flood.1019
The Kaweah River near Three Rivers peaked on January 17: 14,700 cfs. This was the largest flow on that river since record-keeping began in 1903. The Kaweah’s peak natural flow occurred at McKay’s Point on January 17: 15,362 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 10,540 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 8 years for the Kaweah. That is the same recurrence interval as for the 1914 flood.

The Tule River near Porterville peaked on January 17: 6,780 cfs. This was the largest flow on that river since record-keeping began in 1901. The Tule sent significant floodwaters into Tulare Lake.

The Kern River peaked near Kernville on January 17: 9,690 cfs. That was the highest flow since the gage was installed in 1912 and would remain the flood-of-record for several decades.

Total flow for water year 1916 was 177% of the 1894–2014 average for the Kings, 180% for the Kaweah, 254% for the Tule, and 343% for the Kern. This was the highest runoff ever recorded for the Kern.

In 1907, a massive levee had been built around four sides of Tulare Lake to constrain its growth in times of flood. The lake had been harnessed, and the lakebed declared safe for growing orchards. However, the 1916 flood would bring the lake back to life and put an end to those hopes, at least temporarily.

The Tulare Lakebed was dry on January 1, 1916. Inflows from the Kings, Kaweah, and Tule Rivers began in January. Inflow from the Kern River began in March. The Kern River Basin received a huge amount of precipitation during January, some of it falling as rain and some as snow. Apparently the initial runoff from those storms went into filling Buena Vista Lake. Subsequent melting of the snowpack during the March–June period was presumably responsible for the Kern’s contribution to Tulare Lake’s inflows in 1916.

The 1916 flood is worthy of note in that it was a very large flood with especially heavy inflows from the Kern River. As illustrated in Table 39, it was also the first flood in which measurements of inflow were sufficient to determine the contribution from each river basin.

<table>
<thead>
<tr>
<th>Stream</th>
<th>Total Lakebed Inflow (acre-feet)</th>
<th>Percent Contribution</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings River</td>
<td>186,100</td>
<td>18%</td>
</tr>
<tr>
<td>Kaweah River</td>
<td>212,500</td>
<td>20%</td>
</tr>
<tr>
<td>Tule River</td>
<td>96,000</td>
<td>9%</td>
</tr>
<tr>
<td>Kern River</td>
<td>547,100</td>
<td>53%</td>
</tr>
<tr>
<td><strong>Total</strong></td>
<td><strong>1,041,700</strong></td>
<td></td>
</tr>
</tbody>
</table>

This inflow left Tulare Lake about 11 feet deep at its deepest point (elevation 190 – 179 feet). It would be 19 years before the Kern River or Tulare Lake would experience another flood this big.

Paso Robles received 15 inches of rain during the January storm, as much rain as that area receives in an average year. That remains the wettest January ever in Paso Robles.

The 1916 flood was a major event throughout Southern California. We have to look there to understand what happened in the southern part of the Tulare Lake Basin. This storm story begins in the San Diego area in the summer of 1915.

For four years, San Diego had been experiencing below-normal flows in the local rivers. It was feared that the area was entering another prolonged drought. Memories of past droughts were still all too fresh; there seldom seemed to be enough water. To make matters worse, the fall rains didn’t start on schedule. A dry November sharply reduced the supply in the city’s reservoirs.

Charles Mallory Hatfield was a rainmaker — an expert in pluviculture. He was originally from Fort Scott, Kansas, but was by then living near Los Angeles. His supporters believed that Hatfield was the real McCoy.

Under pressure from the San Diego Wide Awake Improvement Club and others, the San Diego Common Council (the equivalent of today’s city council) approached Hatfield and voted 4–1 to pay him to fill the Morena Reservoir.
Reservoir for a $10,000 fee, payable by the inch. They figured that they had nothing to lose; it was a fee-for-
service agreement. The council instructed the city attorney to draw up a contract. By New Year’s Day, 1916,
Hatfield had set up his moisture accelerator near the Morena Dam and was releasing strange and secret vapors
into the atmosphere.

The first of two Pacific storms struck January 14–20. Cuyamaca (in eastern San Diego County, source of the
headwaters of the San Diego River) received 18 inches in that six-day storm, an amount equal to nearly half of
its average annual rainfall. Major flooding resulted downstream in the San Diego area. After the first storm,
Morena Reservoir still wasn’t full, so Hatfield continued producing his vapors; he intended to earn his fee. One
man caught up in the initial flooding in Mission Valley recommended: “Let’s pay Hatfield a $100,000 fee to quit.”

After just a four-day break, the second storm arrived on January 24. It added more than 14 inches of rainfall at
Cuyamaca. Nearby Descanso received a total of 27.79 inches of rain during January 14–28. This represents a
recurrence interval of about 6,400 years. A total of 25 stations recorded their highest-ever rainfalls during the
two storms. There were 8 stations that received 10 inches of rain or more on January 17. The total for the
month at Dorman’s Ranch in the San Bernardino Mountains was 57.91 inches (4.8 feet). The ground was
already saturated from the first storm and could hold no more. When the second storm hit on January 24,
raging torrents of water raced down the rivers and creeks.

The spillway of the Sweetwater Dam was designed with a capacity of 5,500 cfs. At the peak of the first flood on
January 17–18, it had a flow of 45,500 cfs. The dam was overtopped by 3.7 feet of water, causing a large
section of the south abutment dike to fail catastrophically.

On January 26, the city dynamited the dam in Switzer Canyon in the south portion of Balboa Park. It had been
cracked and weakened over the years; blowing the dam was apparently considered preferable to letting it fail
during the height of the flood. Two houses on 16th Street were overturned as the water rushed down to San
Diego Bay.

On the evening of January 27, Lower Otay Dam was also overtopped. Water filled the observation shafts on the
downstream side of that dam’s steel core, and the pressure blew out the rock that provided the dam’s structural
stability. The steel core then swung out like a gate, releasing the full depth of water, which created a flood wave
in the canyon of gigantic proportions. The dam was about 130 feet high, and the depth of the wave in the
canyon a short distance below the dam site was about 100 feet high. As the lower canyon widened, the wave
height decreased to approximately 20 feet, which was still devastating to the people living in the valley below. It
required only 2½ hours for 13 billion gallons to empty out of that reservoir.

The San Diego River reached a crest six feet higher than in any previous flood. The city’s concrete bridge at Old
Town was the first to go. Next was the Santa Fe Railway bridge, even though it was weighed down with loaded
freight cars. In Mission Valley, the San Diego River peaked at an estimated 75,000 cfs. (For comparison, the
more famous 1980 flood peaked in Mission Valley at just 25,000 cfs.) All but one of the large bridges in San
Diego County was destroyed. The Southern Pacific lines were severely damaged, and train service to Southern
California was discontinued. The only way to travel from San Diego to Los Angeles was by boat.

Hatfield talked to the press on February 4, saying that the damage was not his fault and that the city fathers
should have taken adequate precautions. He had fulfilled the requirements of his agreement with the city: the
Morena Reservoir was full.

The city attorney pointed out to Hatfield that although a contract had been drawn up, it had not been signed.
Hatfield had been working without a written contract; he had effectively entered into a gentleman’s agreement
with the Common Council. Hatfield countered that he was then the owner of the water that he had added to the
Morena Reservoir, which was valued at $400,000. It was a nice try, but the city still refused to pay.

Hatfield filed a lawsuit on December 2 to force the city to pay their bill. The city later offered to negotiate;
they’d pay him his $10,000 fee if Hatfield would accept liability for the $3.5 million in claims that had been filed
against the city as a result of the flood. Hatfield declined the offer. His lawsuit never came to trial and was
eventually dropped. The court did rule in related suits that the flood was an act of God and, therefore, the city
was not liable for the damages.

The San Diego Common Council never did pay Hatfield the fee that they had agreed to. But the flood had been
good for Hatfield. His fame continued to grow, and he received more contracts for rainmaking.
Total damage in San Diego County was nearly $8 million and about 15 people died.

Orange County received 11.5 inches of rain during the period January 17–28. The Santa Ana River overflowed, sending a wall of mud through farmland and streets. Almost all the bridges in the county were destroyed. The state highway and virtually all the roads in the foothill and mountain areas were washed out. Four people drowned in the county, two in a cottage floating down the Santa Ana River.1030

The Santa Ana River also flooded upstream in Riverside County as did the San Jacinto River. The cities of Indio, Coachella, and Mecca were completely inundated; 9 inches of rain fell in the Coachella Valley. Lake Elsinore rose very quickly. All rail traffic was halted in the county due to landslides or tracks washing out.1031

Los Angeles experienced flooding January 14–19 and 25–30. The Los Angeles River ran 3 miles wide.1032

San Bernardino County experienced flooding from January 17–28, 40 bridges were destroyed. All roads in Cajon Pass were washed out. The Santa Ana River ran 2 miles wide. It was one of the largest floods ever on the Mojave River. Two drowned in the county.1033

1918–34 Drought

This 17-year megadrought is generally recognized at the state level as having three components:

- 1917–21
- 1922–27
- 1929–34

This drought was statewide, although not all areas were affected equally. The 1918-34 drought occurred in a climatic context that included severe drought conditions over much of the Western U.S. The 1920s–30s were a period of overall dryness that rivaled similar extreme events in the paleoclimate record stretching back at least to the 800s.1034

There were modest breaks between each of the three components of this drought in the Tulare Lake Basin. However, the 2007 EPA report on the Tulare Lake Basin prepared by ECORP Consulting concluded that it was best to view this as a single 17-year-long drought.1035 Peter Vorster, a hydrogeographer with The Bay Institute and a primary author of the EPA report, also recommended taking that approach.

Much drier than average conditions began in the Tulare Lake Basin in 1918. The 1918–34 period averaged just 2,047,511 acre feet, the driest 17-year period in our basin since record-keeping began in 1894. So when this drought is viewed from the perspective of our basin, it was a 17-year megadrought that lasted from 1918–34. It was the first megadrought in our area in over three centuries, since the drought of 1566–1602.

The 1918–34 drought was of epic proportions in the southern San Joaquin Valley. Dave Meko used hydroclimate reconstructions from tree-rings to compare this drought against all the others that have occurred on the upper San Joaquin River Basin at the inflow to Millerton Lake during the 1113-year period 900–2012.1036 As shown in Table 19, average reconstructed runoff for the 17 years that this drought was active on the upper San Joaquin (1917–34) was 68% of the 1113-year average (900–2012).1037

Meko’s reconstruction showed that:

- Water years 1929–31 were the driest 3-years since 983. (However, the average annual combined runoff of the four major rivers in the Tulare Lake Basin for the 3-year period 2013–15 is projected to be 820,902, 34% less than the previous record set in 1929–31 (1,237,573).)
- Water years 1926–31 were the driest 6-years since before 900.
- Water years 1924–33 were the driest 10-years since before 900.

On the San Joaquin River Basin (SJQ4 gage, a summary series defined by CDEC as the total San Joaquin River runoff), water years 1451–1465 were the driest 15-years ever. However, Meko’s reconstruction showed that water years 1920–34 were the driest 15-years ever on the upper San Joaquin at the inflow to Millerton Lake. Likewise, water years 1446–1465 were driest-ever 20-year period on the total San Joaquin River Basin. But on the upper San Joaquin at the inflow to Millerton Lake, water years 1917–1936 were the driest 20-years ever.1038

These two comparisons demonstrate that the 1918–34 drought was more severe in the southern part of the San Joaquin River Basin than even the driest part of the megadroughts of the 900s, 1100s, or 1400s. The 1918–34...
drought was the driest 6-, 10-, 15-, and 20-year period in the southern San Joaquin since at least A.D. 900. These findings by Meko further support the argument that the 1918–34 drought should be viewed as one extended drought instead of three smaller droughts.

A 420-year reconstruction of Sacramento River runoff from tree-ring data was made for DWR in 1986 by the Laboratory for Tree ring Research at the University of Arizona. The tree-ring data suggested that the 1929–34 drought was the most severe in the 420-year reconstructed record from 1560–1980. This indicates that the 1929–34 drought has a possible recurrence interval of more than 400 years.

The 420-year reconstruction also suggested that few droughts prior to 1900 exceeded three years, and none lasted over six years, except for one period of less than average runoff from 1839–46. See the section of this document that describes the 1987–92 drought for a comparison of the multi-year droughts in the Sacramento River Basin.

Table 21 illustrates how severe different droughts were on the upper San Joaquin as measured over different time periods.

Tree-ring reconstruction shows that 1580 is the drought year of record in the Central Valley and the Southern Sierra. The year 2015 will almost certainly be the second-driest. The year 1924 was an extreme drought year. As shown in Table 21, it was the third-driest year on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012. On the upper San Joaquin at the inflow to Millerton Lake, the reconstructed flow for 1580 was only 36% of that of the reconstructed flow in 1924.

There is virtually a three-way tie among 1795, 1924, and 1977 as to which is the third-driest year in the Tulare Lake Basin for the 1115-year period 900–2014. Table 21 shows that flows on the upper San Joaquin River in 1795 (based on tree-ring reconstruction) were about 2% less than in 1924. However, as shown in Table 23, actual flows in the Tulare Lake Basin in 1977 (based on stream gages) were slightly less than in 1924. Based on stream gage data, we know that 1977 was a slightly drier year than 1924. However, we can’t say with any confidence where 1795 falls in this order, especially in the Tulare Lake Basin.

Table 40 shows how the San Joaquin Valley Water Year Index categorized the drought years in that basin. It also shows total runoff for the four major rivers in our basin (Kings, Kaweah, Tule, and Kern) combined during this drought.
First part of drought: 1918–21
The first part of the drought affected the entire state from 1917–21 except the Central Sierra and North Coast and has a possible recurrence interval of 10 to 40 years. It was most extreme in the north. As shown in Table 40, the first part of the drought was most active in the Tulare Lake Basin during the years 1918–19.

Tulare Lake became dry on April 30, 1919 and would remain dry for the next three years (1919–21). Repairs to the lakebed levees, which had been breached in the 1916 flood, could finally be made.

Sometimes a high pressure ridge will form in the northeastern Pacific Ocean, forcing the mid-latitude storm track well to the north of its typical position and preventing winter storms from reaching California. This ridging pattern has preceded some of the worst West Coast droughts.

When the storm track is diverted north of its typical position, it can deliver cold Arctic air to the Midwest and Eastern Seaboard. The Eastern Seaboard received particularly cold air during the winter of 2014–15 while California was experiencing a severe drought. That seems to be what occurred in the winter of 1917–18.

The East Coast experienced exceptionally cold temperatures from December 1917 through mid-February 1918. Three accounts illustrate that effect:

1. The NWS Forecast Office in Louisville, Kentucky described what conditions were like in the Ohio River Valley.\(^{1041}\) The winter of 1917-1918 was the worst winter on record. It began with a ferocious 40 mph blizzard on December 9, 1917 that dropped 16 inches of snow at Louisville in under 15 hours. The temperature in Louisville averaged less than 29 degrees for the three-month period of December through February. Louisville received three feet of snow in January, the highest recorded for that month in Kentucky. The 981-mile-long Ohio River froze over its entire length until January 30, 1918.

2. Most of Arkansas, Tennessee, and Kentucky experienced repeated snows from early December through mid-February.\(^{1042}\) The overall theme was an 8–15 inch snow on December 9, a 4-6 inch New Year’s Eve snow, then numerous back-to-back 4–8 inch snows from mid-January through mid-February. Osceola, Arkansas, about 35 miles north of Memphis, recorded almost 45 inches of snow in a six-week period. We're talking Memphis.

3. North Carolina experienced incredibly cold weather from the mountains to the coast.\(^{1043}\) Just as the new year arrived, the entire Albermarle Sound (15 miles wide at its widest point) froze, perhaps for the first and only time in history. A number of fishing boats were trapped out in the sound and their crews had to walk back to land through very challenging conditions. Coastal North Carolina has never again seen such bitterly cold conditions.

Water year 1918 was exceptionally dry in the Tulare Lake Basin up to the middle of February 1918; less than an inch of precipitation had fallen in Fresno. But then the situation changed dramatically, with more than eight inches of rain falling between mid-February and the end of March. (Likewise, Bakersfield only received a trace of rain during December 1917, but February and March 1918 were both above normal.)

It appears that when the cold Arctic air stopped pouring into the East in mid-February, the winter rains returned to the Tulare Lake Basin. That would be expected if the presumed high pressure ridge in the northeastern Pacific Ocean had broken down, allowing the mid-latitude storm track to return to its typical position and the winter storms to once again reach California. The drought was broken. Gary Sanger concurred with this assessment.

As shown in Table 40, the years 1920–23 were part of the overall drought in the Tulare Lake Basin, but they all had runoff that was at least somewhat above 75% of the long-term average. They provided a modest four-year break before the next severe year of this megadrought.

Second part of drought: 1922–27
The second part of the drought (1922–27) affected much of the state except the Central Sierra and has a possible recurrence interval of 20–40 years. The first year of this part of the drought (1922) had relatively little effect whatsoever in the Tulare Lake Basin. Farmers at the time probably imagined that the drought had ended, but it still had another 12 years to run (1923–34). They had never experienced a drought like this.

Total flow for water year 1924 was 24% of the 1894–2014 average for the Kings, 24% for the Kaweah, 18% for the Tule and 27% for the Kern. Flow for 1931 would be nearly as low as 1924. All four rivers set minimum flows-of-record in 1924, and they were nearly as low in 1931. The Tulare Lake Basin wouldn’t see flows this low again until 1961 (especially on the Tule and the Kern), 1977, and 2014–15 (see Figure 18 on page 111 and Table 23 on page 156).
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

By some measures, the two-year period 1923–24 was the worst Kings River drought on record.\textsuperscript{1044}

In the spring of 1924, rainmaker Charles Hatfield received $8,000 on a rain-producing contract from stockmen on the plains west of Tulare Lake and from lakebed farmers. He set up his plant near Coalinga. However, in the following year, 1925, Hatfield failed to produce the amount of rainfall stipulated in his contract and hence received no pay.\textsuperscript{1045}

W.E. Bonnett, meteorologist with the U.S. Weather Bureau in Fresno, said that only 0.06 inch of rain fell in March 1923; it was the driest March in 42 years. However, April 1923 was the wettest April in 42 years.\textsuperscript{1046}

The drought brought grazing on the ranges almost to a standstill; ranchers were reduced to feeding hay. It was standard practice for cattle from Arizona and Texas to be shipped into Tulare County around the first of the year; 15,000 had already arrived at the beginning of 1924. The foothills had a little grazing left at that time due to recent rain, but the plains were already bare.\textsuperscript{1047}

The drought was simultaneously in effect for the entire state in 1924, and it was particularly severe that year.

1924 was a very small tree-ring year in the Sierra.

The August 1924 national park monthly report said that the drought continued unrelieved. There were fires in the national forests and erroneous reports that travel to the national parks was prohibited or unsafe. Many park streams and springs dried which had never failed in the memory of the oldest inhabitants. Fish were dying in streams. There was ample water at Giant Forest only because of a new water system that had been installed in 1923.

The September 1924 national park monthly report said that virtually no rain had fallen in the park since March, only the lightest showers in April and May. Winter horse feed at lower levels didn’t start until the fall rains, so the park thought that it might have to make emergency purchase of fodder. The September report concluded:

\begin{quote}
But as the hart panteth for the water brook so do all residents of this part of California await the breaking of the Great Drought.
\end{quote}

Table 41 illustrates the runoff in the Tulare Lake Basin during water year 1924.

<table>
<thead>
<tr>
<th>Watershed</th>
<th>Total Runoff (acre-feet)</th>
<th>% of average (1894–2014)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings</td>
<td>391,920</td>
<td>24%</td>
</tr>
<tr>
<td>Kaweah</td>
<td>101,650</td>
<td>24%</td>
</tr>
<tr>
<td>Tule</td>
<td>24,700</td>
<td>18%</td>
</tr>
<tr>
<td>Kern</td>
<td>190,810</td>
<td>27%</td>
</tr>
<tr>
<td>Total</td>
<td>709,080</td>
<td>24%</td>
</tr>
</tbody>
</table>

The combined runoff of the four rivers in the Tulare Lake Basin during 1924 was only 709,080 acre-feet, the second lowest since record-keeping began in 1894, only 1977 would be lower. Runoff in water year 2015 will almost certainly be lower than in either of those years (see Table 72 and Figure 18).

Bill Tweed recalled being told by old-timers in the 1970s that the mainstem of the Kaweah stopped running in 1924; it was only a series of disconnected pools. This may well be true, or it may be a slight exaggeration, colored by the passage of time. There is no way to know for sure. Anecdotal accounts such as this should generally be taken with a grain of salt. For comparison, total runoff in the Kaweah River Basin in water year 1977 was 8% less than in 1924. However, the Kaweah did not quite stop flowing in 1977.
The drought initially appeared to end in October 1924. Sequoia National Park’s monthly report for October began with the statement:

*The principal event of the month was the end of the Great Drought. Heavy rains fell on October 6th, the first for over six months. Rains and snows continued during the month so that precipitations at Giant Forest for the month was 6.30 inches as compared with only 1.07 inches last year. This is an encouraging start for the heavy winter which California so badly needs.*

On November 9–10, 1924, the national parks were deluged with an exceptionally heavy rainstorm; rain came down more heavily than had ever been recorded. December would bring abundant moisture to all of California. For a more complete write-up of that winter, see the section of this document that describes the 1924 Flood. Despite this relief during the winter of 1924–25, the drought would continue for another nine years (1926–34).

One source said that Tulare Lake dried immediately after the April 1923 flood. However, Mae Weis’s account of Tulare Lake history said that a little more water was added from heavy rains during the winter of 1923–24, and that the lakebed went completely dry early in 1924.\(^{1048}\)

The lake did get some inflows from the Kings and the Kaweah Rivers during several of the drought years. However, most of the quantities were small and S.T. Harding believed that much of the water was quickly absorbed by the soil or used directly for irrigation of crops growing in the lakebed. Tulare Lake would not reappear as a large lake until 1937.

Limited floods were welcomed by the grain growers in the Tulare Lakebed since a certain amount of irrigation water was needed, especially during the drought years. On December 6, 1927, practically all of the lakebed reclamation districts joined in forming the Tulare Lake Basin Water Storage District, setting aside 18 sections in the lowest portion of the lakebed for a reservoir.

Yes, you read that right. The drought was so severe that they were creating a reservoir in the lakebed. Since it was being freely predicted that Tulare Lake was a thing of the past due to the increasing use of the water from the tributary streams for irrigation purposes, the idea behind this scheme was to conserve as much as possible of the Kings River runoff for irrigation purposes. The reservoir would afford a certain amount of flood protection; but in 1927, amid the scramble for irrigation water, the flood control angle received scant consideration.\(^{1049}\)

In 1927, California was in the tenth year of the longest drought since 1602. However, the state’s papers would surely have been covering the great flood that was besieging the lower Mississippi River Valley that year. The Mississippi River broke out of its levee system in 145 places, inundating 16 million acres and destroying 130,000 homes. In places, the river swelled to 80 miles wide.\(^{1050, 1051}\) Not until May 2011 would a flood of comparable magnitude come down the Mississippi. That must have seemed like a different world, when viewed from the perspective of a state so mired in drought.

**Third part of drought: 1928–34**

The second part of the drought is considered to have ended in 1927 at the state level, and the last part isn’t generally considered to have begun until 1929. Although 1928 is not considered a drought year at the state level, there are several reasons to view it as a drought year in the Tulare Lake Basin:

- On April 3, 1928, Fresno began a 214-day stretch without measurable rain, the longest such streak on record for that city.\(^ {1052}\)
- As shown in Table 40, the San Joaquin Valley Water Year Index categorized 1928 as a Below Normal year.
- The combined runoff of the four rivers in the Tulare Lake Basin during 1928 was only 53% of the 1894–2014 average.

The 1929–34 drought is considered the longest, most severe drought in the state's history. Parts of the state, especially the northern quarter, were in drought from 1928–37. This drought is sometimes referred to as the drought of 1928–37. Because of the extended duration, the second part of the drought accumulated the largest deficiency in runoff of any drought in the state’s history.

The six-year drought of 1929–34 is unequalled in the historical record of the Sacramento Valley Water Year Index dating back to 1872; this indicates that the drought had a recurrence interval of more than 100 years.

The second and third components of the drought were active somewhere in the state from 1922–34. But within the Tulare Lake Basin, they were primarily active from 1924–34. Based on the combined runoff for the four major rivers in our basin, only two years during that period (water years 1927 and 1932) experienced average
or better flows. The Kern didn’t have average flows even in those years. When Walter Fry wrote Bulletin #8 in November 1931, his editor noted the despondency among valley residents that had been caused by the recent cycle of dry years.

Around 1930, the development of an improved deep-well turbine pump and rural electrification enabled additional groundwater development for irrigation.\textsuperscript{1053} The early wells in the southern San Joaquin Valley were hand-dug pits. In the 1930s, the pits in the Poplar area (northwest of Porterville) began to run dry as the water table dropped.\textsuperscript{1054} See the section of this document that describes the Groundwater Overdraft for a discussion of why the water table was dropping at that time. This marked the transition to the current practice of irrigation using deep-well pumps.

Water year 1931 was an extreme drought year. It was the fourteenth driest year on the upper San Joaquin at the inflow to Millerton Lake during the 1113-year period 900–2012. At a statewide level, water year 1931 ranks as second-driest in 113 years, second only to 1977.\textsuperscript{1055} Water year 1931 was the lowest runoff (184,130 acre-feet) experienced on the Kern River since record-keeping began in 1894. This record would last until 2014.

The winter of 1932–33 was an El Niño event. It was remarkable in the national parks for heavy low-elevation snows. Sequoia National Park’s monthly report for January 1933 began with the note:

\begin{quote}
The only matter of special interest is the weather, which for the last two weeks of January resembled that of Alaska rather than California. All records for snowfall were broken, and old-timers with over fifty years’ experience can recall no such precipitation. The damage to live oak and other trees between the 1,500 and 4,000 foot levels is tremendous; and the damage to park telephone lines, roads, campgrounds, etc. cannot yet be estimated. We shall be hard put to keep the park open and repair damage before summer. But a splendid water supply is piling up in the mountains, and our slight inconveniences are as nothing in the general scheme.
\end{quote}

Following the longest dry winter in the history of the region, a series of storms began on January 15, 1933, that continued with only a two-day break through the end of the month. Precipitation at the national parks’ reporting stations jumped from the least on record to nearly average. Giant Forest received 60 inches (5 feet) of snow in 24 hours on January 19, setting a record for California that would stand until 1982. Giant Forest received 181.5 inches of snow during the month (including 15 feet in 15 days). This huge snowfall was the more remarkable because it occurred during such a severe drought. For the Giant Forest station, both total snowfall and the amount of snow on the ground rose from below average to the greatest ever recorded during January.

The national parks had taken delivery of a new Snow King rotary snowplow just before the storm, and it worked wonderfully. Even so, snow fell faster than equipment could remove it. The Generals Highway closed on the night of January 19 when a slide buried a plow and a pickup. High winds drove the snow into drifts up to 30 feet deep on the steeper sections above Deer Ridge. Equipment continued to be operated 24 hours a day, and additional equipment was rented. Even so, the snow was 6½ feet deep on the roadbed at Granite Springs at the end of the January. The section of road from Deer Ridge to Slide Spring averaged 10 feet deep with many slides 12–15 feet deep. And it was still snowing. It would be February before the storm broke, and the Generals Highway could be reopened.

Although less dramatic, the lower elevations of the national parks also got hit hard in the January storm. Ash Mountain received a total of 23 inches of snow, and snow lay on the ground from January 16–30, a record for that elevation. An emergency purchase of 20 tons of hay was required to feed the parks’ pack and saddle stock for the last half of January. The barn at Clough Cave Station collapsed under the weight of the snow. The water intake at Hospital Rock was demolished by snow slides. Sleet storms wrecked all of the lower telephone lines. Heavy rains at the lower elevations caused considerable damage and slides.

Whatever storm system hit the west side of the Sierra on January 15, 1933 had the power to reach across the Sierra and deliver one of the biggest snow dumps Bishop has ever seen. Waves of heavy wet snow descended on the valley around Bishop from January 15–17 followed by a low temperature of 8 degrees below zero. At Keough’s Place ½ mile north of Keough’s Hot Springs, the snow stayed on the streets until June.\textsuperscript{1056}

Despite this relief during the winter of 1932–33, the drought returned and continued through the fall of 1934.

As illustrated in Table 40, water year 1934 was categorized as a critically dry year in the San Joaquin River Basin.
Water year 1934 was the driest ever in both Fresno and Bakersfield. Fresno received a total of 4.4 inches of rain that year while Bakersfield got only 2.2 inches. That water year had the next-to-lowest runoff experienced on the Tule River since record-keeping began in 1894, only 1977 would be lower.

The weather year is measured from July 1 through June 30; it is very similar to the water year (October 1 through September 30). The weather year is often used for reporting precipitation totals. The 3-year-period July 1, 1931 through June 30, 1934 was the worst 3-year drought in Fresno since record-keeping began in that city in 1878. This record would last until 2011–14.

Not only was 1934 a very low year for precipitation, it was a very warm year. It set the record as the warmest calendar year in California since record-keeping began in 1895. That record would last until 2014.

The national parks’ Ash Mountain headquarters development (originally known as Alder Creek) began operation in 1921. Alder Creek was the sole source of water for that development until auxiliary water sources were developed starting in 1950. With the loss of the Kaweah river pump in 1997, Alder Creek has once again become essentially the headquarters’ sole water supply; there is no longer a significant emergency backup source. Because the flow of that creek is relatively low, there has long been concern about how to get through drought years. In 1934, the park constructed the 75,000-gallon Alder Creek Reservoir; the dam for that reservoir is still largely intact.

Sequoia National Park’s monthly report for October 1934 began with the following note:

A hopeful feeling pervades the park and adjacent Valley because of better than usual early fall rains. The fire hazard has ended and young grass is coming up in the foothills. But the horse ranges have been badly damaged by drought and overuse, and we may have to kill some stock as we cannot get authority to purchase hay and grain.

The period 1927–34 was a serious drought for the Great Basin as well. In December 1934, Lake Tahoe reached its lowest elevation since record-keeping began. A group of stumps was exposed by the receding waters along the south shore of the lake. Since then, these and other submerged stumps in the area have been intensively studied by Susan Lindstrom and others. The stumps measure as tall as 10 feet and up to 3½ feet across. Based on radiocarbon dating of these stumps, the lake reached a low stage on one or more occasions between about 2250–3560 BC. During that period, trees became established and grew for some 100–350 years before the lake raised enough to drown them. The takeaway message is that what we think of as a long-term drought today is mild compared to Megadroughts before the Little Ice Age.

The year 1934 deserves special mention. It was the end of the longest, most severe drought in California’s history (1924–34). But 1934 was also the start of a drought in the Great Plains that would come to be known as the Dust Bowl. That drought lasted 4–8 years, depending on location, from 1933–40.

The Dust Bowl period was the most destructive drought the U.S. has ever experienced. According to NOAA, at least 50 million acres of land were affected. Poor soil management practices made matters worse; without native prairie grasses or cover crops to keep soil in place, the Great Plains quite literally turned to dust and blew away in enormous black dust storms. The drought caused the migration of millions of people from the Great Plains to other parts of the country, especially the West Coast.

The cause of the 1934 drought was the subject of two recent studies:
- A 2014 study from UC Davis led by Richard Howitt.
- A 2014 study led by Ben Cook, a climate scientist at NASA’s Goddard Institute for Space Studies.

Both studies concluded that the 1934 drought was caused by a high pressure ridge over the West Coast deflecting away storms laden with winter precipitation. This ridging pattern has preceded some of the worst West Coast droughts, including 1976 (the first year of the severe 1976–77 drought) and the winter of 2013–14 (part of the 2012–15+ drought).

The Cook study also found that the catastrophic 1934 drought was by far the most intense and far-reaching single-year North America drought on record. It affected about seven times more land area than other droughts of comparable intensity that hit North America during the 1,005-year period 1000–2005, and was almost 30% worse than the 1580 drought, the second most severe drought to hit the continent during that period.
The Cook study used the North American Drought Atlas, a database of drought reconstructions dating back nearly 2,000 years that are based on tree-ring studies. It also analyzed records of air- and sea-surface temperatures and precipitation.

A combination of changes in sea surface temperatures and a lack of rainfall in the Northwest, Southwest, and across the Southern Plains kicked off dry conditions in the fall of 1933 that by the spring of 1934 would spread to the Central Plains and Midwest. Major dust storms — the scale of which had not been seen in North America since the Middle Ages — projected dust from the Central Plains as far east as the Atlantic Ocean.

The drought was likely made even worse by atmospheric effects from human-created dust storms, according to the Cook study. The dust storms may have dried things out further and kicked 1934 into a really extreme event. Based on the Palmer Drought Index, 80% of the contiguous 48 states was experiencing at least moderate drought by the end of June 1934; 63% was in severe to extreme drought. By either metric, 1934 remains the greatest drought year in our country since national record-keeping began in 1895.\footnote{1061, 1062}

**Fire and Drought in the Southern Sierra**

Really large wildfires are relatively uncommon in the Tulare Lake Basin. That is partly because our area is not very susceptible to large wind events. See the section of this document that describes the 1941 Wind Event. Table 42 lists the 32 fires that have occurred in the Tulare Lake Basin since 1910 that have been over 10,000 acres in size.

<table>
<thead>
<tr>
<th>Year</th>
<th>Drought</th>
<th>Fire Name</th>
<th>Watershed</th>
<th>Acres</th>
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<td>1921</td>
<td>1918–1934</td>
<td>Kern</td>
<td></td>
<td>13,172</td>
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<td>1918–1934</td>
<td>Kern</td>
<td></td>
<td>12,523</td>
</tr>
<tr>
<td>1926</td>
<td>1918–1934</td>
<td>Kings</td>
<td></td>
<td>14,969</td>
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<tr>
<td>1926</td>
<td>1918–1934</td>
<td>Kaweah</td>
<td>Kaweah</td>
<td>34,358</td>
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<td></td>
<td>11,993</td>
</tr>
<tr>
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<td>1918–1934</td>
<td>South Fork</td>
<td>Kaweah</td>
<td>22,000</td>
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<tr>
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<td>1918–1934</td>
<td>Kings</td>
<td></td>
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<td>1942</td>
<td>Kern</td>
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<td>1942</td>
<td>Kern</td>
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<tr>
<td>1942</td>
<td>Kern</td>
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<td>Kern</td>
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<td></td>
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<td>1948</td>
<td>1947–1950</td>
<td>Simpson Meadow</td>
<td>Kings</td>
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<td>Kern</td>
<td>McGee</td>
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<td>Kern</td>
<td></td>
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<td>Kern</td>
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<td>1975</td>
<td>Kern</td>
<td>Flat</td>
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<tr>
<td>1984</td>
<td>Kern</td>
<td>Bodfish</td>
<td>Kern</td>
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<tr>
<td>1985</td>
<td>Kern</td>
<td>Deer</td>
<td>Kings</td>
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<td>Kern</td>
<td>Choctaw*</td>
<td>Tule</td>
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<td>2002</td>
<td>1999–2004</td>
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<td>Kern</td>
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<td>Kern</td>
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<td>Kern</td>
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<td>Lion</td>
<td>Kern</td>
<td>20,682</td>
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<tr>
<td>2011</td>
<td>Kern</td>
<td>Breckenridge complex</td>
<td>Kern</td>
<td>25,223</td>
</tr>
<tr>
<td>2015</td>
<td>2012–15+</td>
<td>Rough</td>
<td>Kings</td>
<td>151,623</td>
</tr>
</tbody>
</table>

*Approximately 23,423 acres of the Rankin Ranch Fire and 4,541 acres of the Choctaw Fire crossed the watershed boundary and burned outside the Tulare Lake Basin.
A few of the above fires, especially those in the national parks, were allowed to burn for the benefit of natural resources. Most grew so large despite being aggressively suppressed. Fire behavior is complex; there are many reasons why a fire may grow large. However, it appears that fires are more common during droughts.

Jon Keeley says there is good evidence that large fires are more likely during droughts in Southern California, but it is a bit more complicated in the Sierra; high elevation forest fires burn more area during droughts but lower elevation grasslands and savannas fires are inhibited during droughts due to limited grass fuels.

As discussed in the section of this document that describes the 1918–34 drought, the 17-year-long megadrought of 1918–34 was the driest 6-, 10-, 15-, and 20-year period in the Tulare Lake Basin in at least 11 centuries. According to analysis done by Jon Keeley and Alexandra Syphard, the 1920s was by far the highest decade of burning in the Sierras since 1910. As shown in Table 42, 7 of the 31 large fires in the Tulare Lake Basin occurred during the 1920s. Publicity about fires in the national forests during the 1920s affected visitation to the national parks.

In August 1926, the Kaweah Fire burned 34,358 acres in the drainage of the North Fork of the Kaweah. This fire was originally mapped as 86,000 acres, but was later remapped using GIS. The fire burned in the foothills on the west boundary of the national parks. Of the total acres burned, 11,700 acres were in the parks. This was the largest fire in the history of the national parks.

In August 1928, the South Fork Fire burned approximately 22,000 acres in the drainage of the South Fork of the Kaweah. This fire burned in the foothills just west of the national parks. Although the parks were very involved in fighting the South Fork Fire, it only burned 1,130 acres inside the parks.

The national parks (and the foothills immediately west of the parks) haven’t seen fires of this magnitude since. The next two biggest fires to have occurred in the national parks are:
- 1948 Simpson Meadow Fire — 11,100 acres (occurred during the 1947–50 drought)
- 1977 Ferguson Fire — 10,400 acres (occurred during the 1976–77 drought)

The national parks’ three largest fires (Kaweah, Simpson Meadow, and Ferguson) have all occurred during droughts. These are the only fires that have burned more than 10,000 acres in the parks.

As shown in Table 42, the three largest fires in the Tulare Lake Basin in historic times have all occurred on the Sequoia National Forest during recent droughts:
- The Manter Fire occurred in July 2000, burning 79,223 acres (occurred during the 1999–2004 drought)
- The McNally Fire occurred in July–August 2002, burning 149,475 acres (occurred during the 1999–2004 drought)
- The Rough Fire occurred in August–September 2015, burning a total of 151,623 acres, of which 9,285 acres were in Kings Canyon National Park (occurred during the 2012–15+ drought).

1918 Flood
Flooding in 1918 occurred in September. This flood occurred during the 1918–34 drought.

The 1918 flood occurred during the El Niño of the winter of 1918–19. That El Niño was the subject of a recent study. It was one of the strongest El Niño events of the twentieth century, comparable in intensity to the prominent El Niño events of 1982–83 and 1997–98.

Although a strong El Niño event typically causes flooding in the Tulare Lake Basin, it can have very different effects in other parts of the world. The 1918–1919 El Niño was likely responsible for the severe drought that took place in India in 1918. That was one of the worst droughts that country experienced in the 20th century. There was famine and a lack of potable water, resulting in a compromised population. The drought coincided with an influenza pandemic that was sweeping the globe at that time. The 1918 influenza pandemic killed about 18 million people in India and between 50 to 100 million globally. The authors of the 1918–1919 El Niño study speculated that it might have been linked to the influenza pandemic, especially in India.

In addition to this strong El Niño, 1918 also produced an unusual hurricane. When a Pacific hurricane degrades, it usually makes landfall in Southern California or in Mexico. The year 1918 was the only instance in historic times in which the remnants of a hurricane are known to have come inland as far north as Central California. It isn’t clear what effect this storm had in the Tulare Lake Basin. The Coast Ranges would have gotten heavy rains and flooding. Storms in the Coast Ranges typically spill over into the drainages of western Fresno County and...
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

Kings County. We haven’t found any records to indicate whether that happened in this storm or not. Possibly this storm and flood was just outside the Tulare Lake Basin.

But even if it were, the story of the 1918 hurricane merits inclusion in this document because it serves as a model for two unusual storms that occurred in the Tulare Lake Basin. Both of those were robust cyclic storms which vigorously entered rain shadow areas to the northeast, resulting in a deluge in normally dry areas:

- The February 1978 storm. That storm produced large rainfalls on the windward slopes of Ventura County and then continued over into the rain shadow area in the Buena Vista Lake region.
- The March 1995 storm. That storm produced devastating rainfalls on the windward slopes of the Coast Ranges. It was still quite energetic as it moved into the rain shadow area to create further devastating floods. That was the storm that washed out the Interstate 5 bridges near Coalinga.

On September 11–12, 1918 the remnants of a hurricane tracked to the north-northwest off the coast of Baja California and Southern California, generating only light amounts of rain in the coastal mountains of Southern California.

The storm system apparently moved onshore near Monterey Bay. The town of Antioch, east of Stockton, is on the lee side of the Coast Ranges. Antioch received 6.59 inches during September 12–14 with a recurrence interval of 2,200 years. A total of 12 stations reported rains with return periods in excess of 100 years.

The storm moved north to the Red Bluff area before dissipating. Typically such storms bring a surge of moist warm tropical air that triggers thunderstorms. Red Bluff had 1.19 inches of rain in 30 minutes, 3.72 inches in 2 hours, and 6.12 inches in 24 hours on September 13. Red Bluff received a total of 7.12 inches during the storm event, more than any other station.

1922 Flood

Flooding in 1922 occurred in May and/or early June. This flood occurred during the 1918–34 drought.

The North Fork Kings River near Cliff Camp peaked on June 4, 1922: 6,030 cfs. This was the largest flow on that river since record-keeping began in 1921. This would remain the flood-of-record until the December 1937 flood.

Tulare Lake had dried up on April 30, 1919. One source said that the 1922 flood on the Kaweah was sufficient to leave Tulare Lake eight feet deep at its deepest point. The reliability of that measurement is unknown.

S.T. Harding was aware that the Tulare Lakebed received some inflows from both the Kings and the Kaweah Rivers in 1922. However, he was under the impression that the quantity was relatively small and that the water was quickly absorbed by the soil or used directly for irrigation of crops growing in the lakebed.

Mae Weis’s account of Tulare Lake history from that era said that the floodwaters arrived in May 1922, and that by June a total of 23,680 acres of lakebed cropland was flooded. A little more water was added from heavy rains during the winter of 1923–24, but the lakebed was completely dry again early in 1924.

Total flow for water year 1922 was 130% of the 1894–2014 average for the Kings, 109% for the Kaweah, 104% for the Tule, and 117% for the Kern.

1923 Flood

Flooding in 1923 occurred in April. This flood occurred during the 1918–34 drought.

This was either a rain or a rain-on-snow event. April 1923 was the wettest April in 42 years.

The Kaweah’s peak natural flow occurred at McKay’s Point on April 6: 6,333 cfs. (That was the peak flow; the peak average daily flow was 4,410 cfs.)

The Tule River near Porterville had a maximum daily discharge on April 6: 3,820 cfs. (The term “maximum daily discharge” is presumably the same as “peak average daily flow”). That was the highest flow since record-keeping began in 1901. It would remain the flood-of-record for over three decades. Not even the flood of 1950 would exceed this record.
The Tulare Lakebed was still partially under water in 1923. Possibly it received additional floodwaters during the April 1923 flood, but we haven’t found any records to that effect. The lakebed would eventually go completely dry early in 1924.

### 1924 Flood (2)

There were at least two floods in 1924:
1. October in the Horse Creek Drainage.
2. November across the Sierra

These floods occurred during the 1918–34 drought.

The winter of 1924–25 was a La Niña event.

Frankie Welch said that a cloudburst occurred at Horse Creek in October. It swept a torrent of water four feet deep across the highway in front of the Barney Mehrten ranch, stalling cars and making it impossible to get to and from Three Rivers for a long time.

The November 1924 national park monthly report said that an exceptionally heavy rain occurred over almost the entire park on November 9–10. The rain came down more heavily than had ever before been recorded. Giant Forest received 6½ inches of rain in 24 hours and other points received 2–3 inches in less than an hour. There was severe erosion of park roads, many culverts were washed out, and two small bridges were destroyed. There was a big slide on the Giant Forest Road above Cedar Creek.

The headwaters of Cedar Creek are near the former Colony Mill Ranger Station, but there are several tributary branches of that creek. Bill Tweed said that the Cedar Creek Checking Station was located on the most southwesterly of the branches of Cedar Creek, just below the 4,000-foot contour line. This is probably the Cedar Creek that was being used as a reference for the November 1924 slide.

The December 1924 national park monthly report said that:

> Overshadowing everything else in the park as in California is the abundant rain and snowfall of this winter. There were ten rainy days during December while the month was colder and more gloomy than usual. Six inches of snow were on the ground at Alder Creek park headquarters one morning and it remained for several days. (Alder Creek was the original name for Sequoia National Park’s headquarters development, now known as Ash Mountain.)

By December 31, 1924, Giant Forest had received 108 inches of snowfall, compared with just 31 inches by the same date in 1923. The increase in precipitation was less marked in the valley.

Despite this relief during the winter of 1924–25, the drought would continue for another eight years.

### 1931 Flood

This flood occurred during the 1918–34 drought. Flooding occurred on the Kaweah River at sometime in 1931. Possibly other rivers within the Tulare Lake Basin also flooded.

Little is known about this flood. Two possible clues:
- 1.02 inches of rain fell on Bakersfield on December 8, 1931, setting the record for the wettest December day ever in that city.
- A total of 54 inches (4½ feet) of snow fell in Yosemite Valley in December 1931, setting the record for the snowiest December ever on record there.

### 1932 Flood

Flooding in 1932 occurred in September. This flood occurred during the 1918–34 drought.

This flood was selected by the National Weather Service forecast office in Hanford as one of the top Central Valley weather stories of the 20th century, more noteworthy even than the December 1955 flood.

The subject matter expert on this flood is believed to be Jon Hammond, the editor of the Tehachapi News. Jon has written up this flood in his newspaper. Unfortunately, it has proved impossible to obtain a copy of that
document. The flood is also supposed to be described in detail in the book *Three Barrels of Steam* by James E. Boynton. The following write-up is based on newspaper accounts from the time of the flood and on other sources. 1079, 1080, 1081, 1082

The remnants of an unnamed Pacific hurricane moved up into the Gulf of California and came ashore near Mexicali on September 29. (Pacific hurricanes didn’t start getting names until 1960.)

The storm traveled into the lower desert without much resistance only to break up in the Tehachapi Mountains. Tehachapi received 7.11 inches of rain from September 28 – October 1. The recurrence interval for that event at Tehachapi was 200 to 500 years. Apparently the rains greatly exceeded that in the surrounding mountains.

Reports of this storm are long on details of the effects, but short on quantitative details of the rain amounts. The USACE analyzed the streamflow associated with this storm. They found that the peak runoff rate was 3,815 cfs/sq mi on the 3.5 square miles of the Cameron Creek watershed. This runoff rate (3,815 cfs/sq mi) means that the rain was coming down at an average rate of 6 inches an hour over the entire watershed. That is a rather stunning rate. (The obscure conversion rate used was 1,000 cfs/sq mi = 1.55 inches/hr.)

This is another rare occurrence of extremely large rainfalls on the lee side of an orographic barrier (i.e., the rain shadow of a mountain). For three other examples, see the section of this document that describes the 1918 hurricane.

This is the first time that we know of in historic times in which a Pacific hurricane caused flooding in the Tulare Lake Basin. The 1918 hurricane may have caused some flooding along the northwest side of the basin, but we haven’t found records to document that.

Tehachapi received 4.38 inches of rain in seven hours on September 30, the most extreme rainfall ever recorded in that city. For a time that day, the town of Tehachapi was under three feet of water, with a torrent tearing through the streets and sweeping furniture out of houses. The nearby community of Monolith was also flooded.

Some of the floodwaters flowed north into the Mojave Desert, forming a large lake. The town of Mojave was under two feet of water. However, most of the water poured south down Tehachapi Creek which is the southern fork of Caliente Creek. Caliente Creek drains into the San Joaquin Valley near Arvin, southeast of Bakersfield.

The rain was so intense that it brought Santa Fe Engine No. 3834 to a stop. That train was waiting out the storm atop a new concrete trestle over Tehachapi Creek, ½ mile east of Woodward Station (about a mile upstream from the village of Keene). This was during the Great Depression. In addition to the crew, the Southern Pacific estimated that there would typically have been up to 50 hobos on a freight train such as this. In the middle of that train was a helper locomotive assisting it up the grade to Tehachapi Pass. The Santa Fe train was sitting on a siding. Sitting next to it on the mainline was Sunset freight train No. 829 of the Southern Pacific.

As the floodwaters poured down Tehachapi Creek, they encountered six railroad bridges. At each bridge, debris snagged and created unstable debris dams which held back floodwaters long enough to create temporary reservoirs of runoff. These dams broke apart as water built behind them, creating surges of floodwaters that exasperated the flooding problems. Walls of floodwater, some 40 feet high, raced down Tehachapi Creek as each bridge gave way.

The floodwaters first hit a KAAD service station at Woodford where 15–19 men had taken refuge from the storm. (One source said that this station was at Keene.) Those inside the building were caught up in the floodwaters and some may have drowned. In a separate incident, a family of four in Woodford was drowned when the flood swept away their creek-side house.

Some 20 road camp workers were camped between Keene and Tehachapi. The initial emergency report (before the phone went dead) was that the flood swept down on their camp. Apparently most or all of those workers escaped, but that would not be known until county rescuers could reach the scene a day or two later.

The Kern County Tubercular Sanatorium at Keene was right next to Tehachapi Creek. The flood swept away the pump house, just a few yards from the main building. Three patients were drowned, but the rest survived.
The raging floodwaters piled up 50 feet deep against the trestle that the Santa Fe train was sitting on, undermining it. The trestle gave way directly in the center, collapsing with a roar that could be heard above the deafening noise of the storm. The helper locomotive in the center of the train plunged into the torrent, pulling seven freight cars with it. The Santa Fe locomotive also plunged in, but the Southern Pacific train remained on the mainline track, witnessing the horrifying event. (Southern Pacific passenger train No. 52 had passed only three minutes before the torrent hit the trestle.)

By the time the floodwaters reached Caliente, Tehachapi Creek was flowing at 37,000 cfs. All railroad crossings and 31 miles of track had been undermined, and 600 feet of track were washed out. The cost to the railroad for track repair was $600,000. Huge sections of the state highway through Caliente Canyon (the route that we now know as Highway 58) had been washed out, and at least nine highway bridges were destroyed. In addition to Monolith and Tehachapi, four communities in the Tulare Lake Basin were flooded: Woodford, Keene, Caliente, and Arvin. Flooding in Caliente resulted in the death of a telegraph operator and her two-year-old niece.

High winds accompanying the storm tore down telegraph and telephone lines, isolating the mountain communities. The stricken area was largely cut off from contact with the San Joaquin Valley.

First responders immediately set out from Bakersfield with ambulances to check on the road camp workers and the other areas in the path of the flood. However, they were stopped by the washed out bridges, 20 miles short of the road camp. They had to hike over two mountain ranges to get to the scene of the devastation.

One of the first reports came from Harry W. McGee, a United Air Lines pilot. When he arrived at United Airport in Burbank on October 1, he reported that Tehachapi seemed to have been inundated. He flew over that village in route from San Francisco with 10 passengers, flying out of his way to avoid the worst of the recurrent storms. He reported that mud and debris were visible in the Tehachapi streets.

Total property damage was about $1 million and resulted in 15–26 deaths. Among the dead were the engineer and brakeman of the wrecked train. Two unidentified bodies were assumed to be hobos from the train; there was no way to know how many more remained buried under the mud.

The floodwaters rolled the Santa Fe engine far downstream and buried it under 10 feet of silt and rocks. It took five days to even find the severely damaged engine and a month to free it. All rail traffic over Tehachapi Pass, the inland route between the San Joaquin Valley and the Los Angeles area, was halted by the destruction of the Southern Pacific Railroad track. For the next 14 days, all rail traffic had to be rerouted along the coast.

Following the cloudburst at Tehachapi, another downpour fell near Lebec, near the summit of the Ridge Route (the precursor of the Grapevine or Interstate 5 route over the Tehachapis). Great stretches of that state highway were damaged by the flood and traffic was halted. Most serious was a large rock and mudslide that occurred between Oak Glen and Camp Tejon.

The highway was opened to light traffic the following day, October 1, but guards of highway patrolmen warned motorists that the road was barely passable and prohibited trucking and heavy traffic entirely. County tractors assisted the cars over the stretch of highway between Oak Glen and Camp Tejon.

**1935 Flood**

Flooding in 1935 occurred in April.

The 1918–34 drought had finally ended. The first half of April was stormy. Daily temperatures at both Ash Mountain and Giant Forest were considerably lower than in 1934. Snow fell below Ash Mountain on one occasion. The snowpack in the national park at the end of April was heavier than it had been for many years. However, April was remarkable over and above that. Precipitation in the park for April was much above average. In valley towns, all-time records of rainfall were exceeded in April, Fresno receiving over 16 inches.

The Kaweah River flooded and other rivers within the Tulare Lake Basin may also have flooded. The flood is known from two photographs (on file in the national parks):

- one of Elk Creek flowing across the Generals Highway
- one of the Kaweah in flood adjacent to the Kaweah Hatchery

The culvert at Elk Creek overflowed, washing out 260 yards of the Generals Highway. The Kaweah experienced above-average flows for April, although not nearly as high as April flows would be in the years 1936–38.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

The state had constructed a fish hatchery on the Kaweah River in 1919. It was located directly across from the turnoff for the Mineral King Road. It was originally named the Kaweah Fish Hatchery but was renamed the Hammond Fish Hatchery in 1926. The purpose of the hatchery was to stock the streams of Fresno and Tulare Counties and a portion of Kern County. The hatchery survived a number of floods over the next three decades, some of which caused severe damage. The hatchery would eventually be removed after sustaining major damage in the 1950 flood.

1936 Flood
Flooding in 1936 occurred in February.
This was apparently a rain-on-snow event.
The Kaweah River near Three Rivers peaked on February 13: 8,000 cfs.
The Kaweah’s peak natural flow occurred at McKay’s Point on February 13: 8,360 cfs. (That was the peak flow; the peak average daily flow was 5,366 cfs.)
The Tule River near Porterville peaked on February 13: 12,500 cfs. (This is sometimes incorrectly reported as 12,000 cfs. This was the largest flow on that river since record-keeping began in 1901. Tulare Lake had been dry since 1924. No floodflows reached the lakebed from any river during 1936.
Total flow for water year 1936 was 110% of the 1894–2014 average for the Kings, 115% for the Kaweah, 124% for the Tule, and 109% for the Kern.

1937 Floods (4)
There were four floods in 1937:
1. February (this was the biggest of the four floods on the Tule and Kern Rivers)
2. May/June (this was a snowmelt flood)
3. July
4. December (this was the biggest of the four floods on the Kings and Kaweah Rivers)
Photographs taken during the 1937 and 1938 floods show widespread flooding in the Kaweah Delta. The areas north and east of Visalia looked much as they would in the 1945 flood. Visalia was flooded in one of the 1937 floods, probably the December flood.
The winter of 1936–37 had the heaviest precipitation recorded in the national parks until then. (Record-keeping of weather data began in the parks in 1920. The winters of 1861–62 and 1905–06 were probably heavier, but there was no system for recording precipitation in those years.)
During the last week of December 1936, 75 inches (6¼ feet) of snow fell at Giant Forest, one of the heaviest snowfalls on record up until that time. This was the winter when an avalanche swept away the 125-foot-long Hamilton Gorge Suspension Bridge.
The first flood of 1937 was a rain-flood that occurred in February. In the Tulare Lake Basin, the February flood was much more impressive on the Kern and Tule Rivers than on the Kaweah and Kings. When the biggest floods occur in the southern part of our basin, it usually means that the storm event was centered in Southern California. That was the case with the storm events that caused the 1916 and February 23, 1998 floods.
We haven’t found any specific description of the storm event that caused the February 1937 flood. However, we do have a record of the flood that it caused on Trabuco Creek in Southern California.
Trabuco Creek is a 22-mile-long stream that rises in the Santa Ana Mountains of Orange County and flows toward the city of San Juan Capistrano. The headwaters of that creek are in the large and rugged Trabuco Canyon in the Cleveland National Forest. The last grizzly bear in Southern California was killed in Trabuco Canyon on January 5, 1908. Trabuco Creek once supported one of the most significant steelhead trout runs in Orange County and still has rainbow trout in its upper reaches.
The normally placid Trabuco Creek north of San Juan Capistrano became a river during the February 6–7 storm and changed its course, cutting out over 300 lineal feet of U.S. Highway 101 to a depth of about 25 feet. This would remain the flood-of-record on that stream until the February 23, 1998 flood. The only possible route for a detour of Highway 101 was through the neighboring orange grove, one laden with a fine crop of beautiful fruit.

Fortunately traffic on that section of Highway 101 back in 1937 was relatively light. The owner of the orange grove, a judge, granted permission for the detour to pass through his grove. Provided that fences were erected to protect his orange trees from passing vehicles and to lessen the temptation of passing motorists to sample his fruit. Therefore, the highway department erected eight-foot-high fences made out of chicken wire to create a suitable protective barrier between the beautiful fruit and the slowly passing vehicles.

The Kings River one mile above North Fork peaked on February 6, 1937: 13,400 cfs. This was the largest flow on that river since consistent record-keeping began in 1931. This would remain the flood-of-record for just 10 short months until the much bigger December 1937 flood. The January 1914 flood may well have been bigger, but there was no gage at this location during that flood.

The Kings River at Piedra peaked on February 6: 34,800 cfs. This was the largest flow on that river since the 1916 flood. However, the December 1937 flood would be more than twice as large.

In Giant Forest, 11.96 inches of warm rain fell on six feet of snow between February 5–7, resulting in flood conditions unknown since 1916. There was considerable damage to the Generals Highway and to the Colony Mill Road. The Kaweah River rose 11 feet in 13 hours. Another 7½ inches of rain fell the following weekend, bringing the river to within one foot of its previous high mark. Flood conditions were widespread throughout Central and Southern California.

The Generals Highway was closed for nearly a month (until February 27) by storm damage at a score or more places including major damage at "Deer Creek" (possibly that was a typo and was supposed to say Deer Ridge) in February. Apparently that was referring just to the section of the Generals Highway between Grant Grove and Giant Forest. The section of road below Giant Forest did not reopen until many months later.

Bill Tweed recalled that there was a huge road failure on the Generals Highway just above Deer Ridge. That slide took out more than just the road; the mountainside virtually disintegrated. The CCCs had done a good bit of work in 1933 and 1934 to rebuild the Colony Mill Road. Whether that road was in use at the time isn’t clear. Bill Tweed said that the Colony Mill Road had been closed for at least a few years prior to 1937.

In any case, the landslide of February 1937 was so bad that the Colony Mill Road was reopened and oiled so that it could be used as a detour for much of the following summer. This was the last time that the Colony Mill Road saw significant public use. This may have been one of the last times that the Generals Highway has been closed for more than a couple of weeks. (The highway would be closed for most or all of 1956 and 1967 for replacement of the Marble Fork Bridge near Potwisha.)

Ward Eldredge found a 1937 park monthly report that included the bid for a Deer Ridge bin wall complete with a photograph identifying the location. Manuel Andrade said that this would almost surely have been a wood bin wall, not a metal one. Bill Tweed recalled that a wood bin wall above Deer Ridge was replaced with a metal bin wall about 1980. Perhaps that was the same bin wall that was installed after the 1937 road failure.

The heavy culverts and trash cans that now exist adjacent to the Marble Falls Trail resulted from a failure along the Generals Highway in the vicinity of Deer Ridge and Eleven Range Overlooks, 1,200 feet above. Whether that was from the failure that occurred in 1937, 1952, or 1966 is unknown.

The Kaweah’s peak natural flow occurred at McKay’s Point on February 6: 19,751 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 13,520 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 10 years for the Kaweah.

The USACE said that the 1937 flood was a major flood on the Kaweah and other streams in northwestern Tulare County. It was larger than the 1914 or 1916 floods or any other flood since the turn of the century. Based on the flood exceedence rates in Table 29, the 1914 and 1916 floods each had a recurrence interval of 8 years.

Specific damage in the national parks (including CCC work areas) from the February flood included:

- Salt Creek Truck Trail (outside the national parks but managed by the parks). One section washed out with gulling 2½–3 feet deep (photograph on file in the national parks).
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

- Washout on the Colony Mill Road. Water flowed across oiled surface to shoulder, causing undermining and loss of over ½ mile of road (photograph on file in the national parks).
- Progress Gulch on the Ash Mountain–Advance Truck Trail (what we now call the Shepherd’s Saddle Road). Over 1,000 feet of un-oiled road washed out at one location. Maybe this was where the water tank is near Rattlesnake (photograph on file in the national parks).
- Slide and washout on the Generals Highway at Station 162+79 (photograph on file in the national parks).
- A section of the Generals Highway washed out, tearing out a 36-inch culvert. This happened on the 10 mph curve just below One Shot Rock, about two miles above the 3,000-foot elevation sign (photograph on file in the national parks).

National park CCC crews were employed in the emergency to save life and property in the Three Rivers area.

The Kaweah River near Three Rivers peaked on February 6: 18,900 cfs. This was the largest flow on that river since record-keeping began in 1903.

The river gage at McKay’s Point (three miles below Terminus Beach) had been installed in October 1916. The Kaweah peaked there on February 14, 1937: 16,000 cfs. Two days earlier, the flood had set a record for the highest average daily flow since record-keeping began there in 1916.

The Tule River near Porterville peaked on February 6, 1937 at 12,000 cfs.

The South Fork Kern River near Onyx peaked on February 6, 1937: 3,130 cfs. This was the greatest discharge on that river since record-keeping began in 1911. however, the December 1966 flood (28,700 cfs) would be nine times greater.

The Kern River near Bakersfield peaked on February 7, 1937: 20,000 cfs. Bakersfield narrowly escaped inundation in this event. The levee along the south bank of the Kern came within one foot of being overtopped. An emergency flood-fight helped to protect the levee from overflow. (A similar emergency stand would be required on this levee in the 1950 flood.) The Fruitvale and Fairhaven areas near Meadows Field were flooded, and 16 people had to be rescued by boat in those areas. Over 50 people were evacuated, and all of their homes were destroyed or badly damaged.

The 1937 flood was an impressive event; it was the outstanding early flood in the Kern River Basin. However, when the settlers looked around, they saw high-water marks at much higher elevations. About two miles below the confluence of the North Fork and the South Fork of the Kern, the older marks were 40 feet higher than those of the 1937 flood. Those marks had been left by the 1867–68 flood. That had to be a sobering thought. The reason that the 1867–68 flood had been so high in that area was because of the massive landslide dam failure that occurred on the North Fork in December 1867. See the section of this document that describes the Landslide Dam Failure #4: North Fork of the Kern.

The second flood of 1937 was a snowmelt flood that happened in May and June. In the valley it is sometimes reported as occurring on June 4–7. It did serious damage down in the valley, but there was no damage of note in the national parks.

The North Fork of Kings River below Rancheria Creek peaked on May 14, 1937: 6,510 cfs. This was the largest flow on that river since record-keeping began in 1927. This would remain the flood-of-record for just seven months until the December 1937 flood.

By May 9, 36 sections (approximately 23,000 acres) of the Tulare Lakebed were already inundated by floodwaters. Most or all of this was apparently due to the Kings; runoff from the Kern had not yet reached the lakebed. Cool weather during the first week of May had slowed runoff on the Kings and allowed the lake to drop a little.

Lakebed farmers looked toward the south with apprehension as reports filtered in that the Kern River was slowly making its way toward the southern end of the lakebed. The Kern River hadn’t sent floodwaters into Tulare Lake since 1916. The Kern was flowing at the rate of 1,400 cfs into the sandy country about five miles south of the area that was already inundated. Even though that sand was expected to absorb much of the flow, the river was still predicted to reach the lakebed within the next two or three weeks. Unless the Kern was stopped before then, it might flood another 36 sections or so at the southern end of the lake.
Frank Latta, the San Joaquin Valley historian, viewed the lakebed about the first of May 1937 and declared that a heavy incursion by the Kern could spread the lake out into its dimensions of 1916 and even approach its size in 1890.\textsuperscript{1103}

Some of the wheat and barley was almost six feet tall and promised a record yield. It was perhaps the best grain crop ever grown in the lakebed. In the time remaining to them, the farmers worked to strengthen the lakebed levees protecting their valuable crop.\textsuperscript{1104}

In the early days, only wheat and barley were grown in the lakebed. But by 1937, heavy acreages of cotton and sugar beets were being produced. Gins and cotton oil mills had sprung up in the vicinity. The J.G. Boswell Co. had a feedlot in Corcoran that could accommodate 5,000 head of cattle.\textsuperscript{1105}

The question on the lips of everybody in the communities of Corcoran, Stratford, Lemoore, and Hanford was whether the encroaching lake would get the crops that year before the combines and the cotton pickers. Some observers said that the increased inundation would amount to less than 50 more sections. In the first week of May, it was felt that the next 30 days would tell the story.\textsuperscript{1106}

The second week of May brought hot temperatures, melting the snow and raising runoff levels even higher. The lakebed farmers worked feverishly to strengthen their levees in the hope that they would withstand the crests of water that were then approaching via the Kings, Tule and Kern Rivers. The crests were expected to reach the Tulare Lakebed on the night of May 17 or 18. Lots of water was entering from the Kings River and the lake was rising rapidly, but the ranchers believed the main levees would hold.

The area in the lakebed which was surrounded by the huge dikes was designed to hold a heavy runoff before the reservoir overflowed and flooded the vast farming district. The lake rose three inches on May 16, making an 11 inch rise the previous week. With the Kings, Kaweah and Tule Rivers carrying less water in the mountains due to a diminished melting of snow, the ranchers hoped for continued cooler weather.\textsuperscript{1107}

As of May 17, the Kern River was still being held behind great dikes at the south end of the lakebed. In that area, H.J. Stridde extended his east-west levee a distance of five miles to connect it with the north-south Cohn levee. The hope was that the levee would hold until the valuable grain crops could be harvested. As of May 17, the Kern was flooding across the 23,000-acre Liberty Farms southeast of the lakebed, having reached the headquarters camps that day. It was backing up toward the west across a five-mile front, held back by the levee system. Farmers were constructing other levees in the lakebed in an attempt to save their grain crops if that levee system failed.\textsuperscript{1108}

Farmers in the lakebed spent the week leading up to May 17 strengthening weakened parts of the great dike system. A large crew was rushed to the west bank of the main Kings River channel on May 16 to place brush and sandbags along a curve damaged by crashing action of the water. The brush was cut along the rivers several miles from the lakebed and hauled in trucks. Meanwhile many irrigation and reclamation districts took all of the water their ditches could hold in an endeavor to relieve the pressure on the lakebed levees. Excess water was diverted over rangeland wherever possible. Tulare Lake was a mecca for sightseers on May 17, the water being dotted with boats and levees with automobiles. More than 500 automobiles were parked at the mouth of the Tule River on the afternoon of May 17.\textsuperscript{1109}

Three state highways in Kern County were closed to truck travel on May 17 for parts of their lengths as a result of water spread over them by the overflow of the Kern River. The closing of the roads to trucks resulted from fear that the overflow water might have weakened the roadways to such an extent they would be unsafe for heavy loads. The highways closed to trucks were Pierce Road, a cutoff road connecting the Golden State Highway and the Rosedale Highway; Enas Lane, a cutoff road joining the Wasco-Shafter and Taft Highways west of Old River, and the McKittrick Highway.\textsuperscript{1110}

Four miles of the McKittrick Highway six miles east of Buttonwillow were under water on May 17 and improved county roads into the Wildwood district adjacent to the flooded area were rendered impassable. A considerable area of pasture land was flooded, but low levees kept the water from alfalfa, cotton and potato crops. The Kern River Powerhouse No. 1 reported the river had dropped from a flow of 8,046 cfs feet on May 16 to 7,740 cfs on May 17, and officials believed the peak of the present runoff had passed.\textsuperscript{1111}

Tulare Lake continued to rise into June. We haven’t found any records to indicate how much acreage was flooded, or how successful farmers were at harvesting ahead of the floodwaters.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

July brought numerous thunderstorms to higher elevations of the national parks, a fairly typical situation. However, a cloudburst on the evening of July 24 did considerable damage in the Mineral King area. The road was washed out in one spot and fabric automobile tops were riddled by hailstones. The East Fork Kaweah rose two feet in 20 minutes.

And then came the fourth flood of the year. November was generally dry in the Tulare Lake Basin. Mean discharge of the Kings River at Piedra for November 1937 was about half the 43-year average for that month. That situation was about to change dramatically.

The floods were caused by an exceptionally intense rainstorm of wide extent, which formed over the Pacific Ocean and moved rapidly eastward into Northern California on December 9. It was a well-defined single storm, and most of the precipitation fell within a 48-hour period. The storm was notable for the accompanying warm temperature, which caused precipitation to have the form of rain, rather than snow, up to high elevations in the Sierra. Large amounts of rainfall fell in the middle elevations of the Sierra where normally much of the precipitation during December storms is in the form of snow.

A notable characteristic of the storm was the relatively small amount of rainfall on the floor of the Central Valley. The rainfall was also of only moderate depth and intensity in the coastal areas south of the Salinas Basin, and in Owens Valley on the east side of the Sierra.

In general, there was little snow on the ground at the beginning of the storm period, and contribution from melting snow was not an important factor in the flood runoff.

The storm was a high-elevation event centered in the northeast corner of the state. A total of 21 stations reported their highest-ever two-day rainfall during that storm. The highest intensity part of this storm was in a zone between Inskip Inn (northeast of Chico) to Alturas (northeast of Redding). Alturas had 5.08 inches of rain with a recurrence interval of 22,000 years.

The storm resulted in the highest-ever rainfalls at 80 river gaging stations from the Trinity River in the north to the Kaweah River in the south. Five stations reported over 10 inches of rain on December 11. Hobergs (south of Clear Lake) received a total of 20.50 inches during the two-day storm event. Felton (north of Santa Cruz) and Los Gatos Summit both reported their highest-ever two-day rainfall, with over 14 inches during the storm event.

Table 43 shows the elevation and precipitation for the December 9–11 storm event. The greatest precipitation was received on December 10.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Elevation (approximate)</th>
<th>Storm Total (inches of rain)</th>
<th>Drainage Basin</th>
</tr>
</thead>
<tbody>
<tr>
<td>Auberry</td>
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<td>9.2</td>
<td>San Joaquin</td>
</tr>
<tr>
<td>Crane Valley Reservoir</td>
<td>3,500</td>
<td>14.7</td>
<td>San Joaquin</td>
</tr>
<tr>
<td>Huntington Lake</td>
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<td>San Joaquin</td>
</tr>
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<td>Balch Powerhouse</td>
<td>1,750</td>
<td>10.3</td>
<td>Kings</td>
</tr>
<tr>
<td>Big Creek research facility</td>
<td>1,950</td>
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<td>Tejon Ranch</td>
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<td>1.7</td>
<td>Kern</td>
</tr>
</tbody>
</table>

Flood stages prevailed from the Kaweah River in the south to the Pit and Trinity Rivers in the north, and from the Pacific Ocean to the Sierra.

Maximum discharges were recorded on the Sacramento, upper San Joaquin, Kings, and Kaweah Rivers, as well as on many tributaries. On the Feather River the peak discharge was approximately that of 1928, whereas on the American River it was considerably less than in that record year.
The rivers and creeks over most of Northern California, from the Kaweah on the south to the Pit and Trinity on the north, rose rapidly to very high stages as a result of the December 9–12 storm. The storm was severe from the coast of California to the State of Nevada, and many streams on both sides of the Coast Ranges and the Sierra exceeded previously recorded maximum flood discharges.\textsuperscript{1120}

Record-breaking runoff rates were recorded in the foothill and mountain areas in the Mokelumne, Stanislaus, Tuolumne, Merced, and San Joaquin River Basins but, because of available storage in Lake McClure, Hetch Hetchy, Don Pedro, Melones, Salt Springs, Pardee, and other smaller reservoirs, there was no heavy runoff along the lower reaches of those streams. The discharge of the Merced River at Pohono Bridge near Yosemite was nearly four times the previous maximum recorded in 1922, and about one half of the valley floor in Yosemite Valley was flooded.\textsuperscript{1121}

The discharge of the Kaweah and Kings Rivers and their tributaries far exceeded previous long-period records. However, the storm was much less intense south of the Kaweah, and the peak discharges of the Tule and Kern Rivers did not equal those of the flood of February 1937.\textsuperscript{1122}

The most extreme flood-peak discharges were in parts of the Northern and Central Sierra. The December 1937 flood was widespread over the northern two-thirds of the state. It had a recurrence interval that was greater than 100 years on some rivers. There were several peaks of record in the Northern and Central Sierra. Damage was $15 million.

In coastal streams there was extensive flooding from the Russian River south to the Santa Clara Valley. There was extensive flood damage in the Feather River Basin. Main Street in Chester (east of Red Bluff) was washed away. A new record-high river stage of 31.95 feet occurred on the Sacramento River at Red Bluff on December 11. That was 1.5 feet higher than the previous record high of 30.5 feet set in 1909. The Yosemite Valley Highway was flooded by the Merced River. Extensive flooding occurred in the Tulare Lake Basin.\textsuperscript{1123}

The North Fork Kings River near Cliff Camp peaked on December 11: 14,000 cfs. This was the largest flow on that river since record-keeping began in 1921.\textsuperscript{1124}

The North Fork of Kings River below Rancheria Creek peaked on December 11: 21,000 cfs. This was the largest flow on that river since record-keeping began in 1927.\textsuperscript{1125}

The Kings River one mile above North Fork peaked on December 11: 42,000 cfs. This was the largest flow on that river since consistent record-keeping began in 1931. It was over three times larger than the February 1937 flood.\textsuperscript{1126}

The Kings River at Piedra peaked on December 11: 80,000 cfs. This was the largest flow on that river since record-keeping began in 1895. There was slightly more precipitation in the storm of January 1914 than in December 1937, and it fell on ground previously moistened. However, there was heavy snow at an elevation of about 6,000 feet at the beginning of the 1914 storm, and the snowline lowered about 2,000 feet in elevation during the storm, making the amount of precipitation available in the form of water approximately equal to that in 1937.\textsuperscript{1127}

The Kings River at Piedra peaked on December 11: 80,000 cfs. This was the largest flow on that river since record-keeping began in 1895. This would remain the flood-of-record until the 1950 flood. However, just as on the Kern, there were reminders of an earlier and bigger flood. At Pine Flat, the high-water marks of an early flood (believed to be the 1867–68 flood) were seven feet higher than the December 1937 flood. The 1867–68 flood remains the greatest flood on the Kings since at least the flood of 1805.\textsuperscript{1128}

Farmlands, highways, bridges, and public utility systems were seriously damaged. From Alturas, on the upper Pit River, to Visalia in the Tulare Lake Basin, several towns and cities suffered severe damage from overflow, and large areas of agricultural land were covered by floodwater.\textsuperscript{1129}

In the San Joaquin Valley the damage was largely limited to the foothill and mountain areas and to the lower lands along the San Joaquin, Kings, and Kaweah Rivers where large areas of farmlands were flooded. The Kings River inundated about 15,000 acres in Centerville Bottoms at the edge of the foothills below Piedra, and 30,000 acres in the Burris Park, Laton, and Lemoore areas. The Bottoms acted as a reservoir until the crest of the flood had passed, and then the water drained back into the Kings River and thus increased the flow in the lower reaches.\textsuperscript{1130}
The 1937 flood caused severe damage to the state’s Kings River Hatchery. That hatchery was located on the South Fork of the Kings River, upstream from the junction with the North Fork.

The U.S. Forest Service mapped the high-water line of the 1937 flood in the vicinity of where the Cedar Grove Bridge would later be built. That line coincides reasonably well with the modeled 50-year flood event. That puts it in a category with other Cedar Grove floods of the past 70 years (1937, 1950, 1955, 1966, 1969, 1978, 1982, 1983, and 1997) that rise to the level of the modeled 50-year flood: that is, a flood event that occurs about every eight years on average. See the section of this document that describes Cedar Grove Flooding.

Giant Forest received 16.28 inches during the December 9–12 storm event, 14 inches of which fell during December 10–11. Ash Mountain received 9 inches during the December 10–11 period.

The North Fork Kaweah near Kaweah peaked on December 11: 8,290 cfs. This was the largest flow on that river since record-keeping began in 1910. The Kaweah River near Three Rivers (USGS gage #11-2105) peaked on December 11: 33,300 cfs. This was the largest flow on that river since record-keeping began in 1903. This was believed to be the highest level since 1867. Floodmarks left by the December 1867 flood showed that flood was at least four feet higher than the December 1937 flood. As on the Kings River, the December 1937 flood would remain the flood-of-record until the flood of 1950.

The national parks’ records said that the 1937 flood took out or damaged 13 of the 14 bridges that spanned the Kaweah’s various branches; six of those bridges were in the parks. The Marble Fork Bridge near Potwisha was severely damaged (photograph on file in the national parks).

The records don’t specify which bridge survived the flood unscathed, but it’s tempting to think that it was the Oak Grove Bridge on the East Fork Kaweah. That bridge appears to be relatively flood-proof and we have found no record of it being damaged in any flood.

A concrete arch culvert under the Generals Highway at Dorst Creek was so badly damaged that it had to be replaced. The flood also washed out three trail bridges in the parks.

The national parks were closed for two days until the worst of the road damage was repaired. It took 10 days to rebuild the Marble Fork Bridge so that it was once again passable to traffic.

The approaches to the Pumpkin Hollow Bridge (Bridge #46-29) were destroyed, but the bridge survived. The parks’ approach would be washed out again in the 1955 and 1966 floods, but the original bridge is still in use.

The Kaweah Hatchery near Hammond was severely damaged during the 1937 flood. The hatchery was repaired after that flood and operations continued.

The Dinely Bridge washed away.

Jim Barton recalled that warm rain came down for three days straight. It rained even at Lodgepole, so there was no ice to skate on at the skating rink. The mainstem of the Kaweah crossed the North Fork Drive above the Barton ranch and flooded their pasture. It flowed down North Fork Drive until it turned back toward its original channel across from present-day Flora Bella Farm.

The North Fork Kaweah peaked on December 11: 8,290 cfs. This was the largest flow on that river since record-keeping began in 1910. This would remain the flood-of-record until the 1950 flood.

According to Sophie Britten’s book *Pioneers in Paradise*, the original bridge across the mainstem of the Kaweah in Three Rivers was a trestle bridge that was finished September 10, 1897. For a while, it was known as the River Inn Bridge because the River Inn was located right where the bridge crossed the river. (The hotel was built in May 1910 and burned to the ground in September 1911). According to Sophie’s book, that bridge survived until the December 1937 flood.
By 1937, the bridge at this location was known as the North Fork Bridge or the Three Rivers Bridge. During the 1937 flood, Jim Barton and his family gathered at this bridge to see how it would withstand the flood. Earl McKee, Jr. also witnessed the bridge washing out.\footnote{1138, 1139}

Jim said the bridge was a wooden truss structure built in 1906. The post-and-timber bridge was anchored by four steel cylinders filled with concrete. It was the only public bridge across the mainstem of the Kaweah River. As onlookers watched, two young men — Orlen “Baldy” Loverin and Fred Gimm — drove onto the threatened structure from the North Fork side in Loverin’s 1934 Chevrolet coupe. Upon reaching the Highway 198 side, they realized that the approach was gone, so they backed the car back toward the North Fork side.

But now the water was too deep at that approach, and they were stranded. The men were able to get off the bridge by hanging onto a barbed wire fence while fording the raging water, but the car had to stay. The river rose several more feet, completely inundating the car. Eventually the bridge and car washed away. Earl recalled their car sitting on the bridge as the bridge broke loose from the columns, swung around slowly like a big ship, and headed downstream. Then it slowly rolled over. The bridge and car were found the next morning on the Thorn Ranch below where the present-day North Fork Bridge stands. The car was pretty banged up, but had landed upright back on the bridge, its tires still on the runner planks.\footnote{1140}

The two cylinders on the highway side of the bridge washed completely away and came to rest, along with some other pieces of the bridge, in a swimming hole (across from what would later be Pat O’Connell Towing) but was then known as the “Old Twenty.” That, Jim said, was the end of that swimming hole.

The Airport Bridge survived. At the time, there was no Kaweah River Drive. The road ended at the Taylor Ranch, just beyond the Three Rivers Airport, which had opened just two years before. At the Taylor Ranch, there was a private bridge that connected that area with Highway 198. It was built in two sections, connecting at an island in the middle of the river. That bridge also washed out, leaving North Fork residents stranded. Jack Hill, a county road foreman who lived on what is today the Anjelica Huston ranch, took his bulldozer to the Taylor Ranch and scratched out a road up to Dinely Drive. That route would later become Kaweah River Drive.\footnote{1141}

It took a while to replace the North Fork Bridge so that cars could get across the mainstem of the Kaweah. Initially, people were accommodated by a cable and trolley system. A person would sit on a swing board, place their possessions in an orange box dangling from it, and pull themselves across and over the river. Jim Barton recalled that this system was in place in time for Three Rivers School’s Christmas program as he remembers being on it in the dark with his entire family.\footnote{1142} A cable trolley would again be used at this location after the North Fork Bridge was destroyed during the December 1955 flood.\footnote{1143}

A temporary bridge was put in place for 1½ years until a permanent North Fork Bridge could be completed. That new permanent bridge would be located next to the present-day Three Rivers Market. It consisted of three Bailey Bridges placed end to end on concrete piers. That new bridge remained in use until it washed away in the December 1955 flood.\footnote{1144}

Highway 198 (later known as Old Three Rivers Road) crossed the South Fork Kaweah via a concrete bridge. The South Fork undermined that bridge, causing it to collapse in the center during the flood. A temporary plank bridge was installed until a new bridge could be built in the same location.\footnote{1145}

Sometime after 1938, Highway 198 was realigned and a new bridge across the South Fork was constructed where Kaweah Park Resort is today.

The December flood severely damaged Terminus Beach. The Kaweah at McKay’s Point peaked on December 11: 35,000 cfs. That was the highest flow since the gage was installed in 1916.\footnote{1146} The flood apparently destroyed or otherwise overwhelmed the concrete weir at McKay’s Point.

The Tule River near Porterville peaked on December 11: 11,300 cfs.\footnote{1146, 1147} This was not as big as the flood-of-record which occurred on February 13, 1936 (12,500 cfs).\footnote{1148}

The Kern River peaked on December 11, 1937. It was a significantly smaller flood than the February 6, 1937, flood had been.\footnote{1149, 1150}

The South Fork Kern River near Onyx peaked on December 12: 1,260 cfs. This was less than half as big as the February 1937 flood had been.\footnote{1151}
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The Kern River near Kernville peaked on December 11: 6,800 cfs. The flood-of-record was 9,690 cfs on Jan. 17, 1916. Records had been kept on this stream since 1912.\textsuperscript{1152}

The Kern River near Bakersfield peaked on December 12: 6,859 cfs. The flood-of-record was 20,000 cfs on Feb. 7, 1937. Records had been kept on this stream since 1893.\textsuperscript{1153}

Los Gatos Creek is a short, torrential, intermittent stream near Coalinga. It is in one of a group of foothill streams in semi-arid regions that may have no surface flow for many months in some years. Runoff is largely dependent upon the absorptive condition of the soil of the basin, which in turn may be determined largely by the amount of recent antecedent rainfall. Runoff is typically flashy in these small semi-arid basins where there is no sustained flow.

In the December 1937 storm event, the rain in the Coalinga area fell upon dry ground. Los Gatos Creek and probably other streams in the area flash flooded.

Los Gatos Creek peaked on December 11: 1,530 cfs. It was the biggest flood on that stream since record-keeping began in 1931. (It would remain the flood-of-record for only two months until the February 1938 flood.) The flood no doubt continued downstream on Arroyo Pasajero.\textsuperscript{1154}

The Owens River flow-of-record occurred on December 12, 1937 storm.\textsuperscript{1155} The Owens River may have experienced several floods bigger than this during the 1800s, but that was before gaging began on this river.

Total flow for water year 1937 was 137% of the 1894–2014 average for the Kings, 160% for the Kaweah, 223% for the Tule, and 170% for the Kern.

The Kern River sent floodwaters into Tulare Lake for the first time since 1916. Tulare Lake reappeared on February 7, 1937, for the first time since 1924. American white pelicans, waterfowl, and shorebirds reappeared almost instantly and in incredible numbers. (See the section of this document that describes the Chronology of Tulare Lake for a description of this remarkable biological event.)

After peaking on June 16, Tulare Lake receded until December 14, when it began to receive water due to the December flood.\textsuperscript{1156} By the end of water year 1937, the lake was about 13 feet deep at its deepest point (elevation 191.9 - 179 feet).

**Big Creek Debris Flow**

In addition to flooding, the December storm caused a major debris flow in the lower Kings River Basin. Many debris flows in the Sierra are never recorded. We know about this debris flow in large part because it occurred on a USFS research station: the Pacific Southwest Research Station. That facility was known at the time as the California Forest and Range Experiment Station. The watershed is located immediately north of Pine Flat Reservoir. The event was analyzed and summarized by Jerry DeGraff, a geologist for the USFS.\textsuperscript{1157}

The storm began at 5:00 p.m. on December 9. Precipitation fell mainly as rain and ended at 7:00 p.m. on December 11. The rainfall included two high intensity periods of 2 inches for one half-hour and 1.5 inches for one hour in the Kings River Basin. In the Big Creek Basin, a tributary to the Kings River, the experiment station maintained weir and gage instrumentation on eight small subwatersheds ranging in size between 4 and 15 hectares. Total rainfall in the Big Creek Basin from the storm was 12.3 inches. The rainfall occurred when only eight inches of snow was present on the summits. Structures housing streamflow instrumentation near the mouths of the subwatersheds were destroyed or severely damaged by flooding which carried considerable debris.

While the damage to the dams and instrumentation on the subwatersheds in Big Creek was attributed to flooding, Jerry concluded that it was actually the result of a debris flow. Land slumps occurred in the upper parts of the subwatersheds. Photographs document shallow slope movements which lead into the channels. In the channel above one gaging station, the passage of the debris down the channel appears to have followed its own course rather than remaining strictly confined to the channel banks. Other photographs show the channels scoured to bedrock. The bulk of the debris was described as having originated in the channel bottoms and sides. The small dams were pounded terrifically by large boulders, some weighing as much as 5,000 pounds, which literally battered out the centers of three of the dams. At one of the subwatersheds, the passage of the debris left a four-foot-high debris levee which redirected flow and reduced damage to the dam. The presence of debris levees, the large size of transported boulders, the channel scouring, the variance of the path or track of the flow...
relative to the channel banks, and the shallow landslide sources in the upper watershed are all evidence of a debris flow rather than floodwaters.

1938 Floods (2)

There were two floods in 1938:
1. February–March
2. December

The winter of 1937–38 was a heavy snow year, the second such year in a row. Yosemite recorded 61.09 inches of precipitation in 1938, setting a record that would last until 1983. The winter of 1938–39 was a moderate to strong La Niña event.

Following on the heels of the four 1937 storms, another heavy storm and flood event hit Central and Southern California just three months later on March 2–3, 1938. It was a 50–80 year flood event. The storm of March 2, 1938 produced some of the largest streamflows ever recorded in much of Southern California. (The 1861–62 flood was much worse, but very few Anglos were living in Southern California at the time.) Bakersfield set a 24-hour precipitation record for the month on March 3. Sixteen stations (most of which were in Los Angeles and San Bernardino Counties) reported 10 or more inches of rain on March 2. It resulted in ⅓ to ½ of the average annual rainfall at those stations in that one day. Records were set by the resulting flood that wouldn't be broken until the 1969 flood. The flood totaled $79 million in damages and resulted in 87 deaths.

In February and March, 1938, heavy storms flooded the San Joaquin Valley.

Panoche/Silver Creek west of Mendota flooded sometime in 1938, probably in February or March.

The Kaweah River crested in Visalia on the night of February 26.

The South Fork Kern River near Onyx peaked on March 2, 1938. This was the greatest discharge on that river since record-keeping began in 1911. This flood-of-record probably resulted from less total precipitation than the smaller December 1937 flood, but it occurred later in the rainy season after the absorptive capacity of the ground had been considerably utilized.

The late February part of the 1938 flood was centered in the Tulare Lake Basin. The early March part of the flood was a major event and affected all of Southern California.

Migrant laborers suffered the most from the flooding. John Steinbeck came to Visalia in February 1938 to help relieve their suffering. This experience was apparently the motivating factor in his writing *The Grapes of Wrath*. Horace Bristol photographed the migrant encampment in Visalia and elsewhere. Steinbeck based the central characters in his masterpiece on the farm workers that he and Bristol encountered in Visalia that winter. The book’s climatic flood was based on what he witnessed in Visalia. Bristol’s photographs were used to cast the movie and were later published in *Life* magazine.

In the February 1938 storm event, Los Gatos Creek and probably other streams in the Coalinga area flooded. Los Gatos Creek peaked on February 11: 4,520 cfs. It was the biggest flood on that stream since record-keeping began in 1931.

This greatly exceeded the previous flood-of-record that had been set just two months earlier in the December 1937 flood. The February 1937 flow was so large in part because the rains of that storm fell upon ground previously moistened. The flood no doubt continued downstream on Arroyo Pasajero.

The February 11, 1938 flood may be the flood that caused Los Gatos Creek to flood and severely damage the Coalinga Cemetery. As a result, that cemetery was closed to further burials and the town used the Avenal Cemetery from then on.

The Kaweah’s peak natural flow occurred at McKay’s Point on December 11: 34,799 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 11,232 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 8 years for the Kaweah.

A great flood struck Los Angeles County on March 3, resulting in $45 million dollars in damages and 113 deaths. A total of 5,601 homes were destroyed and another 1,500 were severely damaged. Thousands of people had to
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be evacuated and thousands more were left homeless. Two CCC camps were destroyed and over 300 relief workers had to be rescued. A total of 91 highway and railroad bridges were destroyed or badly damaged. With all the rail lines out of service, there was no mail delivery, so mail was taken by the U.S. Coast Guard between L.A. and San Diego. The peak flow of the Los Angeles River at Long Beach exceeded the average flow of the Mississippi River at St Louis.

Orange County also experienced a great flood in early March. It was the most destructive in the county’s history, resulting in 19 deaths and 2,000 homeless. Eyewitness accounts say that an 8-foot-high wall of water swept out of the Santa Ana Canyon. At the peak of the flood on March 3, the Santa Ana was flowing at an estimated 100,000 cfs. Near the mouth of the river, the Santa Ana overflowed its banks and covered an area 15 miles long and 5 miles wide.

San Bernardino County experienced major flooding from March 1–5 due to a series of storms which resulted in very heavy rainfall. This flooding event seems to have been centered in the upper Santa Ana River Basin. Some areas in that watershed received over 30 inches of rain during the event. Over 100 bridges were destroyed and 800 miles of roads were lost. Over 150 homes were destroyed and many more flooded, leaving over 1,000 homeless. Most USGS gaging stations were destroyed. Cajon Pass was closed to traffic due to miles of road destruction, bridges washed out, rail lines destroyed, and dozens of landslides. All communications were cut off; the only routes left open were by foot or air. The Mojave River experienced a major flood; 22 homes were swept away in Victorville. The peak discharge from the 1938 flood exceeded any flood since the 1861–62 flood which is considered the flood-of-record for this area. Damage in the county exceeded $11 million dollars, and 22 people died.

Riverside County was extensively damaged by the flood of March 1–3. The northern section of Riverside was inundated and many people were forced from their homes. Men, women, and children had to be rescued from trees as they were unable to reach higher ground when their homes became imperiled. Fairmont Park saw great destruction when the dam at Lake Evans was ripped out by floodwaters. Livestock of all sorts was lost to flooding in the Santa Ana River. Damage to roads, bridges, and rail lines in the county was extensive.

December precipitation at both Giant Forest and Ash Mountain was the greatest since record-keeping began in 1920. Between December 9–12, 16 inches of rain fell in Giant Forest, bringing the Kaweah to flood condition and causing unprecedented damage to roads and trails. Six bridges were destroyed. In addition, the Marble Fork Bridge near Potwisha was badly damaged, and the arched culvert over Dorst Creek was so badly damaged that it had to be replaced. Damage and bridge destruction immediately outside the national parks was even greater than damage within the park.

Total Giant Forest precipitation during the calendar year was 66 inches (5½ feet), the greatest since record-keeping began in 1920.

The Kern River near Bakersfield peaked on March 3: 14,600 cfs. Total flow for water year 1938 was 192% of the 1894–2014 average for the Kings, 205% for the Kaweah, 258% for the Tule, and 190% for the Kern. This was the first time since record-keeping began in 1894 that the Kaweah River had two back-to-back years with flows that were over 150% of average.

The 1938 flood caused major flooding in the Tulare Lakebed. When the elevation of Tulare Lake reached 192 feet, one of the main levees in the lakebed broke and the lake spilled over 49 square miles of land. The lake continued rising, eventually cresting at 195 feet. This compared with a maximum elevation of 193.1 feet in the 1906–07 flooding.

By June 1938, 135,600 acres of the lakebed were underwater. That was the maximum acreage flooded since the 1906–07 flooding. Tulare Lake has not been this big since. By the end of water year 1938, Tulare Lake was about 16 feet deep at its deepest point (elevation 195.0 - 179 feet).

In the lakebed, the barley harvest would normally have started in mid-May, but had been delayed by the after-effects of the December 1937 flooding followed by the heavy spring rains. Because of those delays, the harvest got underway just as the rivers went on a rampage, tearing into the levee systems. The farmers found themselves in a race to bring in the harvest before the various lakebed levees failed. As harvesters worked around the clock, earthmovers and an army of shovel-wielding men fought on the levee banks. If the harvesters won and got the grain pulled over the levee in tractors and eased down into the next block, then the levee
would be blown up. But if the levee broke before the harvesters were finished, then the harvesters stayed in the field, working just ahead of the slowly spreading water and, at the last minute, were jerked out of that field and into the next.\textsuperscript{1166} USBR estimated that 126,000 acre-feet of water came into the Tulare Lakebed during water year 1938.\textsuperscript{1167}

While the high lake levels of 1938 were a disaster for the lakebed farmers, others saw their opportunity. Near the height of the flood, Frank Latta and three boys took a 15-foot homemade motor boat from Bakersfield to San Francisco. They left Pioneer Weir on the Kern River on June 18, 1938. The trip ended 14 days later at the wharf on the south end of Treasure Island.\textsuperscript{1168, 1169}

Treasure Island had been built specifically for the Golden Gate International Exposition (aka World’s Fair). The exposition would open to the public on February 18, 1939. Building of the exhibits was well underway when Latta and the boys were there, so they billed their trip as a visit to see the exposition. This was the third of six documented trips between Tulare Lake and San Francisco Bay to occur in historic times. (The other five trips were in 1852, 1868, 1966, 1969, and 1983.) It seems like all the trips after 1868 must have encountered impediments of one type or another; water last flowed out of Tulare Lake in 1878.

**1939 Flood**

Flooding in 1939 occurred in June.

A very intense summer storm struck Fresno on June 14, 1939. At one point, the rain was coming down at a record-setting rate of 0.65 inch in 10 minutes (a rate of 3.9 inches per hour).\textsuperscript{1170} This almost certainly resulted in flooding in the area.

1939 was also the year that the Boyden Bridge (Bridge #42-24) was completed on the newly constructed Highway 180 in Kings Canyon. The Grant Grove approach to that bridge has been washed out on at least two occasions (1955 and 1997), but the original bridge is still in use.

The original Cedar Grove Bridge is believed to have been erected shortly after the Boyden Bridge was completed, but that is based on supposition.

**1940 Flood**

There was at least one, and possibly as many as three, flood events during 1940:

1. flooding at an unknown time on Dry Creek
2. flooding at an unknown time on the Kern River
3. flooding in October in Bakersfield

Sequoia National Park received much more than average precipitation in 1940, but it was largely in the form of rain. Snowfall was less than half of what had occurred in 1939. No flooding was reported in the national parks.

Torrential rains pounded the hills east of the Visalia Electric mainline, flooding Dry Creek and washing away a 45-foot trestle at Dry Creek.

The Kern River flooded enough to damage Highway 178 through the canyon.

A storm dropped 1.51 inches of rain on Bakersfield on October 25, 1940, making that the wettest October day ever in that city.\textsuperscript{1171} Such storms are typically intense and result in street flooding.

**1941 Floods (2)**

There were two floods in 1941:

1. February
2. Sometime during the April–July snowmelt runoff period

A very intense storm struck Fresno on February 24, 1941. At the peak of the storm, rain was coming down at a record-setting rate of 0.48 inch in 5 minutes (equivalent to 5.78 inches per hour).\textsuperscript{1172, 1173} This almost certainly resulted in flooding in the area. This storm is sometimes reported as having occurred on February 24, 1951. However, the NWS forecast office in Hanford researched their files and confirmed that 1941 was the correct year.
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Statewide, the four wettest water years during historic times were 1890, 1941, 1983 and 1995. The heavy rains of 1941 were confined largely to the Sacramento Valley and a narrow strip of land on the South Coast from Santa Barbara to Orange Counties. Both Willows (near Chico) and Santa Ynez (near Santa Barbara) had rain totals for the water year with recurrence intervals in excess of 5,000 years.1174

In the Tulare Lake Basin, runoff in the spring of 1941 was heaviest in the south end of the valley. It was a much bigger flood on the Kern than on the Kings or Kaweah. A photograph on file in the national parks shows the Kern in flood at the historic wooden Bellevue Weir in Bakersfield (just upstream from the present-day Stockdale Highway) in 1941.

Total flow for water year 1941 was 148% of the 1894–2014 average for the Kings, 151% for the Kaweah, 172% for the Tule, and 190% for the Kern. So much water was delivered to Tulare Lake that the lake's elevation rose 12.2 feet (from elevation 184.5 to 196.7). Tulare Lake has not been that high since.

1941 Wind Event

The February 24, 1941 storm wasn't the first vigorous weather system to hit the Tulare Lake Basin during that winter. On January 8, a windstorm of almost hurricane velocity struck Garfield Grove at about 7,000 feet elevation. According to the superintendent's annual report, about 1,000 trees were blown down, including some giant sequoias 20 feet in diameter. The storm then moved north into the drainages of the East Fork and Middle Fork of the Kaweah where hundreds more trees were blown down. Cleaning this up (in crosscut saw days) created "unusual trail maintenance problems."

The national parks have experience with extensive tree failures during the winter, primarily from heavy snow loads and avalanches. Winds have also caused small-scale blowdowns. However, the January 1941 event is apparently the only large-scale blowdown to occur in the national parks in historic times.

Broadly speaking, the Southern Sierra could theoretically experience at least five categories of strong winter winds capable of causing forest blowdowns:

- Winds associated with the passage of a cold front (either with or without embedded thunderstorm cells)
- Mono winds
- Mountain waves (caused by either the passage of a cold front or a low-level jet stream that crosses the crest)
- Jet stream winds protruding from up in the stratosphere and coming down near the surface.
- Low-level barrier jet winds hitting the west slope of the Sierra, resulting in strong upslope winds.

The first category, winds associated with the passage of a cold front, includes both local thunderstorm outflow winds and the cold-front generated winds that accompany those fronts. Cold front storm systems can generate gusty winds and downdrafts that result in the blowdown of a few trees here and there. The following two examples illustrate this type of wind event:

- Frontal winds associated with a storm front (a cold front) on January 1–2, 2006 brought down the Telescope Tree and the second largest limb on the General Sherman Tree. That storm also blew down a number of other trees and power poles throughout the central and southern San Joaquin Valley. Gary Sanger at the NWS forecast office in Hanford said that this event likely was dominated by frontal winds. However, there may have been unreported embedded thunderstorms (and thundersnow) in the cold front's convective band. The National Severe Storms Laboratory has documented a few instances of microbursts associated with heavy showers that did not generate thunder. In those instances, there was apparently drying below the cloud base, and the dry air was caught in the updraft into the cumulus clouds. This caused the rapid cooling of the center of the storm, with the cold, denser air dropping toward the ground (same as a collapsing thunderstorm core). So the winds on January 1–2, 2006, may have had an isolated microburst embedded in the general wind field, but there was not enough evidence to conclusively state this. For more information on this event, see the section of this document that describes the 2005–06 floods.

- The Southern Sierra experienced four days of wind as three storms moved through Central California from January 20–23, 2012. These wind events were a combination of (1) local thunderstorm outflow winds and (2) cold-front generated winds that accompanied the fronts. The second of the three storms moved through Central California on January 21. The passage of this cold front triggered pre-dawn thunderstorms over the region, including Yosemite. There were numerous witnesses to the pre-dawn Yosemite thunderstorm on January 21. Strong winds associated with the passage of that storm caused the failure of four live trees in a small portion of Yosemite Valley shortly thereafter. Brian Mattos, Yosemite’s forester, reported that the tree failures all seemed to radiate from a point near the east end of Stoneman Meadow. One of the trees that failed was a large green ponderosa pine (dominant) which fell on a tent cabin in Curry Village, killing the concession employee inside. That employee was living there while waiting for the Badger Pass Ski Area to
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open so that he could start his winter job. He had previously served as an NPS ranger at Yosemite and at Devils Postpile NM. According to Gary Sanger, the winds that caused the blowdown were probably outflow from a collapsing thunderstorm cell. An extreme case of a collapsing cell would be a wet microburst in which the falling core acts like a piston to force the surface winds rapidly away from the collapsing cell. Gary thought that there wasn’t enough information to determine whether the Stoneman Meadow event was a microburst or a more general, less extreme, thunderstorm collapse. According to Rhett Milne, warning coordination meteorologist at the NWS in Reno, it is very rare for downburst/outflow winds of any sort to cause blowdowns, especially in the middle of winter. The cold front that caused the Stoneman Meadow downburst outflow winds produced strong winds as it moved south. At least two funnel clouds were observed over Fresno County. Winds gusted to 90 mph as measured by the RAWS automated weather station at the Isabella Dam. (The River Kern RAWS near Kernville often reports strong gusts that funnel through that part of the Kern River Canyon into the Lake Isabella area. In Gary Sanger’s opinion, the 90 mph gust was a freak event in terms of its strength.) Steve Bumgardner recalled that the wind was very strong in Lodgepole on January 21, and a few trees were blown down in that area. The Lodgepole weather station reported a thunderstorm in the distance to the east on that day. Gary Sanger thought that the winds at Lodgepole were likely frontal in nature, generated as the cold airmass behind the front pushed under the warmer air ahead of the front. The warmer airmass lifts along the frontal boundary, enhancing instability. Often cold fronts bring moisture and even flooding. However, this was a dry cold front, so it did nothing to break the dry spell that began on November 20, 2011.

The second category of strong winter winds, Mono winds, is capable of much bigger blowdowns. A Mono wind is a type of katabatic wind like the Santa Ana wind in Southern California. A katabatic wind is the technical name for a drainage wind, a wind that carries high-density air from a higher elevation down a slope under the force of gravity. Katabatic winds can rush down elevated slopes at hurricane speeds, although most are not that intense. Not all downslope winds are katabatic. The term does not include rain shadow winds where air is driven upslope on the windward side of a mountain range, drop their moisture, and descend leeward, drier and warmer (e.g., the Chinook wind that occurs along the Front Range of the Rockies).

The Mono wind originates from cold, dry air over the Great Basin. The most pronounced winds cut through a relatively low portion of the Sierra in Mono County, California, hence the name. The wind then spills out of high mountain valleys and streams down the canyons on the west slopes of the Sierra.

As the wind descends, the pressure on the air mass increases. This pressure increase causes the temperature of the air to increase. People in the path of the wind often experience a dramatic temperature increase.

Areas downwind of Mono County are subject to winds of gale force speeds (32–63 mph). Mono winds have knocked down 100-foot-tall trees and have been clocked at 100 mph in Yosemite Valley.

The classic Mono wind pattern is northeast to southwest. Research by Michael Fosberg determined that Mono winds are responsible for most of the trees blown down on the Kings River Ranger District in the Sierra National Forest. That district is located in the North Fork Kings River Basin, a drainage that trends generally northeast to southwest. Mono County is northeast of the North Fork, so that drainage is perfectly aligned for the winds that come from the Mono County area.

There are no significant breaks in the Sierra mountain wall south of Bishop, California. The high mountain crests of the Sierra and Great Western Divide generally prevent Mono winds from reaching the surface within Sequoia and Kings Canyon National Parks.

It’s instructive to look at what a powerful Mono wind event looks like in the Tulare Lake Basin. The east side of the Sierra experienced very strong winds from the NNE from November 30 – December 2, 2011. They were particularly high on the night of November 30. An automated station at the summit of Mammoth Mountain recorded 14 hours of sustained winds over 120 mph with gusts in excess of 150 mph (the limit of the anemometer). (For comparison, a sustained wind of 150 mph is equivalent to an EF3 Tornado or a Category 4 hurricane.)

Rhett Milne analyzed the event. Nothing remotely like this wind event had been recorded on Mammoth Mountain in the past 12 years; the sustained winds in this event were much higher and lasted for a much longer period. Rhett estimated that the top gusts may have been roughly 180 mph (30% greater than the sustained speed of 140 mph).
The strong winds in this event were caused by a huge high pressure system off the West Coast coupled with a huge low pressure in the Desert Southwest. This resulted in incredible pressure gradients, especially east-west across the Sierra. Much of Southern California was windy in this event, but the area around Reds Meadow and Devils Postpile National Monument had the perfect topography and NE/SW alignment to be severely impacted by this downslope windstorm. This wind event affected large sections of the San Joaquin River Basin from Tuolumne Meadows to Mt. Whitney. An estimated 5,000 trees were knocked down just in Devils Postpile National Monument alone.\textsuperscript{1181, 1182}

On the west side of the Sierra, this was perceived as a Mono wind event. Winds gusted to 45 mph at Fresno on December 1, just shy of the record gust for December of 48 mph set on December 28, 1991. Winds gusted to 60 mph at Tioga Pass. Trees were blown down at several locations, including Clovis and Mariposa. Brian Mattos recalled that Yosemite experienced multiple tree failures from elevation 4,000 feet up to over 9,000 feet during this event. Yet despite the strong winds in the Fresno / Yosemite area, the southeastern part of the San Joaquin Valley remained wind-sheltered, allowing areas of dense fog to develop during the night of November 30 – December 1.\textsuperscript{1183} Steve Bumgardner recalled that the wind was quiet in Lodgepole during this event.

Mono winds such as this occur periodically in Yosemite National Park. However, Mono winds are a generally unremarkable phenomenon in Sequoia and Kings Canyon National Parks except along the ridges and peaks. Mono winds strong enough to cause forest blowdowns don’t appear to occur south of the North Fork Kings River.

The third category of strong winter winds is mountain waves. A mountain wave is an atmospheric standing wave formed on the lee side of a mountain range when wind blows over that mountain.

The lee slope of mountains may experience strong downslope winds or many eddies of various sizes which roll down the slope. Within each wave downstream from the mountain range, a large roll eddy may be found with its axis parallel to the mountain range. Roll eddies tend to be smaller in each succeeding wave downstream. The waves downwind of the mountains are referred to as lee waves or standing waves. If sufficient moisture is present, cap clouds will form over the crest of the mountains, roll clouds will be found in the tops of the roll eddies downstream, and wave clouds will be located in the tops of the waves.

Mountain waves occur only on the lee side of mountains. They are often caused by the passage of a cold front. So in our area, mountain waves that are caused by the passage of a cold front occur only on the east side (the lee side) of the Sierra.

But Gary Sanger says that mountain waves often occur when there is a low-level jet (around 700 mb or lower) that crosses the crest. The location of the wave is dependent on the orientation of the jet, and can be on either side of the crest, whichever side is downwind (the lee side).

The Indian Wells Canyon area (15 miles northwest of Ridgecrest) is particularly prone to mountain waves when east-flowing winds funnel through Walker Pass into the Inyokern area. The Bureau of Land Management has a RAWS automated weather station located in the hills there at about 4,000 feet in elevation, downslope of Walker Pass. That station often records very high winds. The mountain waves at Indian Wells Canyon can be indicative of either a low-level jet or a cold front. All that is required is that the winds at 5,000-6,000 feet (around 700 mb or a bit lower) line up perpendicular to the crest in that location to generate standing waves downstream of the crest.

Forest blowdowns due to mountain waves touching down are an occasional occurrence in the Rockies. The largest such forest blowdown ever recorded in the Rockies was the October 25, 1997 event that blew down 20,000 acres on the west side of the Park Range northeast of Steamboat Springs, Colorado.\textsuperscript{1184}

We have not found a forest blowdown in the Sierra that was attributed to a mountain wave, but it is a theoretical possibility. A key feature of such blowdowns is that they occur in the lee of mountains, downwind of the crest.

Gary Sanger researched the Kaweah blowdown event of January 8, 1941. Bakersfield experienced a gust to 47 mph from the east-southeast on that day. (Bakersfield is roughly 100 miles SSW of where the Kaweah blowdown event occurred.) The Giant Forest weather station reported strong winds out of the southwest on January 8. From this we know that the winds were blowing upslope. This eliminates the possibility of the blowdown being caused by a mountain wave touching down. Mountain waves form only on the lee side of a mountain range, so the wind would have had to be blowing out of the east.
The fourth category of strong winter winds capable of causing forest blowdowns is when the jet stream protrudes from up in the stratosphere and comes down near the surface (aka a tropopause fold). This lowering of the jet stream does not occur very often and is most likely to be experienced in a high mountain range like the Sierra. The President’s Day Cyclone of 1979 in the Northeastern U.S. was partially attributed to the lowering of the jet stream. The huge flare-up of the 1988 Canyon Creek Fire in Montana was also attributed to the lowering of the jet stream. We have not found a forest blowdown anywhere in the U.S. that has been attributed to a lowering of the jet stream, but it is at least a theoretical possibility.

The fifth category is low-level barrier jets. A barrier jet is a jet-like wind current that forms when a low-level airflow approaches a mountain barrier and turns to the left to blow roughly parallel to the axis of that barrier. In the Sierra, that results in barrier jets that blow from the south or southwest. Despite the similarity in name, low-level barrier jets aren’t directly related to the jet stream. The airflow is upslope, barrier jets occur on the windward side of the mountain. The strongest winds tend to be elevated off the surface, but top wind speeds can reach up to 100 mph.

Gary Sanger speculated that a low-level barrier jet (around 800 mb) might have been the culprit in the 1941 wind event. Although we aren’t aware of any similar situation, Gary thinks that a southwesterly jet at around 6,000 feet elevation hitting the west slope of the Sierra might have triggered the type of strong upslope winds that occurred in the 1941 event.

This seems plausible, and most of the other explanations have been eliminated. The most likely alternative explanations would be:
- the jet stream protruding from the stratosphere and coming down near the surface
- the passage of a very powerful cold front

We don’t know enough to eliminate the jet stream from consideration. Cold fronts typically don’t produce winds nearly strong enough to cause the damage observed in the 1941 event. In any case, the records from the Giant Forest weather station don’t suggest that a particularly strong cold front passed through during the 1941 event. Looking at the temperatures, the cold front on January 4 can be seen not only in the wind shift, but also in the cold airmass behind the front. (The high on the 4th was only 36, down 13 degrees from the 3rd). For January 8, the temperatures are consistent with warm-sector precipitation, but there was neither a subsequent shift in prevailing wind direction or evidence of a cold (post-frontal) air mass.

Therefore, we’re left with a low-level barrier jet as seemingly the most likely cause. But since there are a lot of unknowns, we really cannot draw any firm conclusions. Regardless of the cause, this was a most unusual event.

1942 Flood
Flooding in 1942 occurred in January.

The winter of 1941–42 was an El Niño event. One source said that it was a strong El Niño, but that could not be confirmed. The NOAA index of El Niño / La Niña events only goes back through 1950.\(^\text{1185}\)

Flooding occurred on the Kaweah and Kern Rivers and possibly other rivers within the Tulare Lake Basin. The national parks’ records make no mention of any flooding that year.

The Kaweah had a peak average daily flow at McKay’s Point of about 11,000 cfs on February 2.\(^\text{1186}\) Visalia was flooded, though not nearly as badly as it would be in the 1955–56 flood. There was a small break in the south-bank levee on the St. Johns River near Miller’s Bridge (Fourth Ave East), northeast of Visalia.

Bakersfield was flooded in 1941, so perhaps this was a December 1941 – January 1942 flood. In any case, Highway 178 through the Kern Canyon was damaged in the 1942 flood.

Total flow for water year 1942 was 119% of the 1894–2014 average for the Kings, 116% for the Kaweah, 99% for the Tule, and 113% for the Kern.

The 1937 flood had brought Tulare Lake back to life on February 7, 1937. Subsequent floods kept the lake generally at an elevation of 190 feet or above through 1944, a level that hadn’t been seen in decades. American
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White pelicans thrived during this period. In 1942, Frances von Glahn made a color movie of them nesting at the lake (video on file in the national parks). So much water was delivered to Tulare Lake in water year 1942 that the lake’s elevation was raised 10.6 feet (from elevation 183.3 to 193.9 feet).

1943 Floods (3)
There were three floods in 1943:
1. January
2. March
3. April

The dry season of the previous year lasted until January 20. There wasn’t enough snow to ski in Giant Forest until January 31. A severe storm occurred from January 20–23, and it rained almost continuously through at least the end of the month. The storm dropped a record 20 inches of rain in Giant Forest and 8 inches at Ash Mountain.

Central and Southern California received a widespread series of storms during the last half of January. Hoegees Camp near Mt. Wilson in the San Gabriels received 26.12 inches of rain in 24 hours on January 22, setting a state record. The recurrence interval for that storm event was 11,000 years.\(^{1187}\)

There was much storm damage in the national parks. Roofs were blown off, and there were heavy landslides and washouts on the road and trail systems. The giant sequoia at Puzzle Corner fell during the first couple of days of the storm.

The Kings River at Piedra peaked on January 21: 46,900 cfs. The Kaweah River near Three Rivers peaked on January 22: 17,000 cfs.\(^{1188}\)

A second round of flooding occurred during March. The floods were concentrated in the Sierra south of the Feather River. During February, there had been occasional periods of light rain as well as warm weather conducive to melting of the mountain snowpack. By early March, the ground was moist and the river stages moderately high. There were a series of light rainstorms from March 4–8. Then heavy rains fell in the mountains and foothills on March 9–10 and March 17–18. The rains on the night of March 9 were particularly heavy and widespread in the foothills. Ten inches of rain fell in Giant Forest on top of a snow base of 29 inches. That cloudburst occurred just as the rivers were nearing crests from the earlier rains.

The Kings River passed flood stage on March 9–10 with only minor damage.

The Kaweah’s peak natural flow occurred at McKay’s Point on March 9: 17,765 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 9,714 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 7 years for the Kaweah.

Flooding on the Kaweah River caused considerable damage. The Kaweah at McKay’s Point peaked on March 9: 14,300 cfs.\(^{1189}\) The USACE said that the 1943 flood was a small but damaging flood in Visalia and elsewhere in northwestern Tulare County.\(^{1190}\)

Flooding also caused considerable damage on the Tule River. The Tule near Porterville peaked on March 9: 15,500 cfs. This was the biggest flood on that river since record-keeping began in 1901. It would remain the flood-of-record until the 1950 flood.

The White River had a major flood: 2,300 cfs. This remains the flood-of-record for that river.\(^{1191}\)

Flooding on the Kern River caused considerable damage, just as on the Kaweah and Tule.

The Kern River near Bakersfield peaked on March 9: 21,700 cfs.\(^{1192}\) This was the biggest flood on that river since record-keeping began in 1896. It would remain the flood-of-record until the 1950 flood. The Kern River was so high in this flood that it overtopped the old Olcese’s Ranch Bridge, a mile downstream from the mouth of the Kern River Canyon.

There was a major flood on Caliente Creek in April 1943, causing extensive flood damage to the Lamont/Arvin area. Presumably this was caused by an intense storm.
The overflow from these streams raised the level of Tulare Lake to near the top of the lakebed levees. Wave action caused levee breaks and the flooding of 28,000 acres. These levee breaks increased the size of Tulare Lake from 46,000 acres to 74,000 acres. By summer, 100,000 acres would be flooded.

Total flow for water year 1943 was 119% of the 1894–2014 average for the Kings, 158% for the Kaweah, 265% for the Tule, and 169% for the Kern. This was one of the very rare years when flows on the Tule River exceeded what could be used for beneficial use by the holders of water rights. Enough water was delivered to Tulare Lake in water year 1942 to raise the lake’s elevation 5.8 feet (from elevation 189.9 to 195.7 feet).

**1944 Flood**

Flooding in 1944 occurred in March.

There was a major flood on Caliente Creek in March 1944, causing extensive flood damage to the Lamont/Arvin area. Presumably this was caused by an intense storm.

Judging from historic photographs, Southern California Edison’s (SCE) Borel hydroelectric facility had a flood canal cut sometime shortly before March 6, 1944 (photograph on file in the national parks). It’s tempting to think that this was due to the same storm that caused flooding on Caliente Creek, but that isn’t known.

**1945 Floods (3)**

Flooding occurred three times in 1945:
1. January–February
2. October (twice)

The storm of January 30 – February 3 dropped a total of over 13 inches of precipitation at Giant Forest. Precipitation in the national parks consisted of more rain than snow below about 7,500 feet. At times it was apparently raining as high as 8,500 to 9,000 feet, but rain at the higher elevations was absorbed by the already good snowpack.

Visalia received 3 inches of rain during February 2–3. Ash Mountain received 6 inches during the storm. Giant Forest received 12 inches during the storm, of which 8 inches fell on the night of February 2.

The American and Sacramento Rivers flooded, as did presumably most of the rivers in the Sacramento River Basin.

The Kings River at Piedra peaked on February 2: 49,300 cfs. Parts of Centerville were inundated when the Kings flooded. There was extensive flooding farther downstream, both north and south of Hanford.

The flood was written up in a special edition of the *New York Times*. According to that account, many houses had to be evacuated in the San Joaquin Valley and farms were inundated.

The Kaweah’s peak natural flow occurred at McKay’s Point on February 2: 18,554 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 9,890 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 8 years for the Kaweah.

Visalia flooded on February 2–3. The USACE said that the 1945 flood was a small but damaging flood in Visalia and northwestern Tulare County. This was described at the time as the most severe flooding ever to hit the town. The flooding in Visalia was big enough news that troops in the South Pacific heard about it on the radio.

The national parks’ annual report made no mention of any flooding in 1945.

The North Fork Kaweah washed out the Airport Bridge, leaving that portion of the Three Rivers community cut off from the highway.
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The causes of the flooding — and the reasons that it was so severe — were described in a series of newspaper articles. The 1906 and December 1937 floods were much bigger events, but Visalia withstood those earlier floods better than it did the 1945 flood. There appear to have been three reasons for this:

1. Visalia was now a bigger and more modern town. Unlike in 1906, it was no longer built for periodic inundations.
2. The primary difference from the earlier floods was that the St. Johns River on the north side of town was no longer able to carry big floodflows. The levee along that river's south bank wasn't being maintained. The channel wasn't being kept clean of debris and trees. Islands and flood deposits had built up in the floodway. A bridge had been built across the channel at the narrowest portion, acting rather like a dam. Most important, a levee had been built on the north side of the St. Johns so that the floodwaters couldn't spread out; the river was effectively restricted to a narrow and very inadequate channel.
3. The 1945 flood came on as a sudden rolling wave instead of a gradual rise. This may have been due in large part to the St. Johns being channelized between two levees rather than being allowed to spread out across its floodplain.

Flooding of Visalia in 1945 resulted from four levee breaks on the St. Johns, all in the vicinity of Miller’s Bridge (Fourth Ave East), northeast of the city (photograph on file in the national parks). The breaks occurred about 10 p.m. on February 1, and the water reached Visalia about three hours later. The Kaweah at McKay’s Point peaked at 10:30 a.m. on February 2, and the depth of water in Visalia peaked about 6 p.m. that afternoon. Downtown Visalia was heavily damaged.

Water was 3–4 feet deep in the northeastern part of the city and more than a foot deep on some of the downtown streets (multiple photographs on file in the national parks). The current coming down Center Street was particularly strong. For the first time since 1906, a rowboat appeared on Visalia’s city streets. It was seen going down Center Street between Court and Church on the morning of February 2, bobbing along the turbulent, muddy current. It then turned up Church Street and continued on to Main.

Main Street was closed to vehicular traffic by 10:00 that morning to stop the wakes that were being thrown into adjacent businesses and homes. Similar problems were happening on nearby streets. Mrs. C.C. Bennett lived at 301 East Mineral King Ave. When she stepped outside to sweep away the water from a wave caused by a passing car, there on her porch was a small, golden-colored river fish which had come to town on the flood.

Over two-thirds of Visalia was flooded. It was a common sight to see a man pick up a lady and carry her across a flooded area. Among the many flooded areas were the Fox Theater, the Tulare County Courthouse (located on Court Street between Oak and Center), and homes on Bridge Street near where the Visalia Convention Center now stands. The flooding was so extensive that it closed TAD’s Drive-In (later renamed Mearle’s) located in what was then considered the far south side of town.

The city water supply remained safe to drink; the floodwaters only reached one well, and that well was isolated from the rest of the system. The State Division of Forestry declared Visalia an area of emergency and sent pumps to help in the city as well as bulldozers to assist with repairing the levee.

The Tule River near Porterville peaked on February 1: 12,600 cfs.

The Warthan Canyon Highway west of Coalinga (Highway 198) was closed on February 2–3, indicating that there was flooding on Warthan Creek and perhaps elsewhere on the west side of the Tulare Lake Basin.

1945 was the first year of use of the new works, built by the USACE, to keep the Kings River out of Tulare Lake. They did not work quite as designed. A break in the bypass occurred on February 3, about 20 miles south of Hanford at the height of the flood. Some ranchers were driven from their homes on the east side of the bypass and considerable grain was flooded on the west side. Nearly 1,000 people were forced to evacuate their homes.

The J.G. Boswell Co. bought the Cousins Ranch in 1946. At that time, the ranch had been under the waters of Tulare Lake for eight years, ever since the 1938 flood.

On October 6, a cloudburst dropped 2.75 inches of rain on the town of Tehachapi in 1½ hours. Rainfall intensity in the nearby mountains was evidently greater. A wall of water estimated to be eight feet high swept down Tehachapi Canyon, killing three people and causing property damage estimated to be $62,000. About half
of this damage was to property in Tehachapi. Several hundred acres of crop land around Tehachapi were heavily damaged. Several hundred feet of railroad track at Keene and near Caliente were washed out. Transportation (both rail and highway) and communication lines were shut down for 24 hours. This was presumably the same storm that caused a major flood on Caliente Creek, causing extensive flood damage to the Lamont/Arvin area.

Apparently there was a cloudburst somewhere in the Kings River Basin on October 29 or 30. The Kings River at Piedra was slightly above flood stage on October 30, but the floodwater was diverted into canals and no damage resulted. The archives at the NWS forecast office in Hanford have nothing to explain where the storm was located, so apparently it was in the mountains east of any of the reporting stations. This is presumably the same event described in the national parks’ monthly report as a late October storm.

One of these October storms is presumably the same storm that caused flashfloods and debris flows that damaged roads and the Los Angeles Aqueduct in several places on the east side of the Sierra in October.

Table 44 gives the precipitation totals for the reporting stations in the national parks.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Grant Grove</td>
<td>4.66</td>
</tr>
<tr>
<td>Giant Forest</td>
<td>4.42</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>1.37</td>
</tr>
</tbody>
</table>

St. Johns Levee — Condition in 1945

At the time that Visalia was founded in 1852, the flow of the Kaweah River was distributed largely through four channels that flowed along the south side of the Kaweah Delta. That all changed thanks to the huge 1861–62 flood and the even bigger 1867–68 flood. One of the legacies of those floods was the creation of the St. Johns River which routed the majority of the Kaweah River floodwaters along the north side of the delta, to the north of Visalia.

This resulted in periodic flooding of Visalia and of the lowlands between the new river and the town. A levee was soon built along the south bank of the St. Johns to protect the town. It didn’t take long for that levee to fail in a flood. The first record that we have of failure of the south-bank levee was in the 1877 flood. It would not be the last.

In 1891, Levee Land District, No. 1 was formed for the purpose of building a levee along the south bank of the St. Johns River north of Visalia to protect the town and the intervening territory from the nearly annual flood hazard. This was going to be a much larger levee than the original one that had been in place and failing. In her history of Tulare County, Kathleen Small said that this was the largest protective measure enacted by the people of the county in the early decades. Upon completion of the levee, the Visalia Delta commented as follows:

"Few people in the city know that a barrier has been erected between the river and Visalia to protect us from the winter floods, and fewer people still have any conception of the magnitude of the undertaking, or the manner in which it was prosecuted. Not a moment was lost. Work was commenced on November 2, and was ended last Saturday, December 5. During that time a small army of men were at work on the embankment during the period mentioned.

On last Saturday the six and a half mile levee, commencing near Cutler’s Bridge and ending at Burrel’s Bridge, was finished. The work was done quickly but thoroughly, and the levee resembles the grade of a new railroad line. It is from ten to twelve feet wide on top and about thirty feet wide at the base. The top of the levee makes a good driveway, although it will not be used for that purpose.

The construction of this levee means nothing less than absolute protection to Visalia by flood. The levee that has been built is strong enough to resist the force of the water. The embankment is three feet higher than the highest water mark. It will be seen, therefore, that the city is guarded from all possible damage from floods. While the levee will protect the people of Visalia from floods, the greatest good will be realized by the land owners between this city and the river, whose property has been subject to inundation for years past."
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This very large 6½-mile-long levee was built in just 33 days at a cost of only $7,210.12. One of the main reasons that the levee could be built so quickly and cheaply was that it was built out of local materials, essentially river sediments. Although it looked very impressive, the levee had relatively little structural integrity for its size. Repairs over the years have generally been made using the same type of materials and engineering.

At the same time that the south-bank levee was constructed, the landowners on the other side of the St. Johns built a two-mile-long levee to protect their land from the floodwaters which would be diverted to the north of the river by the new south-bank levee.

The responsibility for maintenance and repair of the south-bank levee on the St. Johns has generally remained with the local levee district since its construction. Only in emergency situations have the state and federal governments stepped in to assist with levee repairs.

Flooding in the early years (roughly 1875–1940) put only moderate pressure on the levee and caused only moderate flooding problems. That was partly because Visalia was built for periodic inundations. But it was also because the St. Johns had a wide floodplain north of town. Since the floodwaters spread out in a shallow sheet across this floodplain, it put relatively low pressure on the south-bank levee.

But by 1945, the situation had changed dramatically. In the 1945 flood, the south-bank levee failed in four places, and Visalia flooded like never before in its history.

On February 7, 1945, the Visalia Chamber of Commerce hosted a meeting of all the concerned parties to assess the condition of the levees on the St. Johns and determine what could be done to bring them back up to minimum standard. The meeting amounted to a fairly thorough after-action review. It was quickly realized that the problem was bigger than a simple levee repair. It was determined that the channel of the St. Johns in the vicinity of Miller’s Bridge (Fourth Ave East), northeast of the city, was now so deteriorated and constrained that it could only pass 3,000–4,000 cfs. The St. Johns had peaked on February 2: 14,900 cfs.

It wasn’t immediately obvious that channel capacity could be restored or that the south-bank levee could be upgraded sufficiently to protect Visalia in the event of another flood of similar magnitude. The St. Johns River is subject to floods that are much larger than the 1945 flood. The city and county faced four challenges:
1. Remove Miller’s Bridge so that it didn’t act as a constriction on the St. Johns.
2. Restore the natural floodplain and channel of the St. Johns so that it could handle projected floodflows.
3. Strengthen the south-bank levee of the St. Johns, correcting its obvious deficiencies.
4. After restoring the channel and levee, find some way to keep them adequately maintained.

St. Johns Levee — Condition in Recent Years
Over six decades have passed since that after-action review, but there are still major concerns about the ability of the St. Johns channel to safely pass floodwaters. The south-bank levee was partially rebuilt after the 1945 flood, but a portion of it failed again in the December 1955 flood. Using federal emergency funding, the USACE quickly rebuilt the damaged levee. However, federal law required that agency to restore the levee to the same condition that it was in prior to the flood, not fix its obvious deficiencies.

The 2005–06 Tulare County Civil Grand Jury investigated the St. Johns levee and found that it was not constructed to USACE certification standards, it was not being adequately maintained, and there was no adequate source of funds for its maintenance. The grand jury found that after passage of Proposition 13, incoming taxes were insufficient to maintain the levee.

The Tulare County Resource Management Agency surveyed property owners in the levee district in 2002, but those owners were generally uninterested in levee maintenance and did not want to put more of their tax dollars into maintenance. Because the south-bank levee was in such bad shape, $17 million was then needed to bring it up to USACE certification standards. However, no source for those funds has yet been found, and the levee is in approximately the same shape now that it was in when the grand jury assessed the situation.

The National Flood Insurance Program (NFIP) is administered by the Federal Emergency Management Agency (FEMA) to offer flood insurance to properties located in special flood hazard areas (SFHAs). In order to qualify for flood insurance, a community must join the NFIP and agree to enforce sound floodplain management standards.

FEMA is currently updating and modernizing the nation’s Flood Insurance Rate Maps. These digital flood hazard maps provide an official depiction of flood hazards for each community and for properties located within it. In
June 2009, FEMA found that the levee along the south bank of the St. Johns was in such bad condition that it provided essentially no reliable flood protection for Visalia.

On June 16, 2009, Tulare County adopted the new Digital Flood Insurance Rate Maps as part of this FEMA project. This resulted in several thousand properties being moved into the SFHA while some properties moved out of the SFHA.\textsuperscript{1212} A total of about 8,900 additional homes and other properties were designated as being at high risk for flooding. This resulted in some of those homeowners having to purchase flood insurance at relatively expensive rates.

No one disputes the finding by the grand jury and FEMA that the St. Johns levee is poorly maintained and that the levee falls far short of USACE certification standards. The problem is that local agencies and districts don’t have the funds to maintain the levee, let alone bring it up to standard.

On February 3, 2011, 27 U.S. senators sent a letter to FEMA requesting that the agency recognize and try to quantify the degree of protection provided by poorly built and poorly maintained levees such as the St. Johns’ when determining flood risk. The reason given by the senators was that FEMA’s current policy puts American jobs at risk.

1947–50 Drought
This drought is usually referred to as the drought of 1943–51 because it was active somewhere in the state from 1943–51. It affected the entire state from 1947–49. However, as reflected in Table 45, this drought was primarily active in the San Joaquin River Basin from 1947–50.

Precipitation dropped below average in water year 1946, but drought conditions didn’t really set in until 1947. Rivers within the Tulare Lake Basin had about average or near-average flows for the first four water years of the state drought (1943–46). Then the drought moved into our basin. From 1947–50, flows in the Kings and Kaweah Rivers were generally 55%–75% of average. Total flows in the Tule and Kern were generally 35%–60% during that period. So from the standpoint of the Tulare Lake Basin, it makes sense to think of this as a four-year drought, lasting from 1947–50.

The drought was most severe in central and southern coastal areas, where accumulated deficiencies in runoff approached, and in some instances exceeded, those of the drought of 1929–34. Water year 1951 ranks as the driest of record at several gaging stations in Southern Coastal California.

Recurrence intervals for the drought of 1943–51 were about 20 years in the Central and Northern Sierra because of the short duration there. They were about 20–80 years in the rest of the state, where this drought is exceeded in duration and severity only by the drought of 1929–34.

The historical record of the Sacramento Valley Water Year Index indicates that the drought of 1943–51 (recurrence interval of 55 years) ranks second only to the drought of 1929–34.

Table 45 shows how the San Joaquin Valley Water Year Index categorized the drought years in that basin. It also shows total runoff for the four major rivers in our basin (Kings, Kaweah, Tule, and Kern) combined during this drought.

<table>
<thead>
<tr>
<th>Water Year</th>
<th>San Joaquin River Basin</th>
<th>Tulare Lake Basin</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Water Year Classification</td>
<td>Total Runoff (acre-feet)</td>
</tr>
<tr>
<td>1944</td>
<td>Below normal</td>
<td>2,196,030</td>
</tr>
<tr>
<td>1945</td>
<td>Above normal</td>
<td>3,644,000</td>
</tr>
<tr>
<td>1946</td>
<td>Above normal</td>
<td>2,736,810</td>
</tr>
<tr>
<td>1947</td>
<td>Dry</td>
<td>1,867,230</td>
</tr>
<tr>
<td>1948</td>
<td>Below normal</td>
<td>1,654,660</td>
</tr>
<tr>
<td>1949</td>
<td>Below normal</td>
<td>1,521,660</td>
</tr>
<tr>
<td>1950</td>
<td>Below normal</td>
<td>2,082,720</td>
</tr>
<tr>
<td>1951</td>
<td>Above normal</td>
<td>2,718,390</td>
</tr>
<tr>
<td>Drought average (1947–50)</td>
<td>1,781,568</td>
<td>61%</td>
</tr>
</tbody>
</table>

1947 was the driest year ever in Paso Robles, with only four inches of rain.
**Floods and Droughts in the Tulare Lake Basin**

**Specific Floods and Droughts**

After the 1937 flood, the elevation of Tulare Lake was relatively stable for the next 10 years. Then, thanks in large part to the 1947–50 drought, the lake was dry from July 17, 1946, to November 19, 1950.

The Simpson Meadow Fire occurred in 1948. This was the second-largest fire in the history of the national parks, burning 11,100 acres. Only the 1926 Kaweah Fire was larger. The national parks’ three largest fires (Kaweah, Simpson Meadow, and 1977 Ferguson) have all occurred during droughts).

In the national parks, the drought was viewed as beginning in the early winter of 1946–47. Drought conditions of record-breaking proportions prevailed over the parks through February 1948. Livestock by the thousands were shipped from the parched ranges of Central California. Demands arose to open the national parks to commercial grazing as one method of relief for stock growers. Fortunately, above-average rain and snow fell during March, April, and May, 1948.

Table 46 shows total runoff for the four major rivers (Kings, Kaweah, Tule, and Kern) during the 1947–50 drought.

<table>
<thead>
<tr>
<th>Water Year</th>
<th>Total Runoff (acre-feet)</th>
<th>% of average (1894–2014)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1947</td>
<td>1,867,230</td>
<td>63%</td>
</tr>
<tr>
<td>1948</td>
<td>1,654,660</td>
<td>56%</td>
</tr>
<tr>
<td>1949</td>
<td>1,521,660</td>
<td>52%</td>
</tr>
<tr>
<td>1950</td>
<td>2,082,720</td>
<td>71%</td>
</tr>
<tr>
<td>Average</td>
<td>1,781,568</td>
<td>61%</td>
</tr>
</tbody>
</table>

Sequoia National Park’s district ranger Clarence Fry took measurements on July 10, 1949, at three locations along the Middle Fork Kaweah between Potwisha and Ash Mountain. He found only seepage between pools, not enough for fish to swim through or pass over.

By 1951, all the rivers of the Tulare Lake Basin — except the Kern — had returned to approximately average flows.

**1949 Flood**

Flooding in 1949 occurred in March.

On March 7, two thunderstorms hit Bakersfield in the same day. They unleashed heavy downpours that flooded first floors of offices and damaged house foundations, as well as inundating landscaping, streets and storm drains.¹²¹³

**1950 Flood**

Flooding in 1950 occurred in November –December.

The winter of 1950–51 was a weak La Niña event. This association with the 1950 flood was probably a coincidence. Only strong La Niña events have been shown to have any correlation with high precipitation events and floods in California.

The flood affected the Sacramento and San Joaquin Valleys. It had a recurrence interval of 80 years on some rivers. Flooding in the Tulare Lakebed continued for a few months after river flooding subsided.

October 1950 was the wettest October since 1899, with a precipitation total 300% of average. The storms of November and December were general over all of California north of the Tehachapis. Heavy rain fell on both the Coast Ranges and Sierra. However, the flooding was limited to the Central Valley.

The flooding was caused by a series of storms which brought exceptionally warm, moisture-laden air against the Sierra, resulting in intense rainfall instead of snowfall at unusually high elevations. There were four distinct storms:¹²¹⁴

1. November 13–15
2. November 16–21 (this was the big one)
3. December 2–4
4. December 6–8

In the Tulare Lake Basin, most of the precipitation fell as rain. Because it was early in the season, there was relatively little snowpack in place. Rain was relatively light in the valley but heavy in the foothills and Sierra.

The first storm began in the national parks on November 13.

A high-elevation storm passed through Central California on November 18–19, 1950. The rainfall distribution in this storm was quite similar to the January 30 – February 1, 1963 storm. The rainfall in both of these storms was heavy in the coastal mountains as well as in the Sierra. The 1963 storm affected areas south of the wetter zone of the November 18, 1950 storm. The greatest daily total rain for the 1950 storm was 13.16 inches at Giant Forest. Highest-ever daily rainfalls were reported at 30 stations. Nine stations reported rainfall totals in excess of a storm with a recurrence interval of 1,000 years. Seven of these were in the Stanislaus, Merced and San Joaquin River Basins. Highway 140 into Yosemite was washed out near El Portal. Extensive flooding was reported on the lower San Joaquin River. Calaveras Dam in Alameda County received 7.17 inches in one day, which was 33% of its annual average rainfall. The recurrence interval for that event was 23,000 years.\(^{1215}\)

In Three Rivers, the rain was continuous for 20 hours. Long-time residents of that community could not recall such a heavy downpour.\(^{1216}\)

Three days of heavy rain from November 17–19 in the Sierra brought more than 15 inches of rain to some areas as high as 5,500 feet elevation and heavy rain as high as 10,000 feet, which melted the small snowpack. Although the rain was heavy and continuous, the greatest recorded intensity was 0.9 inches per hour at Giant Forest on November 18.\(^{1217, 1218}\)

Table 47 gives the precipitation totals for the reporting stations in the Kaweah River Basin.\(^{1219}\)

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Giant Forest</td>
<td>15.22</td>
</tr>
<tr>
<td>Three Rivers</td>
<td>7.24</td>
</tr>
</tbody>
</table>

As shown in Table 48, this storm broke precipitation records that were set in the 1920s and 1940s.\(^{1220}\)

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>November 1950</th>
<th>All Novembers during period of record prior to 1950</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Total for month (inches)</td>
<td>Greatest daily total (inches)</td>
</tr>
<tr>
<td>Grant Gove</td>
<td>14.51</td>
<td>6.12</td>
</tr>
<tr>
<td>Giant Forest</td>
<td>18.87</td>
<td>9.55</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>6.13</td>
<td>3.19</td>
</tr>
<tr>
<td>Three Rivers</td>
<td>7.86</td>
<td>4.10</td>
</tr>
<tr>
<td>Springville</td>
<td>20.50</td>
<td>10.27</td>
</tr>
</tbody>
</table>

The 1950 flood was so newsworthy that it was written up in at least two issues of the New York Times. The first story described the flooding that was occurring from Sacramento to Bakersfield, including on the Kings and Kern Rivers.\(^{1221}\) The second article also included the flooding on the Tule.\(^{1222}\) Flooding in Three Rivers, Woodlake, and Visalia was featured in a special issue of the Exeter Sun (now The Foothills Sun-Gazette).

The 1950 flood caused major damage in the Central Valley. Damage was estimated to be 33 million dollars, 669,400 acres were flooded, and two people died. The $33 million included $1.2 million in damage to Highway 140 (the All-Year Highway to Yosemite Valley) near El Portal in the Merced River Canyon and $509,000 in damages to park infrastructure within Yosemite. The hardest hit areas in the San Joaquin Valley were Merced, Chowchilla, Centerville, Visalia, Porterville, Oildale, Isabella, and Kernville. Approximately 25,000 people were evacuated from their homes during the entire flood period. Governor Earl Warren proclaimed a state of emergency on November 21.\(^{1223}\)
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Flood crests on the Kings, Kaweah, Tule, and Kern Rivers exceeded all previous records. At the time, this was the biggest flood to occur on those rivers since the 1867–68 flood, an event that occurred before the onset of formal record-keeping. The 1867–68 flood remains the biggest flood to have occurred in historic times in the Tulare Lake Basin.

The Kings River at Piedra peaked at 3:00 a.m. on November 19: 91,000 cfs. That was the highest flow since record-keeping began in 1895.

A recording stream gage was installed in the South Fork Kings from 1950–1957. (See the section of this document that describes the Stream Gages on the South Fork Kings.) Either because of good planning or incredible good luck, that gage began operation on November 16, 1950 just as the second and biggest of the 1950 storms struck.

Thanks to that stream gage, we know that the South Fork Kings peaked at 2:00 a.m. on November 19, 1950: 10,000 cfs. That puts it in a category with other Cedar Grove floods of the past 70 years (1937, 1950, 1955, 1966, 1969, 1978, 1982, 1983, and 1997) that rise to the level of the modeled 50-year flood: that is, a flood event that occurs about every eight years on average. See the section of this document that describes Cedar Grove Flooding.

Work had begun on the Pine Flat Dam in 1947. The dam would be partially operational in 1952, but it wouldn't be completed until 1954. On November 19, 1950 the Kings washed out the newly completed weir, the cofferdam, and foundation work of the dam. It also destroyed a 300-person motion photograph theater that was closed on the day of the flood. Damage at the dam site totaled $900,000; that was the most costly single item of property damaged by the 1950 flood.

Thirty-five people living on an island at Piedra were marooned by the sudden rise of the river, but were saved by rescue workers.

From Piedra to Highway 99 (immediately south of Kingsburg), about 17,000 acres of agricultural land was flooded; most of that land was in Centerville Bottoms. The lower Kings River Bridge in Reedley was washed out. About 500 families were forced to evacuate their homes. Loss of livestock in the area was especially severe. About 30,000 turkeys, valued at $500,000, were lost.

Downstream from Highway 99, the Kings River inundated the following areas during the November 19–21 flood:

- 36,100 acres between Highway 99 and the Crescent Weir
- 13,100 acres between the Crescent Weir and the San Joaquin River along the north distributaries (Fresno Slough and James Bypass)
- 3,000 acres along the south distributaries on the way to Tulare Lake

As the result of breaks in the river levees at numerous points near Laton, that community was virtually surrounded by the floodwaters. Farther downstream, the floods encroached upon Riverdale and fringe communities near Hanford.

A second, somewhat smaller, flood came down the Kings on December 4–6.

This was the last major uncontrolled Kings River flood event.

The flood caused significant damage to roads and trails in Sequoia, Kings Canyon, and Yosemite National Parks.

Flooding occurred in Sequoia and Kings Canyon National Parks from November 18 – December 8. According to the parks’ monthly report, it was not possible to record the maximum high-water level on the Kaweah, as the Potwisha gaging station was under several feet of water. Presumably this was referring to the Marble Fork of the Kaweah.

Rain continued to fall until December 8, causing additional damage by slides and washouts. Slide and washout conditions prevented the use of heavy equipment to keep drainage channels open. In some instances it was necessary to use dynamite to dislodge jams of drifted material endangering structures.
Once upon a time, before there was a Lodgepole Campground as we now know it, there was a big construction operation to complete the Generals Highway from Giant Forest to Grant Grove. As part of that project, the surfacing contractor needed a large borrow pit. It was determined that the best place to locate that pit was in the Marble Fork of the Kaweah; about 200 yards downstream from the bridge we now know as “Log Bridge”. There was an unlimited supply of aggregate in the riverbed.

During highway construction, the river was pushed to the far north side of its channel. The material that came out of the pit had to be processed, plus there had to be an asphalt plant. That operation took place on the south side of the river; that is the location of the present-day big parking lot in the middle of the Lodgepole Campground.

Once the surfacing contractor was through with the river borrow pit, they used their drag line to clean it up in 1934. Then the national park, using CCC labor, made the river borrow pit into a managed recreation area for the Lodgepole campers. The swimming pool was located in the riverbed, adjacent to the old campers’ market (the now defunct Walter Fry Nature Center building). There was no significant dam as far as we know; just a deep pool, a sandy beach, and a diving board.

The pool gradually silted in, but may have survived in some form until 1950. The 1950 flood was likely the end of the Lodgepole swimming pool. You can still see the rows of rock on the south bank of the river that led to the diving board.

Flooding in the national parks was so extensive that personnel were sent from the NPS Washington office, the regional office, and from the Bureau of Public Roads to assist with the damage assessment. There were numerous washouts, including one on the Generals Highway 1½ miles east of the Wye (apparently at Mill Creek). Over 5,000 cubic yards of slide material had to be removed from roads.

The 1950 flood did extensive damage to the national parks’ trails. It appears that the damage to trail bridges — and therefore the extent of flooding — was much worse in the Kaweah River Basin than in the Kings or the Kern River Basins. Some of the major bridge trusses that were washed out were Castle Creek, Middle Fork Kaweah in River Valley (on Route 70, the trail from Bearpaw to Redwood Meadow), Paradise Creek (near Buckeye Flat Campground), East Fork Kaweah (below Atwell Campground), and Clough Cave (near South Fork Campground).

Bob Meadow’s research found that the 1950 flood also destroyed the following bridges in the Bearpaw area: Buck Canyon, Lone Pine Creek, upper and lower trail crossings of Granite Creek, Middle Fork Trail below Redwood Meadow, Lower Buck Canyon, and across the Middle Fork on the Castle Creek Trail.

One source said that this was the flood that washed out the Hospital Rock Bridge. The 1950 flood did damage this bridge and wash out the abutments, but the bridge survived. It is not clear what flood finally destroyed the bridge; perhaps it was the 1955 or 1966 flood. The piers for this bridge were beautiful, as was the bridge itself (photograph on file in the national parks). Those piers were largely demolished by the national parks in about 1974.

The 1949 trails inventory still showed the Board Camp Dome Trail (Route 98, originally constructed as the Hockett Trail) running up the South Fork of the Kaweah past present-day Ladybug Camp to the Hockett Meadow. (Possibly that section of trail was later called the Stakecamp Dome Trail.) The bridge on that trail crossed the South Fork Kaweah just north of Garfield Creek and was one of the few trail bridges in the Kaweah River Basin that survived the 1950 flood. Bill Tweed recalled crossing it in 1960. That bridge would eventually be washed out in the January 1969 flood and was never replaced.

The national parks’ Ash Mountain headquarters development (originally known as Alder Creek) began operation in 1921. The development originally obtained its water solely from Alder Creek. Starting in 1939, there was an extended period of study and debate about how to obtain a better water supply for the Ash Mountain development. Finally, in fiscal year 1950, 4,900 feet of 3-inch pipe was laid to tap into Southern California Edison’s Kaweah #3 flume at Milk Creek.

Shortly after the flume-tapping project was completed, the pipe and pump were damaged during the November 1950 flood. The system was repaired and returned to operation in 1951. The system seems vulnerable to floods because it had to cross the Middle Fork Kaweah River. It’s tempting to think that it might have been damaged again in the big floods of 1955 and 1966. In any case, it was eventually replaced with a system that pumped water directly out of the Kaweah River. That river pump would survive until the 1997 flood.
The peak of the 1950 flood began coming through Three Rivers late on the night of November 18 and continued rising into the early morning hours of the 19th. (One source incorrectly said that it occurred on the night of November 20.)

The North Fork Kaweah peaked at 1:00 a.m. on November 19: 9,150 cfs. That was the highest flow on that river since record-keeping began in 1910. (In the December 1955 and December 1966 floods, the North Fork would experience flows over twice this great.)

The scene was particularly dramatic at Archie and Mary McDowall’s chicken ranch up the North Fork, just above the Kaweah Post Office. (Mary was a teacher at the Three Rivers school for 10 years, from 1930–39. She then served as principal of the school for another 29 years, from 1940–68. McDowall Auditorium is named for her.) Their daughter Bobbie was a junior in high school at the time of the 1950 flood.

When the downpour hit Three Rivers on November 18, Archie realized that it was also hitting the mountains up in the national parks. Even though there wasn’t much of a snowpack, that meant there would be floodwaters arriving at their ranch that night. He thought that the North Fork wouldn’t be affected as much as the mainstem of the Kaweah. All the same, Archie knew that they would be having flooding on their property. Therefore, he and Bobbie went out into the night to move chickens in a low part of their property to a chicken house on higher ground. Sixty years later, Bobbie still has a vivid memory of that night.

The McDowall ranch was ½ mile downstream from the Upper North Fork Bridge (about three miles up the North Fork from Three Rivers). Unknown to Archie and Bobbie, debris (trees, brush, etc.) was building up against the upper side of that bridge. The bridge gave way from the power of the rushing floodwaters against that debris. Archie and Bobbie found themselves in the path of that onrushing wall of water. They were carried downstream to the far side of their property until a barbed wire fence stopped them.

Bobbie had on cape-type rain gear that kept pulling her under. Her dad yelled for her to get that garment off because it kept pulling her under and keeping her from getting firm hold of the fence. Her dad never cussed. But his colorful choice of words on that occasion — words that were unusual for him — sank in. Bobbie was able to get her garment over her head and reach for his hand. He was hanging onto a tree. Archie was able to get to that tree just as he went under for the second time. (Seeing her dad go down was the most frightening part of the whole experience for Bobbie.) How he got back to that tree has always been a miracle to her.

After the two of them got their adrenalin in another gear, they were able to get to a slower portion of the floodwaters and gradually work their way over to the North Fork Road. That took a long time. However, they still hadn’t reached dry land. They found the water running so deep on the road by the Kaweah Post Office that they could hardly walk even there.

The next morning they went back to check on the damage. The fence that they had held onto the previous night was gone; much of the property including a chicken house full of chickens was gone. Archie's pickup truck and much more had also disappeared.

Specific damage in Three Rivers included:

- Most of the bridges in town suffered major structural damage or had their approaches washed away. Only the Pumpkin Hollow Bridge, the Dinely Bridge, and the North Fork Bridge (the one next to the Three Rivers Market that had been built in 1938) remained passable.
- The Upper North Fork Bridge washed out. The Airport Bridge survived, but the approaches were washed out. The North Fork roadbed was badly eroded down to bedrock.
- At least the first three bridges on the South Fork Road (Conley Creek and the first two bridges over the South Fork of the Kaweah) were washed out. Parts of the roadbed were badly eroded when the river rerouted. Huge slabs of asphalt, some as long as 25 feet, were ripped and thrown up by the force of water.
- The Kaweah Park (presumably located at the junction of the South Fork and the mainstem of the Kaweah River) was extensively damaged.
- On the morning of November 19, crowds gathered at the South Fork crossing of Highway 198 (now Cherokee Oaks Drive) to watch that raging river pound and crumble the one highway access for the entire community. By the end of the day, the bridge was knocked out and Three Rivers was cut off.
- At least five homes washed away and many others were damaged or undermined.
- Southern California Edison’s (SCE) #2 flume was heavily damaged: 200 feet of the canal bank was washed away and the upper end was filled with mud. That affected not only power production but left many families
The November 1950 flood did even more damage to the Kaweah Hatchery than the 1937 flood had. The hatchery was shifted off its foundation. Equipment in the interior was greatly disarranged. Pumps, motors, and the entire grounds were covered with tons of sand and debris. This time, the movable property was repaired and transferred to other installations, and the hatchery was permanently closed.

So many water systems were contaminated that the county health department set up a program to inoculate all Three River residents against typhoid.

The mainstem of the Kaweah in Three Rivers peaked in the pre-dawn hours of November 19, 1950. The Three Rivers gaging station (USGS gage #11-2105 Kaweah R. nr Three Rivers) recorded 45,700 cfs at 1:30 a.m. before that gage was destroyed. The 1950 flood was the largest flood on that stretch of the river with respect to both peak and volume since stream gaging began in 1903. It would remain the flood-of-record until the December 1955 flood. The Three Rivers gaging station was rebuilt after the 1950 flood. It was located just above the junction of the Kaweah with Horse Creek. The Three Rivers gaging station would continue operation until 1961 when this site was submerged under Lake Kaweah.

The Kaweah's peak natural flow occurred at McKay’s Point at 3:30 a.m. on November 19: 54,332 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 16,640 cfs.) The 54,332 cfs figure was apparently determined by sloped area measurement: the flood destroyed the gaging station.

Based on the flood exceedence rates in Table 29, this had a recurrence interval of 13 years for the Kaweah. It would have had a recurrence interval of 30 years if calculated using the 54,332 cfs peak flow. (One source reportedly calculated this as having a recurrence interval of 125 years. That result could not be reproduced with either the peak or the daily flows, even when using the now-outdated 1971 flood frequency curves.)

By November 21, the Kaweah River flows were dropping and the danger of flooding in Three Rivers had ended. One source said that eight bridges in the town had been so badly damaged that they remained impassable.

In addition to losing the only highway bridge in and out of town, families on the North Fork were isolated from the main part of town because the Upper North Fork Bridge was gone. More than a dozen families on the South Fork Road were also isolated from the main part of town until the three bridges on that road could be replaced.

One of the bigger challenges was how to get feed to Bob Lewis’s turkey ranch up on the South Fork Road. The 10,000 birds in his flock required bringing in 28 bags of feed every day. By November 20, crews of volunteers had made a plank crossing of Conley Creek, the first washed-out bridge. Men carried the sacks of feed on their backs over the plank crossing, and were met by a private weapons carrier which took the feed over the second washed-out bridge and across the next section of partially washed-out road. At the third washed-out bridge, neighbors had erected a pulley system and the feed bags were taken across in pack animal panniers. People were also being transported in those same panniers.

Upstream from the foothill line, the Kaweah and its tributaries destroyed a total of 7 highway bridges, damaged an extensive stretch of highway, and destroyed or damaged 25 homes.

Terminus Beach was severely damaged, as were two gravel plants and a number of homes in the area.

On November 19, debris carried by the Kaweah lodged against the Visalia Electric trestle near McKay’s Point. This created a jetty, diverting the floodwaters toward Woodlake. A total of about 50 homes in that community were flooded. Six homes were destroyed, and others were extensively damaged. The trestle eventually collapsed, resulting in the destruction of several thousand feet of railroad track and embankment. (A similar event would happen at this trestle in the December 1955 flood.)

The USACE said that the 1950 flood was a large and extremely damaging rain-flood in northwestern Tulare County. It was larger than the 1937 flood or any other flood since the turn of the century. The flooded area was from 2–4 miles in width from Woodlake to Visalia.

It caused serious flooding in Visalia and other valley-floor communities; extensive damage to streets, roads, and bridges; and tremendous agricultural damage. Mill Creek overflowed in Visalia, resulting in extensive flooding of the business section. Sandbags were used to protect residential areas, but it was necessary to evacuate about 100 people from their homes.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Wide areas of cotton, pasture, and grain were inundated along the St. Johns River and Cross Creek and along various distributary channels and canals south of Visalia. Flood damage was generally characterized as extensive. No lives were lost, but many houses and automobiles were swept away. Roads and bridges were extensively damaged and floodwaters extended 3–4 miles wide to the northwest of Visalia. Mill Creek caused serious flooding in Visalia, but not as bad as in the 1945 flood. A lake formed on E. Main St. extending east from Santa Fe St. The water averaged 6–12 inches in depth, although in several places it was 18 inches deep. The St. Johns River flooded extensive tracts of agricultural land on the Kaweah Delta north of Visalia. Wide areas of agricultural lands were flooded south and east of Visalia along the St. Johns River and Cross Creek. The total area flooded by the Kaweah River was about 48,000 acres. Approximately 200 people were evacuated in Woodlake, and 2,000 people were evacuated in Visalia. Damage in the valley reached $5 million, including $500,000 in damage to bridges. The Tule River near Springville (see Station 2032 in Table 63): at 22,400 cfs. This gage had just been installed. This flood would remain the flood-of-record until the December 1966 flood. The South Fork Tule River near Success (see Station 2045 in Table 63) peaked at 7,000 cfs. This was the greatest flow at this site since record-keeping began in 1930. It would remain the flood-of-record until the December 1966 flood. One source said that the Tule River near Porterville peaked at 4:30 a.m. on November 19: 25,500 cfs. Another source said that the Tule peaked at 32,000 cfs. In either case, that would be the highest flow on that river since record-keeping began in 1901. It would remain the flood-of-record until the December 1966 flood. The flooding in Porterville was shallow and was largely confined to a small portion of the residential area. Between Porterville and Highway 99, the Tule spread over agricultural areas to a width of 3–4 miles. The United Concrete Corporation plant on Highway 99 was heavily damaged and was closed for two weeks. Roads and bridges suffered severe damage throughout the Tule River Basin. The total area flooded by the Tule was about 32,000 acres. The North Fork of the Kern River at Kernville (see Station 1860 in Table 63) peaked in November 1950: 27,400 cfs. This was the greatest discharge on that river since record-keeping began in 1912. (This record would be matched in the December 1955 flood. The December 1966 flood would have a discharge more than twice as great (60,000 cfs.) Floodwaters covered portions of the town of Kernville and most of the town of Isabella and forced a mass evacuation of about 1,000 inhabitants. The mainstem of the Kern at the site of the future Isabella Dam (see Station 1910 in Table 63) peaked at 39,000 cfs. This would remain the flood-of-record until the December 1966 flood. That flood would have a computed discharge 2.5 times as great (96,900 cfs) as the 1950 flood. Upstream from the head of the lower canyon near the Isabella Dam site, the river flooded SCE’s Kern No. 3 power plant. It almost completely destroyed the State of California’s fish hatchery. It also inundated summer homes, commercial recreation developments, and USFS recreational developments. Isabella Dam was under construction when the flood occurred and equipment being used to build the dam was flooded or buried, including a big power shovel. Most or all of the damage happened at the lower tunnel. In the lower canyon, floodwaters damaged several commercial recreational establishments, two hydroelectric facilities (SCE’s Kern River #1 and #3), USFS recreational facilities, and the state highway. Five bridges were washed out in the Bakersfield area, including the Kernville Bridge. The flood also washed out the old Olcese’s Ranch Bridge, a mile downstream from the mouth of the Kern River Canyon.
The Kern River near Bakersfield peaked at 4:30 p.m. on November 19: 36,000 cfs. This was the highest flow recorded at that gage since record-keeping began in 1896. The south bank levees protecting Bakersfield almost failed. Heroic efforts by 500 volunteers supported by heavy equipment saved the city from inundation. A similar emergency effort had been required on this levee in the 1937 flood.

The south bank of the Kern wasn’t the only trouble spot during the November flood. Floodwaters behind the historic wooden Bellevue Weir just upstream from the present-day Stockdale Highway Bridge were threatening to overflow on the north bank into the Goose Lake / Jerry Slough system. Historically, floodwaters from the Kern would regularly flow through that slough on their way to Goose Lake (see Figure 14), but that had apparently not happened for a while. If that were to happen now, farms through Rosedale and even 20 miles farther west would be inundated. Given the time available, there was apparently no feasible way to contain the Kern along its north bank. Instead, private interests used heavy equipment to hurriedly throw up a levee to minimize the flooding farther downstream along the slough in the vicinity of the Lerdo Highway. They finished the levee just before the floodwaters arrived. That levee would be tested again in the 1952 flood.

A large portion of the Kern floodwaters entered the Goose Lake / Jerry Slough system, and residents of the Rosedale and Stockdale areas had to be evacuated. Flow was continuous in the Goose Lake Slough until December 7 or 8. The Goose Slough Bridge on Highway 139 was washed away. A total of about 18,500 acres of the Goose Lake / Jerry Slough system was flooded. None of the floodwaters extended past Goose Lake.

The rest of the Kern River floodwaters continued in the main channel to Buena Vista Lake. The total inundated area on the Kern River (including the 18,500 acres of the Goose Lake / Jerry Slough system) was about 37,300 acres. None of the Kern floodwaters made it to the Tulare Lakebed in 1950.

The rivers in the Tulare Lake Basin crested on November 19. However, as of mid-December, 12 Tulare County bridges still remained impassable.

Tulare Lake had dried up on July 17, 1946, thanks in large part to the extended drought of 1947–50. The flood brought the lake back to life on November 19, 1950, if only for a few months. Table 49 details inflows to Tulare Lake during the 1950 flood.

<table>
<thead>
<tr>
<th>Stream</th>
<th>Total Lakebed Inflow (acre-feet)</th>
<th>Percent Contribution</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings River</td>
<td>39,000</td>
<td>57%</td>
</tr>
<tr>
<td>Kaweah River</td>
<td>14,000</td>
<td>21%</td>
</tr>
<tr>
<td>Tule River</td>
<td>15,000</td>
<td>22%</td>
</tr>
<tr>
<td>Kern River</td>
<td>0</td>
<td>0%</td>
</tr>
<tr>
<td><strong>Total</strong></td>
<td><strong>68,000</strong></td>
<td></td>
</tr>
</tbody>
</table>

Emergency levees were constructed as soon as it became apparent that the floodwaters would enter the lake. This initially confined inflow to the lakebed to 12 sections (7,680 acres) of agricultural land that had recently been planted to barley. Maximum depth of water in the lake was 7.4 feet on December 10, 1950. A high wind on that day caused waves that broke the last of the emergency levees. This inundated 4½ additional sections of land. Total flooded area was 10,600 acres.

The lake was dry again by March 10, 1951, and remained dry until the 1952 flood.

The floods of November and December 1950 extended to the east side of the Sierra, causing considerable damage. Almost 0.6 miles of U.S. Highway 395 was washed out in the Walker River canyon and other roads were damaged in Pine Creek canyon and Round Valley.

1951 Flood
Flooding in 1951 occurred in July.

Despite several false reports, there was only one flood in 1951 that we are sure of: a localized event in July in Kings Canyon.

According to one account, there was a major flood on the Kings River in the spring of 1951. That was almost certainly a mistake, and the flood being referred to actually happened in the spring of 1952.
The national parks’ 1952 annual report suggested that 1951 was a heavy snow year. Judging from the runoff in the Kaweah, the winter of 1950–51 was an average year. However, water year 1952 (which reflects the winter of 1951–52) saw nearly twice the average runoff. There is no reliable record of a major flood occurring on either the Kings or the Kaweah River in the spring of 1951.

According to a published government report, a very intense storm struck Fresno on February 24, 1951, resulting in flooding. That was an error; the storm really occurred on February 24, 1941.

There was a record flood on the American River in 1951 (marking this as a record required ignoring the 1861–62 flood). There is no evidence that this flooding extended into the Tulare Lake Basin.

In July 1951, a cloudburst in Kings Canyon caused a significant debris flow to come down across the highway (photograph on file in the national parks). Judging from the photograph, that debris flow may have occurred near Deer Cove.

1952 Floods (3)

There were apparently three flooding events in 1952:
1. January
2. March
3. April–July snowmelt runoff period

The winter of 1951–52 was truly ferocious in California. There were two sets of particularly severe storms: one in January and another in March.

The winter of 1951–52 was a moderate El Niño event. This association with the 1952 floods may well have been a coincidence. Only strong El Niño events have been shown to have a correlation with high precipitation events and floods in California.

Precipitation in both the San Joaquin River and Tulare Lake Basins was consistently greater than normal throughout the winter. Widespread storms began in October and occurred intermittently until the end of March. Most of the storms brought abnormally cold air and produced snow down to and below an elevation of 1,000 feet. Very little of this snow melted, and a very large snowpack accumulated over the entire mountain area.

By New Year’s Day, substantial snow had begun to accumulate from one end of the Sierra to the other. January began with several relatively light storms, but on January 12–13, the first in a series of powerful and cold storms began to move through the state.

Human activity in the Sierra came to a halt as snow fell faster than it could be removed. East of Sacramento, U.S. Highways 40 (now Interstate 80) and 50 closed. Shortly after noon on Sunday, January 13, a huge snowslide west of Donner Pass stalled the westbound City of San Francisco, the Southern Pacific Railroad’s transcontinental passenger train, in the snow. On the train were 226 people.

Highway 40 was nearby, but it was buried in snow, lost in the blizzard.

With heavy snow falling and the wind blowing up to 100 miles per hour, the Southern Pacific set out to free the train from the drifting snow. Within a few hours, the railroad had plowed to the train and attached another locomotive, but the train was already frozen into the rapidly accumulating snowdrifts and could not be moved. The railroad sent out more equipment, but getting to the train proved terribly slow and difficult. Numerous avalanches had to be cut through and still the snow fell. As the afternoon faded, things got worse as the snow deepened and equipment, pushed to the limit, failed. Finally, another avalanche hit the tracks, flipping a huge steam-powered rotary snowplow on its side, blocking the tracks completely and ending any immediate hope of rescue for the snowbound train.

Meanwhile, in the Southern Sierra, things were not much better. Facing the same huge storm, national park and state crews attempting to maintain access to Giant Forest and Grant Grove found their efforts completely frustrated. Highway 180 was closed entirely above 5,000 feet, as was the Generals Highway into Giant Forest from Three Rivers. The storm continued for the next two days.
By January 16, rescuing the stalled train and its passengers on Donner Pass had become a national priority. A wide variety of rescuers were trying to get to it. The Sixth Army tried unsuccessfully to reach the train using their over-the-snow Studebaker Weasels. Pacific Gas and Electric (PG&E) did succeed in bringing in relief supplies with their double-trucked Sno-Cat. A doctor was brought in from Reno, making the last leg of the trek by dog sled. The Coast Guard sent a helicopter, one of the few at that time on the West Coast.

The Southern Pacific continued working their huge rotary snowplows around the clock to reopen the track. But the storm was unrelenting; avalanches and equipment failures gradually knocked the plows out of commission. The railroad called in all of its reserves, reaching far afield. Finally a rotary was brought down from Oregon—the oldest and last one running in the fleet. On January 16, it succeeded in plowing to the Highway 40 overpass west of Emigrant Gap, near the Nyack Lodge. A special rescue train was then brought in to that point.

At Donner Pass, the snow finally stopped at dawn on January 16. By afternoon, the state highway department had managed to open the section of Highway 40 from Emigrant Gap to the stalled train. The passengers aboard the train were led to the highway and driven by a small fleet of private automobiles five miles to Nyack Lodge where they were fed and their needs attended to. By evening, they were on their way to Sacramento and Oakland. The train was finally freed on January 20, seven days after it was stranded.

The Central Sierra Snow Laboratory monitoring site near Donner Pass received 12.8 feet of snow between January 10–17.

The first flood in 1952 occurred in January; it was apparently a rain-on-snow event. The Kaweah’s peak natural flow occurred at McKay’s Point on January 25: 8,851 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 5,918 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 4 years for the Kaweah.

By the end of January, the weather station in Giant Forest had recorded more than 300% of average January precipitation and more than seven feet of snow covered area meadows. At Lodgepole, several national park buildings were damaged and a major equipment shed collapsed under the snow load. Park highways sustained considerable damage. Between Deer Ridge and Eleven Range Overlook, a major section of the Generals Highway slid away into the canyon. The road was closed for three days and travel was restricted for several more weeks as a result. (We assume that the damaged area was eventually repaired with a wood bin wall, but this has not been confirmed.)

In March, another set of similar storms swept the Sierra, again disrupting travel and halting most human activity in the mountains. By the end of March, almost 30 feet of snow had fallen at Giant Forest, one of the wettest seasons ever recorded in the Southern Sierra.

On March 15, a big, late-season snowfall struck the Sierra. Grant Grove received 37 inches of snow in a 24-hour period. This was the second time that month that 30 inches or more of snow was recorded in 24 hours. Grant Grove received a total of 168 inches (14 feet) of snow during March, making it the snowiest month ever at that location. By March 19, the Central Sierra Snow Laboratory monitoring site near Donner Pass had accumulated a snowpack of 20.57 feet, the most ever recorded at that facility. That was reminiscent of the amount that Walter Fry reported at Giant Forest during the record-setting winter of 1905–06.

The second flood of the year occurred in March. It is poorly documented. It was apparently a rain-on-snow event. Evidently so much water was delivered to the Tulare Lakebed that it caused a levee failure within the lakebed, the first of many that year.

On June 27, 1952, Lodgepole received 0.2 inches of snow. This brought the seasonal snowfall to 449.5 inches (37½ feet) (sometimes incorrectly reported as 522.9 inches), making 1951–52 the snowiest winter on record at that location. This record would eventually be broken by the winter of 2010–2011. Snowfall in the winter of 1905–06 was even bigger than this, but that was before a weather station had been established at Lodgepole.
Table 50 provides a monthly record of snowfall at Lodgepole for the winter of 1951–52.

Table 50. Lodgepole snowfall during winter 1951–52.

<table>
<thead>
<tr>
<th>Month</th>
<th>Snowfall (inches of snow)</th>
</tr>
</thead>
<tbody>
<tr>
<td>September 1951</td>
<td>0.0</td>
</tr>
<tr>
<td>October 1951</td>
<td>4.5</td>
</tr>
<tr>
<td>November 1951</td>
<td>36.9</td>
</tr>
<tr>
<td>December 1951</td>
<td>84.1</td>
</tr>
<tr>
<td>January 1952</td>
<td>119.5</td>
</tr>
<tr>
<td>February 1952</td>
<td>30.3</td>
</tr>
<tr>
<td>March 1952</td>
<td>158.2</td>
</tr>
<tr>
<td>April 1952</td>
<td>15.8</td>
</tr>
<tr>
<td>May 1952</td>
<td>0.0</td>
</tr>
<tr>
<td>June 1952</td>
<td>0.2</td>
</tr>
<tr>
<td>Total</td>
<td>449.5</td>
</tr>
</tbody>
</table>

(The national parks began systematically collecting weather data in 1920. The park has operated two cooperating weather stations in the Lodgepole / Giant Forest area in the period since then. One was operated at the Giant Forest Museum in Round Meadow from June 6, 1921 – November 7, 1968. The other was operated at the old Lodgepole Ranger Station from February 22, 1951 through December 31, 1955. Collection of weather data at this station may well have begun much earlier, perhaps since the mid-1930s, but no data from that period have yet been found. After being deactivated for 13 years, the Lodgepole cooperating weather station was reactivated at the new Lodgepole ranger station on November 8, 1968. Lodgepole is only four miles away and 300 feet higher than Round Meadow, but it has significantly different weather. Bill Tweed lived in Giant Forest and Lodgepole from 1978–88 and often marveled at how the weather at Lodgepole differed from that in Giant Forest. Lodgepole sits at the bottom of the Marble Fork Basin and is closer to the Sierra. It is significantly colder and receives about 10% more moisture than the western side of Giant Forest. As a result, it receives much more snow. It often snows in Lodgepole when it is raining or slushing in Giant Forest. Bill thinks that the difference in the snowpack is upwards of 20–40%.)

By April 1, 1952, a huge snowpack had accumulated in both the San Joaquin River and Tulare Lake Basins. That snowpack, in all the sub-basins, exceeded that existing on the same date in 1938, which had been the greatest snowpack on record since the beginning of the California Cooperative Snow Surveys record in 1930.

After the heavy winter of 1951–52, the USGS reexamined available records of snowfall at stations in the San Joaquin River and Tulare Lake Basins from the winter of 1905–06. Based on the limited data available, the USGS study tentatively concluded that the 1952 snowpack appeared to equal or exceed the snowpack that caused the great snowmelt floods of 1906.

Table 51 compares the snowpack for the winters of 1938 and 1952.

Table 51. Comparison of April 1 snowpack for 1938 and 1952.

<table>
<thead>
<tr>
<th>Drainage Basin</th>
<th>Average snowpack in basin</th>
</tr>
</thead>
<tbody>
<tr>
<td>Upper San Joaquin River</td>
<td>170% 180%</td>
</tr>
<tr>
<td>Kings River</td>
<td>155% 190%</td>
</tr>
<tr>
<td>Kaweah River</td>
<td>155% 220%</td>
</tr>
<tr>
<td>Tule River</td>
<td>180% 265%</td>
</tr>
<tr>
<td>Kern River</td>
<td>205% 260%</td>
</tr>
</tbody>
</table>

Thus, in April of 1952, enough snow had accumulated to cause the greatest snowmelt flood on record, with the possible exception of 1906. (The huge 1850 snowmelt flood would have been well before the period of record.) That such a flood did not occur was largely due to the temperature pattern during the snowmelt period. Weather continued to be cold; temperatures from April through July were generally below normal — in June about 5 degrees below normal. The occasional intervals of hot weather that usually cause the peak flows during the period of the snowmelt runoff were short and not as hot as usual.

Based on the data available, the 1952 snowpack appeared to equal or exceed the 1906 snowpack. However, the only way to be certain was to wait until the snowpack melted and ran off. The results turned out to be quite
clear, and rather surprising. As the melt occurred, the volume of the April–July runoff approached closely that of 1938, and on the Mokelumne, Stanislaus, Kaweah, and Kern Rivers, exceeded it. However, in no case where the period of record included the year 1906, did the 1952 snowmelt runoff exceed that of 1906 on any river. That confirmed that all those watersheds had a bigger snowpack in 1906 than in 1952.

The difference was particularly remarkable in the Kaweah River Basin. The winter of 1951–52 set a modern-day snowfall record at Lodgepole, one that would last until the winter of 2010–11. However, as shown in Table 52, the snowmelt runoff in the Kaweah River Basin in 1906 was 38% greater than in 1952.

Table 52. Comparison of snowmelt runoff for 1906, 1938, and 1952.

<table>
<thead>
<tr>
<th>River</th>
<th>Total Runoff April 1 – July 31 (thousand acre-feet)</th>
<th>Maximum daily flow (cfs)</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>1906</td>
<td>1938</td>
</tr>
<tr>
<td>Kings River</td>
<td>2,980</td>
<td>2,320</td>
</tr>
<tr>
<td>Kaweah River</td>
<td>814</td>
<td>562</td>
</tr>
<tr>
<td>Tule River</td>
<td>192</td>
<td>129</td>
</tr>
<tr>
<td>Kern River</td>
<td>1,390</td>
<td>962</td>
</tr>
</tbody>
</table>

The maximum daily discharge during the April–July 1952 runoff period was greater than the corresponding flow in 1938 on the Mokelumne, Kaweah, and Kern Rivers. However, on those streams where the period of record included the year 1906, the maximum daily flow during the snowmelt period did not exceed that of 1906.

The 1952 snowmelt flood affected all the rivers on the east side of the San Joaquin River and Tulare Lake Basins. We’re fortunate that this flood was so well documented by USGS.

Maximum releases from Friant Dam in combination with a maximum inflow of about 4,600 cfs of Kings River water via the Fresno Slough produced a peak of about 8,800 cfs near Mendota on May 29.

One source said that the flood on the Kings River in the spring of 1952 was of a magnitude similar to the big flood of 1914. That seems like a stretch.

Because there was a recording stream gage on the South Fork Kings, we know that stretch of the river peaked on June 5, 1952: 7,460 cfs. That compares with a peak flow of 10,000 cfs in the 1950 flow.

The 1952 flood, which was a combination of rain and snowmelt runoff, produced much more runoff than the rain-flood of November/December 1950.

In May 1952, the Kern River overflowed its north bank west of Bakersfield and entered the natural flood channel known as Goose Slough and Jerry Slough. The floodwaters put pressure on the levee downstream that had been hurriedly thrown up in the 1950 flood. A 40-foot-wide break formed in that levee just north of the Lerdo Highway. Private interests worked into the night, using heavy equipment to repair that break. Larry Frey recalled that the last real Kern River floodwaters coming down Jerry Slough into Goose Lake were in the fall of 1951 and the spring of 1952. That is probably in large part because Isabella Dam began operation in 1954.

Flood crests on the various rivers within the Tulare Lake Basin were not particularly high in 1952. However, moderately heavy flows over a long period caused considerable damage in the valley, particularly in the Tulare Lakebed. The USACE conducted a three-month-long survey to determine the extent of the damages incurred. The results are shown in Table 53.

Table 53. Damages incurred during 1952 snowmelt flood.

<table>
<thead>
<tr>
<th>Water Body</th>
<th>Flooded Area (acres)</th>
<th>Public Institutions and Utility Damage (million dollars)</th>
<th>Agricultural Damage (million dollars)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings River</td>
<td>5,200</td>
<td>.063</td>
<td>.136</td>
</tr>
<tr>
<td>Tulare Lake</td>
<td>72,700</td>
<td>1.418</td>
<td>6.877</td>
</tr>
<tr>
<td>Tule River</td>
<td>300</td>
<td>.017</td>
<td>.04</td>
</tr>
<tr>
<td>Kern River</td>
<td>30,700*</td>
<td>.005</td>
<td>1.21</td>
</tr>
<tr>
<td>Total</td>
<td>108,900</td>
<td>1.50</td>
<td>8.26</td>
</tr>
</tbody>
</table>

*The 30,700 acres flooded by the Kern River included 23,500 acres in the Buena Vista Lakebed and 7,000 acres south of Sand Ridge.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

The Tulare Lakebed had been dry from March 10, 1951 through January 19, 1952. During the 1952 flood, the lake rose by 15.5 feet to a maximum elevation of 194.6 feet (elevation 194.6 - 179.1 feet). The lake has not been this high since, although the 1969 flood came close.

Presumably that was high enough to threaten Corcoran. The last time that the lake had been this big was during the years 1937–44. The New York Times reported on the return of Tulare Lake and that a dike broke, flooding cotton fields. Levee failures in the lakebed occurred from March until June 2, 1952.

As shown in Table 54, all four major rivers contributed water to the Tulare Lakebed during 1952.

<table>
<thead>
<tr>
<th>Stream</th>
<th>Total Lakebed Inflow (acre-feet)</th>
<th>Percent Contribution</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings River</td>
<td>258,000</td>
<td>44%</td>
</tr>
<tr>
<td>Kaweah River</td>
<td>175,000</td>
<td>30%</td>
</tr>
<tr>
<td>Tule River</td>
<td>50,000</td>
<td>9%</td>
</tr>
<tr>
<td>Kern River</td>
<td>100,000</td>
<td>17%</td>
</tr>
<tr>
<td>Total</td>
<td>583,000</td>
<td></td>
</tr>
</tbody>
</table>

Inflows to Tulare Lake would have been even larger, but 44% of the floodwaters were stored elsewhere before they reached the lake:
- The USACE operated the partially completed Pine Flat Dam, storing about 130,000 acre-feet of Kings River floodwater there.
- A local water storage district stored 232,000 acre-feet of Kern River floodwaters in Buena Vista Lake. That flooded 23,500 acres of that lakebed.
- Local interests dammed the Kern River channel where it flowed through Sand Ridge. This created a 7,000 acre lake south of that structure, which stored about 100,000 acre-feet of water. That lake was where the southern extension of Tulare Lake had formerly been, the lake which had been known to the American Indians as Ton Táché.

The flood peaked in the Tulare Lakebed about June 20, 1952. Agricultural land in the lakebed stayed flooded from March 1952 through at least September 1953. The last of the J.G. Boswell Co. land was pumped dry on September 22, 1953.

Flooding also occurred on the west side of the San Joaquin Valley near Coalinga sometime during 1952.

Total flow for water year 1952 was 166% of the 1894–2014 average for the Kings, 194% for the Kaweah, 233% for the Tule, and 206% for the Kern. This was one of the very rare years when flows on the Tule River exceeded what could be used for beneficial use by the holders of water rights.

1955–56 Floods (2)
There were two flooding events in 1955–56:
1. River flooding due to storms that occurred from mid-November through December, 1955.
2. River flooding due to a storm that occurred on January 25, 1956.

The winter of 1955–56 was a strong La Niña event. This would remain the strongest known La Niña until the 2010–11 event.

During December 17–27, 1955, a warm rainstorm melted accumulated snowfall up to an elevation of 10,000 feet. This storm was heaviest in the Central Sierra, the Feather, Yuba, Bear, American, Cosumnes, and Calaveras Rivers, as well as the Russian and Napa Rivers, and the streams of the South Bay Area.

The rainfalls of the higher elevations of both the Coastal Mountains and the Sierra were affected by this storm sequence. A total of 20 stations reported storm intensities in excess of a storm with a recurrence interval of 1,000 years. The Santa Clara Valley was the hardest hit in terms of rainfall events with large recurrence intervals.

This was a relatively high-elevation storm in the Central Valley. Over half the stations reporting a storm with a recurrence interval of 100 years were located at an elevation over 1,000 feet. The highest for the Sierra stations
was 36.57 inches at Strawberry Valley at an elevation of 3,800 feet in the Yuba River Basin. Lake McKenzie, located southwest of San Jose at an elevation of 1,800 feet, received 42.27 inches. Honeydew in the Mattole River Basin received 49.20 inches in 8 days.

A total of 19 stations reported daily rainfall in excess of 10 inches in one day during the December storm. These were located in the Upper Sacramento, Feather, San Joaquin Rivers, and in the Clear Lake area. Lakeshore in Shasta County received 15.34 inches in 24 hours on December 20. This was the heaviest 24-hour rain event ever reported for the Central Valley up to that time. (Hockett Meadow would break this record in December 1966.) This storm did not produce heavy, short bursts of rain, but rather rain continued all week with few breaks. It saturated the soil and filled the surface reservoirs. It resulted in extensive flooding which devastated Yuba City and forced the evacuation of 20,000 people.

Although the river flooding occurred in December 1955 and January 1956, flooding continued in the Tulare Lakebed for about four months thereafter. That was very much how the flooding had played out in 1916.

The December 1955 flood brought large flows to many locations in the Sacramento River Basin. A levee break on the Feather River caused severe flooding in the Yuba City area. The flow in the American River at Fair Oaks was controlled to 70,000 cfs because Folsom Reservoir was nearly empty at the beginning of the event. Had Folsom been up to allowable storage capacity, the project would have exceeded its design outflow and the flow at Fair Oaks probably would have been more than 115,000 cfs. At the Sacramento Weir, 30 gates were opened, and the peak flow reached 48,800 cfs. The peak flow in the Sacramento River at I Street was about 95,000 cfs. Total flow at the latitude of Sacramento, including the Yolo Bypass, was about 380,000 cfs.

Floods in the San Joaquin River Basin reflected those in the Sacramento River Basin. Flows on the San Joaquin River were completely controlled by Friant Dam. Prior to the December 1955 flood, Millerton Lake was well below flood-control pool. If storage had been at allowable flood management levels, uncontrolled flows would have exceeded 37,100 cfs and resulted in extensive damage between Friant Dam and the mouth of the Merced River. The peak flow of 62,500 cfs was a record on the Stanislaus River at Ripon, while the Middle Fork of the Tuolumne River at Oakland Recreation Camp reached a record flow of 4,920 cfs. During the 1955 floods, two of the three forks of the Tuolumne River reached record flows.

Flooding occurred from late November through December 1955, but the worst of the flooding occurred around December 23–24. The flooding affected the northern and central parts of the state as far south as the Tehachapis. It resulted from a family of cyclones originating in the mid-Pacific Ocean. The flood had a recurrence interval of up to 100 years, depending on the river. It caused 76 deaths and $166 million in property damage and was one of the five costliest floods in California’s history.

In the national parks, the first fall storm arrived on November 13. Two more storms came during the next 10 days, resulting in considerable snow at the higher elevations. There was significantly more rain and snow during December. Torrential rains fell in the Kaweah River Basin on December 23–24. Giant Forest recorded 11 inches on the 23rd and 4 inches on the 24th. It kept raining at a lesser rate through December 27.

The valley and much of the Sierra experienced heavy rainfall from December 22–25. As shown in Table 55, the precipitation was particularly intense on December 23.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of moisture)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Fresno</td>
<td>1.72*</td>
</tr>
<tr>
<td>Grant Grove</td>
<td>9.22</td>
</tr>
<tr>
<td>Giant Forest</td>
<td>11.04</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>7.24</td>
</tr>
<tr>
<td>Woodlake</td>
<td>3.3</td>
</tr>
</tbody>
</table>

*This set the record for the wettest December day ever in that city.

Heavy rainfall in the San Joaquin Valley and much of the Sierra from December 22–25 led to flooding across the area. Some rivers and streams reached their highest level on record at the time. It was an unusually warm Christmas in Visalia, even by valley standards.
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

December 1955 is still the wettest December of record, at least in the Northern Sierra and parts of the Tulare Lake Basin. Fresno’s total precipitation during December was 6.73 inches, making it the fourth wettest month ever, and the wettest December on record.¹²⁷⁵ Visalia received 6.06 inches of rain in December, the most since record-keeping began in 1898.

The flood merited three front-page stories in the New York Times. The story in late November talked about flooding from Sacramento to Bakersfield.¹²⁷⁶ There were two stories at Christmastime. One talked about the coast floods starting to recede.¹²⁷⁷ The other story talked about renewed flooding in the Sacramento–San Joaquin Delta.¹²⁷⁸ All three stories apparently addressed conditions in Visalia.

The floods of December 1955 were memorable not only for the magnitude of peak discharge, but also for the duration of rain and the extent of the area affected. Rain fell in coastal areas on 39 of the 44 days between December 15, 1955 – January 28, 1956 as several storms crossed the northern two-thirds of the state. In most areas, the storm of December 21–24 caused the most damage.

Warm, moist air from the southwest released rains that drenched the mountains and melted much of the snow that had accumulated in the Sierra. During December 15–27, extremes of up to 40 inches of rain fell at several locations, and quantities greater than 20 inches were common in the coastal mountains and the Sierra.

The floods of December 1955 produced peak discharges in much of the area that were in excess of any previously recorded. Flooding was particularly notable on the Klamath River on the North Coast, Alameda Creek in the San Francisco Bay area, the San Lorenzo River at Santa Cruz, the Feather River near Yuba City, the Kaweah River at Visalia, and the Carson River east of the Sierra. Peak discharges at these widely separated rivers were generally 1½–2 times the discharge of the previously recorded peak flows. On many streams, the floods ranked among the greatest since the 1861–62 flood.

The Merced River at Happy Isle peaked on December 23: 9,860 cfs. One source said that this flood had a recurrence interval that exceeded 100 years. However, Mary Donahue calculated the recurrence interval as 45 years.¹²⁷⁹

Several rivers, including the Eel, experienced their greatest discharge-of-record during this flood. Overflow of the Klamath River resulted in almost complete destruction of the town of Klamath.

About 382,000 acres of the Sacramento River Basin were flooded. Unusually high tides aggravated the situation by impeding the passage of floodwater through the Sacramento–San Joaquin Delta.

The American River experienced a record flood, the second record flood in seven years.

On December 24, a levee failure on the Feather River flooded more than 3,000 homes in Yuba City, killing 38 people and forcing the evacuation of 12,000 people. The city was inundated with floodwater as much as 12 feet deep.

Marysville was surrounded by the merged floodwaters of the Yuba and Feather Rivers. The entire city, all 12,500 residents, was ordered to evacuate. A photograph (on file in the national parks) shows the city awash even inside the moat-like ring levee and evacuating cars streaming out the only exit.

The Friant Dam contained the entire runoff of the flood on the mainstem of the San Joaquin River and prevented widespread flooding of the agricultural lands from Mendota north to the vicinity of Los Banos. (However, the northeastern part of the town of Los Banos was flooded, forcing the evacuation of 65 people.) Other reservoirs effectively reduced the flooding on the Merced County Stream Group (Burns, Bear, Owens and Mariposa Creeks). The Exchequer Reservoir regulated the flood on the Merced River to a safe outflow of 10,800 cfs.¹²⁸⁰

However, the Fresno and Chowchilla Rivers were not regulated, and the Melones Reservoir on the Stanislaus River and the Don Pedro Reservoir on the Tuolumne River both overflowed; the uncontrolled spill at the peak being 59,400 cfs and 41,700 cfs respectively.¹²⁸¹

This resulted in a huge lake that extended along the San Joaquin River and several of its tributaries from the Fresno-Madera county line north through Merced and Stanislaus Counties and into San Joaquin County, flooding an estimated 300,000–400,000 acres.¹²⁸² Many streams and rivers in this area flooded. Water covered the Chowchilla business district and inundated Highway 99 up to 5 feet deep.¹²⁸³ This forced the closure of Highway
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

99 at Chowchilla at the height of the Christmas travel period. The State Division of Highways and the Highway Patrol worked together to effect a side-road detour around this large water barrier.

This would be one of two great lakes to form during the December 1955 flood. The other would develop around Visalia and be fed solely by runoff from the Kaweah River.

Panoche/Silver Creek west of Mendota flooded at about Christmastime 1955.1284

Because there was a recording stream gage on the South Fork Kings, we know that river peaked on December 23, 1955: 13,900 cfs. That broke the previous record of 10,000 cfs set for this stretch of river on November 19, 1950.1285 This remains the highest flow recorded on the South Fork Kings since record-keeping began in 1950.

The peak day natural flow at Pine Flat Dam occurred on December 23. That is the largest peak day of the year at Pine Flat since the dam was built in 1954. Based on the flood exceedence rates in Table 29, this had a recurrence interval of 100 years for the Kings River at Pine Flat. That puts it in a category with other Cedar Grove floods of the past 70 years (1937, 1950, 1955, 1966, 1969, 1978, 1982, 1983, and 1997) that rise to the level of the modeled 50-year flood: that is, a flood event that occurs about every eight years on average. See the section of this document that describes Cedar Grove Flooding.

Specific damage in Kings Canyon included:
- The Highway 180 Boyden Bridge survived the 1955 flood, but the Grant Grove approach to the bridge was washed out (photograph on file in the national parks). The Grant Grove approach would be washed out again in 1997. The bridge is still standing and in use. It was constructed in 1939 and has withstood many floods.
- The South Fork Kings River caused serious damage to Highway 180 in the Kings River Canyon. Many miles of Highway 180 were washed out or undermined (multiple photographs on file in the national parks). In the vicinity of Boulder Creek, where the highway and the Kings River occupy a steep, narrow canyon, approximately 4,000 feet of road was completely washed out. There were approximately seven or eight other washouts of a serious nature. Even by foot, the road was only passable to a point ½ mile above Boulder Creek.
- The left embankment of the Cedar Grove Bridge washed out. Significant quantities of fill and riprap were required to repair this washout. The damage done to the Cedar Grove Bridge in this flood was similar to the damage that would later occur in the 1997 flood.
- One of the piers on the Roaring River Bridge was reportedly undermined, causing the bridge to tilt.1286

The various branches of the Kaweah River took out many bridges during this flood, one of which was in the national parks. The Visalia City Council hosted what amounted to a miniature after-action review on January 31, 1956. Fred Walker represented the national parks. R.C. Sorenson represented the USACE. Joe Garcia, Jr. represented the Tulare County Road Commission. One of the discussion items at that meeting was the desirability of replacing the washed out bridges with suspension bridges.

Specific damage in the Ash Mountain area included:
- The Marble Fork Bridge near Potwisha was washed out (photograph on file in the national parks). A temporary footbridge was constructed over the bridge piers.
- A washout occurred ½ mile below Potwisha, undermining the Generals Highway (photograph on file in the national parks).
- There was a slide on the Generals Highway ½ mile above the national parks headquarters (photograph on file in the national parks).
- The national parks’ approach to the Pumpkin Hollow Bridge completely washed away (multiple photographs on file in the park). A very tall ladder was placed in the resulting hole (or new riverbed) so that people could climb up to the bridge and then continue on to Three Rivers. Both approaches to this bridge had washed out in the 1937 flood. The 1966 flood would largely wash out the parks’ approach to this bridge. However, the bridge itself has withstood all these floods; it is still standing today.
- About 400 feet of Highway 198 washed out in the lower portion of Pumpkin Hollow (photographs on file in the national parks). Now there is a large curving concrete wall at this point. From the photographs, it is evident that the dirt required to repair the damage was trucked down from the national parks. Perhaps it was brought up Dinely Dr. and through the Riverway Ranch.

It isn’t clear when the section of the Generals Highway between Ash Mountain and Giant Forest was reopened to visitor traffic. Colony Mill Road had been considered unsafe for visitor traffic for many years; that route may not have been used as a detour since 1937. At the least, visitors were probably kept off this section of the Generals
Highway until a new bridge could be constructed across the Marble Fork Kaweah at Potwisha. That presumably took at least a year. This may have been the first time since 1937 or 1938 that the Generals Highway had been closed for more than a couple of weeks.

Jim Harvey recalled that the national parks’ bridge over Yucca Creek washed out in this flood. This was on the road known as the West Boundary Truck Trail.

In Three Rivers, the various branches of the Kaweah rose steadily all day on December 22 due to the heavy downpour. There was no cause for alarm until shortly after midnight, when the rivers surged up with a thunderous roar and swept everything from their path. Many of the town residents were caught up in the battle to save lives that night. One sample from among many in the first post-flood issue of the *Three Rivers Current*:

> Willie Clay was another near-flood victim. He started down the road just before 3 a.m. to lend a helping hand when just past Kath’s, a wall of water spun his car around like a toy. He climbed out and headed for shore about 30 feet away when a second wave hit and tumbled him over and over. He managed to hang on to the first solid thing he felt which was a tree and he climbed up above the water. He spent a terrifying 5 hours wavering above the black rush of water and crashing debris. Shortly after dawn, Fred Walker, John Wollenman, and Leroy Maloy got out to him from the upper side of the road. 1287

In December 2005 or thereabouts, the Kaweah Commonwealth published a 50th anniversary edition featuring the 1955 flood. It had a photograph of the mainstem of the Kaweah at flood stage going around the south side of the Three Rivers Market.

Specific damage in the Three Rivers area included: 1289

- In addition to the damage that occurred in the Pumpkin Hollow area (described above); there were several washouts on Highway 198 between Pumpkin Hollow and Slick Rock.
- The Dinely Bridge (the one that had been built after being washed out in the December 1937 flood) was overtopped and swept away at 3:15 a.m. When that bridge went out, it reduced the height of the floodwaters on the nearby Buchholz’ house. A new bridge was built in 1957.
- The Upper North Fork Bridge washed out. The Airport Bridge survived, but both approaches were washed out. Some of the homes on the North Fork were surrounded by floodwaters, trapping the occupants. Many of the homes and buildings in that area were flooded and badly damaged. The McDowell chicken ranch suffered severe damage. The North Fork Road was badly eroded, isolating residents from the rest of town.
- The mainstem of the Kaweah overtopped and washed away much of the North Fork Bridge, the one next to the Three Rivers Market that had been built to replace the one washed out in the 1937 flood. This bridge consisted of three Bailey Bridges placed end to end on concrete piers. One of those bridges remained on its piers. The other two were washed out and came to rest just downstream of the piers. The Upper North Fork Bridge was replaced after the flood with a Bailey Bridge. Possibly it was the one surviving segment of the North Fork Bridge, especially since that segment was on the same side of the river as the North Fork Road. That Upper North Fork Bridge would later wash out in the 1966 flood.
- Water and sand flooded all three of SCE’s powerhouses, and all were knocked out of operation. The Kaweah #3 hydroelectric complex that is located just inside Sequoia National Park was particularly hard hit (multiple photographs on file in the national parks). One transformer at Kaweah #3 was tipped over, and the other was undermined. The generator floor at Kaweah #3 was flooded to a depth of three inches. There were a number of washouts on the flumes. One of the damaged areas was a major break in the Middle Fork Kaweah concrete flume near Station 60 that was repaired with redwood. Some records associated with SCE’s photographs documenting the damage incorrectly identified the flood as occurring on January 18, 1956. The damage was really sustained in the December 23, 1955 flood.
- The first two bridges on the South Fork Road (Conley Creek and South Fork of the Kaweah) were washed out when these streams merged and created what was described as “a scene of devastation” (photograph on file in the national parks).
- The third bridge on the South Fork Road apparently survived the 1955 flood. It would be replaced with a new Bailey Bridge in 1959.
- Based on Karen Folger’s research, the South Fork of the Kaweah appears to have relocated its channel at some point. This most likely occurred during the December 1955 flood or possibly the November 1950 flood. The old topo maps show the river running north of where the houses are now between the first two bridges on the South Fork Road. However, the river currently runs south of those houses with a split to the north that feeds the ditch intake on the north side at the first bridge.
- The Three Rivers Motel and Trailer Court was demolished.
The Sequoia Hardware Store was undermined and badly damaged (multiple photographs on file in the national parks). That building was located just east of the present Hummingbird Restaurant. Many homes were washed away and many others were severely damaged. Much livestock was lost.

The 1955 flood wiped out the last traces of Conrad Alles’s sawmill that was located on the mainstem of the Kaweah below the present-day Three Rivers Golf Course. At the time of the flood, there were two ways to access Three Rivers from the west, bridging the South Fork of the Kaweah. The primary access was on Highway 198, passing through what is now the Kaweah Park Resort. Farther to the south, the county road known as Old Three Rivers Road followed the former alignment of Highway 198. (The westernmost portion of that road is known today as Cherokee Oaks Drive). The flood caused serious damage to both of those roads, cutting off all access to Three Rivers and to Sequoia National Park.

The Highway 198 bridge over the South Fork of the Kaweah River withstood the flood. (It appears in a photograph from the 1966 flood and is presumably still there today.) However, immediately east of that bridge, the mainstem of the Kaweah River overtopped and washed away 1,600 feet of the highway, and occupied an area approximately 1,000 feet south of the previous location of the route (multiple photographs on file in the national parks). This washed away five houses and the Dunlap Motel (one source incorrectly said that it was the Noisy River Lodge). Because of the magnitude of this washout, the state would decide to abandon that alignment rather than rebuild it.

The South Fork of the Kaweah undermined the center of the county’s South Fork Bridge and caused it to collapse (photograph on file in the national parks). That was the old highway bridge that connected what is now Cherokee Oaks Drive with Old Three Rivers Road. That was the route that had formerly been the alignment of Highway 198. That bridge had been reconstructed by the state after it was destroyed in the 1937 flood.

The mainstem of the Kaweah in Three Rivers peaked on December 23: 80,700 cfs. This was the flood-of-record for this gaging station (USGS gage #11-2105 Kaweah R. nr Three Rivers) during its period of operation from 1903–1961. (The December 1966 flood would have only a 2% greater flow through this stretch: 82,700 cfs. The Three Rivers gaging station was located just above the junction of the river with Horse Creek. This gage location is now submerged under Lake Kaweah.

As soon as the flood passed, the national parks sent equipment and employees to help Three Rivers. The assistance they provided included:

- Worked for days to keep the Buckeye route passable for at least truck traffic through the Riverway Ranch to the North Fork Road.
- Repaired the washed-out approach to the Pumpkin Hollow Bridge.
- Assisted with repairs to the lower portion of Pumpkin Hollow.
- Dozed a route over the Shepherd’s Saddle Road to the North Fork Road.
- Cleared mud and debris from many roads in the Three Rivers area.

This may have been the last time that federal accounting oversight controls allowed the national parks to send such assistance to the Three Rivers community.

The county worked to repair the many bridges that had been damaged or destroyed within Three Rivers. One of the high priorities was to reestablish access across the mainstem of the Kaweah River. Almost immediately after the flood, a cable trolley was rigged at the site of the washed-out North Fork Bridge so that people could be pulled back and forth (multiple photographs on file in the national parks). (A cable trolley had previously been used at this location after the North Fork Bridge had been destroyed during the December 1937 flood.) On December 31, a temporary log bridge was constructed for vehicles.

But a much more permanent temporary bridge was needed. The road commissioner pointed out that it had taken 1½ years to replace the North Fork Bridge after it had been destroyed in the 1937 flood. Therefore, the county authorized the rental of a Bailey Bridge at about $550 per month, to be installed just above where the Chevron Station is today. That may be where the temporary log bridge had been constructed. It was in place by January 18, 1956 (photograph on file in the national parks).

The state worked to restore service to Three Rivers and the national park as soon as possible. This was done by using Old Three Rivers Road (the former state highway) as a detour. This required shoring up and re-decking the county bridge that had collapsed during the flood.
The state then built a new Highway 198 alignment midway between the two bridges that had failed. The new bridge over the South Fork of the Kaweah (Bridge #46-29) was completed in 1957.

As part of building the new bridge, the channel of the South Fork was remade for a distance above and below the bridge. A lot of material was removed from the floodplain, and levees were created on either side of the river. Harry Kulik used to own the Kaweah General Store adjacent to the bridge site. He told Jack Vance that the contractors spent weeks reworking (dredging) the South Fork channel. The effects of that operation are still quite visible today despite many subsequent floods and decades of revegetation.

Harry said that prior to the flood, there used to be an irrigation ditch that sent water from the South Fork to the residents of Pierce Drive. That was also the source of water for the waterwheel landmark visible from Highway 198 in that area. After the dredging, the South Fork was lower than it had been before the flood, too low for water to flow into the head of the irrigation ditch. The ditch has not functioned since the dredging and has been abandoned. Only small traces of it are visible today.

The flood caused several washouts on the Mineral King Road, with the major damage above Faculty Flat.

During the height of the flood, the mainstem of the Kaweah River was very wide at Slick Rock and immediately above. A photograph in the *Exeter Sun* (on file in the national parks) seems to indicate that the river extended to roughly where the new USACE toilets and parking lot are located.

The *Visalia Times-Delta* published a special flood edition on January 20, 1956. That paper presumably contains interesting material about the details of this flood, but we haven’t seen it.

Dry Creek below present-day Terminus Dam peaked at 6,070 cfs. Those floodwaters merged with the Kaweah, adding to the destruction.

The area from Terminus Beach to Highway 99 was hard hit by flooding on December 23. When the Kaweah reached the Lemon Cove-Woodlake Road (Highway 216), it was about a mile wide, stretching from the intersection with Dry Creek Drive almost to the intersection with Highway 198 (photograph on file in the national parks). Both the Terminus Beach and McKay’s Point Resorts were inundated.

The Kaweah overtopped and broke the Friant-Kern Canal south of Woodlake, causing significant damage to that canal (multiple photographs on file in the national parks).

The Kaweah swept away 350 feet of the Visalia Electric mainline trestle and 1,800 feet of track near McKay’s Point, about a mile below the Lemon Cove-Woodlake Road (photograph on file in the national parks). A similar event had happened at this trestle in November 1950.

The flood damage below Terminus Beach had two separate components:
- Damage to agricultural lands because floodwaters were greater than levees were designed to contain.
- Damage to Visalia due to failure of the diversion structure at McKay’s Point.

The Kaweah’s peak flow occurred at McKay’s Point on December 23. The gage at McKay’s Point was swept away that morning, before the peak of the flood. There have been a variety of estimates presented for how big the flood was at McKay’s Point when it peaked.

In a 1956 report, the California Disaster Office reported that the Kaweah River set a record when it peaked at 74,400 cfs during the December 1955 flood.

In 1970, USBR (with assistance from USACE) reported that the Kaweah was flowing at 80,700 cfs when it peaked on December 23. The 80,700 cfs figure was just a reflection of the flow upstream at the Three Rivers gage (USGS gage #11-2105). It did not include the flow that was added from Dry Creek.

In recent years, the USACE at the Lake Kaweah Visitor’s Center reports that the Kaweah peaked at 87,400 cfs at McKay’s Point.

A December 2010 issue of the *Kaweah Commonwealth* reported that the USACE has modeled the December 1955 flood and that it had a recurrence interval of 100 years. The model estimated that at the peak of the flood,
the Kaweah was flowing at an estimated 110,000 cfs in the vicinity of Three Rivers. We haven’t been able to locate this USACE study, and those model results do not appear to have been used when the flood frequency curves were revised for the Kaweah in 2005.

This document uses the estimates used by the USACE when the flood frequency curves were revised for the Kaweah in 2005.

The Kaweah’s peak natural flow occurred at McKay’s Point on December 23: 84,332 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 44,512 cfs.)

Based on the flood exceedence rates in Table 29, this had a recurrence interval of 85 years for the Kaweah. It would have had a recurrence interval of 55 years if calculated using the 84,332 cfs peak flow. (One source reportedly calculated this as having had a recurrence interval of 232 years. That result could not be reproduced with either the peak or the daily flows, even when using the now-outdated 1971 flood frequency curves.)

Since 1870, there had been a weir at McKay’s Point to divert most of the Kaweah River floodwaters into the St. Johns River channel and north around Visalia. Judging from the photographs (on file in the national parks), the 1955 flood cut a channel around the north end of the weir. In addition, the flood plugged the mouth of St. Johns channel with a huge amount of sediment.

This failure of the McKay’s Point diversion resulted in most of the floodwaters going down the Lower Kaweah River channel toward Visalia. The Lower Kaweah River channel ends east of Visalia (just north of the Ivanhoe turnoff (Road 156/158) on Highway 198). The floodwaters rushing down the Kaweah on December 22 and 23 found an outlet to the southwest through Cameron Creek, Packwood Creek, and Mill Creek, none of which had nearly the capacity to handle the volume of water coming their way. The combined capacity of those three creeks was well under 5,000 cfs.

A very big lake was about to form on the east side of Visalia. At its peak, that lake would have a length of 10–15 miles. It flooded an estimated area of 183,000–300,000 acres. The Kaweah Delta was accustomed to flooding, but it had not seen anything like this in historic times (multiple photographs on file in the national parks).

Cameron Creek (a distributary of Deep Creek, southwest of Kaweah Oaks Preserve) caused widespread flooding. Wortman Mill located northeast of Farmersville on the Wortman Road (known today as Road 168) was severely damaged (multiple photographs on file in the national parks). Its entire inventory of 140,000 board feet of dried lumber was swept away as well as 50,000 board feet of logs. The mill office (essentially a small house) was swept off its foundation and was later found over a half-mile downstream where it had come to rest on Farmersville Blvd. The owner of the mill gave the office away, saying that he had no further use for it; the flood had wiped him out completely. Cameron Creek Colony (located just upstream of Farmersville Blvd) and Linnell Labor Camp (located just downstream of Farmersville Blvd) were heavily flooded. Culverts under Farmersville Blvd were inadequate for the flow and plugged. That diverted the flooding creek which turned and flooded the town of Farmersville, ½ miles to the south. The network of roads in the agricultural area around Visalia effectively became a network of canals. One migrant labor camp containing some 200 cottages was entirely inundated, and the workers there escaped with little personal property.

Packwood Creek flooded a large area, including several sections of Highway 198 near the Ivanhoe turnoff (Road 156/158). Packwood Creek caused severe erosion in numerous areas, such as the farmland around the intersection of Lovers Lane and Caldwell (multiple photographs on file in the national parks). Nearly all of the bridges over Packwood and other creeks south of the Mineral King Highway were destroyed during the flood.

The town of Exeter was just outside the flood zone, but the land around Exeter, stretching out toward the communities of Woodlake, Farmersville and Lindcove, was so widely inundated that about 500 families had to leave their homes. Some 50 of those families were from Woodlake. Woodlake and the surrounding area were thoroughly inundated (multiple photographs on file in the national parks). Eight motorboats were initially rushed to Farmersville to help with evacuation of that area, and the sheriff’s office was looking for more. Many of the families in the Woodlake area had to be evacuated by helicopter.

Gary Shadrick was a 10-year-old at the time of the flood. His family lived on N. Valencia Blvd. in Woodlake, four blocks north of the main intersection with E. Naranjo Blvd. He recalled that they had to evacuate their house during the flood. The family gathered and watched as the floodwaters rose higher. He remembers his father using a rowboat to check on their house (and its hardwood floors) during the flood.
The flood swept through Visalia on Christmas Eve, reaching the town just as darkness closed in. Damage was particularly heavy on the southeast side of town. Many residents had to be evacuated from their homes by boat during the night. The National Guard assisted the sheriff’s office and the police department with the evacuations. At 11 p.m. on Christmas Eve, an area covering 21 blocks was sealed off due to the flooding. The National Guard and a detachment of U.S. Marines from Tulare assisted the police with patrolling the town to prevent looting.

The depth of the water in Visalia was not great. While it reached five feet in some places, most areas were only about a foot deep. Because flooding was historically a relatively common occurrence in Visalia, very few buildings had basements. Even so, a lot of silt and mud was left behind in the flooded buildings.

Water was generally flowing in a shallow sheet along the Kaweah Delta, bound for Tulare Lake. South of Mineral King, Mooney Blvd. worked like a levee. The water pooled up when it reached Mooney, flooding the houses on the east side of that road. Houses on the west side of Mooney were grateful for the flood protection.

Pumps were brought in to provide relief for the houses on the east side, and began pumping the floodwaters across Mooney. A particularly large pump was placed at the northeast corner of Meadow Lane and Mooney, where the parking lot for Carrows Restaurant (now Black Bear) is currently located. The round-the-clock pumping achieved the goal of moving the floodwaters across Mooney, but the homeowners on the west side were not happy at receiving all that water.

Most of the campus of the College of the Sequoias was flooded (multiple photographs on file in the national parks). The damage was most severe in the building that housed the college gym because it had a basement. (The building had been designed under the assumption that floodwaters could never reach this part of Visalia.) There was widespread flooding in the neighborhoods around the COS campus.

Farther downstream, Mill Creek broke across the Visalia Airport and flooded the depressed portion of the Visalia Plaza interchange (aka Highway 99 and 198 interchange, since reconfigured but in the same general location). The water came in faster than the pumps could bail the water, and finally the pumps quit working. Traffic had to be detoured around that interchange for the better part of the week of December 25–31.

Largely isolated by the flooding of Highway 99, Visalia and the neighboring communities had to call for food and medical supplies from the outside. The State Office of Civil Defense made those arrangements, with food for the city coming from Holloman Air Force Base near Alamogordo, New Mexico, and medicine being sent from McClellan Air Force Base near Sacramento.

The danger from contaminated drinking water was so great that the Tulare County Health Department embarked on an emergency program to inoculate 35,000 persons. This apparently represented the majority of the county since the combined population of Visalia and Tulare was less than 24,000.

In Visalia, 332 families had to be evacuated from their homes, chiefly in the southeast section of town. Combined with the 500 families from the rural areas around Exeter and some scattered evacuees, an estimated 880 families were driven out by floods in Tulare County. Three mass care centers had to be opened in the Visalia area, and it was eight days before the last of the evacuees could return to their homes.

Immediately after the flood crest passed, emergency work began to reopen the mouth of the St. Johns River at McKay’s Point. Among the first requests made to the State Office of Civil Defense from Visalia was one for five draglines to dredge out the sandbar in the St. Johns River. Tulare County Civil Defense officials were able to locate the necessary draglines, and the California Highway Patrol made arrangements to get the bulky caravan through to its destination. By December 28, 25 pieces of heavy equipment were involved in the effort to reopen the St. Johns channel (multiple photographs on file in the national parks).

This emergency work was overseen by the USACE as was levee repair on the St. Johns more than a mile downstream. The repair work was done using federal funds. Under the terms of the federal law governing such emergency repair work, the USACE could only restore the St. Johns levee to the same condition that it was in prior to the flood, not fix its obvious deficiencies. At best, the levee was returned to the condition it was in when it was brand new on December 5, 1891. When the January 1956 flood hit, portions of some of the newly rebuilt levees failed again. It was a frustrating situation for all parties concerned.
Approximately 183,000 acres of land in and around Visalia was inundated by the Kaweah during the flood. (This had initially been estimated to be more than 300,000 acres.) There was no loss of life in Tulare County, but 100 injuries were attributed to the flood, and there were 24 traffic accidents under flood conditions that involved personal injury.

Private residences in Visalia sustained nearly $1 million in damage, and businesses suffered an additional loss of $600,000. There was $200,000 damage to public property, not including state highways. In the rural areas, there was $200,000 damage to farm homes. Tulare County sustained approximately $10 million in loss to the current agricultural crop plus an additional $10 million in permanent damage to land, orchards and vineyards.\textsuperscript{1307}

According to the California Disaster Office, the Kaweah River Basin sustained the greatest damage of any area in the San Joaquin River Basin during this flood.\textsuperscript{1308} The damage suffered in the Visalia area in the 1955 flood was the impetus for Terminus Dam to be constructed. Flooding in Visalia has been much less of a problem since that dam was completed in 1962.

Tulare was largely outside the flood zone and had only a few blocks flooded. However, roads to that city were closed off in most directions.

Tagus Ranch is located about four miles north of Tulare. Highway 99 was shut down at Tagus Ranch when a large lake covered the roadway there, leaving 25–50 cars stranded (photograph on file in the national parks). Residents of the area around Tagus Ranch had to be evacuated. Trains were also stalled as a result of the area flooding.

There are two powerhouses on the Middle Fork of the Tule River. The Mt. Whitney Power and Electric Co. constructed the first powerhouse in 1909. Intakes for this plant are located on both the North Fork of the Middle Fork of the Tule and the South Fork of the Tule. This powerhouse would later be acquired by SCE.

The San Joaquin Power Company constructed a powerhouse on the North Fork of the Middle Fork of the Tule at the intake of the Mount Whitney plant; it derives its water solely from that fork. After 11 years of construction, this powerhouse opened in 1914. This facility would later be acquired by PG&E.

Water for the San Joaquin Power Company powerhouse is diverted from higher up on the North Fork of the Middle Fork of the Tule. The headworks where the water is diverted are located in Wishon Canyon above Camp Wishon in present-day Giant Sequoia National Monument.

The Doyle Springs Association is a private association of 50 members. It was formed in 1915 or 1916 by people from the Porterville, Tulare, and Visalia areas who wanted a place to get their families out of the valley heat during the summer months. The association leased property from the San Joaquin Power Company to construct cabins. The property is located across the river from the power plant headworks. (The cabin association would acquire the land from PG&E in 1967.)

Smokey McCrea recalled what happened when the Tule flooded through Doyle Springs in December 1955. At that time, several families had cabins near the river. Among these were Frances Barrows (Smokey’s grandmother) and J. G. Boswell, II. These cabins had survived the big floods of the 1930s, but the 1955 flood was different.

Smokey and his family had spent Thanksgiving of 1955 at the cabin, and his grandmother told them where to dig up some brandy and wine that his grandfather (Stanley Barrows) had made and buried during Prohibition. They were both drinkable and they toasted him with their dinner. They had a group of friends who happened to be in camp join them, because they had gone overboard and cooked a suckling pig in the wood stove oven; and a turkey and a goose in the two Westinghouse Roaster Ovens. It was a festive day.

A few weeks later, that cabin was history! The entire main cabin was gone. Nothing was ever found of the cast iron wood stove with the water heater built into the top. The only evidence of the cabin was the foundation of the stone fireplace from the living room. The stone terrace and the large boulders where they used to slide planks in concrete grooves to dam up the river to make their pool deeper — all gone! The only trace was some concrete steps on the far side of the river. The other four bottles of the brandy were lost too.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

The cabins at Doyle Springs get their water supply from springs on the other side of the Tule River. The water distribution system for the entire camp was wiped out because the flood tore out the delivery pipes that carried the water across the river from the springs to the cabins.

Springville is located just downstream from the confluence of the North Fork and Middle Fork of the Tule River. Springville had its water system flooded and sustained some other damage during the 1955 flood. Twenty families had to be evacuated.\(^{1309}\)

The Tule River flowing past Porterville flooded, but its peak flow of 13,900 cfs was considerably less than the 25,500 cfs recorded in the 1950 flood. Still, it was a river to be reckoned with. The Santa Fe Railway trestle bent under the onslaught of the floodwaters, but survived. The highway system was not so fortunate. Traffic coming into Porterville from the west had to be detoured when the east approach to the West Olive Street Bridge caved in, revealing a cavity the width of the highway and 12 feet deep. Homes along the riverbank in Porterville were threatened, and many were evacuated. Thousands of acres were flooded west of Porterville. The highway to Visalia was cut off north of Lindsay.

Damage caused by flooding on the Tule was minimal compared to that caused by the Kaweah. However, Harlan Hagen, the local congressman, used the 1955 flood as an opportunity to push through funding for both Terminus and Success Dams. Success Dam would be completed in 1961.

Over 15 inches of rain over a two-day period caused some flooding along the Kern River. Homes and roads were flooded during this time. The state fish hatchery, rebuilt after being destroyed by flooding in 1950, was washed away again. They lost 683,000 fish this time. The areas near Kernville were evacuated.\(^{1310}\)

The North Fork of the Kern River at Kernville peaked at 27,600 cfs. (That was the peak hourly flow; the peak average daily flow was 12,787 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 13 years for the Kern River.

Isabella Dam had been completed in 1953. That reservoir held virtually all the runoff from the 1955 flood, protecting Bakersfield from flooding. It impounded 55,000 acre-feet during the month of December.\(^{1311}\)

On December 23, President Eisenhower declared Northern California a major disaster area. The following day, he added two Nevada counties to the disaster list.

The North Fork of the Kern River was the southernmost stream listed by the California Department of Water Resources as setting any record in the December 1955 flood.\(^{1312}\) During the period that northern communities were being deluged, the southern part of the state was experiencing a drought.

About one month after the December flood, Central and Southern California experienced flooding as a result of an intense rainstorm on January 25, 1956. The degree of flooding from the January storm varied dramatically by area.

The Kings River flooded on January 25, apparently brought on by a sudden downpour. It was not a major flood. Peak flow at Pine Flat Dam was less than 20% of flows experienced in the 1955 flood.

There was no report of damage in the national parks from the January flood.

Heavy rain fell in the lower Sierra on January 26, 1956, causing rivers to swell. Numerous farms were flooded near Chowchilla and Madera. Flooding was most significant near Visalia.\(^{1313}\) The flows in January 1956 were apparently much lower than they had been in December 1955. With a little luck, the flooding in Visalia should have been relatively minor.

However, luck was not with Visalia. There was apparently no grating at the head of Mill Creek to keep debris from flowing into the aqueduct/conduit that goes under the town. A large tree with a big root wad floated in with the floodwaters. It was some 3 feet in diameter and about 12–14 feet long. The story soon started that it was a sequoia log, but that seems to be just an urban legend.

When the log got into the downtown area (under Garden Street, between Main and Center), it jammed. A major (unseen) debris jam quickly built up behind it. All that water pressure was not to be denied. Geysers started erupting at various places in the vicinity of the debris jam. One of the major clusters of those eruptions was
inside the Harvey Hotel (aka Harvey House). Another major cluster of geysers erupted directly out of Garden Street. A sandbag enclosure was constructed in an attempt to contain where the water was erupting out of the street, but that failed. There was simply too much water. Mill Creek was essentially forced to the surface and started flowing through town on the city streets (multiple photographs on file in the national parks).

Some 40,000 sandbags were used to convert streets and alleys into canals to handle the runaway floodwaters. More than 50% of Visalia was underwater at some point during the flood. Floodwaters inundated up to 72 city blocks for five days. A “glory hole” was cut through the pavement on Church Street, allowing the water to flow back into the aqueduct/conduit.

The USACE brought in a truck-mounted clamshell and started excavating, searching for the log that had to be under the city streets somewhere. Before it was all done, there were a total of four gaping holes in Visalia’s city streets (two on Garden, one on Church, and one on Center). Eventually the huge log was found and extracted. It was rather like a root canal, done on a huge scale.

Along with the key log, at least eight truckloads of logs and timbers were removed from the hole where the debris jam had formed. On January 30, Mill Creek resumed flowing in its underground aqueduct/conduit. The city council decided to leave the sandbags in place because the levee on the south bank of the St. Johns was deemed to be in serious condition, and it was feared that another flood might occur before the season ended. The USACE estimated that damage in Visalia would be on the order of $1.5 million. The Harvey Hotel was so severely damaged by the flood that the building was condemned.

A screen of welded steel pipe to catch floating debris was installed across Mill Creek at Burke Street (near E. Center St.) where the underground conduit begins. This was apparently the first such screen ever installed to prevent blockage in that conduit.

The USACE said that the December 1955 – January 1956 flood was the largest and most damaging rain-flood known to have occurred in northwestern Tulare County since the turn of the century (1900) and prior to completion of Terminus Dam in 1962.\footnote{1314}

In addition to flooding Visalia, the January flood caused damage elsewhere in Tulare County. The Kaweah River washed out a levee where the People’s Ditch takes water out of the Kaweah. This levee had just been rebuilt by the USACE, but they had rebuilt it out of sand because they were required by law to rebuild it out of the same material as before. In the January 1956 flood, the St. Johns washed out a 150-foot section of the south-bank levee northwest of Visalia. The Tule River washed away much of the repair work that had just been done on the Oettle Bridge.

The flooding was so serious in Southern California that it merited an article in the \textit{New York Times}.\footnote{1315} Los Angeles received seven inches of rain in what was described as one of the worst rainstorms in Southern California history. Some 1,500 people had to abandon their homes as a result.

Total flow for water year 1956 was 152\% of the 1894–2014 average for the Kings, 171\% for the Kaweah, 153\% for the Tule, and 125\% for the Kern. Flooding occurred in the Tulare Lakebed in water year 1956. It resulted from runoff from both the December 1955 flood and the January 1956 flood. Tulare Lake had been dry from about July 1, 1953, through December 23, 1955.

The December 1955 rain-type storm on the Kings River Basin was outstanding in both peak and volume of flow, exceeding the previous record runoff of December 1950. However, the entire runoff above Pine Flat Dam was controlled by the reservoir, and no floodflows from the Kings reached Tulare Lake.

The runoff from the Kaweah River was also a record for rain-type storms, exceeding that of December 1950 in both peak and volume. Tule River flows were not as great as the 1950 quantities.

Runoff on the Kern River was also very large. The North Fork of the Kern River at Kernville (see Station 1860 in Table 63) peaked in December 1955: 27,400 cfs. That tied the record set in the 1950 flood. (The December 1966 flood would have a discharge more than twice as great (60,000 cfs.)\footnote{1316} Isabella Reservoir was able to completely control the December 1955 flood.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Inflows to the Tulare Lakebed, in large part from the Kaweah River, are shown in Table 56.\textsuperscript{1317}

Table 56. Inflow to the Tulare Lakebed during water year 1956.

<table>
<thead>
<tr>
<th>Month</th>
<th>Lakebed Inflow (acre-feet)</th>
</tr>
</thead>
<tbody>
<tr>
<td>December, 1955</td>
<td>19,000</td>
</tr>
<tr>
<td>January, 1956</td>
<td>36,800</td>
</tr>
<tr>
<td>February, 1956</td>
<td>10,000</td>
</tr>
<tr>
<td><strong>Total</strong></td>
<td><strong>65,000</strong></td>
</tr>
</tbody>
</table>

Three of the lakebed sumps were flooded. It is unclear whether these portions of the lakebed were included in the calculation of the 183,000 acres of lake flooding that occurred around Visalia in the 1955–56 flood.

The 1955–56 inflow to Tulare Lake was eventually distributed over various portions of the lakebed and absorbed into the ground. The lake was dry by about April 21, 1956 and would remain dry until the 1958 flood.

The principal damages in the lakebed in the 1955–56 flood were the loss of a crop of barley growing on the flooded land, the loss of irrigation equipment, and the erosion of levees and land. Total losses in the lakebed were about $575,000.

1957 Flood

Flooding in 1957 occurred in June. It was a snowmelt flood.

The winter of 1956–57 was a moderate to weak La Niña event. This association with the 1957 flood was probably a coincidence. Only strong La Niña events have been shown to have any correlation with high precipitation events and floods in California.

We know about this flood only from the stream gage record.

The peak day natural flow at Pine Flat on the Kings River occurred on June 4: 13,077 cfs. The Kings remained high from June 2–8. Thanks to the presence of a recording stream gage, we know that the South Fork Kings peaked on June 4: 7,220 cfs.\textsuperscript{1318} That was nearly as big as the much more famous June 1952 flood (7,460 cfs).

The peak day on the Kern occurred on June 5: 3375 cfs. The Kern remained high from June 3–9.

1958 Floods (3)

There were three periods of flooding during 1958:
1. March (west of Mendota)
2. April (near Coalinga and west of Mendota)
3. Flooding occurred in the Tulare Lakebed during February–June as the result of a combination rain and snowmelt flood event.

The winter of 1957–58 was a strong El Niño event.

According to the national parks’ monthly report, March was particularly wet in the parks. There were 22 days of storms during the month. As shown in Table 57, this resulted in more than twice the average precipitation at all three of the parks’ reporting stations.

Table 57. Total precipitation during March 1958.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>1958 Precipitation (inches of rain)</th>
<th>Average Precipitation (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Grant Gove</td>
<td>16.49</td>
<td>6.84</td>
</tr>
<tr>
<td>Giant Forest</td>
<td>17.90</td>
<td>6.51</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>9.72</td>
<td>4.38</td>
</tr>
</tbody>
</table>

Most of the precipitation at higher elevations fell as snow during March. The biggest storm of the month occurred on March 11–17, dropping 38 inches on Giant Forest and 52 inches on Grant Grove. By the end of
March 1958, Grant Grove had received a total of 252 inches (21 feet) of snowfall, compared with 96 inches at the same time the previous year. Giant Forest had received 218 inches by the end of March 1958.

The storm system seemed to break at the end of March. It felt like spring had arrived, giving the national parks time to prepare facilities for normal spring opening. Then April 1–8 brought an almost unparalleled late storm, dropping 72 inches at Grant Grove and 57 inches at Giant Forest. Precipitation and snowfall were the greatest since 1952 and came close to all-time records for a single storm. The weather for the remainder of April was generally clear with below-average temperatures, which greatly reduced the flooding in the valley below.

On March 16, heavy rain triggered debris flows that caused a bridge to wash out 21 miles west of Mendota. A car drove into the raging water, resulting in one boy being killed. A series of storms off the coast with an associated series of fast-moving fronts swept over California during late March and early April, 1958. The San Joaquin Valley experienced several small tornadoes. Thunderstorms were widespread. We know about two of these: one in Stanislaus County and one near Coalinga. We have much better information on the one that occurred in Stanislaus County.

Woodward Dam is located seven miles northwest of Oakdale in Stanislaus County. On April 3, Woodward Dam received 5.72 inches of rain, an amount equal to 45% of its average annual rainfall. That is 8.55 standard deviations above the average maximum daily rainfall with a recurrence interval of almost 300,000 years.

Sometime in April, there was a major flood event near Coalinga (presumably from Los Gatos and/or Warthan Creeks). The flood no doubt continued downstream on Arroyo Pasajero. It mainly affected agricultural lands and public facilities such as roads and bridges. This was one of the three biggest flood events to occur in the Coalinga area during historic times. Panoche/Silver Creek west of Mendota also flooded in April 1958.

Flood releases of 8,000 cfs or greater occurred at Friant Dam on the San Joaquin River during the April–July snowmelt period. Total flow for water year 1958 was 148% of the 1894–2014 average for the Kings, 151% for the Kaweah, 162% for the Tule, and 152% for the Kern.

Flood events in the Tulare Lake Basin in 1958 were a combination of rain and snowmelt. The rains began in February and continued into April. On the Kings River, much of the precipitation above Pine Flat Dam fell as snow, although some intense rain also occurred at low elevations. The rain-flood runoff which occurred in April was well below the record of December 1955. The snowmelt runoff, which began in late May, was well below the 1906 record runoff.

Earl McKee, Jr. witnessed the results of the numerous avalanches that occurred in the backcountry of the national parks during the winter of 1957–58. Two particularly massive avalanches occurred between Grouse and Simpson Meadows in the Devils Washbowl area on the trail along the Middle Fork of the Kings.

Flooding occurred in Tulare Lake in 1958. The lake had been dry from about April 21, 1956 until March 31, 1958. Pine Flat Dam had been completed in 1954. It contained most of the runoff from the Kings River. However, a small amount of Kings water did reach Tulare Lake, largely in June.

Runoff from the Kaweah and Tule River Basins in the 1958 flood was like that from the Kings River in that the runoff from rain was well below the December 1950 record runoff and the snowmelt runoff was less than the 1952 record runoff. However, considerable water did reach Tulare Lake.

Inflows to the Tulare Lakebed are detailed in Table 58.

<table>
<thead>
<tr>
<th>Stream</th>
<th>Total Inflow (acre-feet)</th>
<th>Percent Contribution</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings River</td>
<td>24,000</td>
<td>14%</td>
</tr>
<tr>
<td>Kaweah River</td>
<td>75,000</td>
<td>44%</td>
</tr>
<tr>
<td>Tule River</td>
<td>72,000</td>
<td>42%</td>
</tr>
<tr>
<td>Kern River</td>
<td>0</td>
<td>0%</td>
</tr>
<tr>
<td>Total</td>
<td>171,000</td>
<td></td>
</tr>
</tbody>
</table>
No Kern River water reached Tulare Lake during the 1958 flood. Flows in the Kern were completely controlled by the operation of Isabella Reservoir, which stored about 350,000 acre-feet of runoff between March 1 and June 25.

Maximum depth of water in Tulare Lake was about 9.7 feet on April 20, 1958. However, by August 14, the lake was dry. Presumably much of the water had been used to irrigate lands which had not been flooded in April and May. The lake then remained dry until December 6, 1966.

At some point in 1958 or 1959, a short section of the Colony Mill Road slid off the hillside. That section of the road was immediately downhill of the junction with the Admiration Point Trail. It’s tempting to think this slide was a result of the very wet winter of 1958. In any case, the slide resulted in the Colony Mill Road remaining closed until June 1960 when it had to be hurriedly reopened to support the Tunnel Rock Fire. Rather than constructing a short bypass around the section of road that had failed, a mile or so of new road was constructed (the Over-the-Hump Road) on the other side of the ridge. Bill Tweed recalled that the reason for doing this was to provide a road that was not as exposed to the fire, which was on the south side of the Ash Peaks Ridge.

Earl McKee, Jr. witnessed how high the grass could grow in a year like 1958 that had abundant spring rain. He described a ride he took on a ridge above Greasy Cove across what is now Lake Kaweah. Earl recalled that the wild oats (presumably *Avena fatua*) were high enough that you could tie them in knots over your saddle horn. The grass got so high in places that his horse couldn’t see where he was going and would get panicky and start lunging. Earl would have to get off, knock down the grass, and lead his horse through dense patches like that.1324

### 1959–61 Drought

This drought affected the entire state, but was most extreme in the Sierra and Central Coast. Recurrence intervals were greatest along the Central Coast, in the Sierra, and in the Southern California desert (30–75 years).

Table 59 shows how the San Joaquin Valley Water Year Index categorized the drought years in that basin. It also shows total runoff for the four major rivers in our basin (Kings, Kaweah, Tule, and Kern) combined during this drought. It appears that the drought ended in the Tulare Lake Basin in water year 1962, although that may still have been a drought year in the San Joaquin River Basin.

<table>
<thead>
<tr>
<th>Water Year</th>
<th>San Joaquin River Basin</th>
<th>Tulare Lake Basin</th>
</tr>
</thead>
<tbody>
<tr>
<td>Water Year Classification</td>
<td>Total Runoff (acre-feet)</td>
<td>% of Average (1894–2014)</td>
</tr>
<tr>
<td>1959</td>
<td>Dry</td>
<td>1,282,392</td>
</tr>
<tr>
<td>1960</td>
<td>Critically dry</td>
<td>1,235,021</td>
</tr>
<tr>
<td>1961</td>
<td>Critically dry</td>
<td>882,675</td>
</tr>
<tr>
<td>1962</td>
<td>Below normal</td>
<td>3,011,944</td>
</tr>
<tr>
<td>Drought average (1959–61)</td>
<td>1,133,363</td>
<td>39%</td>
</tr>
</tbody>
</table>

1959 was the driest year on record in Kern County; total precipitation during the year was less than two inches.

Total annual flows for the rivers within the Tulare Lake Basin were generally less than 50% of average for water years 1959–61. In the southern half of the state, 1961 was the driest year of the drought, ranking among the driest years of record at many sites. Total flow for water year 1961 was 33% of the 1894–2014 average for the Kings, 28% for the Kaweah, 18% for the Tule, and 26% for the Kern. The Tule and the Kern both set new minimum flows-of-record, breaking the records set in 1924. The Tulare Lake Basin wouldn’t see flows this low again until 1977 (see Figure 18 on page 111 and Table 23 on page 156).

On January 22, 1962, Fresno experienced its biggest snow in 32 years when 2.2 inches fell. The snow closed schools and caused a rush of people to stores seeking to buy film to photograph this unusual event. Many roads were slippery and some were closed altogether. Five people died on valley roads due to the slick conditions. Other amounts in the valley included 4.0 inches at Madera, 3.0 inches at Wasco, 2.0 inches at Hanford, Avenal, Buttonwillow, and 1.5 inches at Los Banos. The higher elevations were buried in snow, 33 inches was reported at Badger Pass in Yosemite.1325 But this was just an interesting interlude, the drought would continue for another year.
In 1963, a record number of mid-winter foggy and rainless days were recorded at Sacramento associated with high barometric pressures and stagnant winds. This was one of the worst mid-winter droughts of record in Central California.¹³²⁶

National park records indicate that this was a severe and extended drought. In the last week of January 1963, there was so little snow in the parks that three people were able to complete a trip to East Lake, Reflection Lake, and over Langley Pass to South Guard Lake. They made it just in time. The drought ended abruptly with a major storm that began on January 29, 1963.

The similarity of meteorological conditions of the 1860–61 and the 1959–61 droughts are notable. Both were severe droughts that ended with severe flooding.

Tulare Lake was dry throughout the 1959–61 drought. The lake went dry in August 1958, and would stay dry until the 1966 flood brought it back to life, if only for a few brief months.

**1962 Floods (2)**

There were two floods in 1962:

1. a small rain-flood in February
2. a small snowmelt flood in May

We know about these floods only from the stream gage record.

The peak day natural flow at Pine Flat on the Kings River occurred on February 10: 10,236 cfs.

The Kaweah’s peak natural flow occurred at Terminus Dam on February 10: 8,000 cfs. (That was the peak hourly flow; the peak average daily flow at the dam was 3,707 cfs).

The Tule peaked on February 10: 1,337 cfs. The Kern peaked on February 11: 2,438 cfs.

The second flood of the year was a snowmelt flood. In that flood, the peak day natural flow at Pine Flat on the Kings River occurred on May 6: 12,724 cfs. The Kings remained high from May 4–9. Thanks to the presence of a crest-stage gage, we know that the South Fork Kings peaked on about May 6 at 5,600 cfs.¹³²⁷

The Kaweah’s peak average daily flow at Terminus Dam during the runoff occurred on May 5: 2,652 cfs. The Kaweah remained high from about May 3–9.

The runoff came early on the Tule in 1962. The peak day occurred on April 9: 573 cfs. The runoff on the Tule in 1962 really didn’t amount to a flood in any conventional sense.

The peak day natural flow on the Kern occurred on May 6: 3,574 cfs. The Kern remained high from about May 4–10.

**1963 Flood**

Flooding in 1963 occurred in February. It was caused by a storm that came out of the North Pacific.

Most of January was very dry but extremely cold in the national parks. There was little snow on the ground even at the high elevations. The combination of extreme cold and lack of snow caused damage to the national parks’ water systems. This was considered the worst mid-winter drought (or extended dry season, depending on your point of view) in the state’s history.

The dry season of the summer 1962 lasted an unusually long time, well into the winter. It was finally broken by a storm that lasted from January 29 – February 2, 1963. This is often treated as a three-day storm, lasting from January 30 – February 1. The rainfall distribution in this storm was quite similar to the November 18–19, 1950 storm. The rainfall of both these storms was heavy in the coastal mountains as well as in the Sierra.

This 1963 storm resulted in the heaviest-ever three-day rainfalls at 45 stations. These extreme rainfalls were generally at high elevations in the Southern Sierra. The heaviest rainfalls were centered south of Yosemite. Florence Lake received 64% of its average annual precipitation in this storm, which represented a recurrence interval of 33,000 years. Other Sierra stations with a recurrence interval greater than 1,000 years were the South Entrance of Yosemite National Park and Tollhouse.¹³²⁸
Table 60 gives the total precipitation during the January 29 – February 2, 1963 storm event for selected reporting stations.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of moisture)</th>
</tr>
</thead>
<tbody>
<tr>
<td>South Entrance of Yosemite</td>
<td>22.99</td>
</tr>
<tr>
<td>Wishon Dam (near Shaver Lake)</td>
<td>23.25</td>
</tr>
<tr>
<td>Grant Gove</td>
<td>17</td>
</tr>
<tr>
<td>Giant Forest</td>
<td>21</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>12</td>
</tr>
</tbody>
</table>

The snowline associated with this storm was generally over 8,000 feet and at times as high as 11,000 feet. Snowmelt was a major factor in the flooding associated with this storm. Many streams reported record-high flows during this storm. The snowline on the west side of the Great Western Divide was 7,000–9,000 feet. Snow depths were progressively greater to the east and north.

Major flooding occurred to the north of the Tulare Lake Basin, including the cities of Napa, Marysville, and Reno.

Within the national parks, mudslides (or debris flows) occurred in the drainage above Simpson Meadow, suggesting considerable rain before the snow began.

The peak day natural flow at Pine Flat on the Kings occurred on February 1. That is the sixth largest peak day of the year at Pine Flat since the dam was built in 1954. This was a bigger flow than occurred in the much more famous 1983 flood. It seems likely that this was a very high-flow period in Cedar Grove as well.

Thanks to a pair of crest-stage gages, we know that Grizzly Creek peaked on about February 1: 293 cfs. This was the highest flow recorded on that creek between 1960–1973.

The Kaweah’s peak natural flow occurred at Terminus Dam on February 1: 30,900 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 18,405 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 16 years for the Kaweah. Runoff from this storm was very rapid and the Kaweah River reached flood stage before dawn on February 1. Lake Kaweah was able to catch much of this runoff, preventing flooding downstream. Despite the intensity of the storm and the resulting flooding, the national parks received less damage than expected.

In the Kern River Basin, over 14 inches of rain fell in February over a short duration. Forty people were evacuated from their homes in the Kernville area as floodwaters from the North Fork of the Kern River threatened their homes. Once again the state fish hatchery sustained damages, and all the fish were lost — some 225,000 rainbow trout that were about to be released. Everyone staying in low-lying areas was evacuated.

The South Fork Kern River near Onyx (see Station 1895 in Table 63) peaked in February 1963: 3,460 cfs. This was the greatest discharge on that river since record-keeping began in 1911. However, the December 1966 flood (28,700 cfs) would be eight times greater.

There was also flooding on the west side of the valley during 1963. Damage was sustained largely by agricultural lands and public facilities. The area around Coalinga was one of the areas that was damaged. We know this happened sometime in 1963, but we don’t know that it was during the February storm.

**1964 Flood**

Flooding in 1964 occurred in December.

The winter of 1964–65 was a weak La Niña event. This association with the 1964 floods was probably a coincidence. Only strong La Niña events have been shown to have any correlation with high precipitation events and floods in California.

The flooding resulted from meteorological conditions similar to those of the December 1955 flood. An arctic air mass moved into Northern California on December 14, and precipitation on December 18–20 produced large
quantities of snow. Beginning on December 20, a storm track 500 miles wide extended from Hawaii to Oregon and Northern California. Warm, moist air collided with the arctic air and resulted in turbulent storms that produced unprecedented rainfall on Northern California and melted much of the snow from the previous storms. In the Mattole River Basin, nearly 50 inches of rain was reported during December 19–23, with 15 inches observed in 24 hours.

The six-day period from December 19–24 was the wettest ever recorded at 78 Northern California stations. The North Coast had the worst flooding ever experienced in that region. Every major stream in the North Coast produced new high values of extreme peak flows. A total of 34 California counties were declared disaster areas. This storm had three major centers of activity: the Eel River, the Upper Klamath and the Yuba River in the Central Sierra.\textsuperscript{1333}

Branscomb in the Eel Basin received 31.71 inches during the storm event. Most stations in the Eel River Basin reported their highest-ever rainfall during this storm. Gazelle in the Klamath River Basin reported 8.09 inches. That was 7.78 standard deviations above the average with a recurrence interval of over 300,000 years. A total of 35 stations reported daily rainfalls of 10 inches or more on December 22. These stations were located in the North Coast streams as well as in the Central Sierra. The highest-ever rainfalls occurred in the Yuba and Bear River Basins, where Lake Spaulding (east of Grass Valley) received 32.60 inches of rain during the storm event.\textsuperscript{1334}

Floods were widespread across the northern half of the state. The main center of precipitation was in the Feather, Yuba, and American River basins. Runoff from streams of the Coast Ranges, almost without exception, produced peak stages and peak flows that exceeded previous records. Runoff from the Sierra into the Feather, Yuba, and American rivers surpassed all previous records.\textsuperscript{1335}

Bridges on every major stream were destroyed. Several towns along the Eel and Klamath Rivers were totally destroyed. The floods caused $239 million in property damage and 24 deaths statewide. The property damage in north coastal California was about 50% greater than had occurred in the December 1955 floods. The December 1964 flood remains the greatest known flood in the history of Northern California.

Exceptionally large flood peaks were recorded on rivers in north coastal California. Peak discharges of the Eel, Klamath, and Smith Rivers were 30–40% greater than the 1955 peaks and exceeded flood stages of the 1861–62 floods. In Humboldt Redwoods State Park, faded paint marks high up on giant redwoods still record the high stage on the Eel River.

The American River experienced a record flood, the third record flood in less than 15 years. Several rivers, including the Salmon and Klamath, experienced a flood event with a recurrence interval of greater than 100 years. Botanic and geomorphic evidence indicated that floods exceeding the magnitude of the 1964 flood may not have occurred since about 1600.

(During the 1600–1610 time period, a dramatic climatic change was happening across North America. Major precipitation events occurred in the areas near the Sacramento River Basin, Mono Lake, Southern California, Mexico, and elsewhere. See the section of this document that describes the California megafloods for more about that event.)

The Klamath is the second-largest river in California: more than twice as large as the third largest river. It drains a 12,000 square mile watershed that extends more than 260 miles inland to the Klamath Basin in Oregon. The December 1964 flood was the most devastating flood of the Klamath ever. It swept away all of downtown Klamath, destroyed the U.S. Highway 101 bridge, and washed away a great many homes. The river peaked at over 550,000 cfs, 17% greater than the average flow of the Mississippi River.

At the time of the flood, an 800 pound Angus bull named Bahamas was living in the valley of the Klamath. On December 22, the Klamath topped out at 52 feet, covering its valley in a maelstrom of churning logs, brush, lumber and debris. Bahamas was swept up in all the debris and carried downriver and out to sea. He apparently survived the ride by climbing on top of a raft of flotsam.

Bahamas rode there on the open ocean on a constantly disintegrating raft of logs and brush, through huge storm waves. Eventually he and his raft arrived at the Crescent City Harbor, 16 miles up the coast from the mouth of the Klamath River. Somehow he had stayed aboard his accidental raft of slippery, tossing logs and brush.
Bahamas was discovered the next day, 200 feet offshore in the 10 acre mass of floating, churned debris that plugged Crescent City Harbor. He was helped to go from log to log until he reached shore. He was more or less adopted by the town. A novel was written about him: Beloved was Bahamas. He lived out his life in his own grassy paddock in Klamath, his feat commemorated by a large sign on the fence. He was visited by many people over the years who regarded him as a living touchstone of courage and will. He died in the spring of 1983 and was buried in his green pasture.

The peak flow in the American River at Fair Oaks, controlled by Folsom Dam, reached 115,000 cfs. In the remaining watersheds of the Sacramento Valley, peak stages and flows tended to equal those experienced in 1955. At the Sacramento Weir, all 48 gates were opened, and the peak flow reached 84,000 cfs. The peak flow in the Sacramento River at I Street was about 100,000 cfs. Total peak flow at the latitude of Sacramento, including the Sacramento River and the Yolo Bypass, was about 475,000 cfs.¹³³⁶

The December 1964 floods did not extend as far south as those of December 1955. In the Sacramento River Basin, many streams had peak discharges that were greater than during December 1955. However, peak discharges in the San Joaquin River Basin were substantially less than during 1955. In both basins, flood-control operations generally were able to confine downstream flows within flood-control channels. As a result, loss of life was avoided, and damage was less than half that caused by the 1955 flooding.

Although the worst of the storm was in the north, it still brought significant precipitation to the Tulare Lake Basin. The peak day natural flow at Pine Flat on the Kings occurred on December 24, 1964. It was obviously a flood, but it was only half as high as the peak day natural flow for 1963. Thanks to a crest-stage gage on the South Fork Kings, we know that stretch of the river peaked on about December 24: 5,200 cfs.¹³³⁷

As detailed in Table 61, the national parks experienced intense storm activity between December 19–28. This consisted of three separate storms, the most intense of which occurred on December 27–28.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Grant Gove</td>
<td>7.45</td>
</tr>
<tr>
<td>Giant Forest</td>
<td>12.61</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>6.17</td>
</tr>
</tbody>
</table>

The national parks did not report any significant flooding. Most of the precipitation in Giant Forest consisted of rain. The storm lowered the snow on the Wolverton ski area to an unsafe level.

The heavy rains caused some damage to the Generals Highway. On December 27, a portion of the wooden cribbing one mile above Ash Mountain failed, leaving a one-lane roadway for about 60 feet. (That wood cribbing was replaced with a galvanized metal bin wall, constructed by the national parks, in April 1965.) A section of dry rubble retaining wall at the 5,000 foot elevation also failed because of the heavy rains.

The winter 1964 flood is remembered because it did so much damage in Northern California. However, the winter 1963 flood was a more memorable event in the Southern Sierra.

1965 Floods (2)

There were two floods in 1965:
1. March
2. August

The winter of 1964–65 was a weak La Niña event. This association with the 1965 floods was probably a coincidence. Only strong La Niña events have been shown to have any correlation with high precipitation events and floods in California.

Fresno received 1.55 inches of rain on March 12, setting a daily rainfall record. Most of the rain fell in a five-hour window from 4 p.m. to 9 p.m., inundating streets and poor drainage areas with water described as up to hip deep. A number of transformers in the city shorted out, plunging many homes and businesses into darkness.¹³³⁸
Daily computed unimpaired flow (full natural flow) at Pine Flat Dam suggests that the Sierra experienced a series of storms during the week of August 11–17. From August 11–19, Cedar Grove had daily rains with downpours that caused debris flows which blocked the road to traffic until cleared away. Cedar Grove had 2.81 inches of rain for the month of August.

Heavy rain occurred in the national parks during the week of August 11–17. Park records document that this storm affected the area of the South Fork Kings. Thanks to a pair of crest-stage gages, we know that Grizzly Creek peaked on about August 17: 247 cfs.1339

The Marble Fork of the Kaweah flooded through Lodgepole Campground on August 17 (photograph on file in the national parks).

1966 Flood

Flooding in 1966 occurred in December.

A very large storm brought a strong inflow of warm moist Pacific air across Central California from December 3–7. The transport mechanism for the moisture was an atmospheric river.1340 In the Tulare Lake Basin, the most severe effects of the storm were felt south of the Kings River. The storm penetrated deeply inland, bringing significant moisture into the Owens Valley. The heaviest rain was in a narrow band that ranged from SCE’s Kern River Intake #3 in the south to the White Mountains over 100 miles to the northeast.

December 1966 was the wettest five days ever at 58 California stations in an area stretching from the Kern River to the White Mountains, and into Tulare and San Bernardino Counties. A total of 19 stations reported 10 inches or more of rain on December 6. These stations were located mainly in Tulare and San Bernardino Counties.

The heaviest 24-hour rainfall ever recorded in the Central Valley, 17.0 inches, occurred on December 6 at Hockett Meadow. This record would last for 20 years. It would eventually be exceeded by the 17.6 inches recorded at Four Trees in the Feather River Basin on February 17, 1986.

The record downpour that occurred on Hockett Meadow and the surrounding plateau on December 6, 1966, generated unprecedented runoff. Cañon Meadow was predisposed for erosion by years of heavy grazing. It is possible that this was the event that initiated the gullying that we see today in Cañon Meadow. By 2014, the erosion gully was 1200 feet long, up to 17 feet deep, 56–92 feet wide, and had resulted in a 67 foot change in elevation from the headcut to the outlet of the gully.

A total of 42 stations recorded their highest-ever 5-day rainfalls during this storm event. A total of 11 stations reported rainfall totals in excess of a storm with a recurrence interval of 1,000 years. The highest rainfall was reported at Johnsondale with a 5-day total of 30.45 inches.1341

Table 62 shows the elevation and precipitation for the December 2–7 storm event. Springville received its greatest precipitation of the storm event on December 5. The other reporting stations received their greatest amount on December 6.1342

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Elevation (approximate)</th>
<th>Storm Total (inches of rain)</th>
<th>Drainage Basin</th>
</tr>
</thead>
<tbody>
<tr>
<td>General Grant</td>
<td>6,600</td>
<td>23.04</td>
<td>Kings</td>
</tr>
<tr>
<td>Giant Forest</td>
<td>6,358</td>
<td>27.75</td>
<td>Kaweah</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>1,730</td>
<td>15.52</td>
<td>Kaweah</td>
</tr>
<tr>
<td>Three Rivers Powerhouse #2</td>
<td>950</td>
<td>11.85</td>
<td>Kaweah</td>
</tr>
<tr>
<td>Springville Ranger Station</td>
<td>1,050</td>
<td>10.78</td>
<td>Kaweah</td>
</tr>
<tr>
<td>Johnsondale</td>
<td>4,680</td>
<td>30.45</td>
<td>Kern</td>
</tr>
<tr>
<td>Glennville</td>
<td>3,140</td>
<td>8.62</td>
<td>Poso</td>
</tr>
<tr>
<td>Wofford Heights</td>
<td>2,700</td>
<td>11.00</td>
<td>Kern</td>
</tr>
<tr>
<td>Kern River Powerhouse No. 1</td>
<td>970</td>
<td>3.46</td>
<td>Kern</td>
</tr>
</tbody>
</table>

Rain fell as high as 9,000 feet. The rain apparently melted all the snow on the ground at Grant Grove, Giant Forest, and the Wolverton Ski Bowl. However, there may not have been a particularly heavy snowpack to melt. Grant Grove experienced an exceptionally severe rain and wind storm on the night of December 5. That event
brought down a 100-foot forked-top sugar pine, demolishing a visitor cabin that was fortunately closed for the season.

On December 7, the weather turned cold, reducing the amount of flooding. A second storm had been feared, but stayed north of the Tulare Lake Basin. Mountain Home received 23 inches of rain during the storm, and Camp Wishon (northwest of Camp Nelson) reported 36 inches of rain in 69 hours. One source said that the December 2–7 event was the most severe rainstorm on record in the southern San Joaquin Valley.

Paso Robles received 5.25 inches of rain in 24 hours on December 6, setting a record for that city.

Some rivers had a recurrence interval greater than 100 years. It was the first significant storm of the winter. Flooding occurred throughout Northern California, including on the Russian and Eel Rivers. However, this was not to be a replay of the 1964 flood. This time, flooding was most severe in the Kaweah, Tule, and Kern River Basins. San Bernardino and Riverside Counties also sustained serious flooding.

Continuously above-average precipitation from December 1966 through March 1967 resulted in the flooding of 35,000 acres of the northern San Joaquin River Basin. The San Joaquin River above Millerton Lake experienced high runoff during early December. A maximum mean daily inflow of 18,450 cfs was recorded at Friant Dam. However, releases of only 52 cfs were made to the San Joaquin River.

Significant amounts of flooding occurred in both mountain areas and on the valley floor. A total of 141,800 acres flooded, including 122,400 acres of valley floor and 19,400 acres in mountain and foothill areas. These record-breaking floods inundated parts of the towns of Kernville, Springville, Three Rivers, Lindsay, and Lamont.

Panoche/Silver Creek west of Mendota also flooded in December 1966.

The peak day natural flow at Pine Flat on the Kings occurred on December 6, 1966. That is the second-largest peak day of the year at Pine Flat since the dam was built in 1954. (The flood-of-record occurred on December 23, 1955.)

Thanks to a crest-stage gage on the South Fork Kings, we know that reach of the river peaked on about December 6: 11,800 cfs. That put this flood about midway in size between the 1950 flood (10,000 cfs) and the 1955 flood (13,900 cfs). There was major damage to the state’s portion of Highway 180 in Kings Canyon.

Based on the flood exceedence rates in Table 29, this had a recurrence interval of 70 years for the Kings River downstream at Pine Flat. That puts it in a category with other Cedar Grove floods of the past 70 years (1937, 1950, 1955, 1966, 1969, 1978, 1982, 1983, and 1997) that rise to the level of the modeled 50-year flood: that is, a flood event that occurs about every eight years on average. See the section of this document that describes Cedar Grove Flooding.

Pine Flat Reservoir appears to have caught (or at least diverted) the entire flood on the Kings. No floodwaters from the Kings made it into the Tulare Lakebed during 1966.

In the Tulare Lake Basin, the December 1966 flood was generally the biggest flood-of-record on most major streams south of the Kings since the 1867–68 flood. These include Sand Creek draining the area west of the North Fork Kaweah River and Deer Creek draining the area west of the Kern River and south of the Tule River, and Poso Creek draining the area west of the lower Kern River Basin.

Past records of peak flow and 3-day storm-runoff volume in the Kaweah, Tule, and Kern River Basins were greatly exceeded by the floods of December 1966. Extremely high peak discharges occurred at most gaging stations between 11:00 p.m. December 5 and 6:00 p.m. December 6. Snowmelt was not a major cause of the floods, although some snow that had accumulated during minor November and early December storms was melted. Thanks to a USGS report, the December 1966 flood was particularly well documented. Table 63 summarizes the flood discharge data for selected streams in the Tulare Lake Basin.
Severe flooding extended over the Kaweah, Tule, and Kern River Basins in a 60- by 100-mile area in the Sierra northeast of Bakersfield. Moderate flooding occurred in the Kings River Basin and other basins to the north and in streams draining from the Coast Ranges to the west. Flood peaks were the greatest of record at many gaging stations in the Kaweah, Tule, and Kern River Basins. Damage was severe in all headwater areas. Culverts were overflowed or plugged with debris, or usually a combination of both. Most highway bridges were destroyed or severely damaged.1352

The flood caused major damage to roads, trails, and other facilities in the national parks.

Highway 180 was closed at Snowline Lodge by a rock slide, isolating Grant Grove. According to the parks’ monthly report, there was major damage to the state’s portion of Highway 180 in Kings Canyon. However, there was only minor damage to the section of the road that was within the park boundaries.

Table 63. Summary of peak flood discharges for the December 1966 storm event.

<table>
<thead>
<tr>
<th>River</th>
<th>Station ID</th>
<th>Period of record</th>
<th>Discharge (cfs)</th>
<th>Maximum December 1966</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kern River near Quaking Aspen</td>
<td>1853.5</td>
<td>1960–66</td>
<td>4,060</td>
<td>6, 9,360</td>
</tr>
<tr>
<td>Little Kern River near Quaking Aspen</td>
<td>1854</td>
<td>1955, 1957–66</td>
<td>12,200</td>
<td>6, 13,100</td>
</tr>
<tr>
<td>Kern River near Kernville</td>
<td>1860</td>
<td>1912–66</td>
<td>27,400</td>
<td>6, 60,000</td>
</tr>
<tr>
<td>Kern River at Kernville</td>
<td>1870</td>
<td>1905–66</td>
<td>38,700</td>
<td>6, 74,000</td>
</tr>
<tr>
<td>South Fork Kern near Olancha</td>
<td>1882</td>
<td>1956–66</td>
<td>1,280</td>
<td>6, 1,010</td>
</tr>
<tr>
<td>South Fork Kern near Onyx</td>
<td>1895</td>
<td>1911–14, 1919–42,</td>
<td>3,460</td>
<td>6, 28,700</td>
</tr>
<tr>
<td></td>
<td></td>
<td>1947–66</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Kelso Creek near Weldon</td>
<td>1897</td>
<td>1958–66</td>
<td>1,340</td>
<td>6, 5,800</td>
</tr>
<tr>
<td>Kern River below Isabella Dam</td>
<td>1910</td>
<td>1945–66</td>
<td>39,000</td>
<td>30, 2,160*</td>
</tr>
<tr>
<td>Kern River near Democrat Springs</td>
<td>1925</td>
<td>1950–66</td>
<td>40,000</td>
<td>6, 10,100*</td>
</tr>
<tr>
<td>Kern River near Bakersfield</td>
<td>1940</td>
<td>1893–66</td>
<td>36,000</td>
<td>7, 9,290*</td>
</tr>
<tr>
<td>Poso Creek near Oldale</td>
<td>1978</td>
<td>1958–66</td>
<td>2,750</td>
<td>6, 4,300</td>
</tr>
<tr>
<td>White River near Ducor</td>
<td>1995</td>
<td>1942–53, 1943</td>
<td>2,300</td>
<td>6, 1,080</td>
</tr>
<tr>
<td></td>
<td></td>
<td>1958–66</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Deer Creek near Terra Bella</td>
<td>N/A</td>
<td>N/A</td>
<td>N/A</td>
<td>N/A 6, 10,000</td>
</tr>
<tr>
<td>North Fork of Middle Fork Tule River near Springfield</td>
<td>2020</td>
<td>1939–66, 1955</td>
<td>12,400</td>
<td>6, 16,900</td>
</tr>
<tr>
<td>North Fork Tule River at Springfield</td>
<td>2031</td>
<td>1957–66</td>
<td>4,600</td>
<td>5, 24,200</td>
</tr>
<tr>
<td>Tule River near Springfield</td>
<td>2032</td>
<td>1950–66</td>
<td>22,400</td>
<td>6, 49,600</td>
</tr>
<tr>
<td>South Fork Tule River near Springfield</td>
<td>2045</td>
<td>1930–54, 1950</td>
<td>7,100</td>
<td>6, 14,300</td>
</tr>
<tr>
<td>Middle Fork Kaweah near Potwisha</td>
<td>2065</td>
<td>1949–66</td>
<td>46,800</td>
<td>6, 23,300</td>
</tr>
<tr>
<td>Marble Fork Kaweah at Potwisha</td>
<td>2080</td>
<td>1950–66</td>
<td>12,500</td>
<td>6, 6,400</td>
</tr>
<tr>
<td>East Fork Kaweah near Three Rivers</td>
<td>2087.3</td>
<td>1952–55, 1963</td>
<td>2,850</td>
<td>6, 13,000</td>
</tr>
<tr>
<td></td>
<td></td>
<td>1957–66</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Dorst Creek near Kaweah Camp</td>
<td>2090</td>
<td>1959–66</td>
<td>1,540</td>
<td>6, 2,010</td>
</tr>
<tr>
<td>North Fork Kaweah River at Kaweah</td>
<td>2095</td>
<td>1910–66</td>
<td>21,500</td>
<td>6, 23,900</td>
</tr>
<tr>
<td>Kaweah River at Three Rivers</td>
<td>2099</td>
<td>1955, 1963</td>
<td>30,900</td>
<td>6, 73,000</td>
</tr>
<tr>
<td></td>
<td></td>
<td>1958–66</td>
<td></td>
<td></td>
</tr>
<tr>
<td>South Fork Kaweah River at Three Rivers</td>
<td>2101</td>
<td>1955, 1955</td>
<td>10,000</td>
<td>6, 11,600</td>
</tr>
<tr>
<td></td>
<td></td>
<td>1958–66</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Kaweah River at McKay’s Point</td>
<td>N/A</td>
<td>1905–66</td>
<td>84,332</td>
<td>6, 105,000*</td>
</tr>
<tr>
<td>Dry Creek near Lemon Cove</td>
<td>2113</td>
<td>1959–66</td>
<td>1,600</td>
<td>6, 14,500</td>
</tr>
</tbody>
</table>

*Some of the gages in the above table were below dams. Unless noted otherwise in the text, the associated discharges for the 1966 storm generally don’t reflect unimpaired flow (full natural flow) of those rivers because the flows were affected by storage and/or diversion upstream. That is not the case with the discharges for those gages that have been included in Table 28 on page 159. In order for the latter table to show unimpaired flow (full natural flow), those discharges have been adjusted to remove the effects of dams upstream of the gages.
As shown in Table 63, Dorst Creek’s peak natural flow occurred on December 6: 2,010 cfs. A crest-stage gage (USGS 11209000) was located on Dorst Creek near where the Generals Highway crosses that creek. That gage was operated from May 1960 through May 1973. The December 1966 flood is the biggest flow recorded on Dorst Creek during the period of record. It was 30% greater than the 1963 flood, the previous high flood in the area.

Giant Forest was isolated for a week by slides on the Generals Highway between the two national parks; access between Grant Grove and Giant Forest wasn’t restored until December 14.

According to the parks’ monthly report, the section of the Generals Highway between Giant Forest and Ash Mountain was badly damaged:

- The Marble Fork Bridge near Potwisha was washed away (photograph on file in the national parks). This bridge had been badly damaged in the 1938 flood and washed out altogether in the 1955 flood. This time it would be rebuilt in a new location, a hundred feet or so upstream. Manuel Andrade recalled that, as an interim measure, a national park crew constructed a temporary highway bridge at the old location using brace-and-bits and crosscut saws. That bridge was open for administrative traffic in less than two weeks.
- Numerous major rock falls and landslides (multiple photographs on file in the national parks), some of which contained several thousand cubic yards each. From the photographs, one of those appears to have been just below Amphitheater Point.
- Slides / washouts in 10 locations, each of which required cribbing to restore the road to two-lane width. One of those was at Alder Creek (photograph on file in the national parks). Leroy Maloy recalled that one of the 10 sections was a big failure above Deer Ridge that was replaced with a metal bin wall in the summer of 1967. Howard Mancha (oiler on the crane that placed that bin wall) recalled that the Deer Ridge bin wall reconstruction happened after the April 1965 construction of the Ash Mountain bin wall.

According to the parks’ monthly report, park crews had pushed a one-lane road through the slides by December 30. This finally permitted administrative access between Ash Mountain and Giant Forest. All other travel to Giant Forest had to be by way of the Big Stump entrance.

It isn’t clear when this section of the Generals Highway was reopened to visitor traffic. At the least, visitors were probably kept off this section of the Generals Highway until a modern bridge was built across the Marble Fork Kaweah at Potwisha. That presumably took at least a year. This section of highway had probably been closed for the summer of 1956 when the previous bridge had to be replaced.

There are suggestions that the general practice in those years was to keep the Generals Highway open to single-lane visitor traffic even when bin walls were being constructed in the Deer Ridge area. It appears that no concrete barricades were used for such construction; only cones and lane delimiters were used.

As shown in Table 63, the East Fork Kaweah’s peak natural flow occurred on December 6: 13,000 cfs. A stream gage (USGS 11208730) was located on the East Fork at the diversion dam for SCE’s flume, about ¼ mile above where the Mineral King Road crosses the river.1353 It was a complex gage (a water-stage recorder coupled with an acoustic velocity meter) designed to give accurate measurements in variable and low velocity flow situations. The gage was operated from June 1952 through October 2010. It was knocked out of operation during the 1955 flood. The gage apparently wasn’t read from October 1979 – September 1993. The December 1966 flood is the largest flow recorded on the East Fork during the period of record. It was 15% greater than the January 1997 flood.

As shown in Table 63, the North Fork Kaweah’s peak natural flow also occurred on December 6: 23,900 cfs. A stream gage (USGS 11209500) (a water-stage recorder) was located on the North Fork a mile above the Upper North Fork Bridge (the Bailey Bridge). The gage was operated from October 1910 through September 1981. The December 1966 flood is the largest flow recorded on the North Fork during the period of record. It was 11% greater than the December 1955 flood.

As shown in Table 63, the South Fork Kaweah’s peak natural flow also occurred on December 6: 11,600 cfs. A stream gage (USGS 11210100) (a water-stage recorder) was located on the South Fork about ½ mile upstream from where Highway 198 crosses that river, 200 feet upstream from an unnamed tributary. The gage was operated from October 1958 through September 1990, but partial data is available for a longer period extending from December 1955 through January 1997. The December 1966 flood is the largest flow recorded on the South Fork during the period of record. It was 16% greater than the 1955 flood and 36% greater than the 1997 flood.
Specific damage in Three Rivers:

- Ash Mountain was isolated for two days because the national parks’ approach to the Pumpkin Hollow Bridge was so badly eroded that only pedestrian traffic was allowed across. The parks’ approach to this bridge had completely washed out in the 1937 and 1955 floods.¹³⁵⁴
- Both approaches to the Dinely Bridge were badly eroded, but the bridge survived.¹³⁵⁵
- The Upper North Fork Bridge (the Bailey Bridge) washed out. Residents past that point were isolated from the rest of the community. This bridge had also washed out in the December 1955 flood. The Bailey Bridge may have been the surviving segment of the North Fork Bridge across the mainstem of the Kaweah by the present-day Three Rivers Market that was washed out in the December 1955 flood.
- The North Fork Road washed out a short distance past the Upper North Fork Bridge and was apparently badly damaged in one or more other locations.
- The Airport Bridge on the North Fork Kaweah was either badly damaged or washed out.
- Many properties on the North Fork were damaged. The River Isle Trailer Park was heavily damaged (see below for details). The C&D Trailer Park was also damaged. Archie McDowall lost 2,500 chickens and a new pickup. Luther Smeltzer lost his home. B.F. McKinley, C.E. Fairman, and Leroy Maloy received extensive damage to their homes.
- Heavy damage totaling $1.75 million was reported to homes along the South Fork Kaweah. The South Fork flooded as severely as the North Fork, but there were relatively few homes on the South Fork.
- Huffaker’s Candy Shop (now Reimer’s) and Calloway’s Drive-In were flooded. These were the only two Three Rivers businesses that were damaged in this flood. Other businesses were threatened, but these were the only businesses that were actually flooded.¹³⁵⁶
- The Three Rivers Golf Course was heavily damaged. It was cut in two by a re-channeling of the mainstem of the Kaweah. The river’s new course isolated the main clubhouse, destroyed fairways, and washed away green #2 and part of green #6. Many logs were washed onto the fairways.¹³⁵⁷
- Kaweah Public Beach (aka River Park) at Cobble Knoll on the lower side of Three Rivers was washed out. (The exact name of that county park is unclear.) The river was up to the edge of Highway 198 at this point.
- SCE suffered $40,000 in damage to its transmission lines and distribution systems in the Three Rivers area.

The River Isle Trailer Park was located several miles up North Fork Drive. It is reported to still be functioning as a trailer park of some sort in 2012. The trailer park was surrounded and overtopped by the 1966 flood, scattering trailers everywhere. Mobile homes stood on end, upside down, and sideways, completely ruined by the water. The situation was so dire that 19 people had to be evacuated by helicopters from Lemoore Naval Air Station. On one flight, a Lemoore helicopter ran out of fuel and had to make a forced landing with civilians on board. A commentary on the odd things people do in an emergency was the comment from a helicopter crew member that one girl evacuated from the trailer park took with her only hair curlers, hairspray, a comb, and a change of clothing.¹³⁵⁸, ¹³⁵⁹

Bobbie McDowall Harris was living on the North Fork Kaweah near the River Isle Trailer Park at the time of the flood. She has a vivid memory of the floodwaters separating their family (they had five little kids at the time) from the road. As the river started rising, Bobbie and her husband decided to save their cars by moving them up onto the North Fork Road. Her husband went first in the Ford, and she followed him in a VW bus. He got stalled halfway through the water. Knowing that she had an engine in the rear, Bobbie yelled for him to get back in the car and she’d push him. It worked, and they ended up on the road. They rushed back to their children before the water got too high. It was very scary. They were stranded in their home for many days, cut off by the flooding river. They were offered the opportunity to get their family out by helicopter. However, they chose to stay put, feeling that getting kids that young in a hovering helicopter was more dangerous than staying.

The Upper North Fork Bridge was replaced in 1967. The new bridge has low-profile guardrails that can be removed in anticipation of an oncoming flood. The bridge hasn’t washed out since.

Sometime after the 1966 flood, the USACE constructed a levee from the Upper North Fork Bridge down past the McDowall property all the way to the River Isle Trailer Park.

Movie maker Harold Schloss was filming his movie One on Beetle Rock at the time that the flood hit. Floodwaters covered the runway of the Three Rivers Airport, keeping employees from flying in food for his menagerie of characters: mountain lions, a wolf, a hawk, a raven, a badger, and two coyotes. All were stranded on location at the Roping Arena. Shooting of the movie was delayed for several days. The upside was that there was a storm scene in the movie and Schloss got some excellent storm footage.

Fifteen-foot waves were reported to have been common on the mainstem of the Kaweah in Three Rivers at the peak of the flood.¹³⁶⁰ Floodwaters apparently came up several feet on the river side of the Three Rivers Market.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Shortly before noon on December 6, it was announced that some businesses in Three Rivers would have to be evacuated.

The December 1955 flood had been the biggest flood on the Kaweah since record-keeping began. Comparing the size of the 1966 flood in Three Rivers with the 1955 flood is not straightforward.

Floodmarks at two gaging stations in Three Rivers were observed to be slightly higher in the 1966 flood than in the 1955 flood. One of those gages was on the South Fork Kaweah and one was on the mainstem of the Kaweah, upstream from Lake Kaweah.

There was a gage near Three Rivers on the mainstem of the Kaweah (USGS gage #11-2105 Kaweah River near Three Rivers) located just above the junction with Horse Creek (latitude: 36:24:24N longitude: 118:57:12W). The total drainage area upstream of that gage was 519 square miles. That gage was operated from 1903–1961. But by 1966, it was submerged under Lake Kaweah.

Lake Kaweah is roughly comparable to the former Three Rivers gage. It has a total drainage area of 561 square miles. Lake Kaweah is located slightly farther downstream, but it measures essentially the same tributary streams as the former Three Rivers gage. The difference is that the Lake Kaweah gage also measures the inflow from Horse Creek.

The computed maximum bihourly inflow to Lake Kaweah on December 6, 1966 was 82,700 cfs. That was only 2% greater than the peak flow at the former Three Rivers gage on December 23, 1955: 80,700 cfs.\textsuperscript{1361}

Realizing how big the 1966 flood was predicted to be, a temporary sack-concrete barrier was placed on the spillway of Terminus Dam to provide additional storage.\textsuperscript{1362} That proved to be good planning. About 10,250 acre-feet of floodwater was surcharged against that spillway barrier, bringing the reservoir to almost 5½ feet above full-pool level.

The USACE would place these temporary sack concrete barriers on the spillways of Terminus Dam and Lake Success in anticipation of the 1966, 1967, and 1969 floods, removing them after each flood.

Lake Kaweah reached a peak storage of 147,200 acre-feet just before 2 a.m. on December 8, 1966.\textsuperscript{1363} Runoff from the upper basin was completely controlled by the reservoir, and no downstream releases were made until after the first of the year when Dry Creek and other downstream tributaries had subsided.\textsuperscript{1364} Thanks to Terminus Dam (and a break in the weather), Visalia escaped with no flooding other than from surface water. An evacuation center had been set up in the city but wasn’t needed.

The Lake Kaweah marina was ripped loose from its moorings and set afloat in the middle of the lake. Many acres of debris clogged the upper end of the lake following the flood. This included logs, trees, shoes, chickens, dead fish, boats, and general flotsam.\textsuperscript{1365} Most of the heavier woody material was removed somehow the following summer, presumably by burning.

As reflected in Table 28, the peak average daily flow occurred at Terminus Dam on December 6: 53,280 cfs.

Dry Creek below Terminus Dam peaked at 14,500 cfs, the highest flow since record-keeping began on that stream. (For comparison, this was 44% greater than Merced River in Yosemite Valley in the much more famous January 1997 flood.) This flow on Dry Creek had a recurrence interval of approximately 85 years.\textsuperscript{1366} Dry Creek Road was closed from Lemon Cove to Badger during the flood. Culverts were washed out on the Eshom Valley Road. About 48,600 acre-feet of Kaweah River floodwater reached Tulare Lake in December. Dry Creek was the principal source of most of that water, with lesser amounts contributed by Yokohl, Cottonwood, Sand, Lewis and Cameron Creeks.

The Kaweah’s peak natural flow occurred at McKay’s Point on December 6: 105,000 cfs. It was the largest peak day of the year since record-keeping began in 1905; it remains the flood-of-record for this stretch of the river for floods that have occurred since 1905. However, recorded history on the Kaweah Delta began in about 1850. During recorded history, the 1867–68 flood is considered the flood-of-record for the Kaweah; that flood just occurred before any stream gages were installed on this river.

Above Terminus Dam, the 1955 and 1966 floods were roughly equal in size. However, because Dry Creek added so much water, the flood below Terminus Dam was a much more impressive event. The peak average daily flow
at McKay’s Point in 1966 was 20% higher than the more famous December 1955 flood (peak average daily flow of 53,280 cfs in 1966 versus 44,512 cfs in 1955).

Keep in mind that the above flows for 1966 (105,000 cfs and 53,280 cfs) reflect the unimpaired flow (full natural flow) of the river, after adjusting to remove the effects of the dam upstream of the gage. It is how big the 1966 flood would have been if Terminus Dam had not been there. (A much earlier report had estimated the peak natural flow at McKay’s Point as 120,000 cfs. But that has been replaced with the current estimate of 105,000 cfs.)

Based on the flood exceedence rates in Table 29, this flood had a recurrence interval of 170 years for the Kaweah. It would have had a recurrence interval of 100 years if calculated using the 105,000 cfs peak flow. (One source reportedly calculated this as having had a recurrence interval of 140 years. Presumably that was done using the peak flow and the now outdated 1971 flood frequency curves. That result could not be reproduced.)

Yokohl Creek flooded dramatically, undermining the Visalia Electric track adjacent to Highway 198. It also put four to five inches of water into the Yokohl Store and damaged at least one house in the area. Bridges were damaged along Yokohl Creek and at Rocky Hill. Yokohl Creek crosses Highway 245 (Road 204) about 1½ miles north of Highway 198. Just west of that point, Yokohl Creek flows into the Consolidated Peoples Ditch. This is near the Lower Kaweah River but well below McKay’s Point.

Large areas flooded along Cottonwood and Cross Creeks.

Flooding occurred in Lindsay, East Woodlake, Terra Bella, and some isolated areas near Cutler, Orosi, Yettem, and Seville. Flooding was up to three feet deep in Toneyville near Lindsay; 150 people were evacuated from that community.

The flood of December 1966 was one of the largest known to have occurred on the Tule River. Although there are no formal comparative records, historical accounts indicate that the flood of December 1867 was of about the same magnitude as the flood of December 1966.

The Tule was flowing 18,000 cfs at Globe at 10 p.m. on December 5 when the gage was swept away. That was probably well before the peak of the flood.

Doyle Springs is a private recreation cabin complex located on the North Fork of the Middle Fork of the Tule River. It is located in Wishon Canyon above Camp Wishon in present-day Giant Sequoia National Monument.

Smokey McCrea recalled what happened when the Tule flooded through Doyle Springs in December 1966. Several families had cabins near the river, including Frances Barrows (Smokey’s grandmother) and J. G. Boswell, II. The family cabin belonging to Smokey’s grandmother had been destroyed in the 1955 flood. The Barrows’ family rebuilt their cabin in 1959, building on higher ground behind the site of the original cabin.

The December 1966 storm dropped 29 inches of rain in 24 hours onto a moderate snowpack upstream of the Doyle Springs camp. The flood caused a lot of deadfall to flow into jams that created dams in the river. When the accumulated weight of water caused those dams to break, walls of water, logs, and boulders rolled down the streambed, overflowing the banks and destroying anything in their path. The Barrows’ cabin sustained substantial damage, and their driveway was impassable. The just-rebuilt cabin downstream of them was totally destroyed, and the Boswell cabin next to them was mostly destroyed.

The cabins at Doyle Springs get their water supply from springs on the other side of the Tule River. Three 4-inch iron pipes had been welded into a triangular truss that spanned the river below a waterfall, with a crown in the center of the span. The debris in the flood battered that truss and broke the crown, causing the truss to sag and nearly break. After the flood, the present-day cable was added to support the three pipes.

Five cabins washed away at Camp Wishon northwest of Camp Nelson.

Some local residents recall that Highway 190 over the North Fork Tule remained passable throughout the flood. However, the Visalia Times-Delta printed an aerial photo at the time showing that this bridge had washed out. Will Wood and George Costa both confirmed newspaper reports from the time that the lower and upper
bridges on the Globe Drive loop were washed out. The USACE confirmed that three bridges (one of which was the Highway 190 bridge) were destroyed and two others badly damaged.

There was damage to the bridges on the North Fork of the Tule. The Bear Creek Bridge on the road leading to the SCICON school was destroyed. The North Fork Bridge (located about 7 miles up the North Fork of the Tule on the Balch Park Road) was overtopped; the approaches may have been damaged, but the bridge survived.\textsuperscript{1374}

Damage was extensive in Springville with 9–12 homes washed away, 35 homes extensively damaged, and 75–100 people left homeless. Springville’s domestic water supply system was knocked out of commission early in the flood, and it was many days before it could be brought back on line.\textsuperscript{1375, 1376} The Springville sewage treatment plant was severely damaged. The golf course downstream from Springville was also extensively damaged.\textsuperscript{1377}

The Tule Indian Reservation was hit hard. Most of the American Indians live along an eight-mile stretch of the Tule. The floodwaters washed out roads, destroyed all bridges, and swept away telephone and power lines leading into the area. Bulldozers were flown into the reservation to begin construction of a makeshift road. Other equipment went to work from the Porterville end, expecting to meet within four days.

Johnsdale in the southeast part of Tulare County lost power. It also lost all road access, being cut off by the floodwaters of both the Tule and Kern Rivers. Helicopters from Lemoore Naval Air Station were used to fly in food, water, and emergency generators. Later, a temporary access road — rough, but passable — was opened over Parker Pass to California Hot Springs.\textsuperscript{1378}

The December 1966 maximum flow of 49,600 cfs at the gaging station on the Tule River near Springville (see Station 2032 in Table 63) was the greatest flood-of-record and more than double the previous record flow of 22,400 cfs in November 1950. Records at a former gage site inundated by Lake Success in 1961 show that the 1950 flood was the greatest during the period of record (1901–1960). The December 1955 peak discharge was slightly less than that in 1950 and thus was the third-highest recorded flood on the Tule River.\textsuperscript{1379}

The December 1966 peak discharge of 14,300 cfs on the South Fork Tule River near Success (see Station 2045 in Table 63) was also more than double the previous record flow of 7,000 cfs in November 1950.\textsuperscript{1380}

The computed maximum bihourly inflow to Lake Success near Success of 52,800 cfs on December 6 similarly was 1.7 times the peak flow of 32,000 cfs in November 1950 at a former gaging station near the damsite.\textsuperscript{1381} That was the peak bihourly flow. The reservoir had an instantaneous inflow on December 6 of 76,900 cfs and a daily inflow on that date of 40,000 cfs; both of these are flows of record.\textsuperscript{1382}

(The same qualifier applies here as on the Kaweah. This is the greatest flow measured since stream gages were installed on the Tule. However, during recorded history, the 1867–68 flood is considered the flood-of-record for the Tule; that flood just occurred before any stream gages were installed on this river.)

The peak average daily flow at Success Dam, as reflected in Table 28, was 40,085 cfs. It was the largest peak day of the year since that dam was built in 1961; it remains the flood-of-record. Based on the flood exceedence rates in Table 29, this had a recurrence interval of 200 years for the Tule.

Terminus Dam had been able (just barely) to capture the floodwaters of the Kaweah and prevent flooding downstream. Success had a stated capacity in 1966 of 85,400 acre-feet. The reservoir filled, but the floodwaters kept coming. Water began spilling. At one point, the reservoir was holding 101,400 acre-feet (85,400 acre-feet of nominal storage + 16,000 of surcharge storage).\textsuperscript{1383}

About 250 people were evacuated in the Porterville area, mostly from the Doyle Colony. The Tule River broke through levees in the Porterville area. Flooding in the lower Tule continued for several days. Heavy flooding was reported at the Pixley National Wildlife Refuge. However, for all that, Success Dam prevented the bulk of the flooding that would have otherwise occurred. At Porterville, an official put it succinctly: Without Success Dam there would be no Porterville today.\textsuperscript{1384}

Some flooding occurred in agricultural areas downstream from Lake Success during sustained release of floodwater December 6–11.\textsuperscript{1385} Approximately 1,000 sheep were discovered marooned on a levee next to the Tule near Ave 184 and Rd 152 in the Woodville area. They were too heavy to swim because of their full coats of wool. County crews evacuated them by constructing a temporary bridge. Many other cattle, sheep, and other
animals weren’t so lucky and drowned in the flood, contributing to the health problems caused by flooded wells.\textsuperscript{1386}

As shown in Table 64, the Tule sent nearly as much floodwater into the Tulare Lakebed in 1966 as the Kaweah did. That rarely happens.

Deer Creek near Terra Bella peaked on December 6: 10,000 cfs. This is the flood-of-record for that creek. It was a particularly destructive flood. The main road to California Hot Springs follows Deer Creek and crosses it at several locations. All bridges were destroyed or badly damaged. Downstream irrigation diversion structures were washed out and were further damaged by deposition of coarse sediment.\textsuperscript{1387}

White River flooded on December 6, but it was only half as big a discharge as the 1943 flood.\textsuperscript{1388}

There was a major flood on Caliente Creek in December, causing extensive flood damage to the Lamont/Arvin area.

There was a severe rainstorm over the Kern River Basin on December 2–7. Almost 21 inches of rain fell in the area in two days.\textsuperscript{1389} The flooding was most severe in the Kernville area. Flooding there isolated an area of 150 square miles and forced the evacuation of 200 persons. All roads in that area were under water.

Peak flows of the December 1966 flood exceeded previous maximums at most gaging stations in the Kern River Basin except:\textsuperscript{1390}

1. High elevation stream channels. Little storm runoff occurred above 9,000 feet elevation where the precipitation fell during part of the storm as snow and during the remainder as rain, as the freezing level changed during the storm. When inspected on December 9, the channel of Golden Trout Creek was nearly full of ice and frozen saturated snow. There was no evidence of substantial high flow during the storm. Snow already on the ground apparently absorbed and held most of the rain that fell upstream from this gaging station.
2. Stream channels below the Isabella Dam.

The flood damaged many stream gages. It destroyed the water-stage recorder structures and the measuring cableways at two sites:\textsuperscript{1391}
- North Fork Kern River at Kernville
- South Fork Kern River near Onyx

The South Fork Kern River near Onyx (see Station 1895 in Table 63) had a peak discharge of 28,700 cfs. This was eight times greater than the previous maximum discharge of 3,460 cfs recorded in the February 1963 flood. That makes the 1966 flood by far the highest flow on the South Fork since record-keeping began in 1911.\textsuperscript{1392}

The North Fork Kern River near Kernville (see Station 1860 in Table 63) had a peak discharge of 60,000 cfs on December 6, 1966. That was more than twice the previous maximum discharge of 27,400 cfs recorded in the floods of November 1950 and December 1955. That makes the 1966 flood by far the highest flow on the North Fork since record-keeping began in 1912.\textsuperscript{1393}

Isabella Reservoir had a computed maximum bihourly inflow of 96,900 cfs. That was 2.5 times the previous maximum flow of 39,000 cfs which was recorded at the damsite in November 1950 prior to dam construction.\textsuperscript{1394} The reservoir had an instantaneous inflow of 118,600 cfs and a daily inflow of 72,782 cfs; both of these are flows of record.\textsuperscript{1395}

The road from Johnsondale downstream to Kernville (Mountain 99, aka Kern County SM99) is close to the North Fork Kern at many locations. This road was obliterated at the outside bank of many river bends, and the pavement was scoured away in other locations.\textsuperscript{1396}

A hydroelectric plant upstream from the Kernville Bridge washed out, hitting the bridge and collapsing it.\textsuperscript{1397} (Presumably this was SCE’s Kern No. #3 power plant.) Bridge debris, including one 3-foot by 40-foot steel girder, was moved several hundred feet downstream.\textsuperscript{1398}

The state fish hatchery, to the surprise of none, washed away.\textsuperscript{1399} The hatchery had been destroyed in the 1950 flood and washed away again in the 1955 flood.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Sequoia National Forest reported that an aerial survey showed considerable amounts of timber felled by the storm and erosion damage. The Kernville Road (State Highway 155) was badly damaged. Large portions of that road, some sections up to 2½ miles long, were washed away.\textsuperscript{1400}

A trailer court and other buildings along the river at Kernville were badly damaged.\textsuperscript{1401} Many people were evacuated from the Kernville area. Prisoners at a work camp were evacuated and sent to facilities in Bakersfield. A section of the golf course at Kernville was washed away. All of the mountain roads were either washed out or closed by landslides. All trailer parks, motels, lodges and cabins were swept away by floodwaters. The fire station at Lake Isabella was flooded. Highway 178 was closed due to flooding and debris. The historic wooden Bellevue Weir just west of Bakersfield washed out. Two people lost their lives in Kern County.\textsuperscript{1402}

The flooding on the Kern River was covered in the \textit{New York Times}.\textsuperscript{1403}

Lake Isabella was able to fully contain the flood. The only release from the dam for the first 10 days during and after the flood was the 300-500 cfs released to the Borel Canal for power production.\textsuperscript{1404}

The USACE estimated that if Lake Isabella had not existed, flow on the Kern River six miles upstream of Bakersfield at the First Point of Measurement gage would have been approximately 80,000 cfs, resulting in significant damage in the city. Actual flow was only 9,300 cfs and consisted primarily of inflow from tributary streams entering the river below the dam.\textsuperscript{1405}

There was significant flooding on the west side of the San Joaquin Valley near Coalinga in December 1966. Flooding caused extensive road and bridge damage on Los Gatos Creek and Warthan Creeks. Flooding continued downstream on Arroyo Pasajero. East of Coalinga, sewage-treatment facilities and the levees along Warthan Creek were damaged, the Los Gatos Creek channel was severely eroded, and there was extensive damage to utilities and agricultural land. Damages totaled approximately $570,000, and floodwaters inundated 4,500 acres.

Many people in Tulare and Kern Counties were displaced from their homes or otherwise needed assistance. The American Red Cross launched a major relief operation. Evacuation centers were operated in Farmersville, Woodlake, Lindsay, Three Rivers, and perhaps other areas. The Red Cross provided food, clothing and other relief services throughout the affected area. They provided services to remote areas as soon as those areas were reachable. This included organizing a pack train to get food, water, and other supplies into the Tule Indian Reservation. The only other way to access the reservation was by helicopter.\textsuperscript{1406}

On December 9, Governor Edmund Brown declared portions of Tulare, Kern, and Riverside Counties to be disaster areas. There were three deaths and $18 million in property damage. Damage to Tulare County roads and bridges was initially estimated to be $2.5 million, but that was soon deemed to be way too low. Damage to roads and bridges in the Kern River area of the county was particularly bad. The Tulare County civil defense chief estimated damage to ranch property as simply “astounding.”\textsuperscript{1407}

Tulare Lake had been dry since about August 14, 1958. It came back to life on December 6, 1966. The river flood occurred in 1966, but the flooding in the lakebed continued into 1967. From the standpoint of the lakebed, it could be thought of as a flood with two phases. The lakebed flooding of December 1966 wasn’t fully dissipated when the April 1967 flooding arrived.

As detailed in Table 64, the December 1966 rain-flood delivered a total of 87,600 acre-feet to the Tulare Lakebed.\textsuperscript{1408}

<table>
<thead>
<tr>
<th>Stream</th>
<th>Total Inflow (acre-feet)</th>
<th>Percent Contribution</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings River</td>
<td>0</td>
<td>0%</td>
</tr>
<tr>
<td>Kaweah River</td>
<td>48,600</td>
<td>55%</td>
</tr>
<tr>
<td>Tule River</td>
<td>37,900</td>
<td>43%</td>
</tr>
<tr>
<td>Deer Creek</td>
<td>1,100</td>
<td>1%</td>
</tr>
<tr>
<td>Kern River</td>
<td>0</td>
<td>0%</td>
</tr>
<tr>
<td><strong>Total</strong></td>
<td><strong>87,600</strong></td>
<td></td>
</tr>
</tbody>
</table>

Table 64. Inflow to the Tulare Lakebed during water year 1966.
This resulted in the flooding of 26,560 acres in the lakebed. However, that area included 14,750 acres flooded by diverted floodwaters for which there was no flood damage. That is, the land was considered of no value. Flooding first occurred in Sumps #1 and 2, after which agricultural land was flooded. After the heavy flows in these two sumps had subsided, much of this water was transferred to lands in the south and southeastern portions of the lakebed. There it would do little harm. This procedure made Sumps #1 and 2 again available for storage of the snowmelt flood runoff which was anticipated to occur later in the flood season.

The upstream federal reservoirs were all full at the end of the flood, and it was just the beginning of the rainy season. It was imperative to empty those reservoirs as soon as possible to restore their capacity to capture the next potential flood. However, the downstream interests in the Tulare Lakebed warned that this would cause a major disaster. Thousands of acres of farmland were already under water as a result of the flood. However, there was no choice; the reservoirs had to be emptied.\textsuperscript{1409}

Bill Cooper recalled that somebody made the trip to San Francisco in a motorboat in 1966.\textsuperscript{1410} This was the fourth of six documented trips between Tulare Lake and San Francisco Bay to occur in historic times. (The other five trips were in 1852, 1868, 1938, 1969, and 1983.)

The storm of December 1966 caused extensive damage to roads all along the eastern slope of the Sierra as well as to the Los Angeles Aqueduct near Lone Pine.\textsuperscript{1411}

1967 Flood
The 1967 flood was a snowmelt flood, extending from April into early July.

A vast amount of snowmelt from April to July compounded the flood damage already experienced from the 1966 flood. Significant flooding also occurred along the Cosumnes River, in the Morrison Creek and Beach–Stone Lake areas, and in Madera County streams in the lower portions of the Fresno and Chowchilla rivers.\textsuperscript{1412} In addition to the snowmelt, it was a wet spring. Fresno experienced its wettest April ever, with over four inches of precipitation.

The valley floor was already flooding from the 1966 flood, but the reservoirs were able to hold only a portion of the 1967 runoff. Table 65 shows the quantities of snowmelt sent downstream in the spring of 1967.\textsuperscript{1413}

<table>
<thead>
<tr>
<th>Basin</th>
<th>Total Flow (million acre-feet)</th>
</tr>
</thead>
<tbody>
<tr>
<td>San Joaquin River Basin</td>
<td>7.8</td>
</tr>
<tr>
<td>Tulare Lake Basin</td>
<td>3.9</td>
</tr>
<tr>
<td>Total</td>
<td>11.7</td>
</tr>
</tbody>
</table>

The federal reservoirs, acting in concert, made every effort to keep floodwaters out of Tulare Lake in the 1967 flood. However, that goal proved to be more than the system was up to. The reservoirs were full at the beginning of the year because of the December 1966 flood. They would have to make releases at some point in order to provide storage space to catch the predicted large snowmelt runoff that was coming in 1967.

There were sustained high flows on the Kings River from mid-May through late July. The peak day natural flow at Pine Flat on the Kings occurred on July 1. Thanks to a pair of crest-stage gages, we know that Grizzly Creek peaked at 233 cfs, probably sometime during June.\textsuperscript{1414}

The maximum mean daily inflow to Pine Flat was 19,739 cfs, which was controlled to an outflow of 15,034 cfs. The natural flow of the Kings peaked at 20,500 cfs, but PG&E cooperated in manipulating the storage remaining in their upstream reservoirs, reducing this flow at the critical time to assist in minimizing outflows from Pine Flat. Pine Flat Reservoir held all the water that it possibly could. It reached a peak stage of 1.3 feet above full pool. Every effort was made to keep Kings River water out of the Tulare Lakebed. However the runoff was simply too large. It couldn’t all be held in the reservoirs or diverted through the Fresno Slough Bypass. About 62,000 acre-feet of Kings River floodwater wound up in Tulare Lake.

Lake Kaweah started the year full as a result of the December 1966 flood. Like all the other federal reservoirs, it did everything possible to keep floodwaters out of Tulare Lake. However, in April and May, 23,000 acre-feet of water was sent down the Kaweah River to the lakebed in preparation for the snowmelt runoff season.
The USACE created a temporary sack-concrete barrier on the spillway of Terminus Dam. This allowed them to surcharge more than 3½ feet above spillway crest level, reaching 156,700 acre-feet on June 30, 1967. This project prevented 116,000 acre-feet from reaching the Tulare Lakebed.

The USACE did something similar on the Tule River. In April and May of 1967, 9,300 acre-feet of water that had been stored in Lake Success during the December 1966 flood had to be released in preparation for the snowmelt runoff season. That water was passed through to the Tulare Lakebed. In mid-May, a temporary sack-concrete barrier was placed on the spillway of Success Dam to provide an additional 10,000 acre-feet of storage. This extra freeboard turned out to be a good idea; Success Reservoir would eventually peak in the summer of 1967 at about 1.6 feet above full pool. Due to flood inflow into Success Reservoir, eventual maximum outflow reached 8,300 cfs, exceeding channel capacity downstream.

During May and June, large releases were made from Isabella Reservoir for spreading and irrigation use. This made it possible to keep the reservoir from going out of control (that is, to keep it from spilling). The maximum reservoir storage reached was 539,000 acre-feet on July 20, 1967. This was 2.8 feet short of reaching the spillway.

Total flow for water year 1967 was 194% of the 1894–2014 average for the Kings, 241% for the Kaweah, 272% for the Tule, and 227% for the Kern. This was the fourth-highest flow for the Kaweah since record-keeping began in 1894.

As detailed in Table 66, about 94,000 acre-feet of floodwater entered Tulare Lake during the April-July snowmelt period. As a result, about 40,000 acres of the lakebed was flooded, a little more than in the 1966 flood. Two levees within the lakebed were in danger of failing during this period. If that had occurred, the acreage flooded would have been significantly larger.

### Table 66. Inflow to the Tulare Lakebed during water year 1967.

<table>
<thead>
<tr>
<th>Stream</th>
<th>Total Inflow (acre-feet)</th>
<th>Percent Contribution</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings River</td>
<td>62,000</td>
<td>66%</td>
</tr>
<tr>
<td>Kaweah River</td>
<td>23,000</td>
<td>24%</td>
</tr>
<tr>
<td>Tule River</td>
<td>9,300</td>
<td>10%</td>
</tr>
<tr>
<td>Kern River</td>
<td>0</td>
<td></td>
</tr>
<tr>
<td><strong>Total</strong></td>
<td><strong>94,300</strong></td>
<td></td>
</tr>
</tbody>
</table>

#### 1969 Floods (3)

There were three floods in 1969:
1. January (rain-flood)
2. February (rain-flood)
3. April through July (unusually large snowmelt)

The winter of 1969 was very wet in Southern California and the Southern Sierra. Southern California experienced some of the severest flooding since 1938.

Over 200 stations, mainly in Southern California, reported their highest-ever rainfalls for 60 consecutive days. Mount Baldy Notch in the San Gabriel Mountains received 88.50 inches in 60 days from January 13 – March 13. Stations reporting extremely high rainfalls for the 60 days ranged from Cottonwood Creek at 10,600 feet in the Southern Sierra to Death Valley at 194 feet below sea level. A total of 13 stations reported rainfall totals in excess of a storm with a recurrence interval of 1,000 years. The valley floor portion of the San Joaquin Valley also had heavy rainfalls with high recurrence intervals.

The 1969 flood was a major flood in the San Joaquin River Basin, especially in January. It was one of the most damaging natural catastrophes in California’s history. Property damage was about $400 million, and 60 lives were lost. Governor Reagan declared a state of emergency for the January storm; a total of 40 counties were declared disaster areas. President Nixon also declared the State of California a disaster area for the January storm.

In the Sacramento Valley, floodwaters produced by the January storms were largely controlled by major reservoirs, flood channels, and the bypass system. As a result, flows in the mainstem of the Sacramento River
and its major tributaries remained well below project design flows. Peak flow at the latitude of Sacramento was approximately 250,000 cfs.\textsuperscript{1419}

During January 18–27, a series of storms, drawing on a strong flow of warm, moist air from the southwest, moved across Central and Southern California. Massive quantities of precipitation fell on the coastal mountains from Monterey Bay to Los Angeles and in the Southern Sierra. Lytle Creek Powerhouse in the San Gabriel Mountains northwest of San Bernardino received 24.92 inches of rain in a 24-hour period on January 24. The peak discharge on the Santa Ynez River near Lompoc was 78\% greater than during the flood of March 1938.

Fresno recorded 8.56 inches of rain in January 1969, making that the wettest month ever for this city. In all, 22 days of the month recorded precipitation.\textsuperscript{1420} As of January 27, Dinuba had received 21 inches of rain for the season, an all-time record. The USACE said that precipitation during the January storm event varied from slightly more than 8 inches at Terminus Dam to more than 45 inches in some headwater regions. During the most intense period of the storm series, 6\% inches of precipitation occurred in one 24-hour period in the upper reaches of the Kaweah drainage area.\textsuperscript{1421}

As the heavy rains continued in the valley, a snowpack of unprecedented depth and water content accumulated in the mid- to higher elevations of the Sierra. Record after record was broken during the winter of 1968–69:

- The Central Sierra Snow Laboratory monitoring site near Donner Pass received 13.7 feet of snow between January 20–31, the second biggest snowstorm ever recorded at that site. (An even bigger snowstorm would come in March–April 1982.)
- In late February, a series of northwestern cold-front storms moved south along a low-pressure trough that had formed over the California coast. Incredible all-time 24-hour snowfall records were set in parts of the Sierra on February 24–25 with 46.0 inches of snow measured at Lodgepole and 36.0 inches of snow at Grant Grove.\textsuperscript{1422}
- Lodgepole received 187 inches (15.6 feet) of snowfall during the month of February. This is the greatest amount of snowfall ever recorded in one month at that location.
- On February 26, 1969, the snowpack at Lodgepole reached 197 inches (16.4 feet). This is the greatest snowpack ever recorded at that site.\textsuperscript{1423} Despite the similarity of numbers (187 and 197), this is a different record from the preceding bullet. "Total snowfall" refers to new snow falling during a storm event. It is computed by summing the 24-hour snowfalls measured daily during the time period of interest. "Snowpack" refers to the total amount of snow on the ground, including existing snow from previous storms. The snowpack may actually be less than the total snowfall as the snow at the lower depths may be compacted by the weight of the overlying snow.
- On March 13, Grant Grove measured a snowpack of 179 inches (14.9 feet), the greatest ever recorded at that site.\textsuperscript{1424} Much more was to come.
- The Montecito-Sequoia Lodge was damaged from a 20-foot snow-dump.
- As detailed in Table 67, Lodgepole received a total snowfall of at least 440.5 inches (36.7 feet) during the winter of 1968–69. (This only reflects snowfall after November 8, 1968, the date that the cooperating weather station was reactivated. Therefore, the total snowfall for the winter may have been somewhat under-measured.) This is the fourth biggest winter at that location. The winters of 1905–06, 1951–52, and 2010–2011 were all larger.

<table>
<thead>
<tr>
<th>Month</th>
<th>Snowfall (inches of snow)</th>
</tr>
</thead>
<tbody>
<tr>
<td>September 1968</td>
<td>no data</td>
</tr>
<tr>
<td>October 1968</td>
<td>no data</td>
</tr>
<tr>
<td>November 1968</td>
<td>13.0</td>
</tr>
<tr>
<td>December 1968</td>
<td>67.0</td>
</tr>
<tr>
<td>January 1969</td>
<td>93.5</td>
</tr>
<tr>
<td>February 1969</td>
<td>187.0</td>
</tr>
<tr>
<td>March 1969</td>
<td>50.0</td>
</tr>
<tr>
<td>April 1969</td>
<td>27.0</td>
</tr>
<tr>
<td>May 1969</td>
<td>3.0</td>
</tr>
<tr>
<td><strong>Total</strong></td>
<td><strong>440.5</strong></td>
</tr>
</tbody>
</table>

Just as that storm system was starting to form, two State of California snow surveyors, Doug Powell and Murt Stewart, were preparing to embark on a nine-day trip into the upper Kern River watershed. Their planned route would take them through a portion of the Golden Trout Wilderness in eastern Tulare County.
Snow-survey courses are established throughout the Sierra, and the snowpack at these sites needs to be measured four times through the winter. The resulting data helps hydrologists forecast the spring and summer runoff for the rivers that flow into the Central Valley.

Doug and Murt were accomplished backcountry skiers. Doug was also a respected professor of geography at UC Berkeley who loved snow survey work and took time off from his campus work each winter between 1956 and 1984 to ski into the High Sierra. What follows comes from his account of this particular trip and a column that Bill Tweed wrote about it.

What Doug and Murt did not understand as they started off on the morning of February 22, was that they were skiing into one of the 20th century’s most intense Sierra Nevada blizzards. The weather report for the area, obtained from their pocket transistor radio, called for intermittent snow showers, “heavy at times.”

The first obstacle the two faced was to ski over 11,000-foot-high Cottonwood Pass. Deep, soft snow from previous storms made the going slow. As they began the ascent, the cloud cover largely disappeared, and a blast of cold air came downslope. They wondered if perhaps a cold front was passing overhead with following clear weather.

But when they reached the top of the pass and took in the view, their opinion changed abruptly. The cloud cover for many miles to the west was a textbook example of an approaching major storm. This indicated not intermittent showers, but heavy, prolonged snowfall.

They hurriedly measured the snow course just west of the pass. Moving as fast as the difficult conditions allowed, the two skied down to the tiny snow survey cabin at Big Whitney Meadow. They arrived just before dusk and spent the next hour digging out the door from previous snowfalls. A major effort was clearing the stovepipe on the cabin roof.

It began to snow precisely at 6:00 p.m., just as they entered the cabin. Unlike many storms, the rate of snowfall was heavy right at the beginning. For the entire duration of the storm, the snow would come down at a steady rate of three inches an hour.

By morning, when the two ventured outside to check conditions, three feet of new snow had fallen, and the storm was still dumping snow at the rate of about three inches an hour. The snowfall continued all day at that rate without letup, and by dusk the 24-hour snowfall had risen to about six feet.

Their little snow survey cabin was in a dense grove of mature lodgepole pines, so they were relatively sheltered as long as they stayed inside. Although the cabin provided protection from the wind and snow, it was dark inside since the windows were buried under snow and was a chilly 20 degrees. The guys also had to come out into the storm periodically to clear snow off the stovepipe and shovel off the roof so that it wouldn’t collapse onto them.

Tired of being cabin-bound, Doug skied to the edge of the grove to see what conditions were like in Big Whitney Meadow. He ventured only a short distance into the meadow. The snow across that extensive treeless area was propelled by 50–60 mph wind. Visibility was zero and breathing was nearly impossible. He quickly retreated into the shelter of the pines, with gratitude for still being alive.

When the two men emerged from their cabin on the second morning, they could instantly tell that the heavy snowfall had continued all night. After thirty-six hours of continuous storm, about nine feet of new snow now buried Big Whitney Meadow. By this time travel, even on skis, had become almost impossible.

The snow continued all the second day, still at three inches an hour, then quit abruptly at 6:00 pm, exactly 48 hours after it began. The new snow total at the cabin now equaled 12 feet!

By dawn the following morning the sky had cleared, and the snow surveyors spent much of the day laboriously collecting the required ten samples from the Big Whitney Meadow snow-survey course. Each required drilling down through more than 15 feet of snow.

According to their measurements, the two-day storm and its 12 feet of new snow had added up to 16 inches of water to the Sierra snowpack. To put this in context, this snow course’s average annual end-of-the-winter reading is slightly over 17 inches of water. An entire winter’s precipitation had fallen in 48 hours.
Professor Powell would later conclude that at Tulare County’s Big Whitney Meadow in late February 1969, he had witnessed one of the heaviest snowfalls ever measured anywhere on Earth.

Mammoth Mountain Ski Area didn’t begin keeping snow records until the winter of 1969−70, so we primarily have anecdotal accounts of the phenomenal snowfall that occurred in that area in the winter of 1968−69.\textsuperscript{1427, 1428, 1429, 1430, 1431} There were plenty of good snowstorms in December 1968. Then a siege of snowstorms began on January 10, 1969.\textsuperscript{1432} As Peter Vorster recalled, the town of Mammoth Lakes had measurable snowfall for 30 consecutive days.

Up to 25 feet of snow fell in just a few weeks, burying the ski area. The Main Lodge was almost totally buried, and snow tunnels had to be constructed leading down from the surface to lodge entrances and chair lifts. Crews had to dig channels under many of the lift lines so that the chairs could get up the hill. Skiers rode along with their skis touching the snow, even though the towers were high off the ground.

As Peter Vorster recalled, the town of Mammoth Lakes was cut off from vehicle access for many days, and supplies had to be brought in by air, snowcats, and dogsled. When Peter visited Mammoth Lakes in April 1969, he observed snowbanks in town that were still 30 feet high.

LADWP’s February 1, 1969 snow surveys for the Mammoth area averaged 45.0 inches of water, 225% of the 1961−2010 average, breaking the previous record set in 1952 (41.6 inches, 208% of average). That remains the highest February 1 snowpack for this area since record-keeping began in 1940. By the April 1, 1969 snow survey, the snowpack for the Mammoth area held an average of 65.3 inches (5½ feet) of water.

Peter recalled that the sagebrush east of the Sierra was totally buried during the storms. Cold air was trapped in the Owens Valley resulting in lower snow levels there while Lake Tahoe initially had rain.

DWR’s February 1, 1969 snow survey for the Sawmill snow course northwest of Independence held 30.3 inches of water, 245% of the 1961−2010 average, breaking the previous record set in 1967 (29.6 inches, 239% of average). That remains the highest February 1 snowpack for this area since record-keeping began in 1940. By the April 1, 1969 snow survey, the snowpack for Sawmill held 49.3 inches (4 feet) of water.

The town of Bishop received 23 inches in January 1969; that remains the snowiest month ever for that town. Peter recalled that heavy wet snow collapsed roofs in Bishop.

SCE’s snow surveys in the Bishop Creek watershed reportedly averaged 210% of average at some point in spring 1969. A warmer than normal May caused flows in Bishop Creek at Power Plant No. 6 to reach 700 cfs on June 1, 1969. Flows were predicted to reach 1,350 cfs (a recurrence interval of 100 years), but a change in the weather, including temperatures 7 degrees below freezing, slowed the snowmelt runoff in June.

When all that snow on the east side of the Sierra melted, it resulted in flooding of the normally dry Owens Lakebed from 1969−71. An 8-foot-deep lake formed, dissolving 20% of the 6.5-foot-thick salt bed.

Almost the same areas were flooded in February as in January. Peak discharges in Southern California were slightly less than in January, but on February 26 the Salinas River at Spreckels had a new peak discharge-of-record that exceeded the March 1938 peak by 11%.

The flood peak discharges were the largest in 30 years in Central and Southern California and in many places equaled or exceeded those of the March 1938 floods. In the Santa Clara, Santa Ynez, and Salinas River Basins, flood levels may have approached those of 1861−62.

Flood releases of 12,000 cfs or greater occurred at Friant Dam on the San Joaquin River during the April−July snowmelt period.

Panoche/Silver Creek west of Mendota flooded in January 1969.\textsuperscript{1433}

During the period of the January and February floods, storage in Pine Flat Reservoir increased from about 420,000 acre-feet on January 1 to 820,000 acre-feet on March 1. In that time about 223,000 acre-feet of Kings River water was passed through the dam and routed to the San Joaquin River via the Kings River North Channel and Fresno Slough Bypass.
During the preceding dry season, the USACE had fortuitously increased the capacity of the Kings River North Channel from 3,500 cfs to 5,500 cfs. This project had been completed just prior to the onset of the January 1969 flood.

During March, April, May, and June, increasingly large quantities of water were released from Pine Flat for diversion to the San Joaquin River. The total diversion during those four months was about 1,185,000 acre-feet.

No floodwaters from the basin above Pine Flat Dam reached the Tulare Lakebed before June. However, some uncontrolled flows, largely from Mill Creek (a southside tributary of the Kings below Pine Flat Dam) did reach the lake.

Runoff on the Kings during water year 1969 was the second highest since record-keeping began in 1894 (1983 would be even higher). The huge snowpack in the Kings River Basin resulted in the largest-ever releases from Pine Flat: 17,000 cfs.

Runoff for the Kings River at Pine Flat during water year 1969 was 4.2 million acre-feet. This was 253% of the 121-year average (1894–2014) for that river.

One source said that the January 1969 flood on the Kings River was in the same class with the 1914 and 1952 floods. It seems probable that there were other floods (particularly 1861–62, 1867, and 1937) that also belong in this category.

In any case, the peak day natural flow at Pine Flat on the Kings occurred on January 25, 1969. That is the fifth largest peak day of the year at Pine Flat since the dam was built in 1954. Based on the flood exceedence rates in Table 29, this had a recurrence interval of 25 years for the Kings River at Pine Flat.

That puts it in a category with other Cedar Grove floods of the past 70 years (1937, 1950, 1955, 1966, 1969, 1978, 1982, 1983, and 1997) that rise to the level of the modeled 50-year flood: that is, a flood event that occurs about every eight years on average. See the section of this document that describes Cedar Grove Flooding.

Jim Harvey recalled that the flood on the South Fork Kings was very impressive in Cedar Grove. The trail bridge across the South Fork Kings in Upper Paradise Valley was swept away.

No one imagined that Dinuba was susceptible to a major flood. The Kings River was way to the north, the Kaweah way to the south. Flooding was something that Reedley and Visalia worried about, not Dinuba. The January 1969 flood had a big surprise in store for that town.

The Alta Irrigation District gets its water from the Kings River with irrigation releases from Pine Flat Dam. In January, the East Branch of the Alta Canal was running full thanks to the floodflows on the Kings, when an intensive localized rainstorm caused it to overflow. Overflow events had happened in the past on this canal (1937, 1950, 1955, and 1966) and would happen again in the future (1993). However, this overflow would prove particularly memorable. As a result of this overflow, the canal suffered three ruptures near Smith Mountain, one of which was massive, some 60–80 feet long. The floodwaters poured out and flowed cross-country.

Dinuba was eight miles away, but there was nothing to divert the flood before it got there. The downtown area was flooded, as was much of the surrounding ranch land. Flooding was heaviest on the night of January 21. China Town was particularly hard hit. An evacuation center was set up, and a police car with loudspeaker went through China Town urging the residents to evacuate. A spokesman for the American Red Cross said that the China Town residents simply did not want to leave their homes even though they were underwater.

Apparently there was no headgate on the East Bank Canal; the philosophy being to take whatever irrigation releases were available, more is better. But that meant that there was now no way to shut off flow into the canal. As a result, the canal continued hemorrhaging floodwaters into the Dinuba area. Crews from the district and the town worked for a week, struggling to plug the leak. The final leak couldn't be plugged until the Kings River went down. The canal wall was finally repaired on January 27.
Farther to the east, the normally dry Sand Creek flooded Cutler and East Orosi on January 25. Cutler was inundated under a solid sheet of water, several feet deep in places. At least 350 people had to be evacuated, and it was several days before many of those could return to their homes.

In February 1969, Cottonwood Creek had an estimated flow of 4,670 cfs at the Elderwood gage. This was the flood-of-record for that stream and has a recurrence interval of 15 years. Sand Creek had a peak flow estimated at 3,520 cfs at the “near Orange Cove” gage (located 3.8 miles east of Orange Cove) during the 1969 flood. This was the flood-of-record for that stream and has a recurrence interval of 20 years.

Flooding on the Kaweah River washed away the Kaweah Public Beach (aka River Park) at Cobble Knoll on the lower side of Three Rivers. (The exact name of that county park is unclear.)

As on the Kings, runoff on the Kaweah during water year 1969 was the second highest since record-keeping began in 1894 (1983 would be even higher).

The Kaweah’s peak natural flow for the first rain-flood period occurred at Terminus Dam on January 25: 35,200 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 22,437 cfs.)

Based on the flood exceedence rates in Table 29, this had a recurrence interval of 25 years for the Kaweah. It would have had a recurrence interval of 20 years if calculated using the 35,200 cfs peak flow. (One source reportedly calculated this as having had a recurrence interval of 12 years. Presumably that was done using the peak flow and the now outdated 1971 flood frequency curves. That result could not be reproduced.)

The trail bridge over the South Fork of the Kaweah above Ladybug Camp was apparently one of the very few trail bridges in the Kaweah River Basin to survive the 1950 flood. However, one source said that the January 25, 1969 flood washed out this bridge and many others. The loss of this bridge closed the section of the Hockett Trail above this point. See the section of this document that describes the 1950 flood for a description of this section of the trail. Since losing this section of the Hockett Trail, the route has gone through Garfield Grove to get up onto the Hockett Plateau.

Terminus Dam reduced outflow to virtually zero, thus preventing damage to valley-floor communities and thousands of acres of crop and orchard lands.

Inflow to Lake Kaweah was 153,000 acre-feet between January 19–27. Storage rose from 8,300 acre-feet on January 18 to 139,800 acre-feet on January 27, an increase of over 131,000 acre-feet. Most of the subsequent flow, which was released because of flood operating criteria, found its way into the Tulare Lakebed.

Panoche/Silver Creek west of Mendota flooded in February 1969. This may have been a separate flood, or it may have been a continuation of the January flooding.

Peak discharges for the second, and somewhat smaller, rain-flood period on the Kaweah occurred on February 24–28.

Lake Kaweah was drawn down to 82,000 acre-feet on February 23, just prior to the second heavy rain-flood. Storage rose to 117,300 acre-feet by February 28, and eventually reached a maximum of 158,800 acre-feet on June 26.

Runoff for the Kaweah below Terminus Dam during water year 1969 was 1,271,000 acre-feet. This was 299% of the 121-year average (1894–2014) for that river.

The USACE, with cooperation from local organizations and parties, used every possible means to reduce Kaweah River flows into the Tulare Lakebed. Throughout the months of January to July, efforts were made to apply maximum quantities of water in the Kaweah service area, drawing down the reservoir. A temporary sack-concrete barrier was again placed on the spillway of the dam, increasing the storage capacity of the reservoir by 10,000 acre-feet. Even with all these efforts, 430,000 acre-feet of Kaweah floodwaters made it into the Tulare Lakebed in 1969.

Dry Creek below Terminus Dam peaked at 5,710 cfs. This was less than half the flow observed in the 1966 flood.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Over 5,000 acres of pasture and citrus orchards along Yokohl and Mehrten Creeks were flooded, and the floodwaters of those streams flooded neighborhoods in unincorporated Tulare County north of Exeter. About 10,000 acres of agricultural lands were flooded along Deep, Cameron, Outside, and Cross Creeks, and floodflows from Deep Creek entered the southern part of Farmersville. Levees along Cottonwood Creek were breached and 5,600 acres of agricultural land were flooded. About 800 acres of adjacent agricultural land and portions of the communities of Orosi, East Orosi, and Cutler were flooded. Antelope Creek overflowed its natural channel and inundated 350 acres in the Woodlake area. In total, over 25,000 acres of land were flooded in northwest Tulare County and in some areas flooding ranged up to 3 miles in width.

Joe Childress was the manager of the Wutchumna Water Company for many years. He recalled that Antelope Creek produced a large amount of floodwater in the area west of Woodlake’s Presbyterian Church during the 1969 flood. That was particularly remarkable since the Antelope Creek Basin seldom has any flow of note.

As on the Kings and Kaweah Rivers, runoff during water year 1969 on the Tule River was the second highest since record-keeping began in 1894 (1983 would be even higher). Runoff for the Tule below Success Dam during water year 1969 was 504,000 acre-feet. This was 367% of the 121-year average (1894–2014) for that river.

The 1969 storm pattern on the Tule River was similar to that on the Kaweah River. Based on the flood exceedence rates in Table 29, this had a recurrence interval of 67 years for the Tule River. Storage in Lake Success rose from 11,100 acre-feet on January 18 to 77,230 acre-feet on January 28. Gross storage in Lake Success was considered at the time to be 85,400 acre-feet. Available storage was decreased after the January rain-flood to about 63,000 acre-feet. The February rain-flood caused storage to rise to 83,800 acre-feet on February 25. As at Lake Kaweah, USACE and local water-using agencies worked together, making every effort to reduce the amount of water that had to be spilled into the Tulare Lakebed.

A temporary sack-concrete barrier was again placed on the Lake Success spillway, increasing the storage pool above the designed 85,400 acre-feet. During the snowmelt flood runoff season, storage in Lake Success rose to a maximum of 95,300 acre-feet on June 20 and was above 85,400 acre-feet from May 19 through July 15.

The lower Tule flooded in both the January and February floods, and 215,000 acre-feet were passed downstream to the Tulare Lakebed.

Even normally dry Deer Creek was flowing into the Tulare Lakebed in the spring.

The Kern’s peak natural flow was 22,359 cfs. (That was the peak average daily flow.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 56 years for that river. The inflows to Isabella Reservoir during the peak of the rain-flood in January and February were not nearly as great as those during the even more impressive December 1966 flood.

- Maximum mean daily inflow to Isabella Reservoir on January 25, 1969, was 22,200 cfs. The comparable value for the 1966 flood was over three times as great (72,800 cfs on December 6, 1966).
- Maximum 10-day rain-flood inflow to Isabella Reservoir was 132,000 acre-feet from January 19–28, 1969. The comparable value for the 1966 flood was nearly twice as great (254,000 acre-feet for the period from December 5–14, 1966).

There was a major flood on Caliente Creek in February, causing extensive flood damage to the Lamont/Arvin area.

The third flood of 1969 was a snowmelt flood. A great snowpack had accumulated in the Southern Sierra by the beginning of April. It contained over 200% of the average water content. That set the stage for the flooding that was to occur during the April–July runoff period.

The Kings River is controlled by Pine Flat Dam, but the reservoir holds only 60% of the 121-year average runoff (1894–2014) of that river.

See Figure 18 on page 111 to understand why 1969 was not an average year. Pine Flat Dam recorded the largest snowmelt of record during April through July of that year. The 1969 snowmelt exceeded all previous years since record-keeping began in 1895. Pine Flat Dam was operated to control outflow to a maximum of 17,100 cfs.
Flood control releases from Pine Flat Dam in 1969 totaled 1,017,000 acre-feet.\textsuperscript{1445, 1446} Floodwaters were routed both to the San Joaquin River and to the Tulare Lakebed. The James Bypass experienced a maximum daily discharge of 5,570 cfs on June 7, 1969; that remains the flood-of-record for this channel.\textsuperscript{1447} During 1969, the USACE had to pass a total of 4,197,901 acre-feet of Kings River water through to the Tulare Lakebed.\textsuperscript{1448}

Runoff for the Kern River near Bakersfield during water year 1969 was 2,406,500 acre-feet. This was 335% of the 121-year average (1894–2014) for that river. This was the third highest runoff for the Kern since record-keeping began in 1894 (1916 and 1983 were both higher: 2,463,790 and 2,442,500 respectively).

Isabella Reservoir was operated with great care during the 1969 flood. Because of the heavy snowpack, it was known that Isabella would fill for the first time since operation began in 1954. Efforts were made to hold the storage down through March and April so that space would be available for the snowmelt runoff which was expected to be heavy during May and June. Much of the water released from Isabella Reservoir was used in the service areas below the reservoir for irrigation or spreading. Some water was stored in Buena Vista Lake.

Historically, the Kern would fill Buena Vista Lake before spilling over into Tulare Lake. However, in 1969, a giant dike protected two-thirds of Buena Vista Lake from being filled. When the other third of the lake filled, the Kern then spilled or passed through to Tulare Lake. The decision to keep the remainder of Buena Vista Lake dry was not appreciated by those downstream in the Tulare Lakebed.

In 1952, the Tulare Lake Basin Water Storage District had stored Kern River floodwaters in Buena Vista Lake. That was presumably possible because the J.G. Boswell Co., which had a long-term agricultural lease for the Buena Vista Lakebed, was willing to have its land flooded. In any case, no such water storage was allowed in 1969. That created hard feelings among some who were being impacted by the flooding that was occurring in the Tulare Lakebed in 1969. Emotions ran high as did financial losses.

The decision to pass through the Kern River floodwaters was challenged in court. However, in the meantime, the floodwaters continued to come.\textsuperscript{1449} As a result, about 222,000 acre-feet of Kern River water flowed into the Tulare Lakebed. The majority of the Buena Vista Lakebed remained dry, safe behind its giant levee.

The primary reason that this water storage didn’t happen in Buena Vista Lake in 1969 was because the J.G. Boswell Co. stood in the way. They controlled the Tulare Lake Basin Water Storage District and didn’t want their cropland in the Buena Vista Lakebed to be flooded. The resulting lawsuit made it all the way to the U.S. Supreme Court. In a 6–3 split decision, the court eventually found that it was permissible under the U.S. Constitution for the water district to be controlled by the J.G. Boswell Co. to the exclusion of all the other landowners and residents of the Tulare Lakebed.\textsuperscript{1450}

On May 8, 1969, the USACE received approval for a half-million-dollar project to throw up levees to connect the separated segments of Sand Ridge, south of the current Tulare Lake, creating a gigantic holding pond to contain Kern River floodwaters.\textsuperscript{1451}

There was significant flooding on the west side of the San Joaquin Valley near Coalinga. Los Gatos Creek and Warthan Creek experienced extremely high flows in February. Flooding continued downstream on the Arroyo Pasajero. This was the largest flood on the Arroyo Pasajero since record-keeping began on this stream-course.\textsuperscript{1452} It would remain the flood-of-record until the 1995 flood. The resulting floodwaters covered 16,600 acres and caused approximately $4.5 million in damage. Flooding extended from the foothills west of Coalinga to the valley east of the city. Bridges and roads were washed out, agricultural land was eroded, farm and ranch improvements and petroleum installations were damaged and destroyed, areas were isolated, traffic was disrupted, and residential and commercial areas in the northwest and southeast portions of the city were damaged. This was one of the three largest and most damaging flood events to occur in the Coalinga area during historic times.

Tulare Lake reappeared on January 20, 1969; it had been completely dry since August 9, 1967. By the end of March, 125 square miles (80,000 acres) of farmland had been inundated. The total lakebed inflow in 1969 was about 1.155 million acre-feet. This is the second biggest lakebed flood (both by volume and by area flooded) since the federal reservoirs were completed; only the 1983 flood would be bigger.
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Specific Floods and Droughts

In 1969, 88,700 acres (139 square miles) were inundated, significantly more than the 72,700 acres flooded in 1952. The J.G. Boswell Co. had more land flooded in the Tulare Lakebed than any other landowner (almost 50,000 of the total 88,700 acres). Although huge, the 139 square miles inundated in 1969 was just 18% of the 790 square miles that Tulare Lake used to cover when it was at full pool.

During May, the USACE closed the channel where the Kern River flowed through Sand Ridge into the Tulare Lakebed. This caused a lake to form south of Sand Ridge. As a result, about 235,000 acre-feet of Kern River water was prevented from entering Tulare Lake. This essentially recreated the southern extension of Tulare Lake, which had been known to the American Indians as Ton Taché. Along with the South Wilbur Flood Area (located north of Sand Ridge), that is the area known today by Tulare Lake water storage districts and irrigators as the South Flood Area.

As detailed in Table 68, the combined runoff of the four rivers in the Tulare Lake Basin during 1969 was 8,379,585 acre-feet, the second highest since record-keeping began in 1894; only 1983 would be larger (see Table 83 and Figure 18).

<table>
<thead>
<tr>
<th>Watershed</th>
<th>Total Runoff (acre-feet)</th>
<th>% of average (1894–2014)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings</td>
<td>4,197,901</td>
<td>253%</td>
</tr>
<tr>
<td>Kaweah</td>
<td>1,271,328</td>
<td>299%</td>
</tr>
<tr>
<td>Tule</td>
<td>503,856</td>
<td>367%</td>
</tr>
<tr>
<td>Kern</td>
<td>2,406,500</td>
<td>335%</td>
</tr>
<tr>
<td>Total</td>
<td>8,379,585</td>
<td>285%</td>
</tr>
</tbody>
</table>

On June 24, 1969, Tulare Lake reached a peak height of 192.5 feet elevation. (This was the highest the lake had been since 1952, when the lake reached 194.6 feet.)

In 1969, two middle-age ranchers and their three teenage sons took advantage of the high water to boat from Bakersfield through Buena Vista Lake and Tulare Lake to San Francisco Bay. Charlie and Esther Huecker recalled that these men were Ken Wedel from Wasco and Herb Spitzer from McFarland. Charlie said that their daily progress was chronicled in the Bakersfield Californian. It was also put on the AP newswire. We have only found one of those accounts so far.

They traveled in two small motorized fishing boats. The boats had to be light enough to carry around weirs and head gates. Their wives met them at various prearranged stops along the trip route, as they would overnight with friends or motel it up for the night. Ed Nelson recalled reading the accounts of their trip in some newspaper such as The Fresno Bee. He recalled thinking how neat a trip that would be no matter your age. It sort of brought out the Tom Sawyer / Huckleberry Finn in you.

News accounts from the time said that they were attempting a trip that hasn't been made since 1938. But Bill Cooper later recalled that somebody had been made a trip in 1966. Ken and Herb's trip began on about June 13, 1969. They ran into trouble when one of the two boats broke down on their first day in the vicinity of Lost Hills, 45 miles from the beginning of the trip. They apparently got the boat repaired and continued on at a later date. Esther thinks that they reached San Francisco Bay in about mid-July.

This was the fifth of six documented trips between Tulare Lake and San Francisco Bay to occur in historic times. (The other five trips were in 1852, 1868, 1938, 1966, and 1983.)

Bill Cooper recalled that somebody made the trip to San Francisco in a motorboat in 1966, and Ken Wedel (now deceased) then did it in a motorboat with his son.

The 1969 flooding in Tulare Lake caused the complete inundation of sumps #1, 2, 3, and 4, as well as some additional lands outside of those sumps. (That is presumably the area that we now know as the South Wilbur and Hacienda flood units.) Maximum storage in Tulare Lake reached 960,000 acre-feet in June. The difference between total inflow to the lake and maximum storage, 195,000 acre-feet, was lost by evaporation and absorption into the lakebed. A small amount of water was also diverted from the lake for irrigation of lands around the perimeter of the flooded area. The lakebed would remain at least partially flooded through 1971 (see Figure 16). The lakebed was finally dry in calendar year 1972.
The lake threatened the west side of the town of Corcoran during the 1969 flood (multiple photographs on file in the national parks, also see the back cover photographs). An emergency levee was hurriedly built just west of the Corcoran Airport.

The J.G. Boswell Co. took the lead on the levee building with much assistance from Salyer and the smaller farmers in the area (Boyett and Gilkey). Boswell also took the lead on the purchase and movement of junk cars to the levees. These were used as riprap to protect the levee from erosion.

In 1969 Mo Basham was a 12-year-old girl living in Corcoran, so she remembers much of what happened. Tulare Lake was deep enough to cause significant erosion to the emergency levee, even though it was faced with the junk cars. The chop on the water during windy/stormy weather was pretty significant, so the levees were constantly monitored.

Mo recalled going out on rodent patrol to spot/kill ground squirrels and gophers that would dig into the levees. Mo also recalled going on crawdad (crayfish) hunts. The kids would bring back hundreds at a time, and their families would eat them just like lobster. Mo later recalled the adventure:

*My dad, whose ulterior motive was having crawdads to feast on, was the one that got all the neighborhood kids together and loaded us in the back of his pickup (back in the days when it was still legal to do that), and out to the lake’s edge we went. He gave us little or no instruction as I recall but did give us each a rake and a burlap sack. We were to walk along the lake levee, and whenever we saw a crawdad try and sweep it out of the water and onto the levee where it couldn’t get back into the water. And that is exactly what we did. It didn’t take long before we all got the real hang of sneaking up on a crawdad and sweeping it up onto the levee. And some of those crawdads went flying several feet before hitting the ground. We spent several hours each time we went out and with a half-dozen or so neighborhood kids, you can accumulate quite a few crawdads. Now we thought we were having fun and often came home with mud up to our hips but enjoyed the whole process, and after that would always ask Dad when we would be going out again to catch them crawdads.*

The crawdad hunts continued for about two years before the lake receded back behind the larger levees that the J.G. Boswell Co. had built like the El Rico, North Central and South Central levees which were miles out of town instead of just on the other side of the emergency levee by the Corcoran Airport.

The emergency levee was taken down once the waters receded enough to eliminate the threat to Corcoran. As best Mo can recall, that was sometime in early 1971.

Table 69. Inflow to the Tulare Lakebed during water year 1969.

<table>
<thead>
<tr>
<th>Stream</th>
<th>Total Lakebed Inflow (acre-feet)</th>
<th>Percent Contribution</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings River</td>
<td>195,000</td>
<td>17%</td>
</tr>
<tr>
<td>Kaweah River</td>
<td>430,000</td>
<td>37%</td>
</tr>
<tr>
<td>Tule River</td>
<td>215,000</td>
<td>19%</td>
</tr>
<tr>
<td>Deer Creek &amp; other sources</td>
<td>93,000</td>
<td>8%</td>
</tr>
<tr>
<td>Kern River</td>
<td>222,000</td>
<td>19%</td>
</tr>
<tr>
<td>Total</td>
<td>1,155,000</td>
<td></td>
</tr>
</tbody>
</table>

On August 9, 1969, a mature giant sequoia failed in Hazelwood Picnic Area in Giant Forest, triggering the failure of three other mature giant sequoias. That event killed a park visitor, resulting in the permanent closure of this picnic area. The failure likely had multiple causes, one of which may well have been the heavy precipitation of the preceding winter.

The January-February 1969 flood was unusual in being a major event in both the Tulare Lake Basin and in Southern California. The heavy January storm saturated the ground so that the February storm produced particularly high levels of runoff.

San Luis Obispo County experienced flooding in both the January and February storms. Flood damage occurred in places such as San Luis Obispo, Morro Bay, and in state and USFS campgrounds. Considerable debris was
washed up on beaches. The most severe damage occurred in the January storm. Total damage for the county was $5 million.

Santa Barbara County experienced flooding in both the January and February storms. Floodflows were of unprecedented magnitude. The January flows are generally the flood-of-record in the area. Many areas of the county had severe flooding, notably the Solvang and Lompoc areas. The Santa Ynez River near Lompoc peaked at an estimated flow of 100,000 cfs. The spillways of Gibraltar and Cachuma Dams spilled flows exceeding their design flow. Total damage for the county was $4.5 million with 5 deaths.

Ventura County experienced flooding in both the January and February storms. All rivers in the county flooded, and highway damage was heavy. The entire city of Santa Paula was evacuated during both the January and February storms. The February event was the largest flood-of-record in the Simi Valley and Moor Park areas. Total damage for the county was $43 million with 12 deaths.

Los Angeles County experienced severe flooding in both the January and February storms. People who lived in mountain areas such as Topanga Canyon, Mandeville Canyon, and Big Tujunga Canyon were especially hard hit. Landslides, debris flows and overflowing debris basins were a major problem. Total damage for the county was estimated to be $68 million plus $16 million to remove debris; 73 lives were lost.

Orange County experienced flooding from January 18–28 and again in February. The January flood damaged and destroyed bridges, roads, rail lines, and homes. Had Prado Dam not been in place, it was estimated that the Santa Ana River would have flooded in the January flood at 75,000 cfs, resulting in $440 million damage. The county experienced even worse flooding from the storm of February 18–27. The Santa Ana River threatened to breach its levees and emergency work (assisted by the U.S. Marine Corps) was required to prevent disastrous flooding. Total damage for the county from the two storms was $22 million with 7 deaths and 15 serious injuries.

San Bernardino County experienced flooding from January 18–28 and again in February. The January storm was an intense event, Etiwanda (part of present-day Rancho Cucamonga) received 16 inches of rain during this 9-day period. Many people had to be evacuated, and many homes were damaged or destroyed. Along the Santa Ana River, many highway bridges were lost or damaged. Many transportation routes between Bernardino and surrounding areas were closed and impassable until April or later. The February storm generated greater runoff and consequential damage to infrastructure. Places that were damaged in the January event were damaged again in February. Total damage for the county for January and February was more than $54 million with 13 deaths.

Riverside County was struck by flooding in both the January and February storms. The February flooding was even worse than the January flooding. Roads, railroads, and homes were heavily damaged in both floods. The mainline of the Southern Pacific Railroad was washed out. Floodwaters covered the Corona Airport up to 10 feet deep. Total damage for the county was $32 million with 4 deaths.

Floods on Lower Mehrten and Yokohl Creeks
When the Kaweah Delta was first settled, the Kaweah River divided into its distributaries (the Four Creeks, if you will) at a point more or less in the middle of the delta. There was a fifth channel that came directly out of the foothills south of the Kaweah. Presumably that followed the course of Yokohl Creek at least as far as present-day Highway 198. In any case, that fifth channel flowed along the south side of the Kaweah Delta (apparently following a course similar to present-day Outside Creek) and then joined the merged channels of the Kaweah in the marshy ground near where that river flowed into Tulare Lake. Mehrten Creek may have been part of that fifth channel as well.

Floods and man-made ditches have dramatically changed drainage patterns on the Kaweah Delta. Mehrten and Yokohl Creeks are no longer connected to Outside Creek (assuming they once were). Just when and how those channels shifted isn’t quite clear. We have a few clues from past floods:

- In the May 1884 flood, Yokohl Creek was reported to have gone past Merriman Station.
- In the 1906 flood, it is speculated that Yokohl Creek went past SCE’s Venida Substation at the intersection of Highway 65 and Highway 198.

Yokohl Creek has experienced some impressive floods. Even little Mehrten Creek had has had some pretty good floods, especially in February 1969, 1983, and December 2010. It is hard to believe that these two placid little creeks are capable of such big floods.
It is apparent today where Mehrten and Yokohl Creeks flow under Highway 198. However, it can be difficult to imagine where floods occur in their lower reaches.

The area of unincorporated Tulare County that lower Mehrten and Yokohl Creeks flow through is generally located between Woodlake, Lemon Cove, and Exeter. That really doesn’t describe where this flooding occurs. The lower reaches of Mehrten and Yokohl Creeks don’t appear on most maps. After crossing Highway 198, the channels head generally northwest toward Avenue 312 where they turn west. Mehrten Creek then flows on the north side of Avenue 312 while Yokohl Creek flows on the south side.

During floods, Mehrten Creek typically floods the area around the intersection of Road 220 and Avenue 304 as well as orchards in the vicinity.

Yokohl Creek crosses Avenue 304 south of Woodlake. It is not too uncommon for Yokohl Creek to flow over the road at that location, causing the closure of that road.

Mehrten and Yokohl Creeks both cross under Road 204 / Highway 245 about 1½ miles north of Highway 198. During big floods, Yokohl Creek sometimes floods the highway at that point. When this happens, Lort Drive (Avenue 312) also tend to be closed due to flooding.

After Mehrten Creek flows under Road 204 / Highway 245, it spreads out during floods and inundates fields, orchards, and houses.

The Mehrten and Yokohl Creek channels merge and flow into the Consolidated Peoples Ditch just before reaching Road 196 / Highway J27. This is near the Lower Kaweah River but well below McKay’s Point.

1970–71 Floods (2)
Flooding occurred in the Tulare Lakebed in both 1970 and 1971. This lakebed flooding occurred despite the fact that there was no storm event of note in either year. Bakersfield did get hit by a powerful storm on May 26–27, 1971, setting a 24-hour precipitation record for the month. However, that was just a local event during a dry year; none of that moisture made it to the lakebed.

Runoff was below average in both 1970 and 1971, bordering on drought conditions. Runoff during water year 1970 was 82% of the 121-year average (1894–2014) for the four major rivers (Kings, Kaweah, Tule, and Kern) combined. In 1971, it was only 66%.

The flooding in the lakebed in 1970 and 1971 was all left over from the big 1969 flood. As illustrated in Figure 16, the lakebed would remain at least partially flooded through 1971.

Lakebed flooding is a social construct; it is counted based on the number of growing seasons that are missed. The lakebed was flooded for three growing seasons: 1969, 1970, and 1971. Therefore, this is counted as three floods from the perspective of the lakebed farmers, even though the flood event occurred only once.

Something similar happened in the lakebed in 1982–84 and 1997–99. In each of those cases, lakebed flooding continued into a non-flood year.

1972 Floods (2)
There were two floods in 1972:
1. June
2. August

On June 7, an intense thunderstorm centered over the north Bakersfield area caused flooding and damages in its wake. The storm dropped 1.09 inches at Meadows Field in Bakersfield in 45 minutes, making it the wettest June day ever in that city. There was one report of 3.50 inches of rain in an hour in one part of the city. Lightning struck six substations, knocking out the power to most of Bakersfield. The storm produced wind gusts to 50 mph, damaging automobiles and buildings.

Domestic water supply lines were washed out, roads severely damaged, and cars lifted and moved during the high runoff period. Houses were flooded with up to 4½ feet of water, and apartments were flooded and
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

destroyed. Two people drowned, one a high school senior who was returning from his class picnic. Debris flows closed some roads and highways. Highway 178 was closed because of mud and landslides. The cross-town freeway in Bakersfield was closed due to flooding. Kern General Hospital had a flooded basement and first floor, which closed the emergency room services. Memorial Hospital was threatened with evacuation if the flood control canal broke, but it held.\textsuperscript{1461}

Hurricane Gwen formed off the coast of Mexico on August 24. It then spent a few days heading west-northwest and became a major hurricane on August 27. After retaining that intensity for over a day, it rapidly weakened and became a tropical storm on August 29.

On August 30, a cloudburst associated with Gwen off the coast of Southern California dropped 0.99 inches of rain 14 miles southwest of Coalinga in the Bear Canyon Jupiter area resulting in flash flooding.\textsuperscript{1462}

1973 Flood

We don’t have a clear understanding of this flood.

The winter of 1972–73 was a strong El Niño event.
The winter of 1973–74 was a strong La Niña event.

Total flow for water year 1973 was 126\% of the 1894–2014 average for the Kings, 145\% for the Kaweah, 164\% for the Tule, and 133\% for the Kern. Roughly 25\% of years have more runoff than this.

However, flooding still occurred in the Tulare Lakebed. Perhaps this was due to the timing of the runoff rather than to the total quantity of the flow. The lakebed had been flooded for three years, from 1969–71. After three years of below-average runoff, it was dry again in 1972. However, as shown in Figure 16, the lakebed flooded again in 1973.

Judging just from Figure 16, this flooding event was relatively small. It doesn’t look like an event that would threaten any of the valley towns. However, one report said that the lake threatened Corcoran and Alpaugh and stretched toward Kettleman City and Lemoore. That makes it sound like it was big as the 1969 flood, which seems unlikely. It seems almost certain that person was confusing the two floods.

On the other hand, the national parks’ files have a photograph of a valley town (perhaps Stratford or Avenal) with a lot of water in the town. That photograph is identified as having been taken in 1973. Conceivably that photograph was mislabeled and it was actually taken during the height of the 1969–71 flood.

What actually occurred in the Tulare Lakebed in 1973? Was it the small flood documented in Figure 16, or was it the extensive flood described by other reports?

As illustrated in Figure 15, we have an invaluable record of how Tulare Lake changed in elevation for 120 years (1850–1969). If we knew the elevation of the lake in the 1973 flood, we’d have a good measure of how big the flood really was that year. Unfortunately, it has proved impossible to obtain access to the lake gage data for years since 1969. As a result, we really can’t be sure just how big the lake was in the 1973 flood.

1975 Flood

Flooding in 1975 occurred from September 8–12 in Kern County.

In the Isabella area, a high intensity flash flood caused considerable damage in Kern County. One woman was swept from Highway 14 and drowned. High levels of sediment and debris deposits were a clean-up-chore on highways, roads, and on agricultural lands. Agricultural lands saw some damages, mostly to crops waiting to be picked.\textsuperscript{1463}

The South Fork of the Kern also flooded on September 8–12.
Floods and Droughts in the Tulare Lake Basin
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1976–77 Drought
This drought affected the entire state. It was most severe in the northern two-thirds of the state. Although brief in duration, this drought was notable for the severity of its hydrology. Together, the pair of years 1976–77 was the driest two-year period in the state's history. Based on the combined runoff of the four major rivers in our basin (Kings, Kaweah, Tule, and Kern), 1976–77 were also the driest two years in the Tulare Lake Basin since record-keeping began in 1894. The total runoff for those years was 13% less than the total for 2013–14. However, if the current forecast holds, the pair of years 2014–15 will be about 18% less than the record set in 1976–77.

Looking further back in time, the pair of years 1579–80 was the driest ever on the upper San Joaquin at the inflow to Millerton Lake. The flow in those two years was only 53% of the reconstructed flow for 1976–77.

The winters of 1976–77 and 1977–78 were weak El Niño events. The association of these events with the 1976–77 drought was almost certainly a coincidence. Research has shown no relationship between weak and moderate El Niño events and precipitation for any climate region in California.

The April 1, 1977 snow survey showed that the statewide average snowpack was only 25% of the long-term average. This was the lowest level since record-keeping began in 1950. It would remain the lowest April 1 snowpack of record until tied in 2014 and broken in 2015.

Based on 114 years of computed statewide runoff (1901–2014), 1977 occupies rank 114 (driest year) and 1976 is in rank 104. Water year 2015 will almost certainly be drier than any of those years. Statewide runoff in water year 1977 was only about 15 million acre-feet. This represents 21% of the statewide average annual 71 million acre-feet.

Water year 1976 ranks as the second-driest at gaging stations in the central part of the Coast Ranges and among the five driest in the Central and Northern Sierra. Water year 1977 was the driest year of record at almost all gaging stations in the affected area. The two-year deficiency in runoff during the drought was unequalled at gaging stations in the affected area. The recurrence interval was more than 100 years.

In terms of recurrence intervals, the droughts of 1929–34 and 1976–77 are similar; both are of unsurpassed severity among droughts of corresponding duration during the period of systematic record collection. The drought of 1929–34 was longer and accumulated a larger deficiency in runoff. The drought of 1976–77 was more intense and had greater annual deficiencies in runoff. Arguments can be made that either was the most severe drought in the history of the state.

Table 70 compares the 1976–77 drought with other severe droughts of the 20th century.

<table>
<thead>
<tr>
<th>Year Period</th>
<th>Sacramento River Basin Runoff</th>
<th>San Joaquin River Basin Runoff</th>
</tr>
</thead>
<tbody>
<tr>
<td>1929–34</td>
<td>9.8 (56%)</td>
<td>3.3 (56%)</td>
</tr>
<tr>
<td>1976–77</td>
<td>6.6 (38%)</td>
<td>1.5 (26%)</td>
</tr>
<tr>
<td>1987–92</td>
<td>10.0 (57%)</td>
<td>2.8 (48%)</td>
</tr>
</tbody>
</table>

Table 71 shows how the San Joaquin Valley Water Year Index categorized the drought years in that basin. It also shows total runoff for the four major rivers in our basin (Kings, Kaweah, Tule, and Kern) combined during this drought.

<table>
<thead>
<tr>
<th>Water Year</th>
<th>San Joaquin River Basin</th>
<th>Tulare Lake Basin</th>
</tr>
</thead>
<tbody>
<tr>
<td>1976</td>
<td>Critically dry</td>
<td>977,066</td>
</tr>
<tr>
<td>1977</td>
<td>Critically dry</td>
<td>696,572</td>
</tr>
<tr>
<td>Drought average (1976–77)</td>
<td>836,819</td>
<td>33%</td>
</tr>
</tbody>
</table>

Tree-ring reconstruction shows that 1580 is the drought year of record in the Central Valley and the Southern Sierra. Water year 2015 will almost certainly be the second-driest. As explained under the section of this
document that describes the 1918–34 drought, there is virtually a three-way tie among 1795, 1924, and 1977 as to which is the third-driest year in the San Joaquin Valley in the 1115-year period 900–2014. Based on stream gage data, we know that 1977 was a slightly drier year than 1924. However, we can’t say with any confidence where 1795 falls in this order, especially in the Tulare Lake Basin.

Based on tree-ring reconstructions, we know that the reconstructed flow on the upper San Joaquin at the inflow to Millerton Lake in 1580 was only 36% of that of the reconstructed flow in 1795 and 1924. George Durkee recalled that the 1976 dry season was very late in ending in the Yosemite area. In early January 1977, he and others drove over Tioga Pass and ice-skated at Ellery Lake on the east side. One of the first winter storms hit just after, finally closing the Tioga Pass Road.

1976–77 were back-to-back critical drought years for the Kings River Basin. DWR estimated that about 125,000 acres of irrigated cropland were fallowed due to water shortages in 1977, mostly in Fresno and Kern Counties, despite a significant increase in groundwater extraction to compensate for reduced surface water supplies.

1977 is the smallest tree-ring in Sequoia National Park since 1580. It is even smaller than the 1924 tree-ring. On July 25 and 29, 1977, there was no inflow to Pine Flat Reservoir; all three forks of the Kings River had run completely dry. That is the first time this has occurred on the Kings since record-keeping began at Pine Flat in 1953.

Roy Lee Davis recalled that the spring in the draw below Cactus Point in the national parks’ Tunnel Rock Pasture unit ran dry in 1976–77. However, he said that spring and the associated stream were definitely flowing again in 1978.

The 1976–77 drought appears to have been less severe — or at least the effects less noticeable — in the High Sierra than in the lower elevations. Although we don’t have gaging data for that zone, we do have observational data from several individuals. Bob Meadows recalled that none of the lakes that he visited in Yosemite during the summer of 1977 were significantly lower. Based on Bob’s research, none of the daily wilderness ranger logs for Sequoia or Kings Canyon National Parks make any mention of lakes drying up in 1977.

George Durkee was the Crabtree wilderness ranger in 1977. It was his first year there, so he had no baseline to compare conditions to. The signs of drought, while no doubt present, were not particularly obvious. He didn’t observe any lakes drying up. A large multi-day tropical storm occurred in August 1977, bringing significant rain and snow to the Sierra Crest but little to the mid-elevations. That event not only partially recharged the lakes, but caused lots of campers to scurry out of the high country, abandoning their gear. There was one fatality from the storm: a backpacker who died from hypothermia in Lamarck Col.

The spring near the Crabtree Ranger Station continued to flow throughout the summer of 1977. There was a pool above the station that had a stump that had apparently become rooted in a drought during the 1100s (see the section of this document that describes Megadroughts before the Little Ice Age on page 164). That pool did not lower significantly during the summer of 1977, and the spring that fed it continued to flow.

On the other hand, Dave Graber, retired NPS regional chief scientist, recalled that some of the lakes in the High Sierra of the national parks did in fact dry up by the end of 1977; they were nothing but mud.

Total flows for water year 1976 was less than 36% of the 1894–2014 average for each of the four rivers within the Tulare Lake Basin.

Flows for water year 1977 were the lowest experienced on the Kings, Kaweah, and Tule since record-keeping began in 1894. (The Kern had experienced lower runoff in 1931.) Flows on the Kings (386,007 acre-feet) and Kaweah (93,641 acre-feet) during 1977 remain the lowest flow recorded on these rivers during the 1894–2014 period of record. Flows for 2015 are forecast to be 341,000 acre-feet for the Kings and 83,700 acre-feet for the Kaweah.

Water year 1977 was the lowest runoff (15,884 acre-feet) experienced on the Tule River since record-keeping began in 1894. This record would last until 2014. Flows for 2015 are forecast to be even lower, just 11,300 acre-feet.
There have been four water years on the Kern drier than 1977: 2014 (172,946 acre-feet, the driest year of record), 1931, 1961, and 1924. Flows for 2015 are forecast to be lower than any of those years, just 110,000 acre-feet.\textsuperscript{1473}

As shown in Table 72, the combined runoff of the four rivers in the Tulare Lake Basin during 1977 was only 696,572 acre-feet, the lowest since record-keeping began in 1894. That remains the lowest combined flow recorded in our basin during the 1894–2014 period of record. Combined runoff in water year 2015 is forecast to be just 546,000 acre-feet, lower than flows in 1795, 1924, 1931, or 1977.

For a sense of how widely flows vary in our basin, the runoff in 1983 was 8,746,222 acre-feet (see Figure 18 on page 111 for a graph of other years).

Table 72. Runoff in the Tulare Lake Basin during water year 1977.

<table>
<thead>
<tr>
<th>Watershed</th>
<th>Total Runoff (acre-feet)</th>
<th>% of average (1894–2014)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings</td>
<td>386,007</td>
<td>23%</td>
</tr>
<tr>
<td>Kaweah</td>
<td>93,641</td>
<td>22%</td>
</tr>
<tr>
<td>Tule</td>
<td>15,884</td>
<td>12%</td>
</tr>
<tr>
<td>Kern</td>
<td>201,040</td>
<td>28%</td>
</tr>
<tr>
<td>Total</td>
<td>696,572</td>
<td>24%</td>
</tr>
</tbody>
</table>

Dave Parsons recalled that he and Phil Rundel paid close attention to foothills vegetation during the 1976–77 drought. They did not see a significant die-back in that vegetation; nothing like what would occur in 2014.

The 1977 Ferguson Fire occurred during the height of the 1976–77 drought. It burned in the Sugarloaf area south of Kings Canyon. It was ignited by a lightning strike on June 26 and burned until November 9. The Ferguson Fire was the third largest fire in the history of the national parks, burning 10,400 acres. Only the 1926 Kaweah Fire and the 1948 Simpson Meadow Fire were larger. The national parks’ three largest fires have all occurred during droughts.

DWR provided detailed information about the 1976–77 drought in the following reports:

- The California Drought — 1976. May 1976.\textsuperscript{1474}
- The California Drought 1977, An Update. February 1977.\textsuperscript{1475}
- The Continuing California Drought. August 1977.\textsuperscript{1476}
- The 1976-1977 Drought — A Review. May 1978.\textsuperscript{1477}

The following information comes from this report.

The 1976–77 drought was notable for the impacts experienced by water agencies that were unprepared for such conditions. One reason for the lack of preparedness was the perception of relatively ample water supplies in most areas of the state. The SWP’s California Aqueduct had been completed less than ten years before, bringing a new source of water to parts of the San Joaquin Valley and Southern California. Likewise the state-federal joint-use facilities of the San Luis Canal brought new irrigation supplies for CVP contractors on the west side of the San Joaquin Valley. The imported water took some pressure off overdrafted groundwater basins in parts of the valley; growers and irrigation districts took many of their wells out of service with the advent of the new supplies. California was receiving more than its basic interstate apportionment of Colorado River water thanks to supplies unused by Nevada and Arizona and to hydrologic surpluses. There had not been major droughts in the recent past. (Although there had been multi-year dry periods of statewide scope in 1947–50 and 1959–61, those events were far less severe than that of the 1920s–30s.) The 1976–77 drought was a wake-up call for many water agencies.\textsuperscript{1478}

By April 1977, President Jimmy Carter had declared 43 of California’s 58 counties emergency areas. In water year 1977, 47 of the state’s 58 counties declared local drought-related emergencies.

By summer of 1977, Congress had enacted the Emergency Drought Act of 1977 (Public Law 95-18), the Community Emergency Drought Relief Act of 1977 (Public Law 95-31), and had passed the Supplemental Appropriations Act of 1977 (Public Law 95-26) to bolster existing drought assistance programs. These laws provided financial assistance to eligible applicants to mitigate the impact of the drought through such steps as water conservation and improvement to existing water systems. This drought package, including prior legislation approved by Congress, authorized over $800 million in short-term loans and grants nationwide.
All the funds had to be obligated by December 31, 1977. Over $24 million in grants were made to California communities under the Community Emergency Drought Relief Act of 1977. However, as shown in Table 73, only $346,600 of those grants (1%) went to communities in the Tulare Lake Basin and Southern Sierra. Presumably that was because big cities in the larger urban areas had more qualifying packages sitting on the shelf that could be quickly obligated.

<table>
<thead>
<tr>
<th>Community Assisted</th>
<th>Grant</th>
<th>Loan</th>
<th>Description of Project</th>
</tr>
</thead>
<tbody>
<tr>
<td>Yosemite National Park</td>
<td>$100,000</td>
<td></td>
<td>Drill well, construct storage tank, meter system and electrical equipment.</td>
</tr>
<tr>
<td>Fresno</td>
<td>$56,000</td>
<td></td>
<td>Construct water system interties and a conservation program.</td>
</tr>
<tr>
<td>Sanger Public Works</td>
<td>$16,600</td>
<td>$66,400</td>
<td>Construct new wells and install 300 water meters.</td>
</tr>
<tr>
<td>Delano</td>
<td>$36,800</td>
<td></td>
<td>Construct water well and water conservation program.</td>
</tr>
<tr>
<td>Kings County</td>
<td>$41,200</td>
<td></td>
<td>Purchase two tanker trucks with fire pumping capability, 10,000 feet of 3” hose and 2,000 water-saving kits.</td>
</tr>
<tr>
<td>Hanford</td>
<td>$96,000</td>
<td></td>
<td>Two new wells, lower 12 well pumps, and replace leaking pipes.</td>
</tr>
<tr>
<td>Kern County</td>
<td>$5,000,000</td>
<td></td>
<td>Purchase water</td>
</tr>
<tr>
<td>Kern County Water Agency</td>
<td>$859,855</td>
<td></td>
<td>Purchase water</td>
</tr>
<tr>
<td>Total for the Tulare Lake Basin and Southern Sierra</td>
<td>$346,600</td>
<td>$5,066,400</td>
<td></td>
</tr>
<tr>
<td>Total for state</td>
<td>$24,448,810</td>
<td>$54,726,159</td>
<td></td>
</tr>
</tbody>
</table>

Selected Urban Water Supplies

All water users were challenged to meet their needs in 1977. However, most of the urban water systems in the vicinity of the Tulare Lake Basin appear to have gotten through 1977 in reasonable shape. Two systems that were particularly hard pressed were:

- **Mariposa.** That city operated under enforced rationing in 1977 because their main source of supply, Mariposa Creek, was dry that year.
- **Springville.** That city gets its supply from the Tule River. They operated under voluntary conservation for at least part of the year because the flow in the Tule dropped to about 1 cfs in the fall of 1977.

Groundwater Levels and Water Use

In the San Joaquin Valley, reduced imports and depleted carryover storage in local reservoirs combined to put considerable strain on groundwater resources. Pumping capability was insufficient in many areas to maintain the same level of irrigated agriculture in 1977 as in 1976.

In the west side of the valley, from Firebaugh to the vicinity of Kettleman City, hundreds of wells had been abandoned with the advent of surface imports. The reduced pumping capability, together with sharply curtailed imports, reduced applied water in 1977 to about 55% of normal. Groundwater withdrawals increased.

Table 74 shows the shift in water sources that occurred in the entire Tulare Lake Basin during the drought. It shows the amounts of water derived from the various sources in 1975 (a normal year) and in the drought years 1976 and 1977. Also shown are estimates for 1978. The Tulare Lake Basin saw both its local surface water supply and its imported supply cut drastically and, consequently, groundwater withdrawals jumped from 51% to 78% of the total supply. (This percentage is generally reported as 82% instead of 78%.)

In terms of absolute numbers, the first year of the drought (1976) saw a 31% decrease in the amount of available surface water (local plus imported). However, this loss was largely compensated for by a large increase in groundwater withdrawals. As a result, total water supply (applied water) in 1976 was only 4% less than in 1975.

By 1977, however, groundwater withdrawals in the Tulare Lake Basin could not compensate for the large decrease in surface water supply. Despite an overdraft of 3.8 million acre-feet, total water supply in 1977 was 17% less than in 1975. Some of this reduction was accommodated by reductions in demand due to conservation efforts, but the majority represents less than adequate supply to satisfy average use.
Table 74. Estimated water use by source in the Tulare Lake Basin 1975–78. (in 1,000 acre-feet)

<table>
<thead>
<tr>
<th></th>
<th></th>
<th></th>
<th></th>
<th></th>
</tr>
</thead>
<tbody>
<tr>
<td>Deliveries from local rivers</td>
<td>2,462</td>
<td>1,255</td>
<td>850</td>
<td>2,500</td>
</tr>
<tr>
<td>Imported (SWP + CVP)</td>
<td>3,796</td>
<td>3,047</td>
<td>1,307</td>
<td>3,600</td>
</tr>
<tr>
<td>Groundwater</td>
<td>6,420</td>
<td>7,887</td>
<td>7,825</td>
<td>4,100</td>
</tr>
<tr>
<td>Total used</td>
<td>12,678</td>
<td>12,189</td>
<td>9,982</td>
<td>10,200</td>
</tr>
<tr>
<td>Supply (calculated)</td>
<td>11,649</td>
<td>9,230</td>
<td>6,191</td>
<td>11,800</td>
</tr>
<tr>
<td>Groundwater overdraft</td>
<td>1,029</td>
<td>2,959</td>
<td>3,791</td>
<td>-1,600*</td>
</tr>
</tbody>
</table>


The increased dependence on groundwater during the drought is illustrated by the greater activity demonstrated by the well drilling industry.

The number of Water Well Drillers Reports received by DWR increased from 8,687 in 1974 and 8,275 in 1975, to 11,209 in 1976 and 20,115 in 1977. Those figures do not include all new wells drilled, since historically not all work was reported. However, the significant rise in filings supports the conclusion that a large number of new wells were drilled, particularly in 1977.

In the ten San Joaquin Valley counties (San Joaquin, Tuolumne, Stanislaus, Mariposa, Merced, Madera, Fresno, Tulare, Kings, and Kern), about 6,800 reports were filed. It was estimated that approximately 9,000 wells were drilled or deepened in the valley, based on a compliance rate of 75%.

**State Water Project**

During 1977, total deliveries to SWP customers amounted to about 898,099 acre-feet, down from the 1,953,112 acre-feet delivered in 1976.

A total of 940,176 acre-feet were released into the Feather River from Thermalito Complex in 1977 for all downstream purposes, including fish releases, water rights users along the Feather and Sacramento Rivers, Delta salinity control, and delivery to SWP customers. No separate releases were made for power generation, although some power was developed as an incidental use of releases for other purposes.

In carrying out those objectives, Lake Oroville on the Feather River, key storage reservoir for the SWP, was drawn down to 882,395 acre-feet, reaching its lowest level on September 7, 1977. This remains the record low storage for that reservoir, just 26% of capacity. The southern reservoirs, including San Luis, were called upon to furnish the bulk of SWP deliveries for the year and most showed drastic declines in storage. Total December 31, 1977 storage in the seven major project reservoirs amounted to 1,905,133 acre-feet, 52% of average storage for that date.

**Effect of Shortages upon SWP Contractors**

During 1977, the SWP was unable to deliver all the water needed by its contractors. As shown in Table 10, agricultural contractors received only 40% of their contracted amounts and urban contractors 90%.

An example of drought impact is provided by the experience of the Kern County Water Agency (KCWA), whose 19 member districts accounted for 432,500 acre-feet of SWP agricultural water entitlement and 51,100 acre-feet of municipal, or a total entitlement of 483,600 acre-feet. By reason of the cuts, only 218,990 acre-feet were deliverable in 1977. (Actual quantities delivered were somewhat less.) In contrast, 881,400 acre-feet were delivered from the SWP in 1976, including 442,150 acre-feet of surplus water available from reservoir storage resulting from prior years of above-average precipitation.

Thanks to the availability of Colorado River water in excess of the state’s basic interstate apportionment, MWD was able to reduce its use of SWP water, making more water from that source available for other project contractors. To provide a more nearly average quantity for agricultural purposes, DWR arranged for KCWA to purchase water from four Southern California SWP contractors who agreed to forego all or part of their 1977 entitlements. A total of 241,530 acre-feet were purchased.

The low level of deliveries had its effect upon KCWA’s customers, particularly in the economic area. Since annual payment of project costs includes a large fixed component, charges to KCWA (passed on to its member districts)
Floods and Droughts in the Tulare Lake Basin

Specific Floods and Droughts

did not proportionately reflect the marked decrease in deliveries experienced in 1977. As a result, unit costs of the delivered water were considerably higher than in past years.

Until 1977, unit costs to member districts had remained relatively stable at about $15.00 per acre-foot. In 1977, however, average unit prices jumped to $44.27 per acre-foot. This was primarily driven by the increase of SWP water to $54.73 per acre-foot in 1977.

Canal-side unit costs are those prices paid to KCWA by its member districts and are not the prices paid by the farmer. A poll of districts indicates that in 1977 the farmers paid from $55.00–$125.00 per acre-foot, compared to previous year costs ranging from $25.00–$55.00 per acre-foot.

To counteract the decrease in surface water supplies, additional dependence was placed on groundwater. KCWA estimated that groundwater furnished 3,000,000 acre-feet of the county’s 1977 water supplies, up from a normal groundwater withdrawal level of about 2,000,000 acre-feet. The additional draft on the basin produced significant declines in groundwater levels (from 10–60 feet in some areas).

Despite the measures taken to provide the additional supply, some Kern County cropland went out of production. KCWA estimated that 56,000 acres of land, ordinarily irrigated by SWP water, were idled because of lack of project water in areas of limited groundwater pumping capability. This represented 9% of the irrigated acreage and a gross farm income loss of $50 million.

Central Valley Project

During 1977, the CVP, operated by USBR, delivered 3,300,000 acre-feet of water to its users. This was down significantly from the 6,000,000 acre-feet delivered in 1976 and the 7,000,000 acre-feet ordinarily delivered. In addition to deliveries to its customers, water was also released to help maintain Delta water quality, generate power, and provide fish releases. A total of nearly 5,500,000 acre-feet was released for all purposes in 1977 (based on releases from Keswick and Folsom Dams and to the Friant-Kern and Madera Canals).

CVP water users were faced with significant cuts in contract entitlement deliveries in 1977. These ranged from 25% to all users with water rights on the Sacramento and San Joaquin Rivers to 75% for all other agricultural users. Urban and industrial users were cut back 50%. A comparison of diversions for use in each of the project’s main aqueducts during 1975, 1976, 1977 is shown in Table 75. Also shown are projected diversions for 1978.

Table 75. Comparison of diversions for use in each of the CVP’s main aqueducts.

<table>
<thead>
<tr>
<th></th>
<th></th>
<th></th>
<th></th>
<th></th>
</tr>
</thead>
<tbody>
<tr>
<td>Contra Costa</td>
<td>76,752</td>
<td>125,129</td>
<td>95,857</td>
<td>90,000</td>
</tr>
<tr>
<td>Delta-Mendota</td>
<td>1,512,962</td>
<td>1,652,915</td>
<td>983,911</td>
<td>1,290,000</td>
</tr>
<tr>
<td>San Luis</td>
<td>1,375,832</td>
<td>1,425,849</td>
<td>376,678</td>
<td>1,275,000</td>
</tr>
<tr>
<td>Madera</td>
<td>319,651</td>
<td>94,360</td>
<td>31,670</td>
<td>400,000</td>
</tr>
<tr>
<td>Friant-Kern</td>
<td>1,393,977</td>
<td>534,240</td>
<td>258,410</td>
<td>1,200,000</td>
</tr>
<tr>
<td>Corning</td>
<td>33,228</td>
<td>47,864</td>
<td>18,270</td>
<td>39,000</td>
</tr>
<tr>
<td>Folsom South</td>
<td>12,809</td>
<td>22,350</td>
<td>19,530</td>
<td>30,000</td>
</tr>
<tr>
<td>Tehama-Colusa</td>
<td>183,798</td>
<td>267,822</td>
<td>224,878</td>
<td>230,000</td>
</tr>
<tr>
<td><strong>Total</strong></td>
<td>4,909,009</td>
<td>4,170,529</td>
<td>2,009,204</td>
<td>4,554,000</td>
</tr>
</tbody>
</table>

Effects of Drought on Agriculture

Shortages of stock water existed since early 1975 and many herds were maintained by producers hauling water. Forced liquidations of herds began in early 1976 and continued through 1977, particularly as it became apparent that the spring 1977 rains were not forthcoming. Many producers tried to maintain herd sizes in order to keep open lines of credit, but in many instances were forced by creditors to lower inventory levels in an effort to cut overhead costs. Most herds were moved off stressed and overused irrigated pastures early in the season to take advantage of what feed existed at higher elevations.

Ground protection in the Central Coast, Sacramento Valley, and Sierra regions did not exist for new range seedlings in early 1977. Consequently, most new grass died almost immediately after germination because of no moisture. Most counties reported almost 100% loss of grazing capacity during 1977 in these districts. Many sheep and lamb producers either moved flocks or liquidated.

The San Joaquin Valley region comprising the counties of San Joaquin, Stanislaus, Merced, Madera, Fresno, Tulare, Kings, and Kern saw range and pasture conditions averaging only 28% during 1977, the lowest of any
region in the state, and 11 points below the state average. This region suffered some of the most severe effects of the drought and felt the impact over two years, being the first to feel the impact of continued drought. Stock water supplies dried up in early 1976. Dry feed was nonexistent and dry roughage was in short supply until the 1977 hay crop was harvested. Pasture leases were almost nonexistent during 1977. The San Joaquin Valley was the only region in California where the number of cattle being fattened for slaughter market on January 1, 1978 was below a year earlier. Breeding herds were in generally poor condition. Reports indicated that herd reductions in 1976 came about by culling cows and older breeding stock, but in 1977 the numbers disappearing were the younger stock. Producers attempted to hold onto cows and springer heifers in hopes of being in business in 1978.

The number of cattle in the areas affected by the drought was in excess of 3 million head. Losses to the cattle industry in the San Joaquin Valley in 1977 were estimated to have been 206 million.

Besides cattle, there were nearly 900,000 sheep in the areas affected by the drought. Losses to producers in these areas included forced liquidation or major culling, reduced grazing on public lands, loss of irrigated pasture, higher supplemental feed costs, and increased cost of leases. The 1977 loss to the sheep industry in additional costs and reduced production was estimated to be in excess of $6.0 million statewide.

California’s small grain producers suffered from a second year of drought in 1977 with substantial abandonment of planted fields and poor yields on much of the dryland grain brought to harvest. While wheat growers shifted acreage to irrigated ground wherever possible, final estimates showed that 27% of the planted wheat acreage was not harvested, nearly three times the normal loss. Similarly, some 17% of the barley acreage was not harvested. Dryland grain losses in 1977 due to abandonment and yield reduction was estimated at $11.5 million for wheat, nearly $10 million for barley, and $1.5 million for oats. These are all statewide totals.

The sharp curtailment in surface water availability for irrigation also forced field crop producers to leave a substantial acreage of cropland idle. Statewide, nearly 125,000 acres of irrigated cropland was out of production during the 1977 season, with most of the idle land in Fresno and Kern Counties. Using gross values per acre for crops normally grown on this land, it was estimated that field crop producers lost nearly $89 million of income from land idled in 1977.

Cotton acreage in 1977 was increased substantially in the face of sharply curtailed water supplies, since expected crop returns at the time of planting were very favorable. Cotton yields were much better than expected in the San Joaquin Valley.

Total field crop losses due to the drought in 1977, including idle irrigated cropland, were estimated at $112 million statewide. The overall effect of the drought on fruit and nut crops in 1977 was much less than had been expected since producers turned to water-saving techniques such as drip irrigation and more efficient sprinkler systems. Record crops were realized for almonds, plums, and nectarines. Total loss to fruit and nut crops, largely in grapes and walnuts, was in excess of $40 million statewide.

In general, there were no significant overall losses in vegetables. However, canning tomato acreage was reduced in some areas of the San Joaquin Valley, particularly Fresno and Kern Counties, with compensating increases in coastal areas and the Sacramento Valley. There were also significant shifts in lettuce and melon acreage in the San Joaquin Valley, and 1977 melon yields were lower. In contrast, lettuce supplies were larger during the summer months resulting in low returns to producers.

There were other costs to California growers besides those associated with reduced income. The cost of applying irrigation water rose sharply in 1977 with increased power bills to pump groundwater to replace surface water supplies. In addition, there were added bills caused by increased well drilling to provide groundwater.

While well costs varied greatly depending on their depth and the size of the casing, it was estimated that the cost of well drilling for agricultural use totaled $300 million in 1977. Statewide, the extra energy associated with the required lift and additional groundwater extraction was estimated to have required about one billion kilowatt-hours of energy in 1977 at a cost of over $25 million.

**Fresno County.** The eastern 60% of Fresno County, two-thirds of whose growers normally use surface water, had few problems with the drought. Virtually every grower in this part of the county had wells to augment short surface supplies. Fresno Irrigation District’s (FID) irrigation run in 1977 was the shortest ever, just 2 months, extending only from June 1 to August 1. FID later reported that they had only delivered water for six weeks
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts during 1977. In any case, FID wouldn't experience drought conditions like this again until 2014 when they were only able to deliver water for six weeks.\footnote{1481}

Groundwater levels in July 1977 for FID were down an average of just over 7 feet compared to the previous year. Wells in the Consolidated Irrigation District (CID) were also down an average of just over 7 feet when compared to the previous summer. For the second year in a row, CID did not deliver surface water. Approximately 1,800 wells were drilled in Fresno County during the first nine months of 1977.

The most serious problem in Fresno County was experienced in the Westlands Water District where most of the wells had been put out of service with the arrival of CVP water. Part of the reduced federal water supply, cut 75% in 1977, was replaced with water made available from MWD and USBR's water transfer program. Despite these actions, 69,500 acres remained unplanted in the federal service area portion of Westlands. In 1977, a very high acreage of cotton was planted, mostly due to an expectation of a favorable price, but also somewhat due to its lower water requirement. Some Westlands acreage was diverted from sugar beets, which use more water, and from processing tomatoes. 1977 was the first year since the introduction of imported water in the late 1960s that land had to be left out of production because of lack of water.

Estimated pumping in Westlands for 1976 was less than 300,000 acre-feet; it was 408,000 acre-feet for 1977. The latter figure is still less than annual groundwater withdrawals before surface supplies became available.

The reduction in gross farm income in Westlands as a result of the drought was in excess of $100 million. In addition, farmers spent an estimated $7.7 million for new and rehabilitated wells, and more than $7 million for tailwater return systems and sprinkler and drip irrigation systems. Employees were laid off because of the reduced acreage in the district. In 1977, Fresno County had a gross farm income of over $1 billion for the third year in a row. Income was down, however, from the record 1976 level of $1,170,800,000.

Irrigation practices used in Westlands and in the rest of Fresno County to stretch water supplies included skipping of cotton, simply applying less water (in some cases only a preirrigation of cotton), alternate row irrigation of deciduous fruits and vines, and the installation of return flow systems. Many farmers applied for federal financial assistance to convert open ditches to pipelines.

Kings County. Farmers in Kings County were able to make up deficiencies in the surface water supply by pumping. County districts also received some of the water given up by Metropolitan Water District of Southern California (MWD) and other Southern California state water contractors. This was very helpful in the Westside and Tulare Lake Basin areas which normally depend on SWP and CVP water.

Field crop acreages were generally down in Kings County. About 15,000 acres of field crop land was left unplanted. There was an increase in barley acreage and a decrease in wheat acreage which reflects the lower water requirement for barley. Gross farm income in Kings County was reduced about 5% in 1977 below the 1976 level of $403,002,100.

Tulare County. The CVP's Friant-Kern Canal deliveries were decreased to 25% of Class I entitlements in 1977. This was expected to cause serious problems in Tulare County, particularly in the citrus growing area east of the valley floor. USBR, through water transfer programs, and the farmers themselves, by drilling wells and purchasing water from other farmers who could pump into the Friant-Kern Canal, developed a sufficient water supply to allow for nearly normal production.

The rush to drill irrigation wells was most prevalent in the citrus area, and water levels there dropped drastically. For example, the water level in the Exeter Irrigation District dropped from 48 feet in February 1976 to 73 feet in October 1977. In 1950, another drought year, it had declined to 109 feet before rebounding with the advent of Friant-Kern Canal deliveries.

Tulare County experienced a very slight decrease in gross farm income — from $743,327,000 in 1976 to $734,755,000 in 1977. However, there were some drastic shifts in crop acreages. Cotton acreage was up, as it was everywhere. The acreage of alfalfa hay, field corn, grain sorghum, barley, wheat, and sugar beets all showed marked decreases. The aggregate decrease was much larger than the increase in cotton, reflecting a substantial decrease in double cropping.

Little Kern golden trout were salvaged in Tulare County in 1977 when streams became intermittent during the summer months. The rescued fish were returned to their streams following resumption of continuous flow.
Kern County. Cotton acreage in Kern County was greatly increased in 1977 over 1976, but the effects of the drought were still to be seen in the form of reduced yields and quality because of less than optimum amounts of applied water and, in some cases, poor groundwater quality.

Western Kern County, where the entire water supply had to come from the SWP’s California Aqueduct, was most severely affected by the drought. Some of the water service agencies in that area requested deferral of about $6,000,000 in 1977 payments to the SWP. This became unnecessary as federal funds became available to some of those districts.

In the entire Kern County SWP service area, an estimated 56,000 acres of row crops remained unplanted in 1977 because much of the available surface water was allocated to save permanent crops and much of the area had no groundwater supply. There was a reduction in double cropping throughout Kern County, greatly reducing the grain sorghum acreage. Some alfalfa fields were allowed to dry up.

Some Kern County wells had water level drops of 15 feet or more in 1977 instead of the normal 2- to 3-foot drop.

Gross farm income in Kern County was down by $70 million in 1977 from the 1976 level of $873,655,800. In addition, production costs increased due to increased water and power rates.

Contingency Planning Efforts
The state had plans in place for what to do if the drought had continued, and 1978 turned out to be a repeat of 1977. Fortunately that turned out not to be the case; 1978 was a year of abundant rain.

1976 Floods (4)
There were at least four periods of flooding during 1976:
1. February
2. September/October (3)

The winter of 1975–76 was a strong La Niña event.

February 10 was the seventh consecutive day of measurable rain in Fresno, with 4.01 inches falling from February 4–10. Daily precipitation records were set on both February 5 (0.83 inches) and February 9 (1.50 inches). The Fresno City Works department distributed 1,800 sandbags to Fresno residents as several streets and poor drainage areas in that city flooded.¹⁴⁸² Table 76 gives the total precipitation during the February 4–10 storm event for selected reporting stations.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Fresno</td>
<td>4.01</td>
</tr>
<tr>
<td>Sanger</td>
<td>4.27</td>
</tr>
<tr>
<td>Dinuba</td>
<td>3.86</td>
</tr>
<tr>
<td>Batterson</td>
<td>4.15</td>
</tr>
<tr>
<td>Coarsegold</td>
<td>4.89</td>
</tr>
<tr>
<td>North Fork</td>
<td>5.57</td>
</tr>
<tr>
<td>Poison Ridge</td>
<td>5.60</td>
</tr>
</tbody>
</table>

This storm also dropped an inch of snow in San Francisco. That city wouldn’t see any significant snow again until February 2011.

The September/October floods in the Tulare Lake Basin occurred near the beginning of the 1976–77 drought.

Hurricane Kathleen formed off the coast of Baja California on September 9. It was a hurricane for only six hours and was a tropical storm when it made landfall on September 10. Kathleen weakened to a depression after it crossed the U.S./Mexico border near El Centro, but its circulation allowed gale-force winds to be recorded in Arizona and California.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

The storm wasn’t finished yet, as flooding rains continued to plague the Southwest along with gale-force winds. Kathleen’s rapid forward speed allowed it to keep its strength for a long time over land. Kathleen is one of only six recorded tropical cyclones in the eastern Pacific Ocean known to have brought gale-force or hurricane-force winds to the continental United States.

Kathleen moved northward through the deserts of California bringing rain to interior Central California from September 9–11. The heaviest one-day totals were on the 10th at most locations. Two people were swept to their deaths when Interstate 8 was washed out. Lodgepole set an all-time 24-hour precipitation record for the month of September with 5.06 inches of rain. Rainfall totals for the 3-day event were between ½ and 1 inch in the valley, averaging 1–2 inches in the foothills and 3–6 inches in the Sierra. Specific event totals are shown in Table 77.

Table 77. Precipitation during the September 9–11, 1976 storm event.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Yosemite South Entrance</td>
<td>2.08</td>
</tr>
<tr>
<td>Grant Grove</td>
<td>3.96</td>
</tr>
<tr>
<td>Lodgepole</td>
<td>6.21</td>
</tr>
<tr>
<td>Fresno</td>
<td>0.90</td>
</tr>
<tr>
<td>Bakersfield</td>
<td>0.53</td>
</tr>
</tbody>
</table>

Tropical Depression Kathleen dissipated on September 11 while over southern Nevada as it continued accelerating.

Thunderstorms struck the central and southern San Joaquin Valley on September 29 with up to 2½ inches of rain falling in some areas. Dramatic lightning displays were seen from Fowler to Delano, and marble-size hail fell in Visalia and Porterville. The storm knocked out power to several thousand customers and struck two F106 jets operated by the Fresno Air National Guard, causing burn marks on the planes. The heavy rain also caused a roof to collapse at a building under construction as well as flooding homes, businesses and streets. It also caused additional damage to crops that had been seriously affected by the rain associated with Tropical Depression Kathleen.

Heavy rain drenched parts of the central San Joaquin Valley on October 1. Fresno received 1.46 inches of rain, setting a daily precipitation record. Several roads were heavily flooded in that city, temporarily stranding some motorists. Los Banos received ½ inch of rain in just 30 minutes. Many roads and fields in Mendota were flooded.

Flooding occurred on the west side of the San Joaquin Valley near Coalinga in 1976. Presumably this resulted either from Kathleen on September 10 or from the storm system that came through the valley on September 29 – October 1.

1977 Flood
Flooding in 1977 occurred in December. This flood occurred near the end of the 1976–77 drought.

Soaking rains fall in Kern County from December 27–28. Storm totals were 2.05 inches in Lost Hills and 1.11 inches in Bakersfield. Water was two feet deep at some intersections in Bakersfield, stranding some motorists.

1978 Floods (3)
There were three periods of flooding in 1978:
1. February
2. Early summer
3. September

The winter of 1977–78 was a weak El Niño event. This association with the 1978 floods was almost certainly a coincidence. Only strong El Niño events have been shown to have a correlation with high precipitation events and floods in California.
FLOODS AND DROUGHTS IN THE TULARE LAKE BASIN

SPECIFIC FLOODS AND DROUGHTS

Flooding occurred on the west side of the valley in February. Flooding occurred on the east side of the valley and in the Tulare Lakebed in early summer due to high runoff. River flooding occurred in September due to a tropical downpour throughout the Sierra.

Former state climatologist Jim Goodridge rated the February 1978 flood as one of the 10 most damaging floods in the state’s history. Ventura County received over 13 inches of rain in one day, resulting in floods and landslides.

A vigorous winter cyclic storm with widespread flooding and mudslides developed on the windward slopes of the South Coastal Basin on February 10. There was $120 million in storm-related property damage and 18 deaths. This storm was still quite robust as it moved northeasterly into the rain shadow zone of the comparatively dry Buena Vista Lake Basin.

The storm of February 10, based on recurrence interval, was centered in the area around Buena Vista Lake. Blackwells Corner (intersection of Highways 33 and 46) received 3.90 inches of rain on February 10, which was 74% of its average annual precipitation. This was 7.41 standard deviations above the average maximum day with a recurrence interval of 28,000 years. A total of 32 stations reported recurrence intervals in excess of 100 years and 16 stations reported recurrence intervals in excess of 1,000 years.

Bakersfield received 2.29 inches of rain on February 9. That was the wettest day ever in that city. Bakersfield received a total of 5.36 inches of rain during February, making it the wettest month ever in that city. That record would eventually be broken in December 2010.

This was a heavy rain event combined with snowmelt runoff in some areas. Over 6,000 acres were flooded in Kern County, causing extensive damage to agricultural lands. The Lamont/Arvin area was flooded. Mudslides, landslides and debris flows were common. One woman died in Kern County when her car was swept off Interstate 5 by a mudslide. Transportation routes, including rail traffic, were suspended for as long as three days. A total of 91 county roads were closed. The California Aqueduct was damaged. Bridges, culverts and other flood control works were badly damaged. Domestic water supply and sewer lines were washed out. Oil field facilities were also damaged. Total damage in Kern County was approximately $25 million. President Ford declared Kern County a disaster area on February 15.

Within the Tulare Lake Basin, the storm’s effect was felt primarily on the west side of the valley. Panoche/Silver Creek west of Mendota flooded in February. Hanford received 2.4 inches of rain on February 10, the most that city has ever received in any 24-hour period. Flooding occurred along Los Gatos Creek and Arroyo Pasajero from the foothills to the valley floor and damaged agricultural lands, roads and bridges, and utilities. An estimated 4,500 acres were flooded, and damage totaled $160,000.

There is a strong resemblance between this storm and the remnants of a hurricane which came onshore near Monterey Bay on September 11, 1918. Both were robust cyclic storms which vigorously entered the rain shadow areas to the northeast, resulting in a deluge in normally dry areas.

The February 1978 Buena Vista Lake storm also resembled the March 1995 storm. That storm produced devastating rainfalls on the windward slopes of the Coast Ranges. It was still quite energetic as it moved into the rain shadow area to create further devastating floods. That was the storm that washed out the Interstate 5 bridges near Coalinga.

One of the side-effects of the February 1978 storm was the grand display of wildflowers seen in the vicinity of the Tulare Lakebed by mid-March of that year. That was a result of the thorough soaking of the ground at an optimum time of the year.

The February 1978 storm was spectacular south of the Tehachapis. In contrast, the September 1978 storm would be spectacular in the Kings River Basin and to the north. In the Kaweah River Basin, the February storm was a much bigger flood event than the September storm. It produced a peak average daily flow at Terminus Dam of 8,135 cfs while the September storm produced a flow of only 3,890 cfs. The Kaweah’s peak natural flow occurred at McKay’s Point on February 9: 14,700 cfs. (That was the peak hourly flow; the peak average daily flow was 8,135 cfs.)

Flood releases of 8,000 cfs or greater occurred at Friant Dam on the San Joaquin River during the April–July snowmelt period. The peak daily flow at Pine Flat on the Kings occurred on June 9 during snowmelt.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Kerry Arroues said that Cross Creek overflowed its banks in one of the 1978 floods. It isn’t certain which flood that was. It seems like it was probably in early summer due to high runoff. Kerry saw floodwaters one foot deep three miles west of the channel, next to some eucalyptus.

Hurricane Norman was a powerful Category 4 hurricane with a 40-mile-wide eye and sustained winds of 140 mph. It developed in early September well off the coast of Acapulco, and then slowly weakened as it moved over cooler waters west of Baja California. By early on the morning of September 4, moisture from that hurricane had spread north, initiating rains in California. Norman then recurved, turning north toward Southern California. It made landfall as a tropical depression on September 5–6.

When a Pacific hurricane degrades, it usually makes landfall in Southern California or in Mexico. Norman came ashore in the LA area, but its track was aimed straight for the southern end of the Sierra. As the remnants of Norman plowed inland, heavy rains fell across the Sierra, with a maximum amount of 7.01 inches reported at Lodgepole. Rivers rose so quickly that roads closed, and campers found themselves marooned all over the Southern Sierra.

George Durkee (national park wilderness ranger) recalled that the storm was known in the national parks as “Stormin’ Norman.” Tighe Geoghegan was the parks’ fire dispatcher, and one day she reported “rain, rain and more rain” on the weather report. Many backpackers got drenched and sought refuge in the wilderness ranger stations. Two backpackers died of hypothermia at Trail Camp on the east side of Mt. Whitney. They had apparently gotten soaked in the rain down by Crabtree and then encountered cold temperatures and probably sleet at Trail Crest.

Approximately 4 inches of rain fell at Cedar Grove. At Pine Flat, the natural daily flow on the Kings River on September 5 was over 10 times greater than it had been the previous day. It was not quite as high as the natural daily flow that had occurred on June 9 during snowmelt, but it was still a very impressive event.

Labor Day occurred on September 4. Jerry Torres, Kings Canyon National Park’s trails supervisor at the time, recalled that the flood over the Labor Day Weekend covered the North Side Road and spilled over onto Highway 180 in two spots: west of Grizzly Falls and near the Boyden Bridge. Based on the flood exceedence rates in Table 29, this had a recurrence interval of only 8 years for the Kings River downstream at Pine Flat. Still, that puts it in a category with other Cedar Grove floods of the past 70 years (1937, 1950, 1955, 1966, 1969, 1978, 1982, 1983, and 1997) that rise to the level of the modeled 50-year flood: that is, a flood event that occurs about every eight years on average. See the section of this document that describes Cedar Grove Flooding.

Flooding from this storm caused widespread damage to roads and bridges on the east side of the Sierra. This was a negligible flood on the Kaweah and the rivers to the south.

Norman also caused flooding on the east side of the Sierra. At Lake Sabrina, a 10-hour duration rainfall of 1.02 inches was recorded on September 5, producing a peak flow of 940 cfs at Power Plant No. 6 on Bishop Creek. This was the second-largest flood-of-record on that creek and was estimated to have a recurrence interval of 50 years. To prevent flood damage to Bishop and surrounding areas, the LADWP activated the Owens River Canal Bypass.

Total flow for water year 1978 was 203% of the 1894–2014 average for the Kings, 196% for the Kaweah, 199% for the Tule, and 224% for the Kern.

By some measures, 1978 was the wettest water year in Kern County since record-keeping began in 1889. Bakersfield received 12 inches of rain.

Flooding occurred in the Tulare Lakebed; this was the first significant flooding since 1973 (see Figure 16). In order to minimize flooding, 9,000 acre-feet of river floodwater was pumped into the Friant-Kern Canal and routed to the Los Angeles area. Tulare Lake grew to 70 square miles as a result of this flooding.
1980 Flood

There were two floods in 1980:
1. January
2. February

The national parks' records make no mention of any flooding in 1980.

Flood releases of 8,000 cfs or greater occurred at Friant Dam on the San Joaquin River during the winter (prior to snowmelt).

The peak day natural flow at Pine Flat on the Kings occurred on January 13. The flow on that date was 37 times greater than the flow just four days earlier. It was approximately as large as the peak day flow in the 1963 flood. It seems likely that this was a high-flow period in Cedar Grove as well.

The Kaweah's peak natural flow occurred at Terminus Dam (or possibly this was McKay's Point) on January 13: 34,000 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 16,933 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 13 years for the Kaweah and a recurrence interval of 10 years on the Tule.

In the Tulare Lake Basin, the second flood of the year occurred on February 18. The Kern had a much bigger relative response to this storm than the rivers farther to the north. This was the same pattern as in the 1916 flood when the storm was to the south of the Tulare Lake Basin.

The February 14–21, 1980 flooding was most severe in Central and Southern Coastal California. It had a recurrence interval of up to a 50 years on some rivers. Disastrous and record-breaking rainfalls in the South Coastal Basin resulted in the highest-ever rainfall totals over a broad area. Record-high eight-day rainfalls occurred at 133 stations. Recurrence intervals in excess of 100 years were reported at 70 stations. Over 1,500 homes were damaged or destroyed; there was a total of $270 million in property damage, and there were 18 storm-related deaths. Seven counties were declared disaster areas.

Total flow for water year 1980 was 180% of the 1894–2014 average for the Kings, 208% for the Kaweah, 240% for the Tule, and 225% for the Kern. Flooding occurred in the Tulare Lakebed; this was the first significant flooding since 1978 (see Figure 16). In order to minimize flooding in the Tulare Lakebed, 5,000 acre-feet of river floodwater was pumped into the Friant-Kern Canal and routed to the Los Angeles area.

1982–83 Floods (11)

There were at least 10 floods in 1982–83:
1. April 1982 (rain-on-snow event)
2. June 1982 (severe storm)
3. September 1982 (due to remains of Hurricane Olivia)
4. October 1982 (due to remains of Hurricane Sergio)
5. December 1982 (rain-flood)
6. March 1983 (severe storm)
7. Memorial Day Weekend, 1983 (four debris flows)
9. August 1983 (two severe storms caused by monsoonal moisture)
10. September 1983 (severe storm)

The 1982–83 was the strongest El Niño event recorded over the past 50 years. The winter of 1981–82 experienced heavy snowfall in the Sierra. Record after record was broken:
- Echo Summit received 67 inches (5.6 feet) of snow in 24 hours on January 4–5, 1982, breaking the state record that had been set by Giant Forest in January 1933.
- On January 5, 1982, 29.5 inches of snow fell in Yosemite Valley, setting the record for the biggest 24-hour snowfall ever at that location.
- The Central Sierra Snow Laboratory monitoring site near Donner Pass received 15.5 feet of snow between March 27 – April 8, 1982, the biggest snowstorm ever recorded at that site.

That heavy snowpack set the stage for a big spring runoff on all the rivers.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Flood releases of 8,000 cfs or greater occurred at Friant Dam on the San Joaquin River during the April–July, 1982 snowmelt period.

On April 11–12, 1982, a combination of snowmelt and rainfall caused the Merced River to overflow its banks and flood parts of Yosemite Valley. The national park headquarters building was damaged, and parts of a road were washed out. The Merced River at Happy Isle peaked at 4,880 cfs. By Merced River standards, that is a fairly modest flood, having a recurrence interval of 8 years.

Easter Sunday fell on April 11, so this flood is sometimes referred to as the Easter 1982 flood. Table 78 gives the precipitation totals for the reporting stations in Sequoia and Kings Canyon National Parks.

Table 78. Precipitation during the April 11–12, 1982 storm event.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Grant Grove</td>
<td>7.44</td>
</tr>
<tr>
<td>Lodgepole</td>
<td>9.33*</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>4.57</td>
</tr>
</tbody>
</table>

*6.83 inches of this total fell on April 11

Jim Harvey recalled that Elk Creek had a big flood in the early 1980s; this seems like the probable time when that event would have occurred. Jim said the flood damaged the Generals Highway, much as it had in the 1935 flood.

The peak day natural flow at Pine Flat on the Kings occurred on April 11, 1982. The flow that day was 10 times larger than the flow of the previous day. That is the fourth largest peak day of the year at Pine Flat since the dam was built in 1954. That would suggest that there was very high water in Cedar Grove as well, but we have no national park records to substantiate that.

The Kaweah’s peak natural flow occurred at Terminus Dam (or possibly McKay’s Point) on April 11, 1982: 28,800 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 18,514 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 17 years for the Kaweah.

On June 18, 1982, an intense thunderstorm occurred at Forni Ridge in the Eldorado National Forest. This storm and the resulting debris flow were outside the Tulare Lake Basin, but the story merits inclusion in this document as an example of how intense a summer storm can be. The storm lasted only a short time, but is notable because 4.02 inches of rain was measured in 30 minutes (a record-setting rate of 8.04 inches per hour). The storm was centered over a recently burned steep mountain slope adjacent to U.S. Highway 50. The storm was followed by a debris flow that closed the highway.

On June 30, 1982, there were numerous reports of funnel clouds over Clovis, and one touched down near Fresno State University. Thunderstorms caused street flooding in Farmersville and also flooded homes in other parts of the valley. Dinuba received particularly heavy rain.

The June 30 thunderstorm system extended into the national parks. Table 79 gives the precipitation totals for some of the reporting stations in the area.

Table 79. Precipitation during the June 30, 1982 storm event.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Grant Grove</td>
<td>1.15</td>
</tr>
<tr>
<td>Lodgepole</td>
<td>0.65</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>1.30</td>
</tr>
<tr>
<td>Dinuba</td>
<td>1.62</td>
</tr>
</tbody>
</table>

Thanks to a photograph taken by Bill Tweed, we know that the Marble Fork Kaweah flooded through Lodgepole in June (photograph on file in the national parks). This suggests that there was a strong thunderstorm cell in the Tablelands area.
Hurricane Olivia formed about 400 miles south of Acapulco on September 19, 1982. It developed winds of 130 mph, becoming the strongest storm of the season. It gradually weakened as it passed over cooler waters west of Baja California. Olivia then recurved and came ashore as a tropical depression. When a Pacific hurricane degrades, it usually makes landfall in Southern California or in Mexico. Olivia came ashore near the U.S./Mexico border, but its track was aimed for Utah. Olivia’s storm track, as illustrated in Figure 28 followed a fairly typical pattern for Pacific hurricanes that degrade and then make landfall in Southern California.

As the remnants of Olivia plowed inland, heavy rains fell across the San Joaquin Valley and the Sierra from September 23–27. Measurable rain fell from September 24–26 in both Fresno and Bakersfield. Storm totals included 0.70 inches at Bakersfield and 1.10 inches at Fresno, although locally heavier amounts were reported. Far heavier amounts fell in the Sierra with a maximum amount of 7.19 inches reported at Grant Grove.1502 The heavy rain wiped out half of California’s raisin crop, a quarter of the wine crop, a tenth of the tomato crop, and also damaged the almond crop. The rain caused power outages to over 10,000 customers in Fresno County.

Jerry Torres, Kings Canyon National Park’s trails supervisor at the time, recalled that it rained nonstop for at least two days in Cedar Grove, resulting in a major flooding event in the Kings River Basin. Pine Flat experienced a large and abrupt increase in flows on September 25.

The Middle Fork Kings Bridge at Dougherty Creek was constructed in the summer of 1979. Jerry was part of that construction. That bridge was washed out in the September 1982 flood.

Flooding also occurred on the South Fork Kings. In Cedar Grove, the western ¼ mile section of the North Side Road had water over it. The flood damaged a section of Highway 180, 100 yards west of Grizzly Falls. However, the biggest damage occurred two miles west of Grizzly Falls where an entire hillside was washed away, including a section of Highway 180.

Access into Cedar Grove was closed for four days until Caltrans used a dozer to literally cut a road out of the mountain. The September 1982 flood is sometimes remembered in the park as the “Great Trails End Flood” because it occurred during the annual “Trails End” end-of-year celebration in Cedar Grove.

Based on the flood exceedence rates in Table 29, this had a recurrence interval of 35 years for the Kings River downstream at Pine Flat. That puts it in a category with other Cedar Grove floods of the past 70 years (1937, 1950, 1955, 1966, 1969, 1978, 1982, 1983, and 1997) that rise to the level of the modeled 50-year flood: that is, a flood event that occurs about every eight years on average. See the section of this document that describes Cedar Grove Flooding.

The Kaweah’s peak average daily natural flow, as reflected in Table 28, occurred on September 26, 1982: 6,308 cfs. Based on the flood exceedence rates in Table 29, this had a recurrence interval of 4 years for the Kaweah.
The effects of Olivia extended to the east side of the Sierra. The greatest recorded floods in several east-side streams occurred in late September when 6–8 inches of rain fell in two days. Some 38 homes were damaged on the Big Pine Indian Reservation, and more than 40 homes were damaged elsewhere. Bridges were washed out on the Glacier Lodge/Big Pine Creek road and Pine Creek road. The Inyo County Public Works Department estimated the flooding caused more than seven million dollars of damage throughout the county.\textsuperscript{1503}

The fourth flood of the 1982–83 period occurred on October 26, 1982. We know only a little about it. A tropical disturbance was noted southwest of Costa Rica on October 12, 1982. The system organized into a tropical depression late on October 13 and became a hurricane late on the following day. By the afternoon of October 17, Hurricane Sergio was packing sustained winds of 120 mph. Cooler water was reached soon afterwards, and weakening commenced. While slowly moving west, Sergio weakened to a tropical storm by the afternoon of October 21 and to a tropical depression late on October 22. The system dissipated on the afternoon of October 23.

On October 26, $563 million in agricultural damage was reported in Central California when tropical moisture from former Hurricane Sergio spread over the state. Presumably the Tulare Lake Basin was part of that flood area.\textsuperscript{1504}

All that we know about the December 1982 flood comes from the gaging stations at the various dams. The gages all jumped sharply on December 22. Evidently it was a rain-on-snow event. The rise in the rivers was least noticeable in the Kern River Basin, so apparently the storm was located more to the north.

The Kaweah’s peak natural flow occurred at Terminus Dam (or possibly McKay’s Point) on December 22, 1982: 11,100 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 8,325 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 6 years for the Kaweah.

One of the 1982 flooding events, perhaps the September one, was a major flood on the Tule.

Total flow for water year 1982 was 184% of the 1894–2014 average for the Kings, 182% for the Kaweah, 168% for the Tule, and 160% for the Kern.

The heavy spring runoff in 1982 resulted in flooding in the Tulare Lakebed. This was the first significant flooding in the lakebed since 1980 (see Figure 16). In order to minimize flooding in the lakebed, 33,000 acre-feet of river floodwater was pumped into the Friant-Kern Canal in 1982 and routed to the Los Angeles area.

New record-high total annual rainfalls were reported from stations located over a broad range of California during water year 1983 (October 1982 – September 1983). California received a long sequence of storms which left poorly drained areas soaked for many months. This soaking resulted in unusually extensive flooding in several parts of the state. In all regions, the high rainfall totals were associated with a quite noticeably increased numbers of rainy days, rather than with large individual rainfalls.

It had been 93 years since California had as much rain as in water year 1983. The last year with rainfalls as high was 1890. One of the factors which make the 1983 year even more unusual was that 1982 was also one of the wettest years of record. A total of 58 stations reported 100 or more inches for water year 1983. A total of 511 stations reported their wettest year ever during 1983.

During water year 1983, half of the state’s land area had rainfalls in excess of a recurrence interval of 100 years. During that same period, 45 stations reported yearly rainfall totals that were in excess of the 1,000-year amounts. These were distributed from the Klamath River Basin in the north to the Borrego Desert in the south.\textsuperscript{1505}

The winter of 1982–83 was a potent El Niño event. The impact of this El Niño on California’s weather in 1982-83 was complex. The high-pressure ridge between 10 and 20 degrees north latitude was magnified by the heat from the warmer than normal ocean water. Simultaneously, extremely low air pressures developed over the Gulf of Alaska. These contrasting pressure extremes caused the westerly airflow across the Pacific to double. The jet stream that directs storms into California was intensified and displaced to the south so that storms hit the Central California coast fiercely and more often.\textsuperscript{1506}
Northern and Central California experienced flooding incidents from November 1982 through March 1983 due to numerous storms. The melting of the record snowpack then created a second episode of flooding from May through July of 1983.

Statewide, the two wettest water years during historic times were 1890 and 1983. The statewide precipitation for water year 1983 averaged 190% of normal, with many areas well over 220%. New precipitation records were set at 49 locations in the state.

Table 80 summaries the increased precipitation for the three drainage basins from the Upper San Joaquin to the Kern River.

Table 80. Precipitation totals during winter 1982–83.
For the three drainage basins from the Upper San Joaquin to the Kern River.

<table>
<thead>
<tr>
<th>Season</th>
<th>Percent of Average</th>
</tr>
</thead>
<tbody>
<tr>
<td>Fall 1982 (September, October, November)</td>
<td>318%</td>
</tr>
<tr>
<td>Winter 1983 (December, January, February)</td>
<td>183%</td>
</tr>
<tr>
<td>Spring 1983 (March, April, May)</td>
<td>199%</td>
</tr>
</tbody>
</table>

Yosemite recorded 66.39 inches of precipitation during 1983, breaking the record of 61.09 inches set in 1938. The mean yearly precipitation for Yosemite is 35.26 inches. Fresno received a total of 23.57 inches of precipitation during water year 1983, making that the wettest water year on record for that city.

The stage for a disastrous year of flooding had been set in the fall of 1982. In some parts of California, September 1982 was one of the wettest Septembers on record, thanks to subtropical moisture from the remains of Hurricane Olivia. Soils were saturated, and there was less than normal flood control space in many reservoirs. The 1982 and 1983 water years are the wettest pair of years on record.

As shown in Table 81, the snowfall total in Lodgepole in the winter of 1982–83 was a rather awe-inspiring 429.8 inches (36 feet).

Table 81. Lodgepole snowfall during winter 1982–83.

<table>
<thead>
<tr>
<th>Month</th>
<th>Snowfall (inches of snow)</th>
</tr>
</thead>
<tbody>
<tr>
<td>September, 1982</td>
<td>2.0</td>
</tr>
<tr>
<td>October, 1982</td>
<td>trace</td>
</tr>
<tr>
<td>November 1982</td>
<td>46.0</td>
</tr>
<tr>
<td>December 1982</td>
<td>51.3</td>
</tr>
<tr>
<td>January 1983</td>
<td>84.5</td>
</tr>
<tr>
<td>February 1983</td>
<td>74.0</td>
</tr>
<tr>
<td>March 1983</td>
<td>113.0</td>
</tr>
<tr>
<td>April 1983</td>
<td>40.0</td>
</tr>
<tr>
<td>May 1983</td>
<td>19.0</td>
</tr>
<tr>
<td>Total</td>
<td>429.8</td>
</tr>
</tbody>
</table>

Up to four feet of snow fell in the Sierra in less than 24 hours on December 22, 1982. Lodgepole received 27 inches of snow with a storm-total liquid water equivalent of 10.09 inches. The snowpack on some of the higher peaks from this storm was raised to nearly 100 inches. (We tend to think of the incredible winter of 1982–83 as being record-setting at Lodgepole, but there were at least four bigger snowpacks that we know of at that location. Bill Tweed recalled that having been through all that snow, the people who spent the winter at Lodgepole were disappointed that they had not set a new record. The winter of 1968–69 received at least 440.5 inches. In the winter of 1951–52, Lodgepole received 449.5 inches (37½ feet), setting the record for this weather station. That record would eventually be broken in the winter of 2010–2011. The most impressive winter in this area that we know anything about was the winter of 1905–06. In that winter, the snowpack reached a maximum depth of 29 feet on the level in Giant Forest. Even by June 25, 1906, the snowpack in Giant Forest had only melted down to about 12 feet on the level.)

In the winter of 1982–83, snowpack records were set at three-fourths of the Sierra snow courses. On May 3, 1983, snow water content in the Sierra exceeded 230% of normal; the ensuing runoff resulted in four times the average volume for Central Valley streams.
The trans-Sierra highway over Tioga Pass in Yosemite National Park and many Sierra wilderness trails that normally open by early summer remained blocked by snow. Snow survey measurements found the snowpack persisting later into the year leading to an unprecedented July snow survey in the Kings River Basin.\textsuperscript{1515}

Jack Vance recalled being at the Hockett Ranger Station over the Fourth of July Weekend, 1983. Hockett Meadow was one big lake with miniature icebergs floating on it. The high-watermark from that lake can still be seen inside the tackshed.

An intense storm struck the Tehachapi Mountains on March 1–2, 1983. Heavy rainfall of 2–7 inches fell during that two-day period, including 6.50 inches at Frazier Park. This triggered flash flooding on several creeks; Caliente Creek peaked at 15,000 cfs as it flowed into the southeast end of the San Joaquin Valley. Most severely impacted was Lamont, where 1,973 homes were damaged or destroyed — over half of the town. Over 33 roads were washed out in Kern County, and two 100-car trains had to be abandoned after water washed out parts of tracks. The town of Caliente was also flooded, resulting in 77 people having to be rescued by helicopter. Agricultural lands and irrigation works were also damaged and destroyed. Irrigation works were washed out. Total damage from the flood was an estimated $58.7 million. A series of storms resulted in continued flooding through March 13.\textsuperscript{1516, 1517}

Mehrten Creek flooded during one of the storm events in 1983, according to John Hansen and Peter Hickey. The flow on this normally dry channel was greater than the culvert under Highway 198 could handle, and the stream overflowed the highway, forcing its closure. There was apparently extensive flooding of the area between Highway 198 and Foothills Ditch. The store located at the northeast corner of Highway 198 and Mehrten Drive (then known as the Mehrten Market) was flooded.

With such a huge volume of water, the flooding no doubt continued downstream along lower Mehrten Creek all the way to Consolidated Peoples Ditch. See the section of this document that describes \textit{Floods on Lower Mehrten and Yokohl Creeks}. This was apparently the biggest flood on Mehrten Creek since the February 1969 flood. The next big flood on this creek that we have a record of was in December 2010.

During the 1983 Memorial Day Weekend, Kings Canyon experienced three debris flows. Another debris flow occurred in the Redwood Creek drainage in the Mineral King area at apparently the same time. These four debris flows are described at the end of this section along with other mass wasting events that occurred in 1983.

All of the major reservoirs in the Sacramento River and San Joaquin River Basins reached or nearly reached design capacity during the June and July runoff. At least two levees failed in the Sacramento River Basin. Levee breaks caused flooding at four locations along the San Joaquin River. In the Sacramento–San Joaquin Delta, four levees failed, resulting in partial or total flooding of some islands. Damage exceeded $91 million in the Sacramento River Basin and $324 million in the San Joaquin River Basin.

Flood releases of 12,000 cfs or greater occurred at Friant Dam on the San Joaquin River during the April–July snowmelt period.

The peak day natural flow at Pine Flat on the Kings occurred on May 29, 1983. It was a significant flood, but only half of what the peak day natural flow had been during the much less famous 1982 flood. It seems likely that this was a very high-flow period in Cedar Grove as well. That puts it in a category with other Cedar Grove floods of the past 70 years (1937, 1950, 1955, 1966, 1969, 1978, 1982, 1983, and 1997) that rise to the level of the modeled 50-year flood: that is, a flood event that occurs about every eight years on average. See the section of this document that describes \textit{Cedar Grove Flooding}.

Two severe storms occurred in August 1983: one in the southeastern part of the Tulare Lake Basin and the other farther north in the Kings Canyon high country. Both were caused by monsoonal moisture which typically originates in the vicinity of the Four Corners Area. The first storm occurred in the southeastern part of the Tulare Lake Basin. On August 16, more than 1½ inches of rain fell in the Tehachapi Mountains in an hour, washing out portions of Highway 58.\textsuperscript{1518} On August 17, portions of California City were flooded after heavy rain fell in the Tehachapi Mountains and caused Cache Creek to swell. Water was the height of car windows and some houses flooded.\textsuperscript{1519}
Apparently the monsoonal surge continued moving northwest along the crest of the Sierra. Lodgepole received 1.48 inches between August 15–18. Huntington Lake received 1.00 inches between August 15–17. The moisture never reached Yosemite. There may have been multiple severe storms or cloudbursts over the national parks between August 15–18, but only the following account survives.

The second severe storm of August 1983 occurred when a very large black cloud brought intense rain to a section of the Kings Canyon high country. George Durkee witnessed that event while he was the wilderness ranger at the McClure Ranger Station. Ralph Kumano, who was a wilderness ranger on patrol on the Monarch Divide, first spotted the black cloud when it was over Enchanted Gorge, northeast of Tehipite Dome. From there, the cloud moved north to Mt. Darwin near the Evolution Valley.

The cloud settled over Darwin Canyon. It apparently dumped onto the Lamarck and Darwin Glaciers. It resulted in major flooding of Darwin Creek. A lot of debris plugged the creek where it crosses the Pacific Crest Trail (UTM 347825E 4115568N NAD83 Zone 11). It was a very localized event; there was no rain at all at McClure Meadow just two miles down the trail.

At the Pacific Crest Trail, there is a log crossing over Darwin Creek where the water flows under the log. There is a pool just upstream from the trail crossing. The flood filled that pool with sand and rocks, forcing the creek to flow over the log. (The idea is that the creek is supposed to flow under the log; hikers should walk on top of the log.) Both 1982 and 1983 were El Niño years with high runoff, but the flows in those years hadn’t blocked up the log crossing, so it shows what a single event can do. George got out into the pool and dug out under the log until he could clear it a little. He had to do that a couple of times over a week or two because debris kept clogging the opening under the log.

Lower Darwin Lake drains Lamarck Glacier. After the flood, George observed that this lake had turned the milky blue that is associated with suspended glacial silt. That suggests that the flood may have breached an ice dam on Lamarck Glacier. Alternatively the change in the lake could have been caused by the intense rain coming down on the glacier, which drains through a hole in front of the Little Ice Age moraine. In either case, the storm had washed down a lot of fine glacial silt off of the glacier.

After the storm, George also checked out Enchanted Gorge. That gorge is very narrow and had been mostly filled with avalanche snow, much of which had melted. Toward the bottom of the gorge, there was recent scarring on trees about four feet up from the base and there was a lot of sand at the confluence with Goddard Creek. There was no evidence of flooding in Goddard Creek above its confluence with the gorge. Large-scale debris flows usually occur in small, steep stream channels and are often mistaken for floods. The Darwin Canyon event may have been a flash flood that carried a large amount of debris, or it may have been a debris flow. We just don’t have enough data to clearly classify it. The Enchanted Gorge event was probably a flash flood.

The 11th flood of the 1982–83 period occurred on September 30, 1983. We know relatively little about it. Apparently it was caused by a thunderstorm system that spread along the Sierra. Table 82 gives the precipitation totals for the reporting stations in the national parks.

Table 82. Precipitation during the September 30, 1983 storm event.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Grant Grove</td>
<td>1.14</td>
</tr>
<tr>
<td>Lodgepole</td>
<td>1.10</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>1.22</td>
</tr>
</tbody>
</table>

Thanks to a photograph taken by Bill Tweed, we know that the Marble Fork Kaweah flooded through Lodgepole in September (photograph on file in the national parks). This suggests that there was a strong thunderstorm cell in the Tablelands area.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

As shown in Table 83, the combined runoff of the four rivers in the Tulare Lake Basin during water year 1983 was 8,746,222 acre-feet, the largest runoff since record-keeping began in 1894. For comparison, that is 5.4 times the combined current capacity of the federal reservoirs on those four rivers.

<table>
<thead>
<tr>
<th>Watershed</th>
<th>Total Runoff (acre-feet)</th>
<th>% of average (1894–2014)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings</td>
<td>4,286,703*</td>
<td>258%</td>
</tr>
<tr>
<td>Kaweah</td>
<td>1,402,005</td>
<td>330%</td>
</tr>
<tr>
<td>Tule</td>
<td>615,014</td>
<td>448%</td>
</tr>
<tr>
<td>Kern</td>
<td>2,442,500</td>
<td>340%</td>
</tr>
<tr>
<td>Total</td>
<td>8,746,222</td>
<td>297%</td>
</tr>
</tbody>
</table>

*This is measured at the KGF gage (Kings R-Pine Flat Dam). If measured further downstream at the KGP gage (Kings Pre-Project Piedra), total runoff, which included flows from Mill and Hughes Creeks, was 4,473,358 acre-feet.

This was the highest flow year for the Kings, Kaweah, and Tule Rivers since 1894. It was the second highest for the Kern River, only 1916 was higher. This was the last year that the Kaweah River (via the Consolidated People’s Ditch) flooded a significant portion of Kaweah Oaks Preserve.

The 1983 flood had a greater total runoff than the 1969 flood. Unfortunately, some parts of the runoff were not actually measured and can only be estimated. An unknown amount of the Kern River flowed into the Tulare Lakebed, but 759,000 acre-feet of the Kern was diverted into the California Aqueduct and routed to the Los Angeles area. The total estimated lakebed inflow of the other three rivers (Kings, Kaweah, and Tule) was about 1.069 million acre-feet. That was 27% more than the 0.840 million acre-feet from those same three rivers in 1969. The total inflow to the Tulare Lakebed in water year 1969 from all streams was 1.155 million acre-feet. If the 1983 runoff were 27% greater than that, then it would have been on the order of 1.467 million acre-feet if there had been no diversion into the California Aqueduct. That would have been almost as large as the total 1.530 million acre-feet inflow that occurred in the 1906 flood.

The 1983 flood brought the lake to a peak elevation of 191.44 feet, slightly lower than the modern 192.5 foot record set in 1969. In order to protect Corcoran, the USACE spent $2.7 million to construct emergency flood protection levees along Cross Creek and the Tule River. Unfortunately those levees were not strong enough and were breached. Tulare Lakebed inundation began in January and peaked in July. By July 13, 82,000 acres of prime agricultural land were flooded. Based on a comparison of maps, the area flooded in 1983 was slightly greater than the area flooded in 1969.

Mo Basham recalled that the eastern edge of Tulare Lake came to about Avenue 10½. This is about 3 miles west of where the emergency levee was built near the Corcoran Airport during the 1969 flood. Bill Tweed recalled that the lake was so big in the summer of 1983 that you could see it from the High Sierra, shining through the valley haze. To see it was like seeing a ghost, a relic of another time (photograph on file in the national parks, see back cover).

In order to minimize the flooding in the lakebed, an unknown amount of river floodwater was pumped into the Friant-Kern Canal in 1983 and routed to the Los Angeles area.
Table 84 summarizes the damages incurred in the southern end of the San Joaquin Valley during the 1983 floods.\footnote{1521}

<table>
<thead>
<tr>
<th>County</th>
<th>Private Damage (thousand dollars)</th>
<th>Public Damage (thousand dollars)</th>
<th>Road Damage (thousand dollars)</th>
<th>Agricultural Damage (thousand dollars)</th>
<th>Total Damage (thousand dollars)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Merced</td>
<td>$200</td>
<td>$414</td>
<td>$0</td>
<td>$40,000</td>
<td>$40,300</td>
</tr>
<tr>
<td>Madera</td>
<td>$200</td>
<td>$0</td>
<td>$100</td>
<td>$5,648</td>
<td>$13,424</td>
</tr>
<tr>
<td>Fresno</td>
<td>$100</td>
<td>$7,060</td>
<td>$616</td>
<td>$95,000</td>
<td>$97,968</td>
</tr>
<tr>
<td>Kings</td>
<td>$420</td>
<td>$1,998</td>
<td>$550</td>
<td>$97,000</td>
<td>$171,898</td>
</tr>
<tr>
<td>Tulare</td>
<td>$100</td>
<td>$844</td>
<td>$37</td>
<td>$24,731</td>
<td>$24,731</td>
</tr>
<tr>
<td>Kern</td>
<td>$2,750</td>
<td>$356</td>
<td>$1,328</td>
<td>$7,500</td>
<td>$11,934</td>
</tr>
<tr>
<td>Total</td>
<td>$3,770</td>
<td>$10,672</td>
<td>$2,631</td>
<td>$188,971</td>
<td></td>
</tr>
</tbody>
</table>

In 1983, Bill Cooper and John A. Sweetser, Sr. kayaked from the banks of the Kern River just outside of downtown Bakersfield all the way to Richmond Marina on the shores of San Francisco Bay. This was the sixth documented trip between Tulare Lake and San Francisco Bay to occur in historic times. (The other five trips were in 1852, 1868, 1938, 1966, and 1969.)

In a 2014 interview, Bill recalled how he and John did it. Bill was not an experienced kayaker; he had never been in a kayak before.\footnote{1522} They first scouted to Tulare Lake which was the hardest section, and decided that they could make it. Then they threw sleeping bags and a couple tents in their kayaks and put in.

They didn’t expect their trip to be a big deal, but it made the national news. Radio and TV news reported on their progress. They were followed by an airplane for part of their route. Dave Graber, retired NPS regional chief scientist, recalled that their trip was written up in *The Fresno Bee*. Their goal was just to go across Tulare Lake. However, it caught the attention of the news media, so they decided to try to make it all the way to San Francisco Bay.

Bill and John thought they could make it to the south end of Tulare Lake in one day, but they barely made it to Buttonwillow in the first day. Ken Wedel landed his plane and took them up to scout their route. He then returned and re-provisioned them by plane that first day, giving them an air drop of water jugs. After that, they resupplied at farm houses and stores along the way at places like Firebaugh. It took them a full day to cross Tulare Lake, staying close to the western levee. It took them 12 solid days of work, sunrise to sunset, to make it to the Richmond Marina.

Bill said that he and John are apparently the last two people to make it through to San Francisco Bay. He heard that some guys from Reedley tried to get through in canoes in some year after 1983, but didn’t make it; the wind gave them a hard time in the Fresno Slough.

After two years of flooding (1982 and 1983), cotton growers decided to drain their lands. The Tulare Lake Irrigation District applied for a permit to pump the excess water over the top of the Tulare Lake sill. It appears that there was considerable opposition to granting this permit. Under an emergency proclamation issued by the USACE during the spring of 1983, reclamation districts and land companies remade the channel along some 29 miles of the lower Kings River (see Figure 17) to dewater the lake and drain the water north into the Sacramento–San Joaquin Delta region.

A series of pumps were installed with a total lift of 43 feet. The project was designed to remove approximately 2,000 acre-feet of water per day from the lakebed. Pumping began on October 7, 1983 and continued intermittently until the program was terminated on January 19, 1984. Only about 90,000 acre-feet was pumped northward under this program. Pumping was stopped earlier than scheduled due to concern that white bass might be transferred from Tulare Lake to the San Joaquin River. The lakebed would not be fully drained until water year 1985.

An outstanding feature of the 1982–83 storm event was the number of significant landslides and debris flows that resulted. Some of those were in the Central Sierra, north of the Tulare Lake Basin. Those events merit inclusion in this document because they were well studied, and they can inform risk management planning in our area.
Landslide: South Fork American River
This event occurred near White Hall on U.S. Highway 50 about 26 miles east of Placerville and 34 miles west of South Lake Tahoe. At this point, the highway is squeezed in a narrow canyon between the South Fork American River and a steep cliff.

This landslide occurred in the El Dorado National Forest. It is described in several secondary sources.\(^{1523, 1524}\)

The soil was derived from weathered granitic material. It had lots of voids that could hold water. In addition, the landslide included rock with some boulders that measured more than 16 feet (5 meters) in diameter.

The highway and the river had undercut the base of the slope, reducing the overall stability of the hillside. The long period of heavy precipitation during the 1983 water year had raised groundwater levels and increased pore-water pressures within the hillside. That, in combination with the removal of the base of the hill, acted together to trigger the landslide.

At 5:10 a.m. on April 9, 1983, a large section of the hillside gave way. The landslide moved rapidly downhill, across the highway, and dammed the river.

Maximum depth of the lake was 50 feet (15 meters). The river began breaching the landslide dam at 11:30 a.m. the next morning. There were enough large boulders in the dam to prevent rapid breaching and downstream flooding. During the following months, the river gradually eroded the dam down to the original riverbed. By June 1983, the area of the lake had decreased to roughly one-third of its original size.

The landslide had an estimated total mass of about 1,000,000 cubic yards (765,000 cubic meters). It took Caltrans 75 days to reopen the highway.

The Mill Creek Landslide would happen just 0.6 miles west of this location on January 24, 1997.

Landslide: Slide Mountain, Nevada
This event occurred just northeast of Lake Tahoe in the Toiyabe National Forest. It is described in a secondary source.\(^{1525}\)

The winter of 1982–83 was unusually wet and built a record snowpack. A sudden sustained warm period beginning in late May greatly reduced the snowpack and promoted infiltration of water into the subsurface. That increase of moisture content increased local pore pressure in discontinuities and in the unconsolidated surficial deposits covering the bedrock.

At about noon on May 30, 1983, a large section of the hillside gave way. Several types of mass wasting processes were involved, including a rock slump, a rockfall avalanche, and a debris avalanche. The rock slump composed the largest part of the slide and was up to 100 feet (30 meters) thick.

Along the northeastern margin of the landslide, a rapidly moving rockfall avalanche of large boulders and a debris avalanche of gravelly sand entered Upper Price Lake, displacing most of the water in the lake, which breached a low dam. The water then breached the dam of Lower Price Lake and sent a torrent down the gorge of Ophir Creek. This created a debris flow which damaged and destroyed homes, overtopped old U.S. Highway 395, and caused one death.

Total volume of the slide was estimated to be up to 940,000 cubic yards (720,000 cubic meters).

Debris Flow: Camp Creek
The April 10–11, 1982 rain-on-snow storm event was responsible for triggering numerous landslides and debris flows in the Sierra. One of those was a debris flow in Camp Creek, a tributary of the San Joaquin River. That event was analyzed by Jerry DeGraff, a geologist for the USFS.\(^{1526}\) The debris flow began on a 50% slope in a soil composed largely of fine gravel and sand. It flowed about 1½ miles to Mendota Pool reservoir. The Camp Creek debris flow had one of the fastest peak velocities ever recorded for a Sierra debris flow: 16 mph (26 km/hr).

Debris Flow: Garnet Dike
The Garnet Dike debris flow also happened in 1982.\(^{1527}\) It occurred on a tributary of the South Fork Kings River in the Sierra National Forest. It was several miles from Big Creek. That is part of the Kings River Special Management Area. The event was analyzed by Jerry DeGraff.\(^{1528}\)
It was a relatively slow debris flow for the Sierra, only 11 mph (18 km/hr). Despite being a big debris flow, it was able to pass around trees without doing much damage or abrasion to them. It did form a boulder levee on its flanks.

**Debris Flow: Calvin Crest**

On July 5, 1983, a debris flow occurred on the Sierra National Forest adjacent to the Calvin Crest Conference Center near Oakhurst, California. The event was thoroughly analyzed by Jerry DeGraff.1529

The debris flow originated at an elevation of 2,500 feet (762 m). It was located on a 30% slope at the broad head of a small drainage basin. The slope had an open stand of mixed oak and Jeffrey pine with an understory of herbaceous vegetation. Where the 30% slope flattened to 10%, there was evidence of seasonal groundwater seepage at a number of points in the vicinity of the debris flow. The area where the debris flow occurred was underlain by granitic bedrock. A few thousand feet upslope was the contact with remnant meta-sedimentary bedrock capping the hill top. Near this contact, a number of seeps and wet meadows were present.

By July 7, the debris flow was about 600 feet (200 m) long and 72 feet (22 m) wide. The deposit had the consistency of very wet cement. Water discharged from the end of the deposit. Groundwater flowed from the upper scarp and other points along the debris flow track. This made the bottom of the track too soft and muddy to be examined more closely. It remained saturated until the following winter. In succeeding years, grass grew over the debris flow scar and flow path. Despite revegetation, it remains fairly wet to damp throughout most of the year.

Usually debris flows are triggered either by a storm event or by melting of a snowpack. However, that wasn’t the case with Calvin Crest; it was apparently triggered by groundwater conditions resulting from above-average recharge. No precipitation was received during the previous 24 days at the South Entrance of Yosemite National Park, the nearest station to the debris flow. Precipitation totaled only 0.6 inches (15.5 mm) for the 35 days prior to the debris flow. The pre-movement observation of groundwater flow from a depression at the base of the 30% slope suggested high pore-water pressures were present in the slope materials.

**Debris Flow Complex: Kings Canyon National Park**

During the 1983 Memorial Day Weekend, Kings Canyon experienced three debris flows that we know of:

1. Bubbs Creek
2. Unnamed tributary of Lewis Creek
3. Castle Dome Meadow

These debris flows could be thought of as one event that occurred in multiple locations. The event was triggered by a heavy snowpack and an extreme change in temperatures in a 48-hour period.

We know about these events primarily because of the outstanding memory of Jerry Torres, Kings Canyon National Park’s trails supervisor at the time.

The Bubbs Creek debris flow began in two unnamed tributaries of Bubbs Creek, high on the side of Glacier Monument (UTM 365530E 4072600N NAD83 Zone 11, elevation 9,600). From there it flowed south down the steep hillside 1 mile to Bubbs Creek, elevation 6,700. That point was approximately 1½ miles (2½ km) east of the Sphinx Creek Bridge. Aerial photography shows significant scouring all along the flow path. At that point, the debris flow had dropped 2,900 feet in elevation.

That hillside had been burned seven years earlier in the 1976 Sphinx Fire. However, the primary triggering event was probably the large amount of moisture infiltrating into loose soils from the melting snowpack.

Once the debris flow got to the bottom of the hill, it turned and followed the course of Bubbs Creek east. It scoured Bubbs Creek for approximately the first ¾ mile (0.5 km), taking vegetation including large trees, rock, tons of soil, and a good swath of the trail downstream. That section of Bubbs Creek was not any steeper than the sections farther downstream. However, the debris flow had just come off a very steep hill, so possibly that was the cause of the scouring.

The debris flow created a large earthen dam on its eastern flank. This restricted the downstream flow of Bubbs Creek, creating a very large pool which lasted several years. Today a medium-size pool or slow-water area still exists there. Although the Sphinx Creek Bridge was unscathed by the event, it changed the Bubbs Creek bridge channels, drying two of the channels and increasing the flow in one twofold. There were some large tree jams at
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and above these bridges as well as areas of the forest above the fourth and third bridges that were toppled by the event.

The Bailey Bridge crossing on the South Fork Kings (elevation 5,000 feet) two miles east of Roads End marked the western extent of the effects from the Bubbs Creek debris flow. The debris flow had a total length of about 3.6 miles (5.8 km) with a drop of 4,600 feet. The average gradient was about 1,300 feet/mile, but that is misleading. The gradient on the mountain section (2,900 feet in 1 mile) was 2,900 feet/mile while the gradient on the Bubbs Creek section (1,700 feet in 2.6 miles) was only about 650 feet/mile.

The second debris flow occurred in the Lewis Creek Basin. The first tributary of Lewis Creek above its confluence with the South Fork Kings experienced a high-energy debris flow. This debris flow began on a steep, sparsely vegetated slope in a remote trailless area of the park (UTM 352609E 4076728N NAD83 Zone 11). This tributary (now informally referred to as Tsunami Creek) had a massive wall of water and debris come barreling down its channel late on Friday morning (May 27, 1983) prior to the Memorial Day Weekend. The wall scoured the channel, depositing mud fifty feet up the trunks of those trees that survived the onslaught of the “tsunami.” Bill Tweed recalled that huge logs came crashing down in the debris flow, and that there was a lot of silt in the debris. The presence of large logs in the debris suggests that this event has a long recurrence interval.

It has been speculated that the severity of this debris flow might have been attributed in part to the 1980 Lewis Creek Fire. However, this appears to have been just the coincidence of association. The other debris flows that occurred during this storm event were not associated with previous fires. Moreover, there has never been a debris flow in the national parks which was clearly linked to the effects of fire.

There is a more reasonable explanation for what probably triggered this debris flow. The unusually wet spring presumably soaked the ground to depth. The extreme change in temperatures in a 48-hour period then melted much of the heavy snowpack, promoting further infiltration of the soil. That raised groundwater levels and increased pore-water pressures within the hillside. That reduced the soil’s frictional strength, causing the soil mass to begin moving downslope as a flowing mass. That is what triggered the landslide and debris flow at Slide Mountain, Nevada on May 30, 1983.

The Lewis Creek Basin has several areas of loose granitic sand on steep slopes that are sparsely vegetated. A similar high-energy debris flow event would occur in this drainage (but in the main Lewis Creek channel) in the July 15, 2008 flood.

The national parks used to have the main pump house for the Kings Canyon development near Lewis Creek. Maintenance worker Ron Cook was checking the water intake on the bank of the creek on the morning of May 27, 1983, when the debris flood struck. It came upon him so fast that it caught him by surprise, knocking him off his feet. As Ron told the story, he was almost caught up in the maelstrom and killed. He wrapped one arm around a small tree to hold on while he radioed for help with the other.

The debris flood nearly took out the Lewis Creek pump house. (That pump house has since been removed and the Kings Canyon development now gets all of its water from the facility at Sheep Creek.)

The debris flow closed Highway 180 below the Lewis Creek Bridge for most of the day. A bulldozer was used to push Lewis Creek back into its former channel. Remnants of the flood channel are still visible today, just west of the Lewis Creek Bridge on the north side of the road. Once the creek was pushed back, crews removed rocks and several feet of mud and other debris from the roadway.

The third Kings Canyon debris flow of the Memorial Day Weekend occurred at Castle Dome Meadow. It covered a 200-yard section of trail and meadow with decomposed granite and sand.

Debris Flow: Redwood Creek
René Ardesch discovered this debris flow in July 1983 when he was on a cross-country backpacking trip. He later recalled the discovery and the impression that it made on his group:

Back in the late spring of 1983 four of us decided to go on a backpack trip to the Castle Rocks in Sequoia National Park. As we were aspiring wanderers we opted to go on a cross-country course with the destination of Pine Top Mountain for the first night. We started our adventure at the road up to the old Camp Conifer below Atwell Mill, which is where we parked and assembled our gear. One of our
group’s families had a cabin in Camp Conifer long ago and we wanted to see how it looked since all the cabins were removed.

We walked with our heavy packs through the grove and checked out some acorn grinding holes onsite. We then walked on into the forest towards our bivy spot for the night. After sometime of up and down hiking we noticed in the distance an odd scene that looked like a large opening in the heavy forest we were in. As we finally came to the edge we all stood together with a sense of awe, just blown away with what we were looking at. A huge path of destruction up to 100 feet wide with whole trees reduced to logs in big piles gathered along the edges and rocks of all sizes everywhere up and down what appeared to be a creekbed (photograph on file in the national parks). We assumed it to be Redwood Creek as we were headed in a westerly direction and the flow was southerly. The land was fairly level here, and we were at an elevation of around 7,000 feet.

We put our packs down and tried to gather our thoughts about what it all meant and how it happened. Still in the giant sequoias (in the Redwood Creek Grove), the smell in the air was of fresh, moist soil and the ground was damp all around. In one sandy area we moved to later we saw large cat tracks that we compared to our own foot size. We had to circumvent this area for quite a ways as it was in our direct path and the whole time we were talking about what might have caused this catastrophic event. This one sighting was in our minds for many years to come.

Redwood Creek is a tributary of the East Fork Kaweah. The area where they encountered the debris flow was relatively flat, so René inferred that they were near the bottom of the run, and that it had started far above them. Circumstantial evidence suggested that the debris had occurred within the previous couple months, perhaps over the Memorial Day Weekend when the other three debris flows occurred in the Kings Canyon area.

1984 Floods (5)

There were four periods of flooding in 1984:

1. July (twice)
2. August
3. September
4. Lakebed flooding

On July 15-16, a high-intensity, short-duration thunderstorm produced flood conditions in the Goat Ranch Canyon and Long Canyon areas. This storm followed the 26,000 acre lightning-caused Bodfish Fire that began on July 7. Debris flows and debris blocked Highway 178 and many other roads. Uffert Park was covered by debris flows that were about 6 inches deep. Three houses in the Long Canyon area became completely uninhabitable when debris flows inundated them. A small levee in this location was breached and eliminated by the flood. Debris flows threatened homes in the Bodfish Creek area.\footnote{1531}

On July 30, an intense thunderstorm occurred in Scodie Canyon, causing flooding in the community of Onyx. The floodwaters overflowed channels and eroded new channels. Thirty mobile homes were washed away, and nine of these were completely destroyed. Stranded residents had to be airlifted out. Damage was estimated to be $3 million. One man was killed by lightning. A state disaster was declared for Kern County on July 31.\footnote{1532, 1533}

An intense storm occurred in the hills east of Lake Isabella on August 20. Two-thirds of an inch of rain fell in just 40 minutes. Scodie Creek (Sometimes incorrectly listed as Sodie Creek) overflowed its banks, flooding the community of Onyx. Four homes were damaged by mud and one home was washed away.\footnote{1534} This is a different event from the July 30 flood.

Gary Sanger at the NWS forecast office in Hanford researched the two storms in Scodie Canyon. There was abundant monsoonal moisture in Southern California during July 27-30, as reflected by reports of scattered rainfall. However the thunderstorm that occurred in Scodie Canyon on July 30 was considerably more intense than any other storm that was reported during the July 27-30 period. It was a relatively isolated event.

There was also some monsoonal moisture in Southern California during August 19-20, although less than during the July 27-30 period. The storm that occurred in Scodie Canyon on August 20 was considerably more intense than any other storm that was reported during the August 19-20 period. Once again, it was a relatively isolated event. It seems to have been just a bizarre coincidence that Scodie Canyon got hit with two back-to-back huge thunderstorms in this three-week period. You have to wonder what terms area residents used to describe these two events that had been visited upon them.
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The July 30 event was reported to be a thunderstorm; we don’t know about the August 20 storm event. Gary observed that storms such as this don’t necessarily have to be thunderstorms. A nearly stationary storm could also produce very heavy rainfall over a small area, resulting in flash flooding. Conversely, a fast moving, intense storm might spread rain over a much larger area, reducing the impact of the flash flooding.

An intense storm occurred in Lake Isabella on September 19. Over an inch of rain fell in just 45 minutes, washing out ¼ mile of one road, covering others with mud, and destroying two mobile homes.\(^{1535}\)

The Tulare Lakebed flooded in 1984.

This lakebed flooding that occurred in 1984 was completely unrelated to the floods that occurred that year; none of those flood events contributed significantly to the flooding that occurred in the lakebed. Nor did any of the rivers contribute runoff to the lakebed in 1984.

The Tulare Lakebed flooded in the spring of 1982. The floods of 1983 greatly expanded the size of the lakebed flooding. As illustrated in Figure 16, the lakebed was still extensively flooded at the beginning of 1984. The lakebed would not be fully drained until water year 1985.

Lakebed flooding is a social construct; it is counted based on the number of growing seasons that are missed. The lakebed was flooded for three growing seasons: 1982, 1983, and 1984. Therefore, this is counted as three floods from the perspective of the lakebed farmers, even though flood events occurred in only two of those years. Something similar happened in the lakebed in 1969–1971 and 1997–99. In each of those cases, lakebed flooding continued into a non-flood year.

1986 Floods (4)

There were three periods of flooding in 1986:
1. February
2. March (twice)
3. April–July snowmelt period

The first storm event lasted from February 11–24. The actual transport mechanism was an atmospheric river that brought phenomenal amounts of precipitation to a large portion of Northern and Central California and western Nevada.\(^{1536}\) *Rivers of Fear: The Great California Flood of 1986* is supposed to be the most comprehensive source of data available for this flood.

A series of four tropical storms pounded the state between February 11–20. Rains from the first three storms saturated the ground and produced moderate to heavy runoff before the arrival of the fourth storm. The heaviest precipitation from those storms was in a band 200 miles north to 100 miles south of a line from San Francisco to Sacramento to Lake Tahoe.\(^{1537}\)

A total of 200 stations reported their highest-ever rainfalls for 10 consecutive days. Half of the average annual rain fell in the 10 days between February 11–20 at 150 stations in the state. Mono Lake had 95% of its annual average rainfall occur during those 10 days. Bucks Lake in the Feather River Basin had 49.44 inches, which was 71% of its average annual rainfall.

Rains from the first storm started the evening of February 11 and peaked the next day. This storm originated in the Pacific just north of Hawaii and brought up to 6 inches of precipitation to the upper Feather River Basin. On February 13, a second storm developed northeast of Hawaii. A strong cold front generated by this storm moved across Northern California on February 14. Gusty winds and heavy rains hit the entire state. Behind this front, a pattern of overrunning (warm moist air flowing over cold air) produced additional rainfall through much of the following day.\(^{1538}\)

On February 15, a strong, deep flow of warm moist air from Hawaii advanced south of California. On February 16, weather satellites showed enormous development along the jet stream between Hawaii and California. Southwest winds of 210 mph were reported in the jet stream. This storm (the third storm), which entered south of California, began moving slowly north as a warm front. North of the warm front, strong overrunning by a deep moist southwest flow began producing heavy rainfall from the North Bay counties to the Sierra. In many areas, this heavy rainfall continued with only brief breaks through February 17. Rainfall of ½ to ¾ inch per hour was common.\(^{1539}\)
Another Pacific weather system (the fourth storm) approached Northern and Central California on February 18. This storm originated well north of Hawaii, and thus was a much colder front in comparison to the previous three storms. The snow level dropped to 5,000 feet for this storm; during the previous storms, the level was about 7,000 feet.\textsuperscript{1540}

In the Sierra, the storms affected mainly the area from the Feather River Basin in the north to Yosemite on the south. The Sierra stations that received rainfalls in excess of the 1,000-year recurrence interval ranged in a band from Clarks Peak north of Sierra Valley in the Feather River Basin to Calaveras Big Trees in the Cosumnes River Basin in the south.

The heaviest 24-hour rainfall ever recorded in the Central Valley, 17.6 inches, occurred on February 17 at Four Trees in the Feather River Basin, 30 some miles north of Oroville. This broke (just barely) the old record that had been set at Hockett Meadow on December 6, 1966. Four Trees received a total of 56 inches of rain for the month, the greatest February total recorded for any station in the state during 1986.\textsuperscript{1541}

Calistoga, in the Napa River Basin, had 29.61 inches in 10 days. This represented a recurrence interval of 2,600 years.\textsuperscript{1542}

Due to the storms’ tropical nature, snow levels fluctuated between 7,000 and 8,000 feet. Between February 11–20, more than 34 inches of rain fell at Blue Canyon on the American River east of Grass Valley. Above 8,000 feet, storm-total estimates ranged from 15–20 feet of new snow with 20–30 inches of water content.

The widespread drenching rains led to extensive flooding and mudslides. The floodwaters destroyed many bridges and punched through several levees. This was the flood that caused the big levee failure on the Yuba River at Linda, south of Marysville.

Statewide, more than 50,000 people fled their homes, and 13,000 homes and businesses were either damaged or destroyed. Damage was estimated to be $500 million, 13 flood-related deaths occurred, and 96 were injured.

Flooding was widespread with 23\% of streamflow gaging stations in California reporting significant discharges. Flooding was most severe in the northern half of the state. It had a recurrence interval of 100 years on some rivers.

Over much of the area, the precipitation ranged from 100 to 200\% of normal February precipitation for the 9-day period from February 11–19. In many rivers and streams, those storms produced either record or near-record flows. At 16 stream gages, the peak flow recorded either equaled or exceeded the previous maximum. A record flow of 640,000 cfs was estimated at the latitude of Sacramento.\textsuperscript{1543}

The 1986 flood was a record flood on the American River, the fourth record flood in 36 years. The American River dumped more water into Folsom Lake than it was designed to handle. After two days of releases at the maximum design release level of 115,000 cfs, officials were forced to boost releases to 134,000 cfs. Peak discharge-of-record occurred in the Napa River and upper Feather River Basins. Inflow of the Feather River into Lake Oroville reached a high of 266,540 cfs. Record flood management releases of 150,000 cfs made room for this unexpected volume of water.

There was extensive flooding in Plumas County during the January 1986 flood. Two years later, several previously buried beaver dams were discovered on the incised channel of Red Clover Creek in eastern Plumas County, about 60 miles north of Truckee, California and east of the Sierra Crest. This creek is tributary to Indian Creek (via Last Chance Creek), part of the East Branch North Fork Feather River. Presumably these beaver dams were exposed during the January 1986 flood. See the section of this document that describes \textit{Wildlife in and around Tulare Lake} for a discussion of these dams.\textsuperscript{1544}

The Napa River crested near Napa with a peak discharge of 37,100 cfs; it had a recurrence interval of 75–100 years.

The bypasses for the Sacramento River Basin provided much needed storage and flow capacity during the peak of the flood. Before the mid-February storm systems, overflow at each of the weirs had been minor or nonexistent. By February 17, however, all weirs were flowing and all but one of the weirs continued flowing until the last week of March. The peak flow exceeded the project design flow at three of the weirs. System breaks in the Sacramento River Basin included two disastrous levee breaks on the Feather River. Levee breaks along the...
Mokelumne River caused flooding in the community of Thornton and the inundation of four Delta islands. Damages exceeded $172 million and $15 million in the Sacramento River and San Joaquin River Basins respectively.

Much of the San Joaquin River Basin was spared the full impact of the 1986 storms. The major projects for the San Joaquin River Basin did not encroach on their flood-control pool as did their counterparts in the Sacramento River Basin. The exception in the San Joaquin River Basin was Millerton Lake where only 16% of the flood-control pool remained at the end of the February event. \footnote{1545}

A major frontal storm system crossed the Central Sierra in mid-February 1986. The southern edge of this storm triggered three debris flows on the north-facing slopes of Shingle Hill, near Greeley Hill, California, within the Merced River Basin. These three debris flows are described at the end of this section along with two other debris flows that occurred in 1986.

The Kings River experienced a flood from February 13–19. The peak day natural flow at Pine Flat on the Kings occurred on February 18, 1986: 25,060 cfs. This was a slightly bigger flow than occurred in the much more famous 1983 flood. It seems likely that this was a very high-flow period in Cedar Grove as well.

Damage was much greater in Fresno County than in Tulare.

Jerry Torres and David Karplus (Kings Canyon National Park trails supervisors) recalled that there were a large number of avalanches throughout both Sequoia and Kings Canyon National Parks that spring, causing considerable damage. One of those avalanches pushed the Palisade Creek Bridge off its footings. One of the largest and best known avalanches was the Paradise Valley (in Kings Canyon National Park) “logalanche” which scoured a large swath of Middle Paradise Valley below the Kidd Creek headwaters. In addition to the avalanches, the floods created several large log jams along the Kings River as well as other water courses in the national parks.

The Kaweah’s peak natural flow occurred at Terminus Dam (or possibly McKay’s Point) on February 13: 9,852 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 9,428 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 7 years for the Kaweah.

Table 85 summarizes the damage incurred in the southern end of the San Joaquin Valley during the 1986 flood.\footnote{1546}

<table>
<thead>
<tr>
<th>County</th>
<th>Private Damage (thousand dollars)</th>
<th>Public Damage (thousand dollars)</th>
<th>Total Damage (thousand dollars)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Merced</td>
<td>$70</td>
<td>$70</td>
<td>$70</td>
</tr>
<tr>
<td>Madera</td>
<td>$210</td>
<td>$38</td>
<td>$248</td>
</tr>
<tr>
<td>Fresno</td>
<td>$840</td>
<td>$450</td>
<td>$1,290</td>
</tr>
<tr>
<td>Kings</td>
<td>0</td>
<td>0</td>
<td>0</td>
</tr>
<tr>
<td>Tulare</td>
<td>$20</td>
<td>$20</td>
<td>$48</td>
</tr>
<tr>
<td>Kern</td>
<td></td>
<td></td>
<td>0</td>
</tr>
<tr>
<td>Total</td>
<td>$1,140</td>
<td>$488</td>
<td>$1,628</td>
</tr>
</tbody>
</table>

An F0 tornado touched down in Kingsburg on March 7. That tornado was spawned by a strong thunderstorm complex that produced heavy rain over the Southern Sierra, causing flash flooding in Mariposa and Madera counties.\footnote{1547}

On March 10, an intense thunderstorm struck Fresno during the height of the evening commute. About an inch of rain fell in downtown Fresno resulting in widespread flooding, stranding dozens of cars, some with water up to the rooftops. The deluge flooded basements in a number of buildings in downtown Fresno and caused part of the roof to collapse on a store. Hailstones as large as mothballs fell in nearby farm areas and accumulated up to 4 inches deep. In parts of Biola (west of Fresno), up to 3 inches of hail was still on the ground at noon the next day. Locally heavy rain fell farther south, causing the White River to surge over its banks and flood Highway 98 between Earlimart and Delano.\footnote{1548}
Flood releases of over 15,000 cfs occurred at Friant Dam on the San Joaquin River during the April–July snowmelt period. This was the second biggest release since the dam was completed in 1942.

Pine Flat Dam on the Kings River recorded the largest 30-day flood-of-record during late May and early June of 1986. Spring snowmelt was heavy enough to cause flooding in the Tulare Lakebed. In order to minimize flooding in the lakebed, 94,000 acre-feet of river floodwater was pumped into the Friant-Kern Canal and routed to the Los Angeles area. For comparison, that is half as much water as the total capacity of the newly expanded Lake Kaweah.

Total flow for water year 1986 was 192% of the 1894–2014 average for the Kings, 192% for the Kaweah, 180% for the Tule, and 200% for the Kern.

Debris Flow: Shingle Hill
A major frontal storm system crossed the Central Sierra in mid-February 1986. The southern edge of this storm triggered three debris flows on the north-facing slopes of Shingle Hill, near Greeley Hill, California, within the Merced River Basin. The event was analyzed by Jerry DeGraff, a geologist for the USFS.

The storm precipitation fell in the form of rain. No snow accumulation was present on Shingle Hill which ranges from about 2,099 feet to over 3,100 feet. A rural county road along the base of the hill was open the evening of February 17. On the morning of February 18, deposits were blocking the road at two locations. A third deposit was found in an ephemeral channel a few tens of feet up-gradient from the road. The debris flows occurred the night of February 17 or early on the morning of February 18, a time which coincided with a high intensity rainfall period. The slopes ranged from 50%–65%.

Two other significant debris flows occurred in 1986:
- Wolfin debris flow (Tuolumne River Basin). This debris flow began on a 60% slope adjacent to an existing intermittent channel. The mass entered perpendicular to the direction of the channel and immediately began moving down-channel in a clear indication that remolding into a flowing mass took place. It had one of the fastest peak velocities ever recorded for a Sierra debris flow: 15.5 mph (25 km/hr).
- Minarets Highway debris flow (San Joaquin River Basin). It had one of the slowest peak velocities ever recorded for a Sierra debris flow: 6 mph (9 km/hr).

1987–92 Drought
This drought began over most of California in 1987. However, parts of the state were in drought from 1984–93. This was the state’s first extended dry period since the 1920s–30s.

The recurrence interval of this drought in the Sacramento River Basin was approximately 70 years based on the 1906–92 record. On the San Joaquin River, where the drought was more severe, the recurrence interval was approximately 300 years. These statistics reflect both the six-year length and the severity of the drought.

A significant portion of the country experienced drought conditions during the general period that California was in drought: 1987–92. By June 1988, 54% of the contiguous U.S. was in drought condition. The impact was worst in the northern Great Plains, though the West Coast and Northwest were also hit. Particularly memorable were the forest fires that accompanied the drought. In 1988, 793,880 acres of Yellowstone National Park burned, prompting the first complete closure of that park in history.

The drought of 1988 became the worst drought in the U.S. since the Dust Bowl 50 years earlier. Not until 2012 would the U.S. see a drought this extensive. The drought of 1988 remains the costliest U.S. natural disaster ever. Hurricane Katrina ranks second and Hurricane Andrew third.
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Table 86 compares the 1987–92 drought with the other severe droughts of the 20th century. Both the 1929–34 and the 1976–77 droughts had a recurrence interval of more than 100 years, at least by some measures.

Table 86. Comparison of selected 20th-century droughts.

<table>
<thead>
<tr>
<th>Drought Period</th>
<th>Sacramento River Basin Runoff</th>
<th>San Joaquin River Basin Runoff</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Average yearly runoff</td>
<td>Average yearly runoff</td>
</tr>
<tr>
<td></td>
<td>(million acre-feet)</td>
<td>(million acre-feet)</td>
</tr>
<tr>
<td></td>
<td>% of average 1901–2009</td>
<td>% of average 1901–2009</td>
</tr>
<tr>
<td>1929–34*</td>
<td>9.8</td>
<td>3.3</td>
</tr>
<tr>
<td></td>
<td>56%</td>
<td>56%</td>
</tr>
<tr>
<td>1976–77</td>
<td>6.6</td>
<td>1.5</td>
</tr>
<tr>
<td></td>
<td>38%</td>
<td>26%</td>
</tr>
<tr>
<td>1987–92</td>
<td>10.0</td>
<td>2.8</td>
</tr>
<tr>
<td></td>
<td>57%</td>
<td>48%</td>
</tr>
</tbody>
</table>

*This was one component of the larger 1918–34 drought.

DWR provided detailed information about the 1987–92 drought in the following reports:

The 1987–92 drought was notable for its six-year duration. As shown in Table 20 and Table 22, most of the droughts in the Tulare Lake and San Joaquin River Basins last 2–4 years. (A single dry year isn’t generally considered a drought.) As shown in those two tables, we are only aware of 10 droughts in the last 11 centuries that have lasted 6 or more years.

The 1987–92 drought was also notable for the statewide nature of its impacts. In 1991, the single driest year of the drought, the State Water Project terminated deliveries to agricultural contractors and provided only 30% of requested urban deliveries. The federal Central Valley Project provided 25–50% supplies to urban contractors and 25% to agricultural contractors.

At that time (1991), 23 of the state’s 58 counties had declared local drought-related emergencies. Many of the declarations were prompted by economic impacts associated with loss of dryland cattle range, damage to timber resources and associated wildfire damage, and diminution of water-based recreational and tourism activities, rather than by shortages of dedicated and developed water supplies.

Table 87 shows how the San Joaquin Valley Water Year Index categorized the drought years in that basin. It also shows total runoff for the four major rivers in our basin (Kings, Kaweah, Tule, and Kern) combined during this drought.

Table 87. Rating of drought severity during the 1987–92 drought.

This drought was notable in part because all six years were dry, with four of them ranking in the top 10% in terms of driest statewide runoff. Water year 1991 was the driest year at the state level, ranking in fifth place in the statewide runoff record, behind 1977 (driest), 1931, 1924 and 2014.

The Kings River Handbook says that the 1987–92 drought was the worst extended sequential critical drought during recorded history in the Kings River Basin. Perhaps that is so. However, based on tree-ring analysis, the years 1926–31 were the driest six-year period on the upper San Joaquin River at the input to Millerton Lake during the 1113-year period 900–2012.

Bob Meadows was the wilderness ranger at Ranger Lake in 1989 and 1990 and recalled the effect of the 1987–92 drought on that lake. Ranger Lake has an extremely small watershed and almost no regular inflow. However, during the two years that Bob was there, the lake dropped no more than about a foot or so.
Total flows for water years 1987–92 ranged from 22%–64% of the 1894–2014 average for each of the four rivers within the Tulare Lake Basin.

The CVP and SWP met delivery requests during the first four years of the drought, but were then forced by declining reservoir storage to cut back deliveries substantially. In 1991, DWR dropped the SWP allocation to zero for agricultural users and 30% for urban users. That represented the lowest percentage of requested deliveries in the history of that project. That record would eventually be broken in 2014 (see Table 10 on page 108). The CVP delivered 25% to agricultural contractors and 25–50% to urban contractors (see Table 11).

As a result of the 1987–92 drought, the Tulare Lakebed was largely dry from 1987 through 1994. This is the longest period that the lakebed has remained dry since the drought years of 1918–34. During that drought, the lakebed was largely dry from 1924 until February 7, 1937.

Widespread damage to timber resources was reported throughout the Sierra due to bark beetle infestation. The drought’s prolonged duration set the stage for a pattern that would emerge in future extended dry periods — the linkage between severe drought conditions and risk of major wildfire damage in densely populated urban areas located at the wildland-urban interface. The October 1991 Oakland Hills fire was the then-largest dollar fire loss event in U.S. history; 25 lives were lost and more than 3,000 structures were destroyed.1561

In 1991, the “March Miracle” brought abundant snow to the middle and upper elevations of the Sierra. Lodgepole received a total of 147 inches of snow, the third greatest monthly snowfall at that location since record-keeping began.

On February 15, 1992, a strong winter storm lowered the snow level to about 2,000 feet in the Tehachapi Mountains. That was one in a series of storms that dumped a total of 17.32 inches of precipitation on Frazier Park during the month. The cause was an inflow of subtropical moisture that moved over the mountains from the south.1562

This brought relief to portions of Kern County. However, the drought was generally considered to continue through most of calendar year 1992. A series of major Pacific storms brought abundant moisture to the state between December 1992 and February 1993. As a result, Governor Pete Wilson declared the drought to be officially over on February 24, 1993.

Ash Mountain Pasture
The national parks’ Ash Mountain Pasture has experienced seven multi-year drought since it began being used in 1921:
1. 1918–34, a 17-year-long megadrought
2. 1947–50
3. 1959–61
4. 1976–77, the driest two years in the state’s history prior to 2014–15
5. 1987–92
6. 2007–09
7. 2012–15+

The pasture was badly damaged by drought and overuse during the 1918–34 drought. Conditions were so bad by 1934 that the park was seriously considering killing some of its livestock.

The parks’ stock are used to support wilderness operations during the summer. But in the winter, the stock have to be brought back to lower elevation pasture. In the early years, the parks’ stock were kept on the Ash Mountain Pasture during the winter. However, beginning in about 1970, the national parks began sending most or all of their stock outside the parks during the winter whenever they could.

From about 1975 to the present, most of that winter pasture has been on the Horse Pasture Unit at the Pixley National Wildlife Refuge. Pixley manages their Horse Pasture Unit for a particular conservation objective: maintaining average residual dry matter of 800 pounds per acre at the beginning of summer. This is done for the benefit of two threatened and endangered species that live on this pasture: the blunt-nosed lizard and the Tipton kangaroo rat.1563

This partnership between the parks’ stock and the refuge’s conservation objective worked reasonably well until the 1987–92 drought. After the first two dry winters (1987–88 and 1988–89), Pixley’s managers informed the
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national parks that they would have to discontinue putting their stock on the refuge for a while; the refuge Horse Pasture Unit could meet its residual dry matter objective without any grazing. That was because Mediterranean grasses don't grow when the rains don't come.

The parks weren't allowed to put any stock on the refuge during the winters of 1989–90, 1990–91, or 1991-92. As a result, the parks apparently kept most or all of their stock on the Ash Mountain Pasture during those three winters. By the second year (1990–91), the pasture was so depleted that the parks had to purchase a very large amount of supplemental feed to get the stock through the winter.

The drought finally broke in December 1992, and Pixley allowed the parks to bring most of its stock back onto the refuge that winter.

1988 Flood
This flood occurred during the 1987–92 drought.

The winter of 1988–89 was a strong La Niña event.

Flood releases of 8,000 cfs or greater occurred at Friant Dam on the San Joaquin River, supposedly during the April–July snowmelt period in 1988. That is difficult to explain. Total flow for the water year in the Tulare Lake Basin was less than 50% of the 1894–2014 average, and there was no report of flooding in the basin. Possibly the flood release was related to an unusual storm event that occurred in January.

A very intense storm passed through California on January 17, associated with high winds and surf. Several deaths occurred when people became snow-bound in the mountains of Southern California. A 7-foot tide combined with a 15–20 foot surf caused an estimated $50 million in damage to coastal Southern California. Tornadoes were reported in Orange County.

This storm moved out of the Gulf of Alaska and developed into a violent cyclone when it came ashore near Aleva Beach at 1 p.m. on January 17. All-time low barometric pressure was recorded at several Southern California weather stations as the storm moved onshore about 20 miles north of Santa Barbara. Table 88 gives the total precipitation received during the January 17 storm event for selected reporting stations.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Total Precipitation (inches)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Turlock</td>
<td>1.10</td>
</tr>
<tr>
<td>Modesto</td>
<td>1.69</td>
</tr>
<tr>
<td>Newman (northwest of Los Banos)</td>
<td>4.10</td>
</tr>
</tbody>
</table>

Newman's 92-year average annual rainfall is 10.27 inches. The 4.10 inches that Newman received on January 17 was 6.67 standard deviations above the average extreme annual storm. The associated recurrence interval is about 20,000 years. It appears that a local thunderstorm was embedded in the larger statewide storm that hit Newman on January 17.1565

It isn't clear what effect this storm had on the Tulare Lake Basin. Storms such as this often result in localized flooding. The flow of the Kings and Kaweah Rivers roughly doubled on January 18, so there was apparently a strong rain in the northern part of the basin on that day. We haven't found any records to indicate whether this storm caused any localized flooding in that part of the Tulare Lake Basin. However, even if it didn't, the story merits inclusion in this document as an example of how intense a storm can be in Central California.

1991 Flood
Flooding in 1991 occurred in early March. This flood occurred during the fifth year of the 1987–92 drought.

The peak day natural flow at Pine Flat on the Kings occurred on March 4, 1991: 13,078 cfs. This was a sharp peak, 13 times bigger than the flow of the previous day. It seems likely that this was a very high-flow period in Cedar Grove as well.
The March 1991 storm event caused a debris flow at E1 Portal. The event was analyzed by Jerry DeGraff, a geologist for the USFS. Although El Portal is north of the Tulare Lake Basin, it’s worth including this event as an example of how a community can prepare for anticipated debris flows.

This small community was originally established in the Merced River canyon to serve the needs of workers on the local railroad and those working at timber harvest and mining. More recently, it provides residences for employees of Yosemite National Park and its concessionaires. In August 1990, a major wildfire burned parts of the national park and the adjacent Stanislaus National Forest. The burned area included the small, steep watersheds which empty into El Portal. A routine assessment of possible landslide hazard resulting from loss of vegetation was carried out. Previous landslide mapping indicated a large number of past debris flow scars present on the slopes of the burned watersheds. Several closed debris basins were constructed on two drainages where the severity of vegetation loss from the fire, indications of past debris flow activity, and presence of houses at the mouths of these ephemeral channels represented a high risk for debris flow damage.

A major storm occurred in the Sierra from February 27 – March 4, 1991 which triggered debris flows from the drainages above E1 Portal. No precipitation was received in the El Portal area for 21 days prior to this storm. After several days of significant rainfall, the greatest amount of daily precipitation was received on March 3. Interviews with residents of El Portal disclosed that a period of intense rainfall occurred shortly after 11:00 p.m. on Sunday, March 3. The time was firmly established by a number of residents who watch the late television news. At that time water and debris was seen spilling from the closed basin at Chapel Lane. A total of 100 cubic yards of debris was trapped in that basin.

Smaller debris flows occurred on two other drainages leading into E1 Portal at the same time that the basin at Chapel Lane trapped its debris flow. At these drainages, the site conditions had not permitted construction of debris basins. At one location, water, mud, and occasional cobbles flowed against a house and passed between the house and an outbuilding. Fifteen minutes earlier, only water was seen flowing through this location. The debris flow was described by the residents as swift enough to “carry away a small child.” A few tens of meters to the east, residents at another house felt vibrations that seemed greater than the impact of the intense rainfall. Turning on their outside lights, they saw water and debris issuing from a channel and flowing across their backyard. A deposit about one foot thick was formed against the back of their house.

The debris flows at El Portal illustrate the potential for damage to property and threat to life which exists in the Southern Sierra. The Chapel Lane debris basin built as part of burned area emergency rehabilitation, and the small volume of the debris flows from the other two drainages kept costs to a minimum.

1993 Floods (2)
Two floods occurred in 1993, both in January.

January was a very wet month in the Southern Sierra, at least in parts of it. A foot of snow fell in Yosemite Valley on January 28, bringing the monthly total for that site to 175 inches (14.6 feet), a record for any month.

An energetic series of storms swept through Southern California during January 5–19, 1993. High wind and tornadoes were associated with this storm sequence. Extensive flooding occurred in Southern California during this period. The greatest 15-day rainfall totals of record occurred at 132 stations during this storm. Ten stations reported rainfall totals in excess of a storm with a recurrence interval of 1,000 years. Total property damage was $600 million and 20 lives were lost due to flooding. While the great majority of those storms were to our south, some appear to have affected the Tulare Lake Basin.

Horse Creek Dam in Sand Canyon in the Tehachapi Mountains failed on January 9, causing localized flooding.

On January 13, a series of winter storms brought between 1 and 2 inches of rain, flooding numerous farm fields in Fresno County. Several houses between Fresno and Madera were flooded with water up to 3 feet deep and numerous roads were flooded. A debris flow occurred on Highway 33 near Coalinga. A levee collapsed north of Orosi.

Total flow for water year 1993 was 149% of the 1894–2014 average for the Kings, 129% for the Kaweah, 102% for the Tule, and 117% for the Kern.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

1995 Floods (4)
There were four floods in 1995:
1. January (3)
2. March

The winter of 1994–95 was a moderate El Niño event. This association with the 1995 floods may well have been a coincidence. Only strong El Niño events have been shown to have a correlation with high precipitation events and floods in California.

Statewide, water year 1995 had the third-highest rainfall total in historic times. It was exceeded only by rainfall totals for the years 1890 and 1983.

The January and March storms were both events of extremely high one-day rainfall rates concentrated over a relatively small region (i.e., less than 100 miles wide). There were large swaths where 100-year storms occurred with embedded 500- and 1000-year events. These events had over 100 stations which had their highest-ever water year total precipitation. Thirty stations reported over 100 inches for the year; most of those stations were located in the Feather and Yuba River Basins.

Table 89 shows the total precipitation at two mountain stations in the San Joaquin Valley for water year 1995.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Total Precipitation (inches)</th>
<th>Recurrence Interval (years)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Panoche 2 W near the crest of the Coast Range</td>
<td>21.36</td>
<td>1,900</td>
</tr>
<tr>
<td>Florence Lake in the Sierra</td>
<td>50.29</td>
<td>1,300</td>
</tr>
</tbody>
</table>

Historically, years of large rainfall totals were not necessarily years of heavy flood-producing rainfalls. However 1995 was somewhat of an exception as there were numerous periods of robust rainfall activity throughout the state.

The storms of January 1995 extended from Humboldt County in the north to Riverside County in the south. They caused a total of 740 million dollars in damage along with 17 deaths. Extensive debris flows occurred in Santa Barbara County.

The flooding in early January was attributed to a series of two storms originating 500 miles north of Hawaii. The first storm front arrived on January 6. That two-day storm produced moderate precipitation totals in Northern California. The second, and more severe, storm front arrived on January 8 and remained over Northern California through January 10. The evening of January 9–10 brought record rainfall to the already saturated floor of the Central Valley. Sacramento set a new rainfall record, receiving 4.45 inches within a 24-hour period.

Rainfall in December 1994 was just slightly below average, and early January 1995 was well above average for most of the Sacramento River and San Joaquin River Basins. By January 7, Sacramento had received 10 inches of rainfall compared to the average 8 inches.

Record-breaking rainfalls occurred during the 6 days from January 7–12 on the west side of the Sacramento Valley. A total of 50 stations reported their greatest-ever six-day total rainfall. Cobb in the Clear Lake Basin received 35.18 inches in 6 days. Greenville in the Feather River Basin received 30.50 inches in 6 days, which is a recurrence interval of 2,400 years. The main precipitation for this storm series was located in a band extending from Clearlake northeast to the Lake Almanor Region. Another band of high rainfall extended from Whiskeytown north to the McCloud region in the Upper Sacramento River Basin.

The January 10 storm events were embedded in the January 7–12 storm. They occurred almost simultaneously in Sacramento and in Kern County. Needless to say, there were a lot more gages to record the Sacramento event.

On January 10, a major storm event occurred northeast of Sacramento. The peak 24-hour rainfall for this storm was 7.57 inches at the Granite Bay Country Club rain gage. That peak 24-hour storm consisted of three separate rainfall sequences: the first from about 7–11 p.m. on the 9th, the second and heaviest from 4–8 a.m.
on the 10th, and another burst of rain from about 1–5 p.m. Twelve Sacramento area stations reported over 5 inches of rain in one day. Based on the 28-year rainfall record available for Rancho Cordova, the recurrence interval for this storm was 4,000 years. The January 10 storm fell on saturated ground; it was preceded by 8 days of rain. High antecedent rains preceding record rainfalls resulted in devastating flooding in the Sacramento area centered on Linda Creek which flows through Roseville and Rio Linda.\textsuperscript{1575}

On January 10, heavy rain of up to 4 inches caused creeks to swell and washed out several roads in Kern County near Frazier Park and Highway 66 near Maricopa.\textsuperscript{1576}

On January 24, strong thunderstorms moved through the Central California interior, causing flooding in Lamont.\textsuperscript{1577}

On January 25, Kern County was drenched by heavy rain. Up to 5 feet of water surged out of Caliente Creek, washing out roads. Parts of Interstate 5 flooded. Numerous crops were damaged in Arvin, and up to 30 chickens drowned in Loraine.\textsuperscript{1578}

Panoche/Silver Creek west of Mendota flooded sometime in January 1995.\textsuperscript{1579}

A much stronger than normal Pacific jet stream was displaced well south of its normal position during much of the winter and early spring of 1995 due to El Niño conditions in the Pacific. This forced major moisture-laden storm systems directly into California, 15 to 20 degrees south of their normal locations. During January and March, the state was struck repeatedly by very strong storm systems laden with Pacific moisture.\textsuperscript{1580} Flood damages exceeded $498 million for the Central Valley.

Both January and March showed much above-average precipitation over most of the state. Since most of the storms occurred within relatively cool, unstable air masses, much of the precipitation above elevation 5,000 feet accumulated as snow. Water content of snowpack exceeded 150% of average in much of the Sacramento River Basin and Sierra at the end of March.\textsuperscript{1581}

As of January 7, most of the major reservoirs in the Sacramento River and San Joaquin River Basins were less than half full and only 75% of normal after the 1987–92 drought and the relatively dry 1994 water year.\textsuperscript{1582}

None of the major reservoirs in the Sacramento River Basin greatly infringed on their flood-control pool during the January 1995 floods. The major reservoirs in the San Joaquin River Basin experienced similar operations with over 70% of the flood-control pool remaining in all the reservoirs after the January event.\textsuperscript{1583} Runoff from major Sierra rivers during the January flood was mostly stored by reservoirs. Most of the flooding occurred on small streams.

The January storms more severely affected Northern California, while the March storms concentrated more of their impact on Central and Southern California. During March, most locations in the southern San Joaquin River and Tulare Lake Basins received several times their average March precipitation, as illustrated in Table 90.\textsuperscript{1584}

<table>
<thead>
<tr>
<th>City</th>
<th>Percent of Average</th>
</tr>
</thead>
<tbody>
<tr>
<td>Bakersfield</td>
<td>326%</td>
</tr>
<tr>
<td>Coalinga</td>
<td>603%</td>
</tr>
<tr>
<td>Five Points</td>
<td>474%</td>
</tr>
<tr>
<td>Fresno</td>
<td>311%</td>
</tr>
<tr>
<td>Hanford</td>
<td>356%</td>
</tr>
<tr>
<td>Visalia</td>
<td>397%</td>
</tr>
</tbody>
</table>

The heaviest March rainfalls occurred mainly between March 9–10. There was 1.1 billion dollars in property damage attributed to this storm sequence and 16 deaths.\textsuperscript{1585} Three days of soaking rain from March 9–11 resulted in $146.8 million in damage to crops across interior Central California. Mendota was hard hit where many roads and poor drainage areas flooded and gusty winds toppled trees and knocked out power. Highways 140 and 41 to Yosemite National Park were closed due to water, rocks and debris on the roads.\textsuperscript{1586}

The major Central Valley reservoirs had less flood control space to handle the March flood than they had for the January flood. However, runoff from major Sierra rivers was still mostly stored by the reservoirs. Millerton Lake
in the San Joaquin River Basin had less than 5% of its flood-control pool remaining during the peak of the March event.\textsuperscript{1587}

The March storm brought considerable precipitation to the Coast Ranges that borders the west side of the Tulare Lake Basin. It was a major event in the Coast Ranges. Highway 1 was closed by a landslide for approximately a week.\textsuperscript{1588} The highest-ever flood stages were reported on the Salinas River at the Spreckles Highway Bridge. Upstream on the Salinas River, four stations recorded their highest-ever 24-hour rainfall. Paso Robles had a total of 7.40 inches. This event had a recurrence interval of about 1,100 years at Paso Robles.

The March storm on the upper Salinas River spilled over the Coast Ranges into the San Joaquin Valley near Coalinga. Coalinga received 3.74 inches of rain in 24 hours on March 10, breaking that city's previous record of 2.53 inches set in 1914. Since the average annual rainfall for Coalinga is only 7.85 inches, the city received nearly 50% of its average annual precipitation in a 24-hour period.\textsuperscript{1589} The recurrence interval for the Coalinga rain in this storm was 2,400 years.

Kettleman Station and Westhaven also recorded their highest-ever 24-hour rainfalls during this storm event. Fresno tied October 5, 1904 for its wettest calendar day when 2.38 inches of rain fell on March 10.\textsuperscript{1590} High flows occurred on some of the Tulare Lake Basin west side tributaries. Panoche/Silver Creek west of Mendota flooded sometime in March 1995.

Arroyo Pasajero flows from Coalinga east toward Lemoore. The arroyo is fed by four tributaries. From north to south, these are Los Gatos, Warthan, Jacalitos, and Zapato Chino Creeks. On the evening of March 9 (sometimes reported as March 10), extremely high flows in Arroyo Pasajero collapsed the two Interstate 5 bridges near Coalinga, killing seven people. The peak flow in the arroyo was 33,000 cfs, delivering a flood volume of 33,500 acre-feet. That was the flood-of-record in this drainage, larger than the previous record flow that occurred during the 1969 flood event.\textsuperscript{1592} This was one of the three largest flood events to occur in the Coalinga area during historic times.

The next highway in the path of that flood was Lassen Avenue (Highway 269), the road connecting Huron to Highway 198. Before the flood, Lassen Avenue was an elevated highway. Afterwards, it was buried in sediment and other flood debris for hundreds of yards. It was almost as if a lava flow had passed through the area.\textsuperscript{1593}

As described in the section of this document on Gradient Change, significant subsidence has occurred in this portion of the Tulare Lake Basin. This subsidence has resulted in increased gradient for the Arroyo Pasajero and similar stream courses on the west side of the Tulare Lake Basin. This increased gradient results in greater erosion. Many of the soils in this area have silty textures and are cut like butter under these conditions. As a result, the groundwater overdraft and consequent subsidence has increased sources of sediment in the Arroyo Pasajero.\textsuperscript{1594}

Huron is located about 15 miles east of Coalinga on Highway 269 (Lassen Avenue). (Trivia question: In the 2000 Census, Huron had the highest proportion of Hispanics of any city in the United States.) Prior to 1995, Highway 269 was closed an average of 26 days a year due to flooding caused by Arroyo Pasajero. Each time that highway closes, residents of Huron have to drive an additional 20 miles each way when they travel to Fresno. In 1995, Highway 269 was closed for 72 days.

The peak day natural flow at Pine Flat on the Kings occurred on March 11, 1995. The flow on March 11 was nearly five times higher than the flow on the previous day. It was a slightly bigger flow than occurred in the much more famous 1983 flood. It seems likely that this was a very high-flow period in Cedar Grove as well.

The Kaweah’s peak natural flow occurred at Terminus Dam (or possibly McKay’s Point) on March 11: 12,714 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 8,369 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 6 years for the Kaweah.

There is a remarkable similarity in rainfall distribution between the March 1995 storm and the February 1978 cyclic storm which dumped record rainfalls in an area to the south of the area affected by this storm. The 1978 storm produced large rainfalls on the windward slopes of Ventura County and then continued over into the rain shadow area in the Buena Vista Lake region.
The March 1995 storm behaved in a similar manner. It appeared to be a cyclic storm since it produced devastating rainfalls on the windward slopes of the Coast Ranges. It was still quite energetic as it moved into the rain shadow area to create further devastating floods.

A similar cyclic storm came ashore near Monterey Bay on September 11, 1918, resulting in extreme rainfalls at Antioch, again in the rain shadow zone. The 1918 storm was caused by the remnants of a tropical hurricane which originated off the southwest coast of Mexico. As a result of the March 1995 flood, President Clinton declared 39 California counties disaster areas.

1995 was one of the two wettest years ever at Paso Robles, the other being 1941.

Flood releases of 12,000 cfs or greater occurred at Friant Dam on the San Joaquin River during the April–July snowmelt period.

Table 91 summarizes the damage incurred in the southern end of the San Joaquin Valley during the 1995 flood.

<table>
<thead>
<tr>
<th>County</th>
<th>Private Damage (thousand dollars)</th>
<th>Public Damage (thousand dollars)</th>
<th>Business Damage (thousand dollars)</th>
<th>Agricultural Damage (thousand dollars)</th>
<th>Total Damage (thousand dollars)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Merced</td>
<td>$38,854</td>
<td>$38,854</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Madera</td>
<td>$160</td>
<td>$1,300</td>
<td>$10</td>
<td>$829</td>
<td>$2,299</td>
</tr>
<tr>
<td>Fresno</td>
<td>$80</td>
<td>$300</td>
<td>$10</td>
<td>$20,846</td>
<td>$21,364</td>
</tr>
<tr>
<td>Kings</td>
<td>$2,484</td>
<td></td>
<td>$2,484</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Tulare</td>
<td></td>
<td>$48,515</td>
<td>$48,515</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Kern</td>
<td>$10</td>
<td>$1,900</td>
<td>$10</td>
<td>$21,046</td>
<td>$22,966</td>
</tr>
<tr>
<td>Total</td>
<td>$250</td>
<td>$3,500</td>
<td>$30</td>
<td>$132,574</td>
<td>$136,354</td>
</tr>
</tbody>
</table>

Total flow for water year 1995 was 205% of the 1894–2014 average for the Kings, 204% for the Kaweah, 184% for the Tule, and 184% for the Kern.

Flooding occurred in the Tulare Lakebed; this was the first significant flooding since 1986 (see Figure 16). Katrina Young recalled that parts of Manning Road were under nearly two feet of water where it crossed the normally dry lakebed.

In order to minimize flooding in the lakebed, 13,000 acre-feet of river floodwater was pumped into the Friant-Kern Canal and routed to the Los Angeles area.

**1996 Floods (2)**

There were two floods in 1996:

1. April
2. May

An F0 tornado touched down 9 miles west of Fresno on April 16. Small hail that fell in association with this thunderstorm caused $600,000 in crop damage, mainly to grapes. Heavy rain fell in Fresno, stranding motorists in cars. One report had as much as 0.73 inches of rain falling in just 25 minutes.

An intense storm struck Yosemite on May 16. The resulting heavy rain-on-snow event caused the Merced River to flood in Yosemite Valley. Over $2 million in flood damages occurred. On May 26, the Merced at Happy Isle peaked at 5,900 cfs. By Merced River standards, that is relatively big, having a recurrence interval of 15 years. Only the floods of 1937, 1950, 1955, 1964, and 1997 have been bigger.

The Kings River also flooded in May. Peak flow at Pine Flat was 28,705 cfs on May 17. This was more than twice the flow of the day before, suggesting that it was caused by a heavy rain event. This was a relatively minor flood by Kings River standards. There have been 29 bigger floods since record-keeping began.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

1997 Floods (4)
Flooding occurred three times in 1997:
1. January (twice)
2. July
3. September

For a month that was neither an El Niño nor a La Niña, January 1997 saw a lot of damage. The first January flood is sometimes referred to as the New Year’s Day Flood of 1997. By some accounts, this flood was the largest and most extensive flood disaster in the state’s history. It was the second costliest in California’s history.

Statewide, the impacts were:
- 2 deaths, 50 injuries
- 120,000 people displaced by flooding
- $1.6 billion in damages
- 20,000 homes and 1,500 businesses destroyed or damaged
- Disaster areas were declared in 43 counties

Damage to urban and agricultural lands and the cost to replace, restore, and rehabilitate flood damage reached $524 million in the Central Valley.

As a result of severe storms and flooding, a major federal disaster (DR-1155) was declared on January 4, 1997, for the period December 28, 1996 – April 1, 1997. It covered 49 counties including Fresno, Kings, and Tulare.

The flood resulted from a relatively short-duration, high-intensity storm. It was derived from a very warm area of ocean just west of Hawaii. The convection and atmospheric steering resulted in a convergence of cold arctic air and vast tropical moisture. The entire average water year’s precipitation was received by the end of January. The storm lasted from December 29, 1996 – January 4, 1997. The actual transport mechanism was an atmospheric river.

Early winter rainfall was well above average throughout the Sacramento River and San Joaquin River Basins. In the Northern Sierra, total December precipitation exceeded 28 inches, making it the second wettest December of record, exceeded only by the 30.8 inches in December 1955.

The heaviest rainfall fell along the coastal mountains and the Northern Sierra. Over 50 recording stations recorded their historical one-day precipitation totals during this storm. Precipitation was heavy throughout Northern California, with many stations reporting 15–30 inches of precipitation during the nine-day period between December 26 – January 3. Some stations in the Feather River Basins received over 40 inches during that period. Over 14 inches of rain fell in a 24-hour period at Four Trees, north of Oroville.

The flooding resulted from three subtropical storms. Over a three-day period, warm moist winds from the southwest blowing over the Sierra poured more than 30 inches of rain onto the already saturated watersheds. The first of the storms hit Northern California on December 29, 1996. The second storm arrived on December 30. The third and most severe storm hit late on December 31 and lasted through January 2.

Precipitation totals at lower elevations in the Central Valley were not unusually high, in contrast to the extreme rainfall in the upper watersheds. For example, Sacramento received 3.7 inches of rain while Blue Canyon (at elevation 5,000 feet on the American River east of Grass Valley) received over 30 inches of rainfall, thus providing for an orographic ratio of 8 to 1. A typical storm for this region would yield an orographic ratio of about 3.5 to 1 between those two locations. "Orographic ratio" is the contrast between mountain and lowland precipitation, which can be expressed as the ratio of the precipitation at those locations.

In addition to these three subtropical storms, snowmelt also contributed to the already large runoff volumes. Several days before Christmas 1996, a cold storm from the Gulf of Alaska brought snow to low elevations in the Sierra foothills. The low-elevation snowpack that formed had a high water content (five inches at Blue Canyon) and that portion below about 6,000 feet in elevation melted when the three warmer storms hit. The effect of the snowmelt contributed approximately 15% to runoff totals.

George Durkee (national park wilderness ranger) recalled that the rain did not melt the snow very high in the Kings River Basin in the January 1997 event. There was not much evidence of flooding above about 7,000 feet
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

(Cartridge Creek or so). He was skiing in Dusy Basin on really bad sun cups into very early July. (Sun cups are depressions made as snow melts under intense sunshine in a dry atmosphere.)

At the beginning of December 1996, 100% of the flood control space was available. But by Christmas, much of that space was already in use.

Record flows were recorded on rivers throughout California, but particularly on rivers in the Sacramento River and San Joaquin River Basins. Record peak river flows were recorded at 106 gaging stations. Multiple levees on the Sacramento and San Joaquin Rivers broke due to extremely high runoff from melting snow and heavy rainfall.

The North Fork Feather River crested at Grizzly Creek with a peak discharge of 115,000 cfs; it had a recurrence interval that was greater than 100 years. The Feather River fish hatchery was virtually destroyed.

This was a record flood on the American River, the fifth record flood in 46 years. Folsom Lake on the American River experienced a peak inflow of 255,000 cfs.

The Sacramento River crested at Delta with a peak discharge of 62,300 cfs; it had a recurrence interval of 50–75 years.

The Cosumnes River crested at Michigan Bar with a peak discharge of 93,000 cfs; it had a recurrence interval that was greater than 100 years.

As of December 1, 1996, most of the major reservoirs in both the Sacramento River and San Joaquin River Basins were at normal flood control levels (100% of the flood control space was available). Despite this, the San Joaquin River flood management system was pushed beyond its limits during the 1997 flood. Millerton Lake and Don Pedro Reservoir, two of the major projects in the San Joaquin River Basin, exceeded their design capacity.

The January flood caused significant flooding in the San Joaquin Valley as well as the adjacent foothills. Numerous houses adjacent to the San Joaquin River flooded, while agricultural lands near the Merced River were inundated. Flooding also impacted areas in the South Valley, especially Earlimart and Porterville.

The record flows stressed the flood management system to capacity in the Sacramento River Basin and overwhelmed the system in the San Joaquin River Basin. Flood storage behind dams reduced floodflows by half or more. However, levees were overwhelmed in some areas. Levees on Sacramento River tributaries sustained three major breaks. The San Joaquin River levee system failed in 36 places and was extensively damaged throughout its length, resulting in widespread flooding.

In the Tulare Lake Basin, the crest of the flood passed through the mountains and upper foothills very late on the night of January 2. The deltas and valley floor felt the brunt of the flood impact on the following day. The peak flow on the valley floor for many rivers occurred late on the night of January 3. The impacts of the flood were documented by the NWS forecast office in Hanford.
George Durkee and David Karplus (Kings Canyon National Park’ wilderness ranger and trails supervisor, respectively) think that it was the January 1997 flood that altered the streambed of the Middle Fork Kings so markedly. It blew out the many logjams from Grouse Meadow down to the confluence with the South Fork, allowing kayakers to successfully attempt that stretch of the river. It also altered the section of the Kings River below Tehiite Valley, changing the riverbed to more boulders when it was hiked from then on.

Flows were very high on the South Fork Kings in Cedar Grove as well. Significant quantities of fill and riprap were required to repair damage done to the left embankment of the Cedar Grove Bridge during this flood. That damage was similar to what occurred during the 1955 flood. A seven-mile section of Highway 180 from Boyden Cave to the national parks’ boundary was also badly eroded. The Highway 180 Boyden Bridge survived the flood, but about 100 yards of the Grant Grove approach was washed out. This was the same thing that happened in the 1955 flood. There was also bridge damage near Big Creek. Total highway damage in the canyon was estimated to be $1.8 million.

The Kings’ peak natural flow occurred on January 3: 112,000 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 50,217 cfs.) According to the Kings River Handbook, that peak hourly flow of 112,000 cfs was the peak flood-of-record on the river. Flows during the 1867–68 flood would have been greater, but no estimate of those flows has ever been calculated. To give a sense of just how big 112,000 cfs is, the Kings River divides in its lower reaches into distributaries: the South and North Forks. The total channel capacity of those two rivers, measured at Highway 41 north of Lemoore, is about 8,600 cfs.

Based on the flood exceedence rates in Table 29, the 1997 flood had a recurrence interval of 40 years for the Kings River at Pine Flat. That puts this flood in a category with other Cedar Grove floods of the past 70 years (1937, 1950, 1955, 1966, 1969, 1978, 1982, 1983, and 1997) that rise to the level of the modeled 50-year flood: that is, a flood event that occurs about every eight years on average. See the section of this document that describes Cedar Grove Flooding.

Large scale flooding in the Tenmile Creek area in January damaged flood facilities at Hume Lake Christian Camp in Sequoia National Forest.

Heavy rains contributed to high runoff and flooding throughout Sequoia and Kings Canyon National Parks, resulting in significant road, bridge, and trail damage. Eleven inches of rain fell at Hockett Meadow at the 8,500 foot elevation in a 24-hour period.

Kirk Stiltz, the national parks’ roads foreman, recalled January 2 very well. He was the only operator on duty that day at Red Fir. It had been raining steadily through the day, and he was quite busy keeping drains open above snowline. The situation got so bad by early afternoon that two additional operators had to be called in to work. There were many rock, mud, and debris slides on the Generals Highway below Giant Forest. These temporarily blocked the road but did not take it out. The only washout that Kirk recalled on the Generals Highway occurred at Halstead Meadow.

Jim Harvey recalled that the 1997 flood caused a big slide on the old Middle Fork Trail, just west of Elk Creek. That now marks the east end of Tunnel Rock unit of the Ash Mountain Pasture. The parks’ stock come to that slide and can’t go any further.

This was probably the flood that took out the Kaweah river pump at the national parks’ Ash Mountain headquarters. However, Jack Vance recalled that it might have been the smaller February 1998 flood that took out the pump. With the loss of the river pump in 1997, Alder Creek has once again become essentially the headquarters’ sole water supply; there is no longer a significant emergency backup source.

Paul Schwarz recalled that the Ash Mountain river pump was never the primary source of water for Ash Mountain; it was just the parks’ emergency backup in case it was ever needed. However, it was never needed, at least during the time period 1988–1996. Paul recalled that when the Generals Highway construction project went through the Ash Mountain area in 1996, the contractor destroyed the waterline from the river pump to the water plant. They did not tell anyone at the parks about this until it was too late to replace the waterline without tearing up the new road. Then the January 1997 flood destroyed the intake for the river pump system. Therefore, the parks decided not to replace the river pump. Instead, they began drilling a series of test wells to see if they could tap a good groundwater supply.
The 1997 flood was probably the storm event that caused one of the Ash Mountain sewage ponds to lose its integrity. At the time, the ponds were lined with bentonite. As a result of the storm damage, the parks chose to line the ponds with butyl liners to ensure their integrity during high-water events.

Bill Tweed recalled that the 1997 flood caused a rockslide that damaged the piping feeding the Sycamore Creek stock tanks on the Shepherd’s Saddle Road.

The Middle Fork of the Kaweah was high at the Pumpkin Hollow Bridge late on the afternoon of January 2 (multiple photographs on file in the national parks; also see the title page photograph of this document). As shown in Figure 29 the floodwaters kept rising as the last light of day ended.

The title page photograph was taken at about 5:00 p.m. on January 2. The flow shown at that time would be plotted as “17” on Figure 29 using a 24-hour clock. Note that the flow was more than three times as great when the Kaweah peaked at 11 p.m. that night.

Harold Werner, the national parks’ former wildlife ecologist, recalled that streamflows from this storm were so great that the non-native bullfrogs were flushed out of the North Fork Kaweah River Basin within Sequoia National Park and took several years to recolonize it.

The January 1997 flood was apparently responsible for washing away a spare piece of the SCE penstock from Hydroelectric Powerhouse No. 3. That part came to rest against an eight-foot-high boulder 500 yards below the Pumpkin Hollow Bridge. SCE removed that part on December 30, 2013, during the very low-water of the winter of 2013–14. They also removed a steel bridge truss from the upstream side of the Pumpkin Hollow Bridge.

Kirk Stiltz recalled that the Kaweah inundated the low section of the highway between Reimer’s candy store and the Three Rivers school, and that there was a report of a propane tank floating down Highway 198 through that area.

Kirk recalled watching, hearing, and feeling the flood from the Dinely Bridge around 11:00 p.m. on the night of January 2. The river put on a kind of other-worldly lightshow as the rocks crashed together underwater.

Richard Fletcher recalled that the flooding South Fork Kaweah backed up in an unnamed tributary deep into the Cherokee Oaks Subdivision. A large number of rainbow trout moved into this relatively quiet backwater to escape the wild waters of the South Fork. At Richard’s property near Oakridge Drive, this creek was about 60 feet wide (multiple photographs on file in the national parks). When the floodwaters receded, trout up to 16 inches long were left stranded in yards.

Lake Kaweah took on 40% of its total capacity in a 24-hour period. The lake’s elevation went from 620 feet to 670 feet in 36 hours. At that point, the lake was rising nearly five feet per hour.
The Kaweah’s peak natural flow occurred at Terminus Dam (or possibly McKay’s Point) at 11:00 p.m. on January 2: 56,595 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 17,948 cfs.) Based on the flood exceedence rates in Table 29, this had a recurrence interval of 14 years for the Kaweah. (One source reportedly calculated this as having had a recurrence interval of 80 years. However, a USACE Sacramento District hydrologist could not reproduce that result, even using the now-outdated 1971 flood frequency curves.)

Phil Deffenbaugh, park manager at Lake Kaweah, recalled that the flow of Dry Creek exceeded the 5,500 cfs channel capacity of the Lower Kaweah River during the onset of the 1997 flood. Therefore, nothing was released from Lake Kaweah during that part of the flood; the gates in the dam were closed. As the flow in Dry Creek dropped, the gates were opened, allowing the Kaweah River to start flowing through. Lake Kaweah peaked at 115,700 acre-feet on the morning of January 7 with 27,000 acre-feet capacity remaining.

The gates could only be kept shut for a short period, albeit a critical period. Lake Kaweah’s flood-control pool is relatively small compared to the size of its watershed. When there is a moderately severe flood, it is necessary to pass much of the flood through; it just isn’t feasible to keep the dam entirely closed during such a flood. For example, Lake Kaweah filled and emptied twice during the 1997 flood.

The flood left considerable sediment and wood debris in the Lake Kaweah lakebed. Annie Esperanza recalled that some of the trees were obviously giant sequoias. The logs and woody debris were piled and disposed of by burning the following summer.

This was a major flood on the Tule River. It peaked in early January, resulting in significant flooding. Success Dam filled and emptied twice during the flood. A levee broke on the Tule River.

This was also a major flood on the White River; it peaked at the same time in early January as the Tule. State and federal disaster assistance was granted to the town of Earlimart, which suffered millions of dollars of damage to homes and other structures. Highway 99 was closed for over a week due to the flooding. This was the fifth time in 40 years that flooding occurred in the area.

A breach in Poso Creek levees on January 4–5 put water onto the valley floor near Wasco.

The Kern River near Kernville peaked late on January 2 at about 42,000 cfs. (That was the peak hourly flow; the peak average daily flow was 18,780 cfs on January 3.) One mobile home was swept downriver and a couple of others were damaged.

Central San Joaquin Valley agriculture suffered large losses as farmland was inundated from runoff. Uncontrolled small streams and major river flooding caused damage to permanent crops, irrigation equipment, and roads. Agricultural damage was particularly high in Kings County; flooding of the Tulare Lakebed kept acreage from being farmed during the 1997 crop year.

Table 92 summarizes the damage incurred in the southern end of the San Joaquin Valley during the 1997 flood.

<table>
<thead>
<tr>
<th>County</th>
<th>Private Damage (thousand dollars)</th>
<th>Public Damage (thousand dollars)</th>
<th>Business Damage (thousand dollars)</th>
<th>Agricultural Damage (thousand dollars)</th>
<th>Total Damage (thousand dollars)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Merced</td>
<td>$0</td>
<td>$570</td>
<td>$0</td>
<td>$7,610</td>
<td>$8,180</td>
</tr>
<tr>
<td>Madera</td>
<td>$1,400</td>
<td>$270</td>
<td>$20</td>
<td>$2,497</td>
<td>$4,187</td>
</tr>
<tr>
<td>Fresno</td>
<td>$620</td>
<td>$3,400</td>
<td>$0</td>
<td>$1,394</td>
<td>$5,414</td>
</tr>
<tr>
<td>Kings</td>
<td>$1,500</td>
<td>$770</td>
<td>$500</td>
<td>$38,857</td>
<td>$38,857</td>
</tr>
<tr>
<td>Tulare</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Kern</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Total</td>
<td>$3,520</td>
<td>$5,010</td>
<td>$520</td>
<td>$56,424*</td>
<td>$65,474</td>
</tr>
</tbody>
</table>

*Another source put total agricultural damage for these counties at $70.7 million.

Following the heavy rain and snowmelt floods of early January and another storm passage around January 20, another period of heavy rain occurred from the afternoon of January 24 through the evening of January 26.
Storm totals in the Southern Sierra were generally 3–4 inches of rain. Storm totals in the valley were about an inch but rather variable. Table 93 shows the precipitation totals for some of the reporting stations during this storm event.

Table 93. Precipitation during the January 24–26, 1997 storm event.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Oakhurst</td>
<td>4.06</td>
</tr>
<tr>
<td>Fresno</td>
<td>0.83</td>
</tr>
<tr>
<td>Visalia</td>
<td>1.4</td>
</tr>
<tr>
<td>Bakersfield</td>
<td>0.31</td>
</tr>
</tbody>
</table>

This heavy precipitation induced a second round of flooding in the San Joaquin Valley and the foothills. In the valley, small streams swelled and poor drainage roads were covered with water.

In Sequoia National Park, the rain brought rock falls and debris flows. Damage to the Generals Highway occurred just below Giant Forest at 4:30 p.m. on January 23.

The USACE had to make large releases from Lake Success in anticipation of the forecasted heavy mountain rains. This resulted in the Tule River running quite high downstream. The event was anticipated, and at-risk structures were closed. A 10-foot section of the Jaye Street Bridge in Porterville was washed away on January 24.

The second storm of 1997 occurred on July 23. We know this storm from two places. While these were caused by two separate storm events, they could be thought of as one event that occurred in multiple locations:

1. Alder Creek flash flood in the Ash Mountain area. Bill Sullivan recalled that Alder Creek flash flooded on this date. The water was so high and carried so much sediment that the Ash Mountain water plant had to be shut down for a while. The sediment/turbidity situation cleaned itself up by early/mid-evening, and the water plant was producing water again by that time.

2. There was a flash flood on the creek west of Silver City in the Mineral King area. This creek is locally known as Silver Creek or Silver City Creek. This flash flood resulted from an estimated 2 inches of rain that fell within a 45 minute time period over this creek’s small watershed. Water washed over the Mineral King Road, but only minor road erosion occurred. The impacts of the storm and resulting flood were documented by the NWS forecast office in Hanford. This creek would have a much more severe flash flood in 2006.

The third storm of 1997 occurred in September. Monsoonal moisture over the Southwest supported thunderstorm activity over the desert portions of Kern County and the northern Kern County Mountains in early September:

- On September 2, thunderstorms brought 1 inch diameter hail to Mount Mesa (near Lake Isabella) and dropped 1.11 inches of rain in 30 minutes at Ridgecrest. The heavy rain in Ridgecrest caused numerous intersections in that town to flood and some were covered with up to 6 inches of mud. An automated station just west of Ridgecrest recorded 0.90 inches of rain in just 8 minutes (a rate of 6.8 inches per hour).

- On the evening of September 3, a particularly large thunderstorm cell produced 4.5 inches of rain in a little over an hour in Red Rock Canyon State Park. The resulting flash flood brought 28,000 cfs down Red Rock Creek, across Highway 14, and on into Koehn Dry Lake. A 12-foot wall of water swept over Highway 14 at 7:10 p.m. Several highway bridges were damaged, and four cars were swept into the floodwaters. The highway had to be closed until repairs and clean-up could be made. Nearly 100 motorists were stranded by the flooding. A related thunderstorm on the same evening occurred just west of Mojave. The flash flooding associated with that storm produced flooding four feet deep at the intersection of Highways 14 and 58, floating cars. (The Ridgecrest and Red Rock storms and floods were just outside the Tulare Lake Basin, but the story merits inclusion in this document as an example of how intense a summer storm can be.)

- On September 3, a thunderstorm east of Lake Isabella resulted in flash flooding in Scodie Creek, causing water to flow over Highway 178 at Onyx. Hall was also reported with this thunderstorm up to ½ inch in diameter.

Total flow for water year 1997 was 154% of the 1894–2014 average for the Kings, 180% for the Kaweah, 260% for the Tule, and 174% for the Kern.
About 48,000 acres of agricultural land were submerged in the Tulare Lakebed, returning the lake to 1983 levels. Apparently the western edge of the lake came to about the intersection of 10th and Pueblo on the west side of Corcoran. In order to minimize flooding in the lakebed, 87,000 acre-feet of river floodwater was pumped into the Friant-Kern Canal and routed to the Los Angeles area. For comparison, that is half the total storage capacity of the newly expanded Lake Kaweah. The Tulare Lakebed wouldn’t be completely drained until 2000.

An outstanding feature of the 1997 storm event was the number of significant landslides and debris flows that resulted. The large ones that we know about were in the Central Sierra, north of the Tulare Lake Basin. These events merit inclusion in this document because two of them were well studied, and they can inform risk management planning in our area.

**Mill Creek Landslide: South Fork American River**

This event occurred near White Hall on U.S. Highway 50, about 25 miles east of Placerville and 35 miles west of South Lake Tahoe. At this point, the highway is squeezed in a narrow canyon between the South Fork American River and a steep cliff. Another major landslide had occurred just 0.6 miles east of here on April 9, 1983.

This landslide occurred in the El Dorado National Forest. The event was thoroughly analyzed by Robert Sydnor, an engineer geologist for the California Division of Mines and Geology.\(^{1624}\)

The soil was a very wet sandy colluvium, a sandy mud. It contained some silt, but lacked a cohesive clayey matrix to bind it together. It had lots of voids that could hold water. The December 31 – January 1 rain-on-snow event saturated the ground, filling it with water. This was followed two weeks later by a week-long period of sustained heavy rainfall. That was the proverbial straw that broke the camel’s back and triggered the landslide.

At 11:20 p.m. on January 24, 1997, a large section of the hillside gave way. The head of the landslide began on a 44% slope. The landslide moved rapidly downhill, across the highway, and dammed the river. The headscarp was about 1,100 vertical feet above the elevation of the river (elevation 3,420 feet).

The landslide dam was breached at 4:30 a.m. the following morning. The river took five hours to erode the dam down to the original riverbed. This occurred slowly enough that no downstream flooding occurred.

The landslide had a total estimated mass of about 2,000,000 cubic yards (1,500,000 cubic meters).\(^ {1625}\) The highway was buried under 75 feet of debris. It took Caltrans 27 days to remove approximately 275,000 cubic yards of debris and reopen the highway.

**Sourgrass Debris Flow: North Fork Stanislaus River**

This event occurred on Sourgrass Creek, a small tributary of the North Fork Stanislaus River. The debris flow crossed U.S. Highway 4 six miles east of Dorrington and 14 miles west of Ebbetts Pass. The flow ended in the North Fork Stanislaus, five miles upstream of Calaveras Big Trees State Park.

The debris flow occurred in the Stanislaus National Forest. The event was thoroughly analyzed by Jerry DeGraff, a geologist for the USFS.\(^ {1626, 1627, 1628}\)

This debris flow began in glacial till and eroded into other unconsolidated material.

The event began at about 6:30 p.m. on January 1, 1997. It started as a debris slide on a 40% slope. The headscarp was at an elevation of about 5,960 feet. The debris slide almost immediately disaggregated into a debris flow and continued as such all the way to the North Fork Stanislaus River, elevation 3,960 feet. Total distance from the headscarp to the river was 2.4 miles with an elevation drop of 2,000 feet.

The debris flow had an average gradient of about 830 feet/mile. In the gentler sections, it was moving at only 2–3 mph and not eroding much material. However, when the debris flow entered the steeper sections of Sourgrass Ravine, it picked up speed and began scouring the creek channel to bedrock. This would have been a very impressive event to witness. Here and there in the trees, there are large cobbles 12–18-inches in diameter that were hurled out of the channel as the flow passed by. (Imagine a moving catapult, and you get the picture.)

In addition to the scouring, the debris flow cleared everything from its path. Because the flow was cohesive, it left a surprisingly clean swath behind it. About 10% of the flow was deposited on the sides as levees. However, except for a few large boulders, scattered rocks, and tree pieces, the debris flow carried everything else to the end.
The debris flow was approximately 500 feet wide and 20 feet high when it mobilized over Highway 4. Because the force of the mass was spread over a wide area, damage to the highway was minimal. No drivers were there that night to witness this phenomenon.

By the time the flow reached the North Fork Stanislaus, it had attained a maximum speed of over 12 miles per hour. This nearly equals the average peak velocity of past debris flows in the Sierra.

The debris flow began with an initial mass of approximately 65,000 cubic yards. The volume of the debris flow increased due to erosion of material along its path. By the time it reached the river, it was about 300 feet wide and 35 feet high and its mass had tripled to about 190,000 cubic yards. There was a campground located at this point, but fortunately it was closed for reconstruction. Otherwise, there could have been a disaster. (Particularly since this happened in the middle of the night.)

The debris flow poured into the North Fork Stanislaus, completely filling its channel and damming the river. The river was experiencing a major flood at the time, flowing at 28,000 cfs compared with a seasonal average of 250 cfs. It took the river’s floodwaters about one hour to overtop and erode the dam enough to restore unimpeded flow. About 200 acre-feet of that debris would be washed eight miles downstream and deposited in McKays Reservoir, causing that reservoir to lose 10% of its capacity.

**Other 1997 Debris Flows: Central Sierra**

Three USGS geologists conducted an aerial reconnaissance of potential landslide activity in the Central Sierra on January 8, 1997. In addition to the Sourgrass debris slide, their reconnaissance detected the following large debris flows:

- They observed several large debris flows in the Royal Gorge canyon of the North Fork American River, south of Snow Peak. Those debris flows had fallen from the top of the canyon over a thousand feet down into the river.
- About two miles west of Strawberry Lodge (19 miles west of South Lake Tahoe), they observed a large debris flow that had covered U.S. Highway 50 with mud and granite boulders for a distance of several hundred feet. Several other large debris flows crossed the highway in that vicinity.
- Several large debris flows dropped tons of mud and woody debris into Salt Springs Reservoir on the Mokelumne River. That reservoir is located north of Calaveras Big Trees State Park in the El Dorado National Forest.

**1998–99 Floods (7)**

There were six flooding events in 1998:

1. February, 1998 (twice)
3. April, 1998
4. May, 1998
5. September, 1998 (remnants of Hurricane Isis)
6. 1999 Tulare Lakebed flooding

The winter of 1997–98 was a strong El Niño event. It is because of the publicity surrounding this El Niño that Californians came to associate El Niño events with high precipitation and floods.

February began with a strong jet stream (170+ mph) oriented perpendicular to the Sierra. The transport mechanism for the moisture was an atmospheric river. The accompanying storm lasted from February 1–3. The storm’s greatest impact was felt on the Central Coast and the Santa Cruz Mountains.

Weather stations reported 14 inches of rain falling in 45 hours over the coastal mountains. Various creeks and rivers reached flood stage throughout the Santa Cruz Mountains:

- Pescadero Creek crested at Pescadero with a peak discharge of 10,600 cfs; it had a recurrence interval of 25–50 years.
- The San Benito River crested at Hollister with a peak discharge of 34,500 cfs; it had a recurrence interval of 50–75-years.
- Tres Pinos Creek crested at Tres Pinos with a peak discharge of 27,200 cfs; it had a recurrence interval of 75–100 years.
- San Lorenzo Creek crested at San Lorenzo with a peak discharge of 10,300 cfs; it had a recurrence interval of 75–100 years.
Although the storm was focused on the coastal mountains, it had a significant impact on the Tulare Lake Basin. The storm began on February 1. By the next day, flooding was being reported from the far south end of the San Joaquin Valley. On February 2, heavy rainfall led to flash flooding and water over Highway 166 southwest of Bakersfield.

The storm peaked on February 3. It resulted in the lowest barometric pressure ever recorded during a February in Fresno, and the lowest barometric pressure ever recorded in any month in Bakersfield. Southerly winds increased throughout the morning of February 3, blowing down trees, power lines, fences, and damaging buildings. Bakersfield experienced near record-setting wind gusts. Many roofs were damaged in Tulare, Kings, and Kern Counties. One woman was injured near Lemoore when a tree fell on her.

Fresno County was hit by a series of storms that brought heavy rainfall to the Coast Ranges to the west and high wind and heavy rainfall to the San Joaquin Valley floor. Runoff from the Coast Ranges caused flooding in west Fresno County affecting agricultural areas around Mendota, Firebaugh, and Cantua Creek. Approximately 9,300 acres of farmland were flooded.

Some of the worst flooding was about 15 miles southwest of Mendota. The estimated flow in Panoche Creek at Interstate 5 (northwest of Mendota, between Belmont and Nees) was 17,000 cfs on the morning of February 3. Cantua Creek and Arroyo Hondo combined to flood 240 acres of farmland.

Over 100,000 chickens died near Gustine (northwest of Los Banos) when two chicken farms were inundated by the floodwaters of Garzas Creek just before dawn on the morning of February 3. The Los Banos area received 3.16 inches of rain in the previous 24 hours by 10:00 a.m. on February 3. Los Banos Creek peaked at midnight on February 2 with a flow of 14,480 cfs into the Los Banos Creek Reservoir. This set a new record, breaking the 11,500 cfs record flow set in 1955.

That was only the beginning of one of the wettest months on record in the Tulare Lake Basin. Measurable rain fell in Fresno on 21 of February’s 28 days, resulting in the third-wettest February on record for that city. With 5.36 inches in rain, Bakersfield experienced the wettest month ever recorded since record-keeping began in 1889. (This record would be broken in December 2010.) The near constant rains kept the ground in the foothills and the San Joaquin Valley floor saturated, and runoff caused persistent problems through the month.

In Bakersfield, significant rain led to ponding water and flooding on many secondary roadways on the evening of February 7.

The same storm system impacted the west side of Fresno County that evening and the following day. Streamflow from Panoche/Silver Creek crested at 13,000 cfs at 10:00 p.m. on February 7. The resulting flooding downstream in Mendota peaked on the evening of February 8. Flooding affected Highway 198 west of Interstate 5 in far western Fresno County.

A second major storm struck on February 23. The impact of that storm was apparently focused on the southeast side of the Tulare Lake Basin from the vicinity of Lindsay south to the Tehachapis. This storm event also affected Southern California; Trabuco Creek in Orange County had its flood-of-record on February 23, 1998. This was similar to the 1916 and February 1937 floods.

Lewis Creek near Tonyville and Frazier Creek near Strathmore both overflowed early on February 24, causing an estimated $1.5 million in damage to area homes and businesses. Rainfall in the 24 hours prior to the flooding was estimated to be 1–1½ inches in the lower Tulare County foothills.

The White River had a 700 cfs flow in its shallow channel by midnight on February 23. All the tributaries of the White River (Speas, Chalaney, Coho, and Tyler Gulch Creeks) were flowing heavy. The river breached a levee at 1:30 a.m. on February 24 and flooded the town of Earlimart. Highway 99 had to be closed through the town and remained closed for a week. Up to 250 homes in the town were impacted by the flooding, with 50 homes having 3 feet or more of water in them and 220 people forced to evacuate. Damage in Earlimart was estimated to be $13.7 million.

Poso Creek breached its banks late on the night of February 23 with a peak flow estimated to be 7,000 cfs. The creek flooded 112 homes in the town of McFarland; damage was estimated to be $2.5 million. Poso Creek floodwaters also threatened some rural homes downstream near Wasco later on February 24.
Evacuation of Lamont was begun at 8:00 p.m. on February 23 in anticipation of flooding due to the steady accumulation of rain. Caliente Creek progressed from nuisance flow to flooding before dawn on the morning of February 24. Caliente Creek at Bena (in the Tehachapi foothills upstream of Lamont) peaked by 2:30 a.m. on February 24: 6,000 cfs. Farther downstream of Lamont, water from Caliente Creek flooded and closed the northbound lanes of Highway 99 at Herring Road (8 miles south of Bakersfield) by 8:00 a.m. on February 24.

February precipitation for the Sacramento 8-station index was 265% of average. For the 8 reference stations in that index, average precipitation is 7.9 inches, but February 1998 had 20.9 inches. The statewide snowpack water content by the end of February was running at 160% of average. The impacts of the heavy February rains and resulting flooding were documented by the NWS forecast office in Hanford as shown in Table 94.

<table>
<thead>
<tr>
<th>City</th>
<th>1998 Rainfall (inches of rain)</th>
<th>Average Rainfall (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Fresno</td>
<td>5.10</td>
<td>1.80</td>
</tr>
<tr>
<td>Hanford</td>
<td>4.26</td>
<td>1.45</td>
</tr>
<tr>
<td>Bakersfield</td>
<td>5.36*</td>
<td>1.03</td>
</tr>
</tbody>
</table>

*The 5.36 inches of rain recorded by Bakersfield made that the city’s second wettest February on record.

Seasonal rainfall 10 miles northeast of Springville was 28.86 inches by February 23.

As a result of severe winter storms and flooding, a major federal disaster (DR-1203) was declared on February 9, 1998 for the period February 2, 1998 – April 30, 1998. It covered 41 counties including Fresno, Kern, and Tulare.

Table 95 summarizes the damages from the various February storms.

<table>
<thead>
<tr>
<th>County</th>
<th>Property Damage (million dollars)</th>
<th>Agricultural Damage (million dollars)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Merced</td>
<td>$2.0</td>
<td>1.4</td>
</tr>
<tr>
<td>Fresno</td>
<td>$1.6</td>
<td>1.8</td>
</tr>
<tr>
<td>Kings</td>
<td>$0.02</td>
<td>1.0</td>
</tr>
<tr>
<td>Tulare</td>
<td>$13.9</td>
<td>1.5</td>
</tr>
<tr>
<td>Kern</td>
<td>$12.5</td>
<td>5.4</td>
</tr>
<tr>
<td>Total</td>
<td>$30.02</td>
<td>$11.1</td>
</tr>
</tbody>
</table>

1 In Kings County about $1.0 million in flood protection costs was expended to try to protect agricultural land.
2 In Kern County, the areas most severely impacted were Arvin-Lamont and McFarland, although flooding also occurred in the Lebec-Frazier Park-Cuddy Valley area and in the Kern River Valley.

Flood releases of 8,000 cfs or greater occurred at Friant Dam on the San Joaquin River during the April–July snowmelt period.

The peak day natural flow at Pine Flat on the Kings occurred on June 17. There were sustained high flows throughout June and July. These were clearly above-average flows, but there had been about a dozen floods greater than this in the previous 45 years. Although the river didn’t reach a particularly impressive height, it delivered a tremendous amount of water to the valley floor.

On March 25, a band of quasi-stationary thunderstorms deluged Merced with 3 to 6 inches of rain in a 12- to 18-hour period. One gage in the northern part of the city of Merced had 6.8 inches in a 48-hour period from late March 23 to the 25th. The Merced Airport recorded 3.25 inches of rain on the 24th alone. Bear Creek reached a crest of 19.3 feet on the morning of March 25, resulting in 1,000 people being evacuated. A total of 65 homes
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

and 19 apartments were flooded. Damages totaled $9.6 million to property with agriculture suffering a $1.5 million loss.1636

Panoche/Silver Creek west of Mendota flooded sometime in March 1998.1637

On April 1, Lewis and Frazier Creeks swelled due to heavy rains and snowmelt, resulting in flooding in Lindsay, Strathmore and Tonyville, damaging 32 homes.1638

On May 2, thunderstorms unleashed locally heavy rain over rugged terrain northeast of Bakersfield. A spotter reported 1.5 inches of rain falling in about an hour. The rapid runoff from those storms resulted in flash flooding on several streets in the Bakersfield area, including Highways 178 and 58. Many vehicles became stalled in the high water, and some residences and apartments were flooded with as much as 3 feet of water.1639

On September 4, moisture associated with remnants of Hurricane Isis brought rain to parts of interior Central California. Frazier Park received 1.53 inches of rain and Bakersfield received 0.27 inches of rain, setting a new daily precipitation record. Trace amounts were reported in the valley as far north as Madera.1640

Total flow for water year 1998 was 180% of the 1894–2014 average for the Kings, 220% for the Kaweah, 336% for the Tule, and 225% for the Kern. This was the fifth-highest year ever recorded on the Kaweah and the fourth highest on the Tule.

The combined runoff of the four rivers in the Tulare Lake Basin during water year 1998 was 5,990,549 acre-feet, 204% of average. As shown in Figure 19 on page 123, that allowed the groundwater aquifer to be recharged slightly that year. Such a recharge event rarely happens.

Tulare Lake had reappeared in 1997, and the 1998 flood exacerbated the problems being experienced by landowners in the lakebed. In order to minimize flooding there, 202,000 acre-feet of river floodwater was pumped into the Friant-Kern Canal and routed to the Los Angeles area in 1998. For the same reason, 130,000 acre-feet of Kern River water was also routed to Los Angeles. An additional 984,000 acre-feet of Kings River water was routed through the James Bypass to the San Francisco Bay.

That helped to minimize the lakebed flooding, which was beneficial for the farmers there. However, it meant that the Tulare Lake Basin lost 1,316,000 acre-feet of water that could have been put to productive use or used to recharge our groundwater aquifers. For comparison, that 1,316,000 acre-feet of water is equivalent to 81% of the combined current capacity of all four of the federal reservoirs in the Tulare Lake Basin.

The Tulare Lakebed flooded in 1999. This flooding occurred despite the fact that there was no storm event of note that year.

Runoff was also below average in 1999. Runoff during water year 1999 was only 74% of the 121-year average (1894–2014) for the four major rivers in the Tulare Lake Basin (Kings, Kaweah, Tule, and Kern) combined.

About 32,000 acres of agricultural land were submerged in the Tulare Lakebed in 1998. The flooding in the lakebed in 1998 and 1999 was to some degree left over from the big 1997 flood. The floods of 1998 exacerbated the lakebed flooding. As illustrated in Figure 16, the lakebed would not be fully drained until 2000.

Lakebed flooding is a social construct; it is counted based on the number of growing seasons that are missed. The lakebed was flooded for three growing seasons: 1997, 1998, and 1999. Therefore, this is counted as three floods from the perspective of the lakebed farmers, even though flood events occurred in only two of those years.

Something similar happened in the lakebed in 1969–1971 and 1982–84. In each of those cases, lakebed flooding continued into a non-flood year.

1999–2004 Drought

As reflected in Table 106, this drought was active from 2001–2004 in the San Joaquin River Basin. Tree-ring reconstruction of flow on the San Joaquin River showed that the drought didn't begin anywhere in that basin until 2001.1641
Table 106 also shows total runoff for the four major rivers in the Tulare Lake Basin (Kings, Kaweah, Tule, and Kern) combined during this drought. Those flows show that this drought started two years earlier in our basin. Here it was active from 1999–2004. Average flow over the six years of the drought was 70% of the 1894–2014 average.

As described in the section on Megadroughts since the Little Ice Age, California’s 1999–2004 drought is part of a longer-term megadrought that has impacted most of the Western U.S. since 2000. Because we are on the edge of that huge drought system, we tend to only be aware of it when it reaches out to encompass our area.1642

The San Joaquin River Basin has only been affected by that megadrought for 12 of the years that it has been active: 2000–04, 2007–09, and 2012–15 (based on the San Joaquin Valley Water Year Index and/or total runoff for our four major rivers). From our perspective, we tend to think of those three episodes as individual droughts of relatively average duration instead of being part of the larger megadrought.

Two very large fires occurred on the Sequoia National Forest during the 1999–2004 drought:
- The McNally Fire occurred in July and August 2002. It burned 149,475 acres.

These are the largest fires to have occurred in the Tulare Lake Basin in historic times.

The drought affected all of the Central Valley including the Klamath Basin. Some of the most heated water battles in the West have taken place on the Klamath. In 2001, federal officials shut off irrigation to thousands of acres of farmland in Oregon and California to protect endangered fish during this drought. In the aftermath, federal marshals had to be called in to stop angry farmers from reopening locked irrigation gates.1643

The 2001-02 rain season in Southern California was the driest since record-keeping began in 1877. San Diego recorded only 2.99 inches compared to the annual average of 10.34 inches.

In 2002 the drought affected much of the Western U.S., including the Southwest, the Midwest, the Rocky Mountains, and the prairie provinces in Canada. Exceptional drought conditions were accompanied by hot temperatures and wildfires.

During 2002, more than 50% of the coterminous U.S. was under moderate to severe drought conditions. In a report published in 2009, a team led by Edward Cook used the Palmer Drought Severity Indices (PDSI) to compare the 2002 drought year with other droughts going back to 1900.1644 They found that 2002 was second in intensity only to the Dust Bowl drought year of 1934.1642

The single driest year of record for inflow to Lake Powell was 2002 (the prior dry year record had been set in 1977). The decade of the 2000s (2000–2009, inclusive) was the driest decade in the historical record. During these prolonged dry conditions, total system storage dropped to just below half of capacity.1646

**2000 Floods (2)**

Flooding occurred twice in 2000:
1. January
2. October

From January 23–25, a three-day storm brought locally heavy winter rains to the valley, foothills and lower elevations of the Sierra as shown in Table 96.1647

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Bass Lake</td>
<td>6.78</td>
</tr>
<tr>
<td>Shaver Lake</td>
<td>5.69</td>
</tr>
<tr>
<td>Northeast Fresno</td>
<td>2.29</td>
</tr>
<tr>
<td>East Visalia</td>
<td>2.20</td>
</tr>
</tbody>
</table>

Valley urban areas had significant ponding of water and mountain streams exhibited moderately large amounts of flow. There was some flooding along the valley floor and foothill interface.1648
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

An early season storm brought several inches of snow to the Central and Southern Sierra on October 10. Table 97 shows precipitation totals during that storm event.

Table 97. Precipitation during the October 10, 2000 storm event.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of snow)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Lodgepole</td>
<td>10</td>
</tr>
<tr>
<td>Mount Tom</td>
<td>8</td>
</tr>
<tr>
<td>Huntington Lake</td>
<td>5</td>
</tr>
<tr>
<td>Tuolumne Meadows</td>
<td>4</td>
</tr>
</tbody>
</table>

Over an inch of rain fell in some areas in the valley, including Fresno, resulting in the closure of the Fresno Fair for the first time since 1922. The rain caused numerous flooding problems in Fresno and ceilings to collapse in buildings in Tulare.\textsuperscript{1649}

**2001 Floods (2)**

Flooding occurred twice in 2001:
1. March
2. August

These floods occurred during the 1999–2004 drought.

Although not a flood, an unusual snow event occurred on February 12–13. A low-pressure system came onshore near Point Conception and brought unusually heavy snow to the mountains at the southern end of the San Joaquin Valley. The snowline in the Tulare County foothills dropped to 3,000 feet. Greenhorn Summit received 18 inches of snow, and Frazier Park received 30 inches. Some 500 motorists were stranded on Interstate 5 over the Grapevine.\textsuperscript{1650}

Gary Sanger at the NWS forecast office in Hanford speculated that a system such as the above storm might have accounted for the phenomenal snowfall that the Southern Sierra experienced in January and February 1906, especially if a deep atmospheric river were entrained.

Western Fresno County experienced heavy rain from late on March 4–6; Coalinga received 2.99 inches. Several roads in the area were washed out, including Highway 33/198.\textsuperscript{1651}

An intense thunderstorm struck a large portion of the Rock Creek Basin on August 9. This caused Rock Creek to quickly flash flood, sending a large quantity of water out onto the Kern Valley floor about two miles north of Kern Hot Spring. Erika Jostad was on patrol and witnessed both the deluge and the flood. (Tony Caprio, the national parks’ fire ecologist and his fire effects crew showed up shortly thereafter.) In a place that does not typically flow water, a wall of milky water, gravel, and woody debris flowed across the High Sierra Trail and into the Kern River. It is unclear whether this event was a debris flow or a flash flood that carried large quantities of debris.

Tony recalled that the volume of water coming out of Rock Creek was so great that the Kern River ran milky white for some distance downstream. The flood debris obliterated the High Sierra Trail for a stretch of a few hundred feet (multiple photographs on file in the national parks). In the photographs, tree trunks appear to have a couple feet of gravel piled against them. Some of the images show damage to the standing trees a couple of feet above the level of the gravel bank, giving some indication of the depth of water and debris flowing over the area at peak flow. The High Sierra Trail was not rebuilt but gradually reestablished through use.

A similar flash flood would occur in the Rock Creek Basin in July 2011. Although that flood may have been even larger than the 2001 flood, it did not put any debris onto the High Sierra Trail. We don’t know why these two floods differed in this way. Tony Caprio speculated that the 2001 flood might have cleared accumulated material from the stream channel, and there had been insufficient time to accumulate a similar quantity of material before the 2011 flood.
2002 Floods (2)
Flooding occurred twice in 2002:
1. May
2. November

These floods occurred during the 1999–2004 drought.

Thunderstorms dropped 1.01 inches of rain at the Hanford Airport in just 21 minutes on the afternoon of May 30 (a rate of 2.9 inches per hour), resulting in street flooding in that city. Some 260 lightning strikes were recorded in just an hour in the central and southern San Joaquin Valley.\(^{1652}\)

Hurricane Huko (the Hawaiian equivalent of the name Hugo) formed in the central Pacific and became a hurricane on October 28. Huko then became a wanderer, impacting all three North Pacific basins (east, central, and west) and eventually morphed into an extratropical cyclone.

Tropical moisture from Huko combined with a major trough from the eastern Pacific to bring copious amounts of precipitation and gusty wind to the Tulare Lake Basin from November 7 until early on the 9th. The impacts of that storm and the resulting flood were documented by the NWS forecast office in Hanford.\(^{1653}\)

The flood covered a number of locations in the foothills and mountains of Tulare and Kern counties. Flooding problems were most pronounced in the Tulare County mountains and the higher foothills. Table 98 provides precipitation totals for some stations during the storm event. Numerous foothill locations received 5–10 inches of rain during the three-day period.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Merced</td>
<td>1.80</td>
</tr>
<tr>
<td>Fresno</td>
<td>1.76</td>
</tr>
<tr>
<td>Lodgepole</td>
<td>11.60</td>
</tr>
<tr>
<td>Ash Mountain</td>
<td>5.89</td>
</tr>
<tr>
<td>Hanford</td>
<td>1.44</td>
</tr>
<tr>
<td>Glennville (NE of Bakersfield)</td>
<td>6</td>
</tr>
<tr>
<td>Bakersfield</td>
<td>1.29</td>
</tr>
<tr>
<td>Johnsondale</td>
<td>16.38</td>
</tr>
<tr>
<td>Tehachapi</td>
<td>4.67</td>
</tr>
</tbody>
</table>

Many valley locations set new 24-hour rainfall records on November 8. For example, Fresno’s old record for the 8th was 0.23 inches; the new record set during this storm was 0.98 inches.

Snow levels were relatively high at 9,000 feet. Chagoopa Plateau received 80 inches (6.7 feet) of new snow during the three-day event.

Gusty winds associated with the storm caused 23 pole fires, resulting in 102,000 valley residents losing power.

There was very heavy rainfall in Sequoia National Park on the morning of November 8. After a few hours of such rain, the Kaweah River was rising at a rate comparable to the January 1997 storm. If that rain had continued for a few more hours, it was conceivable that the national parks’ approach to the Pumpkin Hollow Bridge would have washed out as it had in 1937, 1955, and 1966. As a precaution, the Ash Mountain Entrance was closed and a partial evacuation of the national parks was begun. Fortunately the rain stopped before the approaches to the bridge sustained any damage.

The November 2002 flood was apparently responsible for washing away a 500-pound, 16-foot-long steel channel beam from Hydroelectric Powerhouse No. 3. This part came to rest behind the Buckeye Tree Lodge. It was removed by kayaking volunteers during the low-water of the winter of 2013–14.\(^{1654}\)

According to records provided by the USACE, the Kaweah’s peak natural flow occurred at Terminus Dam on November 8: 30,273 cfs. (That was the peak hourly flow; the peak average daily flow, as reflected in Table 28, was 9,436 cfs.)
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Based on the flood exceedence rates in Table 29, this flood event had a recurrence interval of 7 years for the Kaweah River. It would have a recurrence interval of 14 years if calculated using the 30,273 cfs peak flow. (Another source reportedly calculated this as having a recurrence interval of 18 years. Presumably that was done by using the peak flood flow and the now-outdated 1971 version of the flood frequency curves.)

Bill Sullivan and Cal Kessner recalled that Alder Creek flooded with water that was the color of coffee and cream. The water was so high and carried so much sediment that the Ash Mountain water plant had to be shut down. (The floodwaters and debris washed on downstream to where they clogged the culvert on the Generals Highway and caused significant erosion problems there.) After the flood, the sluice gate in the Ash Mountain dam was opened, and much of the sediment was worked downstream. It was two or three days before the water plant could be brought back on line.

During the storm event, a landslide dam apparently formed on a very small stream on the north side of Shepherd’s Saddle in the national parks. When that dam failed, a huge wall of water came down this small drainage, washing out the Shepherd’s Saddle Road and placing three huge rocks at a surprisingly high elevation (photograph on file in the national parks). Those rocks didn’t fall down from a higher elevation. However improbable, the floodwaters from this very small stream lifted the rocks up on high. The road was washed out, but the parks were able to repair it the following year for about $90,000. (This was the cost to repair all the damage to the Shepherd’s Saddle Road, only about $30,000 of this went to repairing the big washout near the saddle.)

That storm also caused Sycamore Creek, a normally very small stream, to wash away two stock tanks that sit beside the Shepherd’s Saddle Road.

With numerous rock falls and debris flows, flooding, and road erosion problems, the Generals Highway and Mineral King roads were closed. Campers were evacuated from Potwisha on the morning of November 9. The Mineral King Road and Generals Highway sustained significant damage and required 1.25 million dollars to repair.

An additional $150,000 was required to repair damage to the Crystal Cave Road in the national parks. A large culvert plugged and allowed a creek to flow down the road washing a deep trench along the uphill side of that road until it finally crossed and washed out the road and fill slope. The road was left impassable due to the sinkhole. The parks used gabion baskets to stabilize the shoulder and used boulders from a slide to fill the sinkhole.

Kirk Stiltz, the national parks’ road foreman, recalled that many of the culverts that plugged during the November 2002 storm, plugged as a result of debris flows. The parks experience a lot of small debris flows when there are heavy rain events.

The total cost to repair all the national park roads damaged during this storm event was approximately 1.49 million dollars. That included the Generals Highway, the Mineral King Road, the Crystal Cave Road, and the Shepherd’s Saddle Road.

Numerous other roads flooded and debris flows occurred in the foothills of the Southern Sierra. Several roads were flooded in Kern County. Three roads were washed out in southeast Tulare County:
1. The Parker Pass Road
2. The road below the Durwood Resort
3. The road that leads from Johnsondale southward to Kernville along the North Fork Kern (Mountain 99, aka Kern County SM99)

Rock falls and debris flows occurred on Highway 168 and Highway 180 in the Southern Sierra foothills.

The McNally Fire had burned about 150,000 acres of Sequoia National Forest in July and August 2002. Some of that area burned quite hot. When the intense November storm hit that area a few months later, some erosion problems resulted. Debris was spread across many mountain roads in the area as well as contributing to a fish kill in the Kern River.

Peak flow into Lake Isabella from the Kern River of 26,500 cfs occurred on the night of November 8. (That was the peak hourly flow; the peak average daily flow was 10,306 cfs on November 9.) The lake storage increased
from 82,000 acre-feet to 109,000 acre-feet, and the lake rose 5 feet in elevation during the two-day period from November 8–9. Flooding and debris flow problems occurred along Highway 178.

### 2003 Floods (3)
Flooding occurred at least three times in 2003:

1. February
2. August (several times)
3. December

These floods occurred during the 1999–2004 drought.

An intense storm struck northwest Fresno on February 13, dropping 3.40 inches of rain in just 2 hours. Up to 3 feet of water flooded parts of the area.°

There were monsoonal influences over the Central and Southern Sierra throughout August, resulting in periods of heavy rain, localized flooding, and brief road closures. Multiple flash floods occurred around Kernville, Tehachapi, Johnsondale, and along the Sherman Pass Road. Estimated rainfall rates of 3–4 inches per hour occurred in an area from near Lake Wishon (near Shaver Lake) south to near Lodgepole on the afternoon of August 2. The Mineral King Road flooded in Mineral King from heavy rain.

Heavy thunderstorms drenched parts of the Kern County mountains in the morning hours of August 21. Piute received 1.78 inches of rain.

Monsoonal moisture generated thunderstorms late on August 25 that produced over an inch of rain in some areas. Cottonwood Creek in the Sierra received 1.68 inches, and 1.48 inches fell at Lost Hills in the valley. Roads were closed in parts of Sequoia National Park and in the Kern Plateau.

On December 25, locally heavy rain on the Southern San Joaquin Valley floor and adjacent foothills led to flooding at several locations over the South Valley due to runoff, including in Bakersfield.

Fresno received 0.99 inch of rain on December 25, setting a daily precipitation record; the old record of 0.53 inch had been set in 1946. Bakersfield received 0.91 inch that day, also setting a daily precipitation record; the old record of 0.76 inch had been set in 1931. A storm total of 1.54 inches was reported just east of Fresno. Large amounts of snow fell in the neighboring Sierra. There were 4 indirect deaths in vehicles caused by the heavy rainfall southeast of Bakersfield.

In the fall of 2003, several wildfires raged across the mountains and hills of Southern California. Two of those fires — the Old and Grand Prix — occurred in areas adjacent to Cajon Pass. When winter rains fall on steep Southern California land that has burned, water runoff is greater than on land that has not burned. The possibility of debris flow events also increases.

During a 24-hour period from December 24–25, 2003, more than 4 inches of rain fell on an area that had burned in the Old / Grand Prix complex. This torrential rain resulted in excessive runoff, sending a debris flow through the small town of Devore on December 25. This debris flow was captured in a dramatic video as it came through town (multiple videos on file in the national parks).°

Devore is on the south side of the Tehachapi Mountains. Although it isn't in the Tulare Lake Basin, it merits inclusion in this document because it was nearby and it is one of the few videos that we have to illustrate what a debris flow looks like.
Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

2005–06 Floods (8)
Flooding occurred eight times in 2005–06:
1. April 2005
2. May 2005 (twice)
3. August 2005
4. October 2005
6. April 2006
7. July 2006

Severe thunderstorms struck in the afternoon hours of April 28, 2005, dropping hail as large as 1¼ inches in diameter in Kings County, damaging crops, including 20% of the cherry crop. A number of streets were flooded in Fresno, and 3.57 inches of rain fell in Parlier. A funnel cloud was spotted and photographed north of Visalia.

Heavy rainfall from thunderstorms on the afternoon of May 5, 2005, triggered flooding in the Madera area. A roof collapsed on a building in downtown Madera, and several roads were flooded. Parts of Highway 99 and Interstate 5 experienced flooding in Kern County. All the roads in Coalinga were flooded with some described as impassible.

On May 16, 2005, a rain-on-snow (actually, a rain-through-snow) flood occurred on the Merced River in Yosemite Valley. Although milder in degree, this was similar in nature to the more famous flood that occurred on that river May 16, 1996.

This was the first major flood observed by the new high-country hydroclimatic network in Yosemite. That network was developed by scientists from the USGS, Scripps Institution of Oceanography, California DWR, National Park Service, and other institutions, and consists of streamflow and air-temperature loggers, plus snow-instrumentation sites.

A storm drew warm, wet subtropical air into the Sierra, bringing moderate rain to the Southern Sierra on the night of May 15–16. That resulted in flooding on a number of rivers on May 16. Like so many cool-season floods in the Southern Sierra, the flood was mostly due to unusually warm temperatures and large catchment areas that received (moderate) rainfall rather than snowfall. In Yosemite, temperatures were above freezing up to about 10,000 feet elevation. Rain fell and streams filled up to 10,000 feet compared to typical freezing levels of about 5,000 feet.

With snow levels so high, rainfall amounts averaging 1.75 inches in the mountains combined with a snowmelt runoff contribution of about 1 inch water equivalent caused river flooding on the valley floor in Yosemite. The Merced River rose 3 feet on the morning of May 16, prompting the park to evacuate campers in Yosemite Valley. The Merced River crested at 12.5 feet in Yosemite Valley, forcing the closure of roads into the valley.

Warm storms — past and future — can unleash floods when rain falls over unusually large catchment areas. (In this flood, the area receiving rainfall may have been as much as five times normal.) Warmer temperatures in the future may increase the frequency and severity of these floods, even as snowpack volumes decline.

The storm was less intense in the Tulare Lake Basin. The flow on the Kings and Kaweah Rivers on May 16 increased 81% and 74% respectively compared to the previous day. The flow on the Tule and the Kern Rivers only increased 47% and 22%.

A large and severe thunderstorm swept over the southeast part of Kern County on the evening of August 15. The California City Fire Department rain gage measured 5 inches of rain in just one hour from the deluge. This led to flash flooding in that town from Cache Creek and extensive sheet flow through the area. Portions of Highway 14 and Highway 58 flooded. (This storm and flood was just outside the Tulare Lake Basin, but the story merits inclusion in this document as an example of how intense a summer storm can be.)

Weak low pressure off the Southern California Coast entrained tropical moisture that resulted in an intense storm striking the Tehachapi Mountains from the evening of October 17 into the morning of the 18th. The impacts of the storm and resulting flood were documented by the NWS forecast office in Hanford.
There were numerous rainfall reports of 2–3 inches of rain from throughout the Tehachapi Mountains during the storm event. Among the various reporting stations, Bear Valley Springs had 2.31 inches and the Piute Forest Service RAWS automated weather station reported 2.31 inches of rain for the event. Numerous locations from Tehachapi to Taft experienced flash flooding, but the Frazier Park area was especially hard hit. Cuddy Creek overflowed in that community, flooding some areas 4-feet deep and resulting in the evacuation of at least twenty people.

The December 2005 – January 2006 flooding was most severe in Northern California, but did come as far south as the Tulare Lake Basin. Statewide, the storm lasted from December 29, 2005 – January 2, 2006. The transport mechanism for the moisture was an atmospheric river.

Damages totaled $300 million. As a result of severe storms, flooding, debris flows, and landslides, a major federal disaster (DR-1628) was declared on February 3, 2006 for the period December 17, 2005 – January 3, 2006; it covered 30 Northern California counties. The impacts of the storm and resulting flood on the Tulare Lake Basin were documented by the NWS forecast office in Hanford.

Rainfall totals for December 24 – January 3 exceeded 20 inches throughout the Northern Sierra. Coastal Range stations in the Russian and Napa River Basins received 18–30 inches. On December 31, there were widespread 24-hour rainfall totals in excess of 5 inches. As a result of this heavy rain, several rivers throughout Northern California came above flood stage. Recurrence intervals for peak discharges generally ranged from 10–25 years.

Major flooding in the state was concentrated in the Napa and Russian River Basins. The Russian River crested near Guerneville with a peak discharge of 85,800 cfs; it had a recurrence interval of 10–25 years. The Napa River crested near Napa with a peak discharge of 29,600 cfs; it had a recurrence interval of 10–25 years. Approximately 1,000 homes were flooded in Napa.

Sonoma Creek crested at Agua Caliente with a peak discharge of 17,600 cfs; it had a recurrence interval that was greater than 100 years. The Klamath River crested near Klamath with a peak discharge of 416,000 cfs; it had a recurrence interval of 25–50 years.

Heavy rainfall fell in the San Joaquin Valley from January 1–2. Continuous rain, heavy at times, brought an abnormally high 2.84 inches of rain to Fresno during the January 1-2 storm event; 3.19 inches in Selma, and 2.25 inches at Coalinga in a little over 24 hours. Fresno set a new daily precipitation record of 1.88 inches on January 2. Flooding occurred in the city of Fresno as 15 ponding basins overflowed. Over 150 houses were damaged within the Fresno County.

Rainfall in excess of 2.5 inches in just over 30 hours on January 1–2 led to water-covered roadways in Kings County. Hanford measured 2.82 inches of rain in a 30-hour period while Lemoore and Corcoran received just over 3 inches. Ponding basins overflowed in Lemoore, and flooding occurred in Huron and Corcoran.

Consistent rains led to more than 3 inches of rain in a 30-hour period from mid-day on January 1 to the evening of the 2nd around Visalia and over 3.5 inches of rain in the city of Tulare. Over 2 feet of water flooded portions of West Visalia as well as flooding just east of Tipton. Tulare County had 45 homes damaged to some extent by flooding. Some of the flooding in Visalia was due to detention basins and pumps failing to perform as expected.

Over 7.5 inches of rain was reported during the storm event at the 2,000 foot elevation in the Tulare County foothills. Strong winds in the Tulare County mountains felled several large trees on January 2. Among those were:

- The second-largest limb on the General Sherman Tree
- The Telescope Tree, a large, hollow giant sequoia on the Congress Trail in Giant Forest
- A large tree that fell onto the Runciman cabin in East Mineral King

Strong wind events were reported elsewhere in the San Joaquin Valley during the January 1–2 storm event:

- During the late morning and early afternoon of January 1, gusty southeast wind commonly hit over 40 mph in the northern portions of the central San Joaquin Valley.
- On the afternoon of January 2, strong wind in the eastern part of Fresno County blew down trees and power lines, leaving over 60,000 customers without power.
- During the mid-afternoon hours of January 2, strong southeast wind in the Taft area in southwest Kern County resulted in downed power poles.
- On the evening of January 2, gusty wind caused damage around the community of Oakhurst and blew down numerous trees.
Gary Sanger at the NWS forecast office in Hanford said that the January 1–2, 2006 events likely were dominated by frontal winds. However, there may have been unreported embedded thunderstorms (and thundersnow) in the cold front's convective band. See the section of this document that describes the 1941 Wind Event for more detail about strong winter winds capable of causing forest blowdowns.

Rainfall totals of almost an inch in less than 24 hours at Bakersfield resulted in significant water flows from nearby mountains onto the south San Joaquin Valley floor. Flooding occurred on Highway 33 north of Highway 46 on the morning of January 2 on the west side of the valley. The storm dropped significant snow at higher elevations. Lodgepole and Mineral King reported 36 inches of new snow between January 1–2. Charlotte Lake received 41 inches of new snow, and Farwell Gap had 72 inches (6 feet) during the same period.

Table 99 summarizes the damages incurred during the January 1–2 storm event.

<table>
<thead>
<tr>
<th>County</th>
<th>Property Damage (million dollars)</th>
<th>Agricultural Damage (million dollars)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Fresno</td>
<td>$1.5</td>
<td>not assessed</td>
</tr>
<tr>
<td>Kings</td>
<td>$0.1</td>
<td>$1.0</td>
</tr>
<tr>
<td>Tulare</td>
<td>$5.72*</td>
<td>not assessed</td>
</tr>
<tr>
<td>Kern</td>
<td>$0.025</td>
<td>not assessed</td>
</tr>
<tr>
<td>Total</td>
<td>7.345</td>
<td>$1.0</td>
</tr>
</tbody>
</table>

*$220,000 of the property damage in Tulare County was attributed to falling trees, including one at Mineral King.

In April, there was another round of storms which combined with snowmelt to create more significant flooding over a greater area. As a result of severe storms, flooding, landslides, and debris flows, a major federal disaster (DR-1646) was declared on June 5, 2006, for the period March 29, 2006 – April 16, 2006. It covered 17 Northern California counties. Minor flooding occurred throughout the San Joaquin River and Tulare River Basins.

The peak day natural flow at Pine Flat on the Kings occurred on April 5. The flood surge lasted three days, but was unremarkable by Kings River standards. This was followed by sustained high flows from May through June. Although the river didn’t reach a particularly impressive height, it apparently delivered more than the average amount of water to the valley floor.

A cloudburst occurred above Silver City in the Mineral King area on or about July 20, 2006. It resulted in a flash flood on the creek west of Silver City. That creek is locally known as Silver Creek or Silver City Creek. (This is the same drainage that had a much smaller flash flood in 1997.) The 2006 flood damaged some of the cabins in Cabin Cove and washed out the Mineral King Road (multiple photographs on file in the national parks). Tony Caprio and Joel Despain (the national parks’ fire ecologist and geologist, respectively) recalled that the flood eroded these low-gradient stream channels down nearly to bedrock. This level of erosion implies a very high stream discharge far from the statistical norm, an event that would occur only rarely. The effect of the storm was very localized. There was little effect east of Silver City (for example, High Bridge) or west of Deadwood Creek, with the latter having only a moderate increase in flow.

Total flow for water year 2006 was 173% of the 1894–2014 average for the Kings, 167% for the Kaweah, 149% for the Tule, and 152% for the Kern. In order to minimize flooding in the Tulare Lakebed, 29,000 acre-feet of river floodwater was pumped into the Friant-Kern Canal and routed to the Los Angeles area in 2006.

**Debris Flow: Cement Table**

The national parks do not know precisely when this event occurred. It may or may not have occurred during the floods of 2005. David Karplus, Kings Canyon National Park’s trails supervisor, discovered it in 2008, and at that time it appeared to have been there for a couple of years at least. He knows that it was not there in 1999.

The debris flow began on the east side of Cloud Canyon, and flowed across the trail and into the creek. It started over 1,000 feet up the canyon wall (UTM 36S105E 4060830N NAD83 Zone 11). The debris flow was about 100 yards wide and consisted of two channels where it crossed the Colby Pass Trail, each channel about 15 feet deep. It eroded down to bedrock in many places. Rather than fill in the eroded trail, David simply had his crew dig a trail across the new mud slope.
2007–09 Drought

As described in the section on Megadroughts since the Little Ice Age, California’s 2007–09 drought is part of a longer-term megadrought across most of the Western U.S. since 2000. Because we are on the edge of that huge drought system, we tend to only be aware of it when it reaches out to encompass our area. The San Joaquin River Basin has only been affected by that megadrought for 12 of the years that it has been active: 2000–04, 2007–09, and 2012–15 (based on the San Joaquin Valley Water Year Index and/or total runoff for our four major rivers). From our perspective, we tend to think of those events as individual droughts of relatively average duration instead of being part of the larger megadrought.

Table 106 shows how the San Joaquin Valley Water Year Index categorized the drought years in that basin. It also shows total runoff for the four major rivers in our basin (Kings, Kaweah, Tule, and Kern) combined during this drought. Average flow over the three years of the drought was 62% of the 1894–2014 average.

Water years 2007 and 2008 clearly constituted a multi-year drought; that is obvious from looking at Table 106 or Figure 25. The runoff was so low in those years that the state’s water year index rated those years as critically dry. From a hydrologic standpoint, the drought lasted only through 2009, the last year of below-normal flows.

California experienced three consecutive dry years during 2007–09. Those years also marked a period of unprecedented restrictions in State Water Project (SWP) and federal Central Valley Project (CVP) diversions from the Sacramento–San Joaquin Delta to protect listed fish species. Exports from the Delta in recent decades have adversely affected the health of the Delta and the health of the San Francisco Bay.

San Francisco Bay is a huge and valuable estuary that depends on inflow of freshwater. Because of exports from the Delta, it has experienced a 350% increase in the frequency of “very dry” years in inflows. The bay is now effectively in a persistent, man-made drought. Decades of monitoring and scientific research have shown that reduced freshwater inflows are a major cause of habitat degradation and declining fish populations in the estuary: since the 1970s, populations of many of the most common species have plunged by 66-98%.1671, 1672

Statewide hydrologic conditions overall were not as severe during 2007–09 as compared to prior droughts of statewide significance. Water years 2007–09 were the 12th driest three-year period in the state’s measured hydrologic record, based on DWR’s 8-station precipitation index. That means that the state experienced 11 three-year periods during the 20th century that were more severe than the 2007–09 drought.

Table 100 compares the 2007–09 drought with more severe droughts of the 20th centuries.

<table>
<thead>
<tr>
<th>Drought Period</th>
<th>Sacramento River Basin Runoff (million acre-feet)</th>
<th>% of average 1901–2009</th>
<th>San Joaquin River Basin Runoff (million acre-feet)</th>
<th>% of average 1901–2009</th>
</tr>
</thead>
<tbody>
<tr>
<td>1929–34</td>
<td>9.8</td>
<td>56%</td>
<td>3.3</td>
<td>56%</td>
</tr>
<tr>
<td>1976–77</td>
<td>6.6</td>
<td>38%</td>
<td>1.5</td>
<td>26%</td>
</tr>
<tr>
<td>1987–92</td>
<td>10.0</td>
<td>57%</td>
<td>2.8</td>
<td>48%</td>
</tr>
<tr>
<td>2007–09</td>
<td>11.2</td>
<td>64%</td>
<td>3.7</td>
<td>63%</td>
</tr>
</tbody>
</table>

DWR provided detailed information about the 2007–09 drought in the following report:


A system deposited 0.08 inch of rain in Bakersfield on May 27, 2008. That was the only measurable rain to fall in the entire March-May period in that city and tied 1992 for the driest meteorological spring on record. The 2007–09 drought appears to have been focused on the Tulare Lake Basin. The only counties that proclaimed a local emergency during the drought were Fresno, Tulare, Kings, Kern, and Riverside. A DWR report said that the 2007–09 drought impacts were most severe on the west side of the San Joaquin Valley.1676

A 2011 Pacific Institute report summarized the impacts of the 2007–09 drought on agriculture in Fresno, Tulare, Kings, and Kern Counties. Drought-period impacts in those counties were most apparent in acreage, but also to some degree in yield and production values. Observing changes in acreage, production, and values in those four counties over the entire drought period, and in comparison to average and wet years (2000 and 2006), supported the conclusion that agriculture in the Tulare Lake Basin suffered short-term losses (moderated in part by crop shifting), yet managed to keep acreage, yield, and gross revenues relatively steady overall.
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The analysis of San Joaquin Valley counties that stood to be most affected by drought-induced water supply reductions showed that the valley agricultural sector experienced declines in terms of acreage, gross revenue, and some crop yields between 2007–09. Acreage declines were most dramatic in 2009, yet yields of top crop commodities and total production values remained steady or increased during the drought in the Tulare Lake Basin. Overall, gross revenues for all four counties were higher between 2007–09 than between 2000–06. Even in 2009, Kern and Kings Counties’ gross revenues were only 2% less than values in the most recent wet water year of 2006; Fresno and Tulare’s were both 4% higher in 2009 than 2006.

This was the first drought since the 1920s–30s during which locally significant impacts due to economic recession and drought resulted in emergency social services response actions (food banks and unemployment assistance). Governor Arnold Schwarzenegger proclaimed a state of emergency on June 12, 2008, recognizing the onset of the drought. The 2007–09 drought was California’s first drought for which a statewide proclamation of drought emergency was issued. That turned out to be critical. When precipitation returned to above-average conditions (see Table 106), it was hard politically for the governor to declare an end to the drought. There clearly wasn’t enough water to go around.

The meteorological drought had ended, but the drought had turned into a socioeconomic drought. Our basin relies on a great deal of supplemental water imported from the San Joaquin River (via the Friant-Kern Canal) and from the Delta via the state and federal canals. Reduced imported supplies can stimulate a socioeconomic drought even when precipitation and runoff in the Tulare Lake Basin is not in a meteorological drought. It is more challenging to mark the end of this type of drought. Precipitation can return to average or even above-average conditions, but there still isn’t enough water to meet our needs (the amount of water we choose to apply). See the section of this document on What Constitutes a Drought for a description of the different types of droughts.

Other droughts had been declared over when water conditions returned to near-average conditions. That didn’t seem possible in this drought. It wasn’t until March 30, 2011, after an incredibly wet winter, that Governor Jerry Brown issued a proclamation rescinding the state of emergency. The drought had finally met an official end. That was long after the end of the hydrologic drought.

However, there still wasn’t enough water to go around. There was still a significant groundwater overdraft. There were still signs posted along some of our highways expressing the opinion that the drought was caused by Congress. The signs are still there, and the groundwater overdraft persists. For more about these issues, see the section of this document that describes the Groundwater Overdraft.

As explained in the section on Groundwater Overdraft, water users have come to rely on the groundwater aquifer more and more, especially during droughts. During the four year period between April 2006 and March 2010, water users in the Tulare Lake Basin used a huge amount of groundwater. Not all of those four years were drought years. The San Joaquin Valley Water Year Index categorized water year 2006 as a wet year; the drought didn’t move into the valley until 2007.

Apparently there are two separate estimates of the size of that water withdrawal:

- In February 2011, the University of California Center for Hydrologic Modeling at UC Irvine, estimated that based on satellite data, the groundwater loss was more than half the size of Lake Mead (19.5 million acre-feet), the third largest decline in 50 years.

- In 2012 Bridget Scanlon and her colleagues at the University of Texas published what appears to be a separate analysis of the same four-year period. According to news accounts, Scanlon found that water users in the Tulare Lake Basin had used enough groundwater to fill Lake Mead, the nation’s largest man-made reservoir.

Role of the Endangered Species Act in Reducing Delta Exports

The major difference between the 2007–09 and prior droughts was the severity of SWP and CVP delivery reductions, which began immediately in the first year of the drought. During the drought, there was considerable controversy around the role that environmental protections, particularly the federal Endangered Species Act, played in the reduced exports to south-of-Delta water users.

The Sacramento-San Joaquin Delta ecosystem has been in declining health for a number of decades. One of the major causes of this is the amount of fresh water removed from the Delta and exported south. The impacts of those exports are greatest during droughts. Many of the animals that live in the Delta (salmon, steelhead, sturgeon, striped bass, Delta smelt, sea lions, etc.) rely on fresh water flowing through the Delta.
The Delta smelt has been listed as a threatened species since 1993. In 2008, the USFWS issued a biological opinion on the Long-Term Operational Criteria and Plan (OCAP) for coordination of the SWP and CVP.\textsuperscript{1681} The USFWS determined that the continued operation of those two water projects, as described in the plan, was likely to jeopardize the continued existence of the Delta smelt and adversely modify its critical habitat.

The biological opinion found that there were reasonable and prudent alternatives to the proposed operating plan, and the water agencies accepted those alternatives. After the USFWS opinion, what followed were restrictions on pumping of water from the Delta to farmers and cities south of the Delta, especially during droughts. These restrictions on exports ensured minimum flows of fresh water in the Delta to maintain salinity standards and support the ecosystem including the critical habitat of the Delta smelt.

South-of-Delta water users objected to these pumping restrictions. They generally felt that it was more important to send water to the arid south than to protect the Delta smelt and the ecosystem that it lives in. Farmers and cities in the San Joaquin Valley and Southern California generally tried to simplify the argument to the needs of a small fish versus the needs of people. If only the question were that simple.

The debate was reported as “farms vs. fish,” but the actual role of regulation in affecting water supplies was far more complex. Several pieces of federal and state environmental legislation affect Delta pumping by providing a network of protections for human and environmental health. In 2009, the final year of the 2007–09 drought, Delta exports were reduced by about 40%. Analyses from the California Department of Water Resources and the Congressional Research Service showed that over three-quarters of the reductions in Delta exports (1.6 million acre-feet) was due to drought conditions and that less than a quarter (0.5 million acre-feet) was due to environmental protections such as protecting endangered fish and maintaining Delta salinity standards.\textsuperscript{1682, 1683}

Water that flows out of the Sacramento-San Joaquin Delta has sometimes been characterized as “wasted to the sea.” The Public Policy Institute of California analyzed SWRCD data to identify the components of that water for water year 2014, a major drought year. There was 11.4 million acre-feet (maf) of water available within the Delta that year. Cities and farms that diverted water upstream of the Delta, along with Delta farmers, used 5.4 maf, well below average. Just 1.9 maf of Delta water was exported, the lowest volume in decades. Roughly 4.2 maf flowed to the sea, a near-record low. Table 101 shows the components of that water.

<table>
<thead>
<tr>
<th>Components of outflow</th>
<th>Outflow to the sea (acre-feet)</th>
<th>% of total outflow</th>
<th>% of total water in Delta</th>
</tr>
</thead>
<tbody>
<tr>
<td>Required for salinity control for urban and farm water</td>
<td>3,000,000</td>
<td>71%</td>
<td>26%</td>
</tr>
<tr>
<td>Required for fish habitat</td>
<td>750,000</td>
<td>18%</td>
<td>7%</td>
</tr>
<tr>
<td>Uncaptured storm flows (pumps lacked capacity to export)</td>
<td>450,000</td>
<td>11%</td>
<td>4%</td>
</tr>
<tr>
<td>Total</td>
<td>4,200,000</td>
<td>100%</td>
<td>100%</td>
</tr>
</tbody>
</table>

Although the protection of endangered fish has been pointed to as the cause of reduced water deliveries, there are actually a variety of water quality and statutory obligations that contributed to restricting Delta exports. The Congressional Research Service concluded that even if one piece of environmental legislation were waived or overridden (e.g., the federal Endangered Species Act), federal and state agencies would still be required to comply with other state and federal laws and directives that protect the environment, including the federal Clean Water Act, the state Porter-Cologne Act and its implementing directive D-1641, the California Endangered Species Act, the California Fish and Game Code, and the Central Valley Project Improvement Act.\textsuperscript{1684}

The history of these obligations pre-dates the federal Endangered Species Act. Delta salinity standards, for example, arise from the original water rights provided by the State of California to the SWP and CVP to divert water upstream of the Delta, thereby raising the salinity of water used by in-Delta users. The 2007–09 drought issues serve as an expression of the nature of this extreme conflict, and its problematic, unresolved status.\textsuperscript{1685}

Several water agencies brought a lawsuit challenging the U.S. Fish and Wildlife Service (USFWS) over protections for the endangered Delta smelt and the larger question of water flow through the San Francisco Bay and Sacramento-San Joaquin Delta.\textsuperscript{1686, 1687} They questioned whether limits on pumping water to the southern part of the state were required under the Endangered Species Act. They said that the restrictions were particularly harmful to consumers, farmers, and other water users during the drought.
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That lawsuit was eventually heard by the U.S. Court of Appeals for the Ninth Circuit. In its March 2014 decision, the 9th Circuit reasoned that the USFWS’s duty is “to halt and reverse the trend toward species extinction, whatever the cost.” The court concluded that the USFWS could curtail water deliveries to farms in order to protect the Delta smelt under the Endangered Species Act, without regard to the human or economic cost.

“The law prohibits us from making such fine utilitarian calculations to balance the smelt’s interests against the interests of the citizens of California.” The court concluded that the USFWS’s actions were “reasonable and prudent.” In addition to protecting the Delta smelt, the 9th Circuit has upheld limits on pumping to protect several species of salmon which migrate through the San Francisco Bay and the Delta.

The water agencies appealed this Delta smelt decision to the U.S. Supreme Court. On January 12, 2015, the justices turned down that appeal. The Supreme Court did not rule on the merits of the case, and no oral arguments were heard. The court’s action had the effect of upholding the decision of the 9th Circuit.

The Delta smelt represents a large number of other fishes that live in the Delta such as chinook salmon, steelhead, longfin smelt, and green sturgeon, all of which require a functioning estuary. The three-inch-long Delta smelt is not more important than any of those other species. It is not a commercial fish or even a sport fish. However, it is considered a key indicator of the health of the entire estuary; it functions as the canary in the Delta coal mine. If the Delta smelt population were to collapse, that would be an indication that the estuary is not working very well to support the other fish species.

Different views have been expressed about the value of the Delta smelt and protecting endangered species from going extinct. Congressman George Radanovich described the Delta smelt as a “worthless little worm that is going extinct.” Congressman Devin Nunes described it as “their stupid little fish, their little Delta smelt that they care about.” Former Alaska governor Sarah Palin said “where I come from, we call that bait.”

DWR conducts a Kodiak trawl survey each spring designed specifically to catch Delta smelt in the reaches where they gather during spawning. The March 2015 survey caught just six fish and the April survey caught just one. UC Davis biologist Peter Moyle said that the 2012–15+ drought has stressed the species to the brink of extinction. This collapse was part of a long-term decline that has been made worse by the drought.

Water exports from the Delta are limited by two factors: the amount of available water and whether the pumps can run. The pumps are sometimes required to shut down because endangered fish species or exotic water hyacinth plants are too close. The water hyacinths only affect the federal pumps, not the state pumps.

Moyle said that even if the Delta smelt were declared extinct in the wild, it would be unlikely to significantly affect the amount of water exported to south-of-Delta users. Actions taken to protect the Delta smelt are also needed to protect other endangered species or needed to prevent the intrusion of salinity. The release of large amounts of fresh water into the Delta during the summer to prevent salinity from intruding is done primarily for economic reasons: to protect irrigation water for farms in the Delta and to protect water being pumped out for cities. Those water releases would largely be required even if there were no Delta smelt.

The main effect of the extinction of the Delta smelt is that the pumps in the south Delta would no longer have to shut down when the smelt get too close. However, the number of Delta smelt have been so low in 2013–15 that they have not caused any of the pumps to shut down.

2007 Flood
Flooding in 2007 occurred during October. This flood happened during the 2007–09 drought.

The National Weather Service rated this event as one of the largest severe weather outbreaks on record in interior Central California. It occurred during the afternoon and evening hours of October 29. An upper-level low moving inland across Central California interacted with a surge of tropical moisture, triggering thunderstorms that produced hail in many places as large as one inch in diameter and gusty winds as well as locally drenching rains. Hardest hit was the northwest side of Fresno where rainfall totals of one to two inches were reported and a number of streets flooded quickly during the evening rush hour resulting in a good many stalled vehicles. Some streets in northwest Fresno were still covered with several feet of water nearly four hours after the thunderstorms had ended. In addition, hail up to an inch in diameter fell. The combination of the heavy rain and hail resulted in the collapse of the roof on an 80,000 square foot warehouse. Thunderstorm winds also knocked out power to 18,000 customers in Fresno.
Two houses in Visalia had trees fall on them, and about 200 boats were damaged at a boat dock on Lake Kaweah. Downed trees were reported in the valley from Merced County to Tulare County and eastward into the Sierra at Yosemite and Sequoia National Parks.

**2008 Flood**

Flooding occurred in July, primarily from July 12–15. This flood occurred during the 2007–09 drought.

It was caused by a series of individual storm events, but could be thought of as one event that occurred in multiple locations.

These storm events were caused by the North American Monsoon. From July 7–11, high pressure centered over the Four Corners area dominated the weather over the Southern Sierra. By July 12, a major low pressure area had formed in the Southwest. In addition, moist air influenced by Hurricane Bertha east of Bermuda began to reach the Southwest. Conditions were also changing in California. An upper-level high pressure ridge moved inland off the Pacific on July 11, with a low pressure area moving along the coast. This pattern set up a southerly wind pattern over California, drawing up monsoonal moisture from the southeast on July 12. As the air flow from the southeast brought in this moist air, thunderstorms formed over the Southern Sierra. Thunderstorms formed over the Tulare County mountains by early on the afternoon of July 12 and remained in the region through July 15.

Following is a sample of the storm events that occurred during the July 12–15 period.

**Mud flows: Tioga Pass Road**

On July 14, the Tioga Pass Road in Yosemite National Park was closed due to mud flows across that road. Thunderstorms dropped a lot of hail, resulting in pea-sized hail covering about a two-mile section of the road.

**Debris Flow: Oak Creek**

This debris flow was generated within Oak Creek, an east-flowing drainage near Independence. Oak Creek is within the Inyo National Forest. That is outside the Tulare Lake Basin. However this event merits inclusion in this document because of the similarities of this debris flow to the one that would occur in the nearby Lewis Creek Basin just two days later.

Like Lewis Creek, the bedrock in the Oak Creek Basin is granitic. Perhaps there are some soils in the upper elevations that are composed largely of granitic sands. However, most of the bedrock is apparently overlaid with deposits of alluvium, basaltic lava flows, glacial outwash and moraines, landslide deposits, alluvial fan deposits, and colluviums. Whatever the specific soil mix, the rapid rainfall apparently saturated this granular soil, reducing its frictional strength, and causing the soil mass to begin moving downslope as a flowing mass.

Late on the afternoon of July 12, a large convective cell centered over Oak Creek Canyon produced a brief period of intense rainfall. While debris flows are sometimes triggered by a prolonged rain or snowmelt event, the Oak Creek debris flow was triggered by this short, intense rainfall event. The debris flow started within an hour or so of the onset of that cloudburst. The rain started sometime after 4:00 p.m. We don’t have any direct precipitation data for that thunderstorm because the RAWS automated weather station located upstream from the South Fork Oak Creek junction was swept away by the debris flow just before its scheduled transmission at 4:42 p.m., less than an hour after the onset of the rain.

The Oak Creek debris flow was initiated in the headwaters of the North and South Forks of Oak Creek at about elevation 10,825 feet (3,300 m). It then traveled about 11 miles (18 km) before coming to rest on the floor of Owens Valley at about elevation 3,840 feet (1,170 m).

The debris flow that came down the North Fork of Oak Creek deeply eroded the existing creek channel through the Oak Creek Campground and severely damaged the nearby road. In the campground, a camper escaped his collapsing motor home when it became entangled in a grove of trees. He jumped into the muddy debris flow and “surfed” several hundred yards until he eventually reached stable ground where he could walk to help. This fortunate individual reported experiencing three distinct waves or surges of material with the larger one being between 6–12 feet (2–4 m) high. The velocity of that larger wave would have been about 12.5 mph (20 km/hr).
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The debris flow that came down the South Fork of Oak Creek destroyed the Bright Ranch, including the main house. Fortunately the inhabitants were away at the time that the debris flow struck. That ranch had been occupied since 1872 and had never been impacted by a debris flow. That suggests that this debris flow had a long recurrence interval.

The watersheds serving as the source for the Oak Creek debris flow were burned during the Inyo Complex Fire in July 2007. There is circumstantial evidence that this might have created the conditions necessary for this debris flow. Presumably equally severe thunderstorms had occurred in the Oak Creek Basin since 1872. Yet this was the first time that the Bright Ranch had been impacted by a debris flood. That raises the possibility that the effect of the Inyo Complex Fire on this watershed contributed to this unusual event.\(^{1699}\)

The South and North Forks converge at about 5,250 feet (1,600 m). Just east of the junction of these two tributaries, the debris flow struck the historic Mt. Whitney Fish Hatchery. The buildings survived, but all the fish were wiped out.

In the reach below the fish hatchery, 17 homes were damaged or destroyed. By about 5:30 p.m., most people along Oak Creek had abandoned their homes and only one person had to be rescued from a rooftop. Fortunately no one was seriously injured or killed.

The debris flow blocked the drainage structure under U.S. Highway 395, forcing the flow to mobilize over the top of the roadway. This resulted in closure of the highway for five hours, from 6:30 p.m. – 11:45 p.m. It took Caltrans nearly a week to fully restore the highway.

After crossing the highway, the debris flow damaged another 25 homes on the Ft. Independence Indian Reservation.

The debris flow had a total length of about 11 miles and a drop of 6,985 feet in elevation. It had an average gradient of 635 feet/mile. That averages the steeper mountain section and the flatter section lower down. The peak velocity measured was about 12.5 mph (20 km/hr).

The debris flow had a total estimated volume of about 2.04 million cubic yards (1.56 million cubic meters). This is the largest Sierra debris flow ever definitively measured.

**Debris Flow: Erskine Creek**

There were multiple bouts of flash flooding in the Kern County foothills and mountains during the July 12–15 time period.\(^{1700}\)

The Piute Fire was a major wildfire that began south of Lake Isabella on June 28, 2008. Precipitation information was available for this area from RAWS automated weather stations because the Lake Isabella area had experienced flash floods in the past and also because it is near a major reservoir. In addition, the firefighting efforts directed at the Piute Fire placed a temporary RAWS at Piute Peak to assist with fire weather forecasting. An NWS meteorologist was assigned to the firefighting incident, giving us a good record of the events that would unfold.

Heavy rain hit the Piute Fire area for three days in a row: July 12–14. On July 15, heavy rain occurred just south of the fire area. The rain was sometimes quite intense. For example, on July 15, a Claraville weather station (due south of Lake Isabella) reported 2.15 inches of rain in 90 minutes (a rate of 1.43 inches per hour).

The intense rain caused at least seven debris flows and flash floods in the area plus one in the town of Tehachapi. The most impressive of those was the July 12 Erskine Creek debris flow. Erskine Creek would also experience what the NWS characterized as debris flows on July 13 and July 14.\(^{1701}\) Erskine Creek is within the Sequoia National Forest. The July 12 Erskine Creek debris flow was analyzed by Jerry DeGraff and others.\(^{1702}\)

When thunderstorms began to develop late on the afternoon of July 12, the NWS weather specialists on the fire team and at the NWS forecast office in Hanford recognized the high likelihood of flash flooding. The Piute Fire Unified Command issued a flash flood warning and recommended evacuation notice for Erskine Creek. As the storm developed, a helicopter operated by the Kern County Fire Department was dispatched for aerial observation.

The July 12 debris flow didn’t start from a prolonged rain or snowmelt event. It was triggered by a short, intense rainfall event. The debris flow started within an hour or so of the onset of that cloudburst.
Whenever a debris flow occurs after a fire, there is a tendency to assume that the debris flow was caused primarily by the fire. That is, to assume the fire was necessary to create the conditions for the debris flow to occur. The Piute Fire did contribute to the progressive bulking of the Erskine Creek debris flow. But otherwise, there was no clear-cut relationship between that fire and the debris flow.

Debris flows are rarely seen, let alone photographed. However, this one happened to occur where an incident command post was already established; one that had both good meteorological data and helicopter support. It also occurred in daylight. That combination made all the difference.

The debris surge had multiple surges. The first surge was the largest and darkest one, containing abundant ash from the burned slopes. The helicopter spotted the debris flow in the upper watershed and followed the leading edge through the town of Lake Isabella, capturing dramatic footage (video on file in the national parks).

The Erskine Creek debris flow was generated by flows from within the South, Middle and East Forks of Erskine Creek.

Like Lewis Creek and Oak Creek, the bedrock in the Erskine Creek Basin is granitic. Perhaps there are some soils in the upper elevations that are composed largely of granitic sands. However, most of the bedrock is apparently overlaid with deposits of landslide deposits, colluviums, and alluvial fan deposits. The town of Lake Isabella is built on the alluvial fan deposits.

The longest tributary, the South Fork, flows about 6.5 miles (10.5 km) from its headwaters at about 8,185 feet (2,495 m) to the junction with the other two tributaries at an elevation of 4,395 feet (1,340 m). From there, Erskine Creek flows another 8.3 miles (13.4 km) to its junction with the Kern River at an elevation of 2,445 feet (744 m).

The debris flow had a total length (including the South Fork tributary) of 14.8 miles (23.9 km) and a drop of 5,740 feet in elevation. It had an average gradient of 388 feet/mile. That averages the steeper mountain section and the flatter section lower down.

The portion of the debris flow that was on the South Fork tributary had a length of 6.5 miles and a drop of 3,740 feet in elevation. That section had an average gradient of 583 feet/mile.

The gradient along the flow path was about 23%. The peak velocity measured was about 12 mph (19 km/hr).

Because the debris flow entered and dispersed within the Kern River, there is no accurate estimate of its volume. On the basis of the affected area, it would seem comparable in size to the Oak Creek debris flow which had a total estimated volume of about 2.04 million cubic yards (1.56 cubic meters).

There was an initial and natural temptation to associate the three Erskine Creek debris flows / flash floods with the Piute Fire. That is, to assume that the debris flows were caused in some way by the fire. However, as Jerry DeGraff and others showed in their analysis, that was not the case with the Erskine Creek debris flows.

However, in 1984 the Lake Isabella area did experience a major debris flow that might have been due in part to a wildfire. That was in the Goat Ranch Canyon and Long Canyon areas (see the section of this document that describes the 1984 flood).

Erskine Creek flash flooded on three consecutive days, July 12–14. When the July 12 debris flow swept into the town of Lake Isabella at about 6:30 p.m., Erskine Creek overflowed Lake Isabella Blvd (video on file in the national parks). Because of the advance notice, crowds had gathered on each side of the barricaded area to watch this dramatic event. This was probably the most viewed and photographed debris flow ever in the Tulare Lake Basin.

Erskine Creek would overflow Lake Isabella Blvd. again on July 13 and briefly on July 14. Each time, the Piute Fire Unified Command issued a flash flood warning and recommended evacuation notice far enough in advance so that the road could be closed to traffic.

The debris flood of July 13 was very powerful. In places, the floodwaters were 100 yards wide and 18–24 inches deep. Many people had heeded the flood warning and evacuation notice. Up to 80 homes in the Erskine Canyon area were evacuated. But still, despite the warning — and the experience of the previous day — some people
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were caught unprepared. Helicopter crews from the Kern County Fire Department rescued two families, consisting of a total of seven people and two dogs, from their homes along Erskine Creek. One of those families and their dog had to be plucked off their roof.\footnote{1708}

The three debris flows / flash floods in the town of Lake Isabella resulted in $1.5 million in property damage.\footnote{1709}

Erskine Creek empties into the Kern River; so much of the debris from the three days of debris flows / flash floods became dispersed into the flow of the Kern and carried downstream. At a point 17 miles (27 km) downstream is Democrat Dam, a low diversion dam to divert water to the KR1 hydroelectric plant operated by SCE. The KR1 powerhouse is located another 10.2 miles (16 km) downstream. Near that powerhouse is the Kern Canyon diversion dam to direct water to the PG&E-operated Kern Canyon powerhouse at the mouth of the Kern River Canyon. Both of those plants were taken off-line to avoid damage to generating equipment and remained closed for a number of days. Additionally, habitat supporting a genetically pure strain of Kern River rainbow trout was severely damaged within the Erskine Creek watershed.\footnote{1710}

On July 15, the Bakersfield water supply was threatened by dirt and silt washing down the Kern River. A portion of the water treatment facility had to be shut down. The city had only a three-day supply of clean emergency water, and that was beginning to run out at one of the treatment plants.\footnote{1711}

**Other Debris Flows and Flash Floods: Kern County**

In addition to the three debris flows / flash floods on Erskine Creek, there were a number of other flash floods and debris flows in the Kern Mountains during the July 12–15 period. The ones that we are aware of were:

- On July 12, a debris flow passed down Thompson Creek, a tributary of Walker Basin Creek. This event was analyzed by Jerry DeGraff.\footnote{1712} Thompson Creek Basin adjoins that of Erskine Creek. Like the South Fork of Erskine Creek, the head of this debris flow was near Piute Peak where the thunderstorm cell was centered. Sediment from the Thompson Creek debris flow was evident a distance of about 12.5 miles (20 km) downstream where Kern County Highway 483 crosses Walker Basin Creek. A number of residential structures were impacted by this debris flow.

- On July 12, a debris flow passed down Clear Creek, a tributary of Havilah Canyon Creek. This event was analyzed by Jerry DeGraff.\footnote{1713} The Clear Creek Basin adjoins that of Erskine Creek. The head of the debris flow was near Piute Peak where the thunderstorm cell was centered. While the Clear Creek debris flow did not appear to pass the entire 10½ miles (17 km) to where the creek is crossed by Kern County Highway 483, the sediment from this event was visible from that point. There are no roads or other infrastructure in the bottom of Clear Creek, so there was no damage from this event.

- On the afternoon of July 14, a flash flood / debris flow occurred on Johns Rd between Caliente Creek and Walsher Rd.\footnote{1714}

- On July 14, flooding occurred in the town of Tehachapi. An apartment complex in that town sustained significant damage.\footnote{1715}

- On the afternoon of July 15, Thompson Creek Road was washed out, stranding 40 homes about 10 miles south of Lake Isabella.\footnote{1716}

**Debris Flow: Charlotte Lake**

This debris flow resulted from a cloudburst that occurred on July 14, 2008. A total of 1.73 inches of rain fell between 2:00–6:00 p.m. The intense part of the storm began at about 3:45 p.m. The first 20 minutes or more was heavy hail followed by heavy rain. George Durkee, the wilderness ranger at Charlotte Lake, looked up at about 4:20 after hearing the rain intensify for about 15 minutes and saw a flood of brown water and debris (small rocks, mud, and pine needles) coming down the hillside behind the ranger station.

The cloudburst resulted in several small debris flows and numerous gullies up to a foot deep in a 30 square mile area between Charlotte Lake and Bullfrog Lake (UTM 372965E 4072095N NAD83 Zone 11) (map and multiple photographs on file in the national parks). George attributed the huge runoff to the heavy warmer rain melting the recent hail and the combined water coming down all at once.

**Debris Flow: Lewis Creek**

According to Mel Manley and Ken Hires, Kings Canyon experienced a lot of rain over the weekend of July 12–13. A couple of those storms dropped 1½ inches of rain in 1 hour.\footnote{1717} The storms continued into the following two days as well. These were evidently localized storms. The Cedar Grove gage (CGR) recorded only 0.44 inches of rain for the July 13–16 period.

On July 14, a particularly intense thunderstorm cell appears to have been centered near Kennedy Mountain along the Monarch Divide. This date is based on a ranger weekly report which recorded the storm as occurring...
on July 14. There is some confusion about this date. A photograph taken of the ensuing flood in the Kings River has the date stamp of July 15. In addition, Bill Templin recorded two weeks later that Ken Hire, the Cedar Grove lead interpretive ranger, said that the storm happened on July 15.

Possibly some of the July 12–13 storms had included the Lewis Creek Basin; we have no way of knowing. The ground may have been relatively dry. In any case, the July 14 thunderstorm triggered a small- to moderate-size debris flow on the north side of Kennedy Mountain. We know about that debris flow because it crossed the Kennedy Pass Trail.

In addition, that storm triggered a major debris flow on the south flank of Kennedy Mountain, just west of Kennedy Pass at about 10,200 feet elevation (UTM 351797E 4081928N NAD83 Zone 11). The following description refers to this southern debris flow.

No soils map exists for this hillside. However, judging from the material that came off in this debris flow, the soil appears to be largely sand with some silt and perhaps some clay. Such granular soils depend on grain-to-grain friction for soil strength. Rapid rainfall can exceed the infiltration capacity of the soil so that it becomes saturated. This results in water filling the pore spaces of the soil and reducing the contact between individual grains. The decreased frictional contact temporarily reduces soil strength at which point the soil mass can begin to move downslope as a flowing mass. Debris flows in the Sierra are initiated when these conditions arise from intense rainfall, rain-on-snow events, or rapid snowmelt. A large debris flow can be triggered in as little as one hour from the onset of an intense rainfall event.

This southern debris flow started on an unburned slope, well above any of the fires that have burned in that area, including the 2005 Comb Fire. From there, the debris flow traveled down Lewis Creek 5–6 miles (9 km) to the valley floor. Several smaller tributary debris flows joined the main flow along the way (PowerPoint on file in the national parks).

Whenever a debris flow occurs after a fire, there is a tendency to assume that the debris flow was caused primarily by the fire. That is, to assume the fire was necessary to create the conditions for the debris flow to occur. That was the case with the 2008 Lewis Creek debris flow; it was initially attributed to the 2005 Comb Fire.

However, by using a set of pre-fire and post-fire/post-debris flow aerial images, the parks’ fire ecologist Tony Caprio was able to ascertain that the flow started above and outside the area burned by the fire. That is clearly shown in the PowerPoint that Tony created to document his research (see the PowerPoint referenced above).

The debris flow scoured the hillside, removing vegetation and eroding deeply into the ground (multiple photographs on file in the national parks). Although bedrock was exposed in some places, that doesn't necessarily mean that this event had a long recurrence interval. Jerry DeGraff, a geologist for the USFS, said that sometimes debris can fill in pretty fast after one of these events.

In any case, the debris flow scoured out rock, decomposed rock (granitic sand), and ash from the 2005 Comb Fire. It crossed the Kennedy Pass Trail in two places, causing major damage (multiple photographs on file in the national parks). A California Conservation Crew worked for 1½ months the following summer repairing that damage at a cost of $80,000.

This dramatic event was similar to a debris flow that occurred in an unnamed tributary of Lewis Creek on this hillside on May 27, 1983. The high energy debris flows that occur in the Lewis Creek Basin are somewhat reminiscent of the power and ferocity of avalanches.

No one witnessed the 2008 debris flow as it came off the hillside, so we don’t know how fast it was moving. Debris flows in the Sierra have an average peak velocity of 12.4 mph (20 km/hr). From looking at the aftermath of the 2008 event, it’s tempting to think that it was a much faster than average debris flow. In areas of very steep slopes, debris flows can reach speeds of over 100 mph.

All of the material that was scoured off the hillside in the July 2008 event was delivered to the valley floor. When the debris flow reached the gentler slopes at the canyon mouth, the event changed from scouring mode to depositional. This transition point occurred at about elevation 5,200 feet, about one mile upstream from the Lewis Creek highway bridge. That was 5–6 horizontal miles (9 km) from where the debris flow had begun. By this point, the debris flow had dropped about 5,000 feet in elevation.
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The elevation at the mouth of the Lewis Creek canyon is 5,100 feet.

Significant deposition of coarse alluvium is occurring just below the mouth of the Lewis Creek canyon, about 1,000 feet upstream from the Lewis Creek highway bridge. This deposition is the result of large woody debris (tree trunks) being caught between live trees and catching the large alluvium emanating from the upper watershed during the significant flow events over the years. The depth of this deposition has accumulated as high as 10 feet above the low point of the channel at the downstream extent, virtually filling the main channel above this location.

A log jam / debris dam 20–30 feet high formed about 200–300 feet downstream from the canyon mouth (multiple photographs on file in the national parks).

Bill Templin, Rick Hartley, and the Kaweah Flyfishers visited the lower Lewis Creek channel on July 26, 2008. Bill noted that the event had covered up all of the benthic invertebrate habitat in Lewis Creek with a heavy deposit of sediment. This material was presumably a mixture of granitic sand, silt, and possibly clay.

By the time of their visit, bear tracks had been pressed into a soft, ashen-colored sediment in the bed of Lewis Creek that appeared to be silt or clay (photograph on file in the national parks).

Bill also observed that there was a layer of fine, gray material that seemed to have been blasted onto the sides of the creek channel. It clung to the rocks and trees like a coating of shotcrete would adhere to the sides of a swimming pool (photograph on file in the national parks). Jerry DeGraff said that such fine, gray material is a common feature of debris flows when you are there soon enough after occurrences — before rainfall and wind essentially scrub it off the trees and rocks. It is not ash from fires because it appears in both burned and unburned areas. It is mineral in nature, presumably finely ground rock.

David Karplus, Kings Canyon National Park's trails supervisor, recalled that Lewis Creek experienced some very big channel changes as a result of this debris flow. The national parks have no record of what the channel was like prior to the debris flow.

David and others noticed how black or chocolate colored the flow in Lewis Creek and the Kings River was when this event occurred. Ned Kelleher (Kings Canyon district ranger) captured this in a time-lapse sequence that he took of the Kings River during the event (multiple photographs on file in the national parks). The color was presumably due to the ground rock and the mysterious silt/clay/ash component of the debris.

The 2005 Comb Fire occurred three years earlier, and the remaining ash would have been a relatively small component of the total debris that washed off the hillside. However, that ash could very well have been mixed in the early parts of the flow; it is floatable material and tends to be part of the leading edge of a debris flow.

Ned’s log recorded the debris flow on Lewis Creek and the Kings as occurring between about 3:00–5:00 p.m. The lower portion of Lewis Creek came up 3–4 feet and flushed an enormous amount of debris and sediment into the Kings River. The Kings came up over a foot within 10 minutes.

A large amount of the debris (some cobblestones, but consisting primarily of granitic sand and silt/clay/ash) washed down the Kings River as part of the debris flow. Bill Templin recalled that a thick covering of sediment completely covered the benthic invertebrate habitat as far as the old USGS gaging station below Grizzly Falls. Silt deposits were observed all the way down to Boyden Bridge and presumably continued below that. Jeff, an employee at Boyden Cave, told Bill that a large amount of woody debris floated downstream past the cave.

David Karplus recalled that deep holes in the Kings River were filled with up to about 4 feet of sediment, and a thick covering was spread over the entire riverbed. The effect was to generally level out the entire riverbed.

Rick Hartley said that the bottom of the Kings was covered with roughly 8 inches of sandy sediment, and that there were many new sandbars (multiple photographs on file in the national parks). In addition to the sand, the sediment in the riverbed contained a fine gray material that could easily be stirred up. It wasn’t clear whether that was silt or clay (photograph on file in the national parks). Since the ash is floatable, it wouldn’t have settled out with the sand and silt in the riverbed; it would have continued on downstream.
Bill Templin and others caught several fish in the Kings River above Deer Cove Creek. Those fish were ashen-colored and full of this fine, gray material inside. They contained no benthic invertebrates, only terrestrial insects in their stomachs. The female’s eggs were brown instead of orange.

Bill estimated that the sand and silt/clay debris covered the Kings riverbed thickly for about 10 miles from the junction with Lewis Creek to just below the old USGS gaging station at about elevation 4,100 feet. That would be 11 miles below where the debris flow had changed from its scouring to depositional phase. The debris flow had dropped about 1,100 feet in elevation in the 11 miles since it changed to its depositional phase.

Bill documented the effects that he observed in a July 28, 2008 email to Sequoia National Park, Sequoia National Forest, California Department of Fish and Game, and others. He also wrote it up in a newsletter for the Kaweah Flyfishers. The total length of this debris flow was about 17 miles, 6 miles in the scouring phase and 11 miles in the depositional phase. It dropped a total of 6,100 feet in elevation, most of which was in the high-energy scouring phase. The gradient during the scouring phase (5,000 feet in 5–6 miles) was between 833–1,000 feet/mile. No debris flow like this had previously been recorded in the Tulare Lake Basin.

For comparison, the nearest equivalent debris flow that we are aware of was the Oak Creek debris flow that occurred just two days earlier on July 12, 2008. The Oak Creek debris flow had a total length of about 11 miles and a drop of 6,985 feet in elevation. It had an average gradient of 635 feet/mile. That averages the steeper mountain section and the flatter section lower down. The Oak Creek debris flow occurred just east of the Lewis Creek debris flow and was caused by the same storm system. It had an equivalent length and drop. However, it had a much greater volume.

By the end of the spring 2009 runoff, the Kings River depths and cobble sizes apparent to a casual swimmer (David Karplus) had returned to the pre-event 2008 levels.

While the big thunderstorm was triggering the Lewis Creek debris flow in the north part of the national parks in 2008, another storm event was occurring farther south. At 4 p.m. on the afternoon of July 14, the Atwell / Cold Springs area in Mineral King was receiving heavy rain, resulting in flood damage to the Mineral King Road. Mud and rocks washed onto that road, forming deposits up to three feet deep, leaving the road impassable.

Lewis Creek forms a delta where it exits the canyon; that delta extends down to the Kings River. The main highway cuts across the lower portion of that delta, approximately 1,500 feet downstream from the mouth of the canyon. The primary channel of Lewis Creek flows from the mouth of the canyon straight across the delta, under the highway bridge, and into the river.

The July 2008 debris flow left an unstable channel above the highway bridge. A large amount of sediment and debris was deposited on the delta, especially along the general course of the primary channel. In addition, a log jam formed on the primary channel a short distance below where Lewis Creek exits the canyon. Within the next year or two, about 100 feet of the primary channel above the log jam filled in. During roughly this same period, a side stream developed along the northwest side of the delta, along the road to the parks’ wastewater treatment plant. That side stream gradually came to be called the Lewis Creek overflow channel. That side stream splits off at elevation 5,080 feet.

On June 6, 2010, a small debris jam formed in the primary channel of Lewis Creek upstream of the highway bridge. That debris jam resulted in diverting water flow outside the primary channel into the overflow channel. The 16 inch culvert where the overflow channel goes under the highway could not carry the resulting flow, so Lewis Creek (via its overflow channel) overflowed the highway for a couple of days and caused some minor road damage. It also caused some erosion along the road to the wastewater treatment plant. This channel-shifting happened again from about June 20–30, 2011.

In the fall of 2011, the national parks’ road crew installed two more culverts (an 18 inch and a 24 inch) to help carry the flow that results when a portion of Lewis Creek moves into the overflow channel.

Where Lewis Creek emerges from its canyon, there is very little impediment to keep it from changing course from the primary channel to the overflow channel. If the primary flow of Lewis Creek were to move to the overflow channel, then a different solution would have to be found. The three highway culverts on the overflow channel are not sufficient to carry the full flow of Lewis Creek. Either a bridge would be required, or the creek would have to be pushed back into its primary channel as was done after the May 1983 debris flow.
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The situation on this portion of the Lewis Creek channel was investigated by Bart Bergendahl, a Federal Highways Works Administration engineer. "

2009 Flood
Flooding in 2009 occurred in October. This flood occurred during what might be considered the ending months of the 2007-09 drought.

On October 8, Super-typhoon Melor struck Japan. Four days later, rain from the remnants of that powerful typhoon encountered the Sierra. Melor itself didn’t travel across the Pacific, but water vapor from the typhoon moved via an atmospheric river. Figure 30 shows that atmospheric river channeling water vapor from the decaying Typhoon Melor over the western North Pacific, across nearly the entire width of the ocean basin to the Sierra on October 14, 2009.

![Figure 30. An atmospheric river channeling water vapor from the decaying Typhoon Melor across the Pacific Ocean to the Sierra on October 14, 2009. Source: Michael Dettinger, USGS/Scripps.](image)

Moisture from the remnants of Melor arrived over the Central California interior, bringing high-elevation snowfall and large amounts of rainfall. A deep low pressure trough tapped into the moisture from the remnants of Melor on October 13-14. The impacts of the storm and resulting flood were documented by the NWS forecast office in Hanford. Many valley and Sierra locations set new record-high precipitation amounts.

At the onset of the storm during the morning of October 13, quite a bit of snow fell over the crest of the Sierra. However, once the precipitation began in earnest by the afternoon of October 13, snow levels rose to elevations of over 10,000 feet. By the end of the event on October 14, most of the precipitation was falling as rain.

Maximum rainfall totals were near 19 inches in 24 hours along the Central Coast and greater than 10 inches along the Southern Sierra. Several roads and highways along the Central Coast were closed due to flooding. Landslides were reported in the Santa Cruz Mountains.

The Dinkey Creek RAWS automated weather station southwest of Shaver Lake received the most precipitation of any Sierra site. That station received a storm total of just over 13 inches, about 9 inches of which fell in 12 hours on October 13. Dinkey Creek experienced a flash flood.
Grant Grove received 7.7 inches of rain during the October 12–13 storm event. That was significantly more precipitation than had been recorded during any two-day period of October at Grant Grove or Giant Forest in the previous 80 years. The resulting flooding caused significant erosion in the newly restored Halstead Meadow and elsewhere in the national parks.

At Three Rivers (the TRR gaging station), the Kaweah River had been flowing at 27 cfs. Within a matter of just five hours, it peaked at 20,937 cfs. (That was the peak hourly flow; the peak average daily flow was 7,360 cfs.) One source said that Lake Kaweah rose 30 feet in less than 24 hours. René Ardesch captured footage of the flood as the Kaweah River passed under the Pumpkin Hollow Bridge (video on file in the national parks).

The flood went over the top of the SCE bridge leading to the Kaweah #3 hydroelectric complex that is located just inside Sequoia National Park.

This fall flood was reminiscent of the floods that occurred in September, 1976, September 1978, September 1982, and November 2002. All of those floods were caused by the remnants of Pacific hurricanes.

The October 2009 storm caused many small-scale debris flows in the Mineral King Valley which clogged a number of the culverts on the Mineral King Road.

The storm caused significant trail damage, including to the High Sierra Trail west of Bearpaw. For example, the bank gave way under a large boulder at the Buck Creek crossing, causing that boulder to fall into Buck Creek, leaving a lot of rocks and debris on the trail.

The flood also washed out several hundred yards of the Cliff Creek Trail below Pinto Lake. The damaged sections were mainly below the area known as the Waterfalls where the trail follows the creek, and also at the trail crossing to Timber Gap. The damage was caused in large part by floodwaters running across and beside the trail. The flood deposited rocks and debris on the trail and caused some bank erosion. In some areas, the trail crew had to dig into the cut bank to reestablish the trail. The amount of trail damage from this flood was rather impressive. The trail crew repaired this section of trail in 2010.

**Debris Flow Complex: Sequoia National Park**

The worst damage on the High Sierra Trail occurred at Hamilton Gorge (UTM 359334E 4047933N NAD83 Zone 11). That section of trail was the target of one of the largest of several debris flows that impacted the Sequoia National Park wilderness during this storm event. The trail in that area passes through a granite gorge. The debris flow picked up a large mass of rocks and deposited them on the trail several hundred yards downstream. It took a trail crew a good bit of effort to repair the damage to the trail. That is the largest debris flow to occur in the Sequoia National Park wilderness in recent memory.

The Middle Fork Trail in the Kaweah River Basin was also damaged in a number of places during this storm event. Moderate-sized debris flows occurred in several places along the Middle Fork Trail.

**Debris Flow: Black Rock Pass**

The following year (2010), Tony Caprio, the national parks’ fire ecologist, observed where several large high-elevation debris flows had occurred on the west-facing slope of Black Rock Pass (from UTM 360400E 4038080N NAD83 Zone 11 south on the slope to UTM 360530E 4037950N NAD83 Zone 11). Those flows were about 15–20 feet across and had scoured from 4–6 feet deep. They were deep enough that they were an obstacle, and it took some looking to find the right place to cross. These flows looked fresh and were presumed to have been from the October 2009 storm event.

The coordinates listed above were for the points where Tony observed the debris flows, relatively near the bottom. The 2010 NAIP aerial imagery shows that the flows began roughly ½ mile upslope from there. The imagery also shows that there are several other old debris flows in the area (multiple photographs and aerial imagery on file in the national parks).

**Debris Flow: Tablelands**

Tony also observed where a large high-elevation debris flow had occurred on the Tablelands above Pear Lake (UTM 351900E 4052080N NAD83 Zone 11). The size of this debris flow was hard to judge since it just lifted mats of soil and vegetation, which were a foot or so thick, off of the bedrock and deposited them down on flatter areas (multiple photographs and video on file in the national parks).
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2010–11 Floods (4)
There were at least four periods of flooding in 2010–11:
1. January 2010 (localized flooding from five back-to-back storms, treated as 1 flood)
2. December 2010 (2 severe storms)
3. July 2011 (multiple severe storms caused by monsoonal moisture, treated as 1 flood)

The winter of 2009–10 was a moderate El Niño event. This coincided with the January 2010 flood. This may well have been a coincidence. Only strong El Niño events have been shown to have a correlation with high precipitation events and floods in California.

The winter of 2010–11 experienced one of the strongest La Niña events ever. It was stronger even than the 1955–56 La Niña event. In addition, December experienced the most extreme jet stream pattern on record for that month. These conditions coincided with the December 2010 flood and the heavy snowpack that accumulated during the remainder of the season.

Tulare County proclaimed a state of emergency on December 20, 2010, as winter storms created widespread flooding, debris and mud flows, and numerous road washouts and closures. The following day, on December 21, Governor Arnold Schwarzenegger proclaimed a state of emergency for Kern, Orange, Riverside, San Bernardino, San Luis Obispo, and Tulare Counties.

As a result of winter storms, flooding, and debris flows, President Barack Obama declared a major federal disaster (DR-1952) on January 26, 2011 for the period December 17, 2010 – January 4, 2011. It covered 10 Southern California counties including Tulare, Kings, and Kern. This was just one of 99 major federal disasters declared in the U.S. in 2011, breaking the record of 88 disasters that had been set in 2010.

The Tulare Lake Basin experienced a major precipitation event in January that consisted of five distinct storms. The storm series started on January 17 and continued for a week. This line of storms had similarities to the atmospheric river event of January–February 1998 that brought flooding to the central and southern San Joaquin Valley, but there were some significant differences.

While both events occurred during El Niño/La Niña events, the 1998 event brought warm, subtropical moisture that eroded the mountain snowpack, which appreciably added to the runoff. Although the event of January 2010 was similarly moisture-laden, surges of cold air kept snow levels low, and snow even fell on the Southern Sierra foothills on January 21–22. In addition, the spacing between the storms in the 2010 event allowed for several-hour breaks between the first few storms, enabling the ground to absorb some of the moisture before the runoff from the next storm hit. There was still some flooding in the Tulare Lake Basin in 2010, but most of the flooding events were either due to clogged storm drains or occurred in normally flood-prone areas.

The first storm moved rapidly through the area late on January 17. It brought rain but no flooding.

The second storm arrived on January 18. Strong winds over the Tehachapi Mountains ahead of that storm felled a tree onto a house, killing the occupant. This would be the first of two fatalities in this week-long storm event. Later that day, a brief tornado formed southwest of Fresno.

The third storm arrived on January 19, and an upper-level disturbance rotating around the low moved into Southern California, bringing snow to the Tehachapi Mountains. Snow levels with this storm were down to about 4,000 feet. Runoff resulted in some road flooding, and creeks in rural areas ran high. The second storm-related fatality occurred that evening when a man drove around barriers in an attempt to cross a flooded road near the Merced County-Stanislaus County line. That road had been flooded by Orestimba Creek, and the driver was swept away by the fast current.

The heavy rain associated with the third storm brought flooding to portions of Fresno, Kings, and Kern Counties on the afternoon of January 19. Among the flooded roads were:
- Highway 33 and Merced Avenue southeast of Coalinga
- The southbound lanes of Interstate 5 from Highway 269 (Lassen Avenue) to the Fresno County line (or so this section of highway was described)
- The Herring Road exit on Highway 99
- Highway 33 from Highway 46 to the Lerdo Highway
The fourth storm passed through the region late on January 20. It dropped snow on the Grapevine and triggered isolated early evening thunderstorms.

The fifth and final storm of the series arrived on January 21, bringing very cold air to the region. Snow levels dropped to about 2,200 feet, with snow falling in the towns of Mariposa and Oakhurst. Both the Grapevine and Tehachapi Pass were closed for several hours by the snow. This was a very deep low pressure system, and all-time low pressure records were set in both Bakersfield and Fresno; both reported a barometric pressure of 28.94. This storm also brought very strong winds to the Kern County Mountains where 10-inch diameter tree branches were downed in the Grapevine area due to wind gusts approaching 70 mph. Arroyo Pasajero flooded, closing Highway 269 (Lassen Avenue) north of Huron at 8 p.m. on January 21.

A cold pool of air moved over the area on the afternoon of January 22, triggering strong convective showers, one of which blanketed parts of the city of Clovis with about 2–3 inches of pea-size hail on the ground. A funnel cloud was reported near Clovis Ave and Highway 168 in Clovis.

By the time the last storm moved east of the region, the total rainfall amounts in the central and southern San Joaquin Valley were mostly between 1.5 and 2.5 inches, with a few locations around 3 inches. Snowfall amounts in the Southern Sierra and Tehachapi Mountains were measured in feet, with the heaviest snowfalls reaching around 10 feet of new snow.

Winter 2009–10 was much colder than average for the U.S., and it delivered a string of record-breaking snowstorms that began on the winter solstice. (This was touted by some as proof that global climate change was a myth.) The snow and cold didn’t linger far into the spring, however. By the end of April 2010, North American snow cover had retreated to the lowest extent since satellite records began in 1967.

However, conditions were very different in the Southern Sierra. Here a snowy winter and a cool spring contributed to a snowpack that lasted later than normal. This was followed by an extended period of high water that lasted from the spring into the summer.

The winter of 2009–10 was a moderate- to strong-El Niño event. Because of this, Hawaii experienced severe drought, but Sequoia and Kings Canyon National Parks had a moderately wet year. Cedar Grove experienced the highest water levels since the 1997 flood. Flow in the South Fork of the Kings River peaked on June 6–8, 2010. That high-water event brought ponding/flooding and lapping of water along roads at various locations throughout Kings Canyon, including the Motor Nature Trail, the North Side Road, and the main highway.

A large (50± inch dbh) ponderosa pine fell perpendicular to the flow of the South Fork Kings, creating an obstruction in the river just east of the Zumwalt Meadow parking lot. The resulting obstruction elevated the water level upstream, creating ponding along both sides of the main highway. Water was nearly up to the shoulder on both sides of the main road immediately upstream of the Zumwalt Meadow parking lot. Immediately adjacent to the root wad where the tree once stood, an eddy developed in the river, causing scouring of the river bank. About three days after the failure, that eddy location increased in size moving toward the parking lot. Scouring/erosion of the bank slowed by the end of June, but remained vulnerable.

On June 6, 2010, a small debris jam formed in Lewis Creek upstream of the main highway bridge. This resulted in diverting water flow outside the primary channel into the overflow channel. The 16-inch highway culvert for the overflow channel could not carry the resulting flow, so Lewis Creek overflowed the highway for a couple of days and caused some minor road damage. There was concern that heavy equipment might have to be used to push Lewis Creek back into its primary channel. (A bulldozer had been used for that purpose on Memorial Day Weekend, 1983.) However, the debris jam broke on the evening of June 6–7, redirecting all water flow back into the primary channel.

This debris jam was presumably caused in large part by the July 15, 2008 flood which resulted in a huge debris flow that deposited a large amount of sediment and debris on the Lewis Creek delta, created a log jam in the primary channel, and generally destabilized that channel.

This channel-shifting happened again from about June 23–30, 2011. In the fall of 2011, the national parks’ road crew installed two more culverts (an 18 inch and a 24 inch) to carry the flow that results when Lewis Creek moves into the overflow channel. Where Lewis Creek emerges from its canyon, there is very little impediment to keeping it from changing course from the primary channel to the overflow channel. Should the main flow change to the overflow channel, a bridge would be required to handle the flow under the highway. Otherwise, the creek would have to be pushed back into its primary channel as was done after the May 1983 debris flow.
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Flow in the East Fork Kaweah peaked about June 6–7, 2010. The river was very close to the bottom of the Disney Bridge. Monarch Creek jumped its banks for a few days, but resulted in no road damage (photograph on file in the national parks). This appeared to be due at least in part to debris resulting from an avalanche. (Monarch Creek would do exactly the same thing in June 2011.)

The mainstem of the Kaweah crested in Three Rivers (the TRR gaging station) on June 6, 2010 at 5,129 cfs. Total inflow into Lake Kaweah during the snowmelt period peaked on June 6 when the combined flow from all forks of the Kaweah reached 7,546 cfs. It had been years since the Kaweah had been this high during snowmelt at either of those reporting stations.

The Southern Sierra stayed unusually wet throughout the summer of 2010. Even at the end of August, meadows throughout both Sequoia and Kings Canyon National Parks continued to be reported as quite wet.

Contrary to long-term forecasts that called for the winter of 2010–11 to be a relatively dry La Niña-pattern winter, December 2010 turned out to be quite wet in the San Joaquin Valley and Southern California. Precipitation records were broken and flooding occurred. Although moderate La Niña events are often dry in our area, strong La Niña conditions are generally wet. The winter of 2010–11 developed into one of the strongest La Niña events ever.

The December 2010 flood was marked in the Tulare Lake Basin by two pulses of moisture; the first came on December 16–20 followed by a smaller one on December 28–29. The first pulse was very intense, and the media quickly dubbed it the “Wallop”.

Virtually the entire state was affected by that first pulse, from coastal cities to the Central Valley, the Sierra, and the southern deserts. Large snowfalls occurred in the Sierra. Very heavy rainfalls hit Southern California.

The stormy weather began hitting the northern part of the state late on Thursday, December 16, and the southern areas on Friday after a large storm front moving out of the Gulf of Alaska met with moist subtropical air coming across the Pacific Ocean.

Writing in the Visalia Times-Delta, Bill Tweed described the uncommon combination of events that brought so much precipitation to the Tulare Lake Basin. First, a powerful low-pressure zone stalled off the Oregon/Washington coast and remained stationary for several days. Low-pressure areas circulate internally in a counter-clockwise manner, and since we were geographically at the bottom of the clock, this set up and held in place a strong southwesterly flow aimed at Central California.

The second critical factor in our big rain was that this southwesterly flow into Central California tapped into a huge mass of very moist tropical air from the central Pacific Ocean beyond Hawaii. The effect was to turn on a strong stream of water-saturated air, aim it directly at the Tulare Lake Basin, and then hold it in place for several days.

Finally, and this was also significant, the flow of wet air arrived from a direction that pushed it straight at the Sierra and forced the moisture to rise over the mountains at just the right angle. Put another way, our mountains run from northwest to southeast, and the wind in this storm blew at the mountains from the southwest, a perfect fit for maximum precipitation.

Many valley floor weather stations recorded 3 or 4 inches of rain during the period that began Friday, December 17, and ended Monday morning, December 20. (The storm continued through the 22nd in the southern part of the San Joaquin Valley and through the 23rd in Southern California.)

The foothills received about twice as much rain as the valley floor. Three Rivers received about 8.5 inches of rain during the storm event.

Sunday, December 19, was the date of the Christmas Bird Count in the national parks. The participants in that count were all willing to testify to the intensity of the downpour.

On December 19, flooding, landslides, rock falls, and debris flows forced the national parks to close the inbound lane of the Generals Highway at the Ash Mountain Entrance Station. To oversimplify the situation, rocks and debris were falling on the road faster than the road crews could remove them. Therefore, the entrance station
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was closed as a safety precaution to avoid potential accidents. In addition, the road between the two national parks had to be closed that day with more than 10 feet of snow on the roadway.

On December 21, two soil scientists from the U.S. Department of Agriculture, Kerry Arroues and Phil Smith, documented about 15 small landslides that had occurred on the Generals Highway for several miles below Hospital Rock. They concluded that most were probably related to the angle of repose. Kerry and Phil had planned to work their way farther uphill to where the big problem areas had been, but the road was still blocked by landslides and mass wasting events of various sorts.

The USACE automatic weather station in Giant Forest at 6,600 feet elevation received a total of over 15 inches of moisture during the December 17–20 storm event. Satellite data from an automatic sensor south of Mineral King at Wet Meadows (elevation 8,900 feet) also reported 15 inches, while another automatic station above Mineral King recorded almost 19 inches of precipitation.

Above 7,000 feet elevation, almost all of the precipitation fell as snow. At least 200 inches (nearly 17 feet) of snow was received at the automatic snow sensor at 9,600 feet elevation at Farewell Gap. That sensor abruptly flat-lined on December 19, apparently because it was hit by an avalanche. (When spring came, the tower was not to be seen.)

That represented an amazing amount of precipitation. In just three days, upwards of 1.5 feet of water had fallen in the Sierra above 8,000 feet elevation. For comparison, the total average annual rainfall in Visalia is about 0.75 foot.

Because the moisture flow was tropical in origin, the snowline remained above 7,000 feet elevation during most of the storm, and a great deal of foothills and middle elevation precipitation ran off into streams and rivers.

Considering the huge amount of moisture delivered by that storm event, there was relatively little runoff in the rivers of the Tulare Lake Basin. Because nearly all of the precipitation that fell above 7,000 feet elevation was captured as snow, a heavy runoff event did not occur. Had the snowline been one or two thousand feet higher, as had been forecast, mountain rivers would have risen as they did in similar storm events in 1955 and 1966.

At Three Rivers (the TRR gaging station), the mainstem of the Kaweah River had been flowing at about 229 cfs. The river began rising just before midnight on December 17. It climbed at a dramatic rate in the middle of the day on Sunday, December 19, peaking at 3:00 p.m. at 15,831 cfs, bringing the river to near flood stage in the town.

Carole Combs recalled that the North Fork of the Kaweah peaked at about the same time as the mainstem. She said that the North Fork was terrifying. Their driveway flooded so that they couldn’t have escaped if they had wanted to, which she did.

Terminus Dam did the job that it was designed to do, catching the floodwaters of the Kaweah River. Valerie McKay said that Lake Kaweah rose 40 feet in 48 hours during that storm event. The USACE staff was kept busy moving stuff out of the way of the rapidly encroaching waters.

Because Terminus Dam caught the entire flood on the Kaweah, Visalia was expecting only minor flooding impacts from that storm. However, that is not exactly how events played out.

The biggest challenge Visalia officials faced during the December 17–20 storm event was overflowing ponding basins. Basins built to accept stormwater filled to capacity and spilled over, flooding nearby areas. That flooding prompted officials to call a local emergency on the evening of December 19. Areas that were flooded by overflowing ponding basins were:

- A park at Pinkham Street and Mary Avenue, near the Annie R. Mitchell Elementary School in southeast Visalia
- Pinkham Street at Cherry Avenue in southeast Visalia
- Constitution Park near Tulare Avenue and Akers
- Mooney Boulevard and Cameron Avenue in south Visalia
- Sierra Village in west Visalia
- Walnut at Roeben Road in west Visalia

Those were comparatively minor overflow situations, relatively easily dealt with. However, there was one overflowing ponding basin that proved much more consequential.
Highway 198 is below grade through much of Visalia. During rainstorms, water collects in the low spots, and Caltrans’ pumps lift that water out and transfer it to city ponding basins to prevent highway flooding. Those pumps worked as designed, but the rain was so intense that the ponding basin at Linwood Street and Mineral King Avenue began overflowing about 3 a.m. on December 20, threatening adjacent homes and businesses. At the city’s request, Caltrans turned off their pumps to protect those properties.

With the pumps turned off on that section of highway, water eventually started flooding the lanes (photograph on file in the national parks). As a result, the freeway was closed between Akers and Demaree starting at 10:30 a.m. on December 20. Later that day, city crews were able to fix the problem by digging a channel to divert water from the overflowing ponding basin into the Persian Ditch. Caltrans then used two pumps to drain the highway and reopened it to traffic at 10:00 p.m. that night.

The December 17–20 weekend storm caused only minor problems in Fresno and Fresno County. The brunt of the storm was felt in the southern end of the valley. As detailed in Table 102, Visalia and Bakersfield received almost twice as much moisture as Fresno did. In the Central Valley, the storm event lasted from December 17–20. In Southern California, the storm continued through December 23.

Bakersfield shattered several precipitation records during this storm event. Among those were:
- December 18 — record rainfall for the date of 1.37 inches. The old record was 0.30 inch, set in 1921.
- December 19 — record rainfall for the date of 1.53 inches. The old record was 0.48 inch, set in 1984. This was also the wettest day on record for December at Bakersfield. The previous wettest day in December was December 27, 1936 with 1.02 inches of rain.

The December 17–20 storm caused numerous, mostly minor, flooding problems across the valley, including:
- Several Fresno streets flooded. A portion of Palm Avenue near Clinton Avenue flooded. Jameson Avenue, south of Church Avenue, flooded. The northbound lane of Reed Avenue, north of Floral, flooded.
- Several streets in and around Visalia flooded. Road 64 at Avenue 308 was closed due to flooding.
- Highway 180 was closed about 7 miles west of the junction with Highway 63 for several hours due to a mud or rock slide.
- Many of the ponding basins in Tulare filled or overflowed, including Live Oak Park, Del Lago Park, and the Mission Oak High School ponding basin.
- Several streets and roads in the Tulare area flooded. The westbound lanes of Prosperity Avenue between Laspina Street and Mooney Boulevard flooded. Laspina south of Prosperity flooded. San Joaquin Avenue was closed between J Street and I Street. Highway 137 (the Tulare-Lindsay Highway) flooded west of Road 168.
- Several streets and roads in the Porterville area flooded. Highway 190 flooded at Bourbon Drive and at Westwood Street. Lots of water flooded Highway 190 at Road 284. The Eagle Mountain parking lot at Avenue 136 and Westwood Street flooded with vehicles almost submerged.
- About 30 roads in Tulare County were closed by flooding, and 1,000 acres were inundated, including farmland planted in wheat or barley.
- Ten people were evacuated on December 20 from three homes in Weldon (near Lake Isabella) due to creek flooding.
- Extensive areas of farmland in the Lamont area (southeast of Bakersfield) flooded, possibly due to the failure of a dike.
- The California Highway Patrol reported flooding, rocks, and mud on various foothill and mountain roads into the Sierra.
- Highway 59 between Merced and Los Banos was closed when Mariposa Creek overflowed it. That highway didn’t reopen until December 23. Mariposa Creek overflowed it again on December 30.

Deer Creek flooded in south Tulare County. Among the roads it flooded were:
- A segment of Avenue 56 near Road 88 between Earlimart and Alpaugh
- Several segments of Highway 43, just west of Road 88

Deer Creek flooded part of Pixley NWR in December 2010 or January 2011; Nick Stanley said that the flood came on all at once and was very loud when it blew out the wall of the creek bank. It flooded the portion of the Pixley NWR immediately west of Road 88.

In McFarland, high water in Poso Creek caused the evacuation of about 2,000 people on December 20. Between 400 and 500 homes were in danger of flooding. Santa Fe Railway crews worked to keep that creek free of debris, helping to ensure that it didn’t overflow.
The storm brought very heavy rains to Southern California. Some locations received more than 12 inches of rain during the December 17–23 event. A number of rainfall records were broken. For example, 3.45 inches of rain fell in Pasadena on December 19, shattering the old record of 1.5 inches set on the same date in 1987. Los Angeles received 70% of its annual rainfall in just seven days. It was the most rainfall from one storm event since 2005.1737

The storm also brought very heavy rain to the Southern California deserts. The normally dry Mojave River flooded portions of the Apple Valley / Victorville area, peaking at 17 feet deep on December 21 (truly “ginormous” for those who had never seen that river in its magnificence).

The rain in the Mojave Desert was so intense that the resulting flooding wasn’t restricted to the riverbeds. On the evening of December 22, Shauna Austin encountered a flash flood flowing across seemingly open desert, flooding U.S. Highway 395 a few miles north of Adelanto. The water was so deep that passenger cars that tried to push through it were being swamped.

Table 102 gives the total precipitation for that storm event for selected reporting stations.

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Nature Point near Bass Lake</td>
<td>18.56</td>
</tr>
<tr>
<td>Wishon Dam (near Shaver Lake)</td>
<td>18.89</td>
</tr>
<tr>
<td>Fresno</td>
<td>2.47</td>
</tr>
<tr>
<td>Visalia</td>
<td>4.49</td>
</tr>
<tr>
<td>Three Rivers</td>
<td>8.5</td>
</tr>
<tr>
<td>Tulare</td>
<td>2.8</td>
</tr>
<tr>
<td>Bakersfield</td>
<td>4.02</td>
</tr>
<tr>
<td>Camp Nelson</td>
<td>18.6</td>
</tr>
<tr>
<td>Wofford Heights (Isabella Lake)</td>
<td>15.78</td>
</tr>
<tr>
<td>Near Crestline (west of Lake Arrowhead)</td>
<td>26.16</td>
</tr>
<tr>
<td>Tanbark Flats north of Pomona</td>
<td>19.22</td>
</tr>
</tbody>
</table>

December 2010 still had one more pulse of moisture in store for the Tulare Lake Basin. As detailed in Table 103, a fast-moving storm swept through the region on December 28–29.1738

<table>
<thead>
<tr>
<th>Reporting Station</th>
<th>Storm Total (inches of rain)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Fresno</td>
<td>1.54</td>
</tr>
<tr>
<td>Visalia</td>
<td>1.24</td>
</tr>
<tr>
<td>Three Rivers</td>
<td>2.63</td>
</tr>
</tbody>
</table>

The brunt of this storm’s effects was felt in Tulare County. A section of the Mineral King Road ¼ mile above the Hammond Fire Station collapsed on the morning of December 29 due to erosion and undermining. That road was closed for several hours until the roadbed could be rebuilt. Anne Birkholz recalled that she couldn’t get to work on the 29th because of flooding on several small tributaries along the South Fork of the Kaweah.

The community of Seville (northwest of Woodlake) was hit particularly hard by flooding on the night of December 28. By the next morning, the surrounding area was described as looking like the Nile River. At least five homes, a business, and the Stone Corral School in Seville were flooded. Rushing water covered Road 156 just south of Avenue 385 in Seville.1739, 1740, 1741

Yokohl Creek at Highway 198 was flowing at more than 1700 cfs during the storm. Yokohl Creek broke its levee on December 29, flooding hundreds of acres of orange groves and causing the closure of Avenue 304 south of Woodlake. Efforts to shore up the levee failed. The sight of the overflowing creek caused people to stop and stare. It is not too uncommon for Yokohl Creek to flow over the road at that location, but this particular flood caused it to cover the road to an unusual depth. Yokohl Creek crosses Highway 245 (Road 204) about 1½ miles north of Highway 198. Just west of that point, Yokohl Creek flows into the Consolidated Peoples Ditch. This is near the Lower Kaweah River but well below McKay’s Point.
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This area of unincorporated Tulare County is generally located between Woodlake, Lemon Cove, and Exeter.

According to two reports, a levee at Mehrten Creek also broke on December 29 and flooded hundreds of acres of oranges and nearby homes and businesses. John Kirkpatrick said that was probably immediately downstream of Highway 198, right after Mehrten Creek passes over the top of the Foothills Ditch. A home and a business flooded at Road 220 and Avenue 304 in 2 to 4 feet of water, and 300 acres of orange orchards were flooded. Mt. Whitney Pest Control on Road 220 was flooded. Highway 245 (south of Woodlake) was closed due to flooding from Avenues 304 to 312. Lort Drive (Avenue 312) in the same area was also closed due to flooding. It’s hard to believe Mehrten Creek could do all that. Mehrten Creek also flooded in February 1969 and sometime in 1983. There have surely been other floods on this stream, but these are the only ones that we have records of.

Rocky Hill Drive near Exeter was closed because of flooding.

Lewis Creek overflowed in Lindsay.

A lot of woody debris came down Cottonwood Creek during the flood. After the flood, emergency crews had to use heavy equipment to clear the debris off of the Highway 63 bridge over Cottonwood Creek, about 5 miles north of Visalia.

The wet winter of 2010–2011 caused a number of debris flows in the Kings River Special Management Area along the Garnet Dike road. There was both a December 2010 and a March 2011 event along that road. Some of those debris flows rivaled the 1937 event in the Big Creek Basin.

December 2010 was one of the wettest on record in the Tulare Lake Basin, in the Southern Sierra, and in Southern California. The December 28–29 storm pushed several communities into record-setting territory:

- Fresno received a total of 5.92 inches for the month, making it the second-wettest December in that city’s history. The wettest was December 1955 with 6.73 inches.
- Three Rivers received more than 13.5 inches for the month, making it one of the wettest Decembers on record for that community.
- It was the wettest December in Visalia’s history. The previous record had been 6.06, set in December 1955.
- The total precipitation for Bakersfield for the month was 5.82 inches — nearly eight times the average amount (0.76 inches) for December. That was the most rain recorded in any month since record-keeping began in 1889, and broke the record set in February 1998 during a very strong El Niño.

The winter of 2010–2011 was a La Niña season, and the long-range forecast had been for less-than-average precipitation. However, as of December 30, the season total for Three Rivers was 18.52 inches of rainfall, well above average.

At the end of December 2010, the Southern Section Sierra snowpack was reported at a phenomenal 284% of average for the date. Three of the snow sensors in the Kern River Basin — Pascoes, Tunnel Guard Station, and Casa Vieja Meadows — had recorded more snow in three months than the average amount for the entire six-month season.

To put the amount of snow in perspective, there was substantially more snow at Pascoes at the end of December 2010 than there was at the beginning of January during either the big El Niño events of 1982–83 or 1997–98.

Mammoth Mountain received 209 inches of snow in the month of December, the greatest for that month since record-keeping began in the winter of 1969–70. (The previous December high was 139.8 inches set in 1971.) Mammoth Mountain received a total of 668.5 inches of snow during the winter of 2010–11, breaking the season-total record of 578.5 inches set during the winter of 2005–06.
The snowpack continued to build and lasted through the spring. As shown in Table 104, the snowpack in the Tulare Lake Basin at the end of April 2011 was well above average.\textsuperscript{1751}

<table>
<thead>
<tr>
<th>Watershed</th>
<th>Predicted Runoff</th>
<th>% of average (1956–2005)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kings</td>
<td>2,050,000</td>
<td>167%</td>
</tr>
<tr>
<td>Kaweah</td>
<td>490,000</td>
<td>171%</td>
</tr>
<tr>
<td>Tule</td>
<td>115,000</td>
<td>180%</td>
</tr>
<tr>
<td>Kern</td>
<td>860,000</td>
<td>187%</td>
</tr>
<tr>
<td>Total</td>
<td>3,515,000</td>
<td>173%</td>
</tr>
</tbody>
</table>

As detailed in Table 105, the winter of 2010–11 broke the snowfall record at Lodgepole that had been set in the \textit{winter of 1951–52}. Snowfall in the \textit{winter of 1905–06} was even bigger than this, but that was before a weather station had been established at Lodgepole.

<table>
<thead>
<tr>
<th>Month</th>
<th>Snowfall (inches of snow)</th>
</tr>
</thead>
<tbody>
<tr>
<td>September 2010</td>
<td>0.0</td>
</tr>
<tr>
<td>October 2011</td>
<td>4.0</td>
</tr>
<tr>
<td>November 2010</td>
<td>52.5</td>
</tr>
<tr>
<td>December 2010</td>
<td>116.7</td>
</tr>
<tr>
<td>January 2011</td>
<td>44.7</td>
</tr>
<tr>
<td>February 2011</td>
<td>89.0</td>
</tr>
<tr>
<td>March 2011</td>
<td>127.4</td>
</tr>
<tr>
<td>April 2011</td>
<td>23.5</td>
</tr>
<tr>
<td>May 2011</td>
<td>17.0</td>
</tr>
<tr>
<td>June 2011</td>
<td>1.8</td>
</tr>
<tr>
<td>Total</td>
<td>476.6</td>
</tr>
</tbody>
</table>

2010 didn’t just set records in the Sierra. The Global Historical Climatology Network announced that 2010 was the wettest year that the world has seen since at least 1900. The La Niña conditions that brought so much precipitation to the Tulare Lake Basin and Southern California in December 2010 were also responsible for catastrophic flooding in Australia that month. All in all, it was a bang-up way to close out the year.

The large amount of snow caused havoc on some of the national parks’ trails. The parks’ trail crews expect to spend the first part of each spring clearing (logging) trees that have fallen during the preceding winter. However, significantly more trees came down in parts of the parks during the winter of 2010–11 than average.

Over 1,000 downed trees had to be cleared from trails in the Kaweah and Kern River Basins in Sequoia National Park; that is roughly five times the average. The areas that were most affected were:

- Giant Forest/Wolverton in the Kaweah River Basin
- High Sierra Trail to Buck Creek in the Kaweah River Basin
- Redwood Meadow/Cliff Creek in the Kaweah River Basin
- Tar Gap/Hockett Plateau in the Kaweah River Basin
- Lower Kern Canyon in the Kern River Basin (this area was particularly hard hit)
- Chagoopa Plateau in the Kern River Basin

Many of the above trees were tree-top failures. Tree-top failures are typically caused by a combination of heavy snow loads and wind.

There were also many up-rooted trees. Up-rooted trees are an indication of ground saturation. Based on his experience, Tyler Johnson, Sequoia National Park’s trails supervisor, thinks that these events seem to occur when winter precipitation exceeds about 150% of average. Tyler recalled that up-rooted trees were also a significant problem on the parks’ trails in 2006. In that year, the May 1 snowpack for the Kern was 152% of the long-term (1956–2005) average.\textsuperscript{1752} As shown in Table 104, precipitation in the Kern in the winter of 2010–11 was 187% of average, which helps to explain why the fallen tree count was so much higher.
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Kings Canyon National Park also experienced about five times the average number of trees falling across trails during the winter of 2010–11. Some trails had many trees down, scattered along the entire trail (Roaring River Basin, Woods and Bubbs Creek Basins, and South Fork Kings below Upper Paradise Valley). Some trails had about the average amount (Middle Fork Kings River Basin and Monarch Divide). Some had an average amount interspersed with large avalanches that had a lot of trees (San Joaquin River Basin). Upper Basin between Pinchot and Mather had no trees down at all — although there are typically one or two there.

One very noteworthy event was that the Middle Fork of the Kings Trail was closed to stock travel into the first half of September 2011 due to a large avalanche snow deposit that had not melted. That trail was closed just below the 7,200 foot elevation level. That was approximately mid-way between Grouse Meadow and Simpson Meadow in the Devils Washbowl area. When last reported in late September, there was still a large snow patch there, but the trail had melted out and was passable. Based on the experience of the Kings Canyon trail crew, this is the first time since at least the mid-1960s that a snow patch has lasted until late September at such a low elevation. Partly this was because the big winter of 2010–11 created the conditions necessary for a big avalanche. However, equally important was the cool spring and summer of 2011 that allowed that snow to persist into the fall. It’s tempting to think that conditions such as this haven’t existed since at least the winter of 1951–52.

The Mineral King Road was unusually late melting out in June 2011 due to the heavy winter of 2010–11.

On September 30, 2011, two adjoined giant sequoias failed along the Trail of 100 Giants in Giant Sequoia National Monument. One of those trees was 17 feet in diameter and 300 feet tall. Upon investigation, a forest pathologist found no rot in either tree. A suspected primary cause of the failure was lingering wet soil due to the winter of 2010–11.

In the Sierra, the majority of the snowpack usually melts in May so that there is little snow remaining by June 1. However, in 2011 the snowpack lingered well into June. This was the result of an above-average amount of snowfall during the winter, followed by an exceptionally cool spring, which helped keep the snow in place much long than normal. As of the first of June, the amount of snow still on the Sierra was nearly six times greater than average.1753

The snowpack at the Central Sierra Snow Laboratory monitoring site near Donner Pass lasted until June 30, 2011. That was the latest date for melt-off observed at that site since record-keeping began in 1946. It tied the record set in 1967, another big winter in the Tahoe area. (Older Southern Pacific Railroad records suggest that this might have been the latest melt-off dating back to 1879. However, those measurements were taken at a slightly different location and were not recorded in a rigorous manner, so no reliable conclusion can be drawn.) In any case, the June 30 date was a rather astounding five weeks later than the average May 23 date for melt-off at this location.

Because of the near-record amount of snowfall, there was a large amount of runoff. Since the spring was cool, that runoff didn’t result in peak flooding events, just large amounts of water delivered to the valley floor. Lake Kaweah reached peak storage (714.83 feet elevation, equivalent to 185,264 acre-feet) on July 7, 2011.1754

As a measure of the size of the runoff, the plan for the operation of Pine Flat Reservoir was to end the irrigation season with a full reservoir.1755 Normally Pine Flat would be drawn down to low-pool by then.

Flooding in 2011 occurred in July. It was caused by a series of individual storm events, but could be thought of as one event that occurred in multiple locations.

On July 28, southeast winds aloft began to bring mid-level moisture from northern Mexico and the Desert Southwest. Isolated thunderstorms developed over the Sierra crest around Kings Canyon and points just to the north and east. The surge of monsoonal moisture continued through the end of the month, with the strongest thunderstorm activity on July 30 and the morning of the 31st. The storm system extended south as far as Edwards Air Force Base in Kern County. A few thunderstorms over the Southern Sierra had rain rates of an inch or more in an hour.1756

Thunderstorms generated by this storm system deluged the Rock Creek Basin in the national parks from noon on July 29 through the evening of July 30. It was a severe and sustained event with flash floods on the afternoon of both July 29 and 30. We know about this storm event because the Rock Creek wilderness ranger station was staffed in 2011 by ranger Dave Alexander and Elizabeth Curry, a volunteer in the parks (VIP).
July 29 was one of Dave’s scheduled days off. (Wilderness rangers are always on call in case of an emergency.) Dave and Elizabeth were near the ranger station on July 29 because it was cloudy and threatening to rain.

It began to rain heavily around noon on July 29 with periods of thunder and lightning. (There was only light hail that day; storms the next day would bring heavy hail.) After a few hours, it cleared somewhat. At about 3:00 p.m., Rock Creek, which runs just in front of the ranger station, turned from clear to muddy and rose rapidly over the next two hours even though it was only raining lightly (multiple pictures and video on file in the national parks, also see cover photograph of this document).

Concerned for the safety of park visitors, Dave and Elizabeth hiked down to the commonly used camp area where the Pacific Crest Trail (PCT) crosses Rock Creek. On the way there, it was obvious from the rumble of unseen boulders being swept along and large, unearthed logs floating by, that Rock Creek had become impassible. The meadow below the ranger station was flooded, and the ranger station trail was under water.

Dave and Elizabeth contacted one commercial group of pack-supported hikers who decided to stay on the ranger station side of Rock Creek. (That group spent the following two nights there without being able to safely cross. They decided to forgo their trip to Mt. Whitney and returned the way that they had come.) There was also a group of commercial packers that was stuck on the opposite side of Rock Creek. It was impossible to shout over the roar of the creek, but they indicated with hand gestures that they had decided to wait overnight and would try to cross the following morning.

Early the next morning, July 30, Rock Creek had receded somewhat from the high of the previous afternoon, but it was beginning to rain again. The commercial pack group from the opposite side was able to cross the creek downstream where it split around a small island, but the water was above the horses’ bellies, and the packers were very relieved that they were able to get across without incident.

Throughout the early afternoon, there were a series of thunderstorms that covered the Rock Creek Basin with an inch of hail and heavy rain. The area around the ranger station was completely covered with hail; it looked like it had snowed. Mt. Langley, at the head of the watershed, was left blanketed with hail and snow (multiple photographs on file in the national parks). Rock Creek rose considerably higher than the previous day. It again became impassible to stock and hiking parties. The creek was so loud that Dave and Elizabeth had to yell to be heard when talking near the ranger station.

The rains ended on the evening of July 30, and Rock Creek receded to near-normal levels by late on the afternoon of July 31.

The flash flood did significant damage to the trails in the Rock Creek Basin. It swept away most of the upstream log crossings, depositing some of them in a large snag near the regular crossing below the ranger station. The section of the Rock Creek trail from Soldier Lakes to the Rock Creek ranger station was particularly hard hit. Along that section, Rock Creek jumped its banks and seriously eroded the trail, leaving it obscured, difficult to follow, and impassable to stock. The flood also caused considerable erosion of the trail near the Army Creek crossing (east end of the Rock Creek Basin, near Lower Soldier Lake). This left the trail too deep to walk in, so hikers created new trails parallel to the old trail.

Managing the logjam of stock and hiking parties unable to cross Rock Creek due to the flooding was a dangerous situation that required a great deal of intervention by Dave and Elizabeth. The heavy downpours and cold temperatures caught a number of hikers unprepared. Some who were suffering from near-hypothermia came to the ranger station to get warm and dry out. Two of them used the station as shelter for the night.

The flooding was so intense that it made significant changes to the channel structure of the creek. Most noticeable, a sand and log dam was created below the ranger station which split Rock Creek and formed what appeared to be a permanent additional channel through the meadow.

A somewhat similar storm event occurred in August 2001. In that event, an intense thunderstorm struck a large portion of the Rock Creek Basin, causing Rock Creek to flash flood and send a large quantity of water out onto the Kern valley floor about two miles north of Kern Hot Spring. The 2011 flood may have been a larger flood, but Erik Frenzel (national park meadow monitor) reported that it did not put any debris onto the High Sierra Trail. We don’t know why these two floods differed in this way. Tony Caprio speculated that the 2001 flood might have cleared accumulated material from the stream channel, and there had been insufficient time to accumulate a similar quantity of material before the 2011 flood.
The 2011 Rock Creek storm and flood lasted for nearly three days, from July 29 through the morning of July 31. On the second day of that event, July 30, a number of isolated strong thunderstorms developed at various points elsewhere in the Southern Sierra and in the desert region of Kern County. The July 30 storm events that we know about included:

- Mono Hot Springs (30 miles northeast of Shaver Lake in Fresno County) received 0.89 inches of heavy rain from a thunderstorm at 1:00 p.m.
- Whiterock Creek (5 miles northeast of Tehachapi Pass in Kern County) received 1.44 inches of rainfall from a heavy thunderstorm at 4:30 p.m.
- A RAWS automated weather station located 10.5 miles south of Onyx in Kern County received 0.56 inches of rain in only 33 minutes (a rate of 1 inch per hour) from a heavy thunderstorm at 4:00 p.m.
- Several storms caused small-scale debris flows that flowed onto mountain highways. One near Johnsondale in Tulare County caused damage to Salmon Creek Highway and Mountain Highway 99.
- A severe thunderstorm moved over Edwards Air Force Base in Kern County during the afternoon.

Another of the July 30 thunderstorms occurred in the Cedar Grove / Canyon View area of Kings Canyon National Park. This was some 25 miles northwest of the Rock Creek Basin in Sequoia National Park. The Cedar Grove storm began at about 4:00 p.m. and lasted only an hour or so. The 24-hour rainfall total was 1¾ inches. A number of small to moderate debris flows occurred toward the latter part of the storm. This storm struck in the same area where the October 2008 Cedar Bluffs prescribed burn had occurred. A small part of the storm extended into the area where the 2010 Sheep Fire occurred.

Whenever a debris flow occurs after a fire, there is a tendency to assume that the debris flow was caused primarily by the fire. That is, to assume the fire was necessary to create the conditions for the debris flow to occur. See for example when the 2008 Lewis Creek debris flow was initially attributed to a fire that had occurred three years earlier.

That was the case with the 2011 Canyon View debris flows; they were initially attributed to the Cedar Bluffs burn. However, investigation by Tony Caprio, the national parks’ fire ecologist, showed that the Cedar Bluffs fire was incidental to the Canyon View debris flows. It was a low-severity burn that had occurred three years prior to the debris flows. It had contributed little to creating the conditions necessary for those debris flows.

Damage from the Canyon View runoff, erosion, and debris flows included:

- The Heliport Road suffered one washout. Otherwise, the damage to that road was mainly a lot of material and debris deposited on top of the roadbed. A lot of that is now new road grade.
- There were 7 or 8 culverts overwhelmed and plugged on Highway 180 and the Heliport Road. One of the debris flows that came out onto Highway 180 was 4–18 inches deep and prevented vehicle traffic movement until it could be cleared. There were 7–8 significantly smaller debris flows that only blocked one lane of traffic.
- There were punctures to several vehicle tires as a result of driving over the debris flows.
- Lots of debris including rocks and trees was washed down the hillsides. The main waterline feeding Moraine and Canyon View Campgrounds runs along the side of the highway and is buried at least four feet deep. At one spot, about 100 yards before entering Canyon View Campground, this line crosses a natural gully. At that point, the line had originally been buried, but it had been exposed prior to the storm and had not been reburied. There was nothing to protect the line from the onrushing force of the floodwater and rock and wood debris. The combination of the debris and the rushing water broke the line, draining much of the water out of the main Cedar Grove water tank before the valve could be shut off. After the flash flood, the gully was about three times the size of what it was prior to the event. The drainage that caused the pipe break flows to the southeast of the Cedar Grove Bridge; it does not flow through Canyon View Campground.
- A moderate-sized debris flow came into Canyon View Campground. When the storm started, children along with their driver jumped into a school bus to seek shelter. Rumor (probably unfounded) was that the school bus was moved sideways by the debris flow. What we do know is that this debris flow made a really big mess of the campground with the decomposed granite (DG) and other debris. This debris flow came from a drainage just east of the one that broke the main waterline. In addition to the debris flow, this drainage caused erosion within the campground.
- One of the bridges on the foot/bike path to Canyon View Campground between the highway and the river was pretty well buried with some large rocks washed up to it.
- Sheep Creek deposited a moderate amount of sand and ash in the waterhead for the Sheep Creek water plant, filling it to the top of the dam. This sediment had to be removed by hand because no backhoe could access the area. This problem was compounded by the fact that the waterhead had not been cleaned out in...
recent years, so there were layers of old sediments that also needed to be removed. A total of 20 people worked for two days to flush all this material through the system.

- When the Sheep Creek water plant later tried to produce water from this surface intake, the turbidity was too high. That was because there was so much fine ash in the water from the 2010 Sheep Fire, and the ash couldn’t be filtered out.

The storm cell that caused the July 30 Cedar Grove / Canyon View event extended south into the Roaring River area. Cindy Wood was the Roaring River wilderness ranger and was caught up in that storm event. When the storm hit, she was riding out from her station leading a string of four pack animals. Intense rain and lightning continued from Ferguson Creek through the Sugarloaf Valley (about 1½ hours by horseback).

As Cindy rode, she listened on the park radio to all of the happenings in Cedar Grove. She kept riding through the storm because stopping under a tree was not really an option due to the lightning. The rain was so intense that Cindy was soaked through in five minutes, even with a coat and long rain slicker. The Sugarloaf Creek trail crossing was very high, but she made it across okay. It was just a very long, cold ride out to the trailhead.

Total flow for water year 2011 was 200% of the 1894–2014 average for the Kings, 203% for the Kaweah, 202% for the Tule, and 202% for the Kern. The combined runoff of the four rivers in the Tulare Lake Basin during water year 2011 was 5,910,342 acre-feet, 201% of average. This was nearly as large as the runoff in 1998. This may have allowed some recharge of the groundwater aquifer.

2012–15+ Drought
As described in the section on Megadroughts since the Little Ice Age, California’s 2012–15+ drought is part of a longer-term megadrought across most of the Western U.S. since 2000. Because we are on the edge of that huge drought system, we tend to only be aware of it when it reaches out to encompass our area.  

According to the U.S. Drought Monitor, about 56% of the contiguous U.S. experienced moderate to exceptional drought at the end of June 2012. The last time drought was this extensive was in December 1956 when about 58% was in moderate to extreme drought based on the Palmer Drought Index.

The San Joaquin River Basin has only been affected by that megadrought for 12 of the years that it has been active: 2000–04, 2007–09, and 2012–15 (based on the San Joaquin Valley Water Year Index and/or total runoff for our four major rivers). From our perspective, we tend to think of those events as individual droughts of relatively average duration instead of being part of the larger megadrought.

Table 106 shows how the San Joaquin Valley Water Year Index categorized the drought years in that basin. It also shows total runoff for the four major rivers in our basin combined during this drought. Conditions for 2015 are projected. Average flow over water years 2012–15 is forecast to be just 34% of the 1894–2014 average.

<table>
<thead>
<tr>
<th>Drought</th>
<th>San Joaquin River Basin</th>
<th>Tulare Lake Basin</th>
</tr>
</thead>
<tbody>
<tr>
<td>1999–2004</td>
<td>Above normal</td>
<td>2,103,599</td>
</tr>
<tr>
<td>2000</td>
<td>Above normal</td>
<td>2,453,729</td>
</tr>
<tr>
<td>2001</td>
<td>Dry</td>
<td>1,703,125</td>
</tr>
<tr>
<td>2002</td>
<td>Dry</td>
<td>1,881,093</td>
</tr>
<tr>
<td>2003</td>
<td>Below normal</td>
<td>2,557,853</td>
</tr>
<tr>
<td>2004</td>
<td>Dry</td>
<td>1,747,866</td>
</tr>
<tr>
<td>2005</td>
<td>Wet</td>
<td>4,435,498</td>
</tr>
<tr>
<td>2006</td>
<td>Wet</td>
<td>4,883,910</td>
</tr>
<tr>
<td>2007–2009</td>
<td>Critically dry</td>
<td>1,156,074</td>
</tr>
<tr>
<td>2008</td>
<td>Critically dry</td>
<td>2,122,906</td>
</tr>
<tr>
<td>2009</td>
<td>Below normal</td>
<td>2,177,485</td>
</tr>
<tr>
<td>2010</td>
<td>Above normal</td>
<td>3,563,153</td>
</tr>
<tr>
<td>2011</td>
<td>Wet</td>
<td>5,910,342</td>
</tr>
<tr>
<td>2012–2015+</td>
<td>Dry</td>
<td>1,551,604</td>
</tr>
<tr>
<td>2013</td>
<td>Critically dry</td>
<td>1,091,371</td>
</tr>
<tr>
<td>2014</td>
<td>Critically dry</td>
<td>830,549</td>
</tr>
<tr>
<td>2015</td>
<td>Critically dry</td>
<td>546,000</td>
</tr>
</tbody>
</table>
Two guys, Shawn Forry and Justin Lichter, took advantage of the drought to make the first wintertime through-hike of the 2650-mile-long Pacific Crest Trail. They left the Canadian border on October 21, 2014 and arrived at the Mexican border on March 1, 2015, just 132 days later. They encountered less than three feet of snow in the upper basins of the High Sierra. The relative lack of snow had its advantages, but it also complicated travel in many ways. Stretches where they would have liked to ski faster were covered with exposed boulders and brush.

The following narrative is informed in large part by Daniel Swain’s very informative California Weather Blog. 1769

Floods and Droughts in the Tulare Lake Basin
Specific Floods and Droughts

Water years 2013, 2014, and 2015 were very dry. Tulare Lake Basin has experienced a number of very dry years during historical times. For a comparison of such dry years, see Table 23 and Table 24. Table 23 compares the ten water years since 1894 that have experienced the runoff levels. It puts the two most severe years of the 2012–15+ drought (2014 and 2013) in context of those other very dry years. As shown in Table 23, 2015 is projected to have the lowest runoff in historic times, 22% less than the previous record set in 1977.

As Bill Tweed observed in his May 9, 2014 Visalia Times-Delta column, the low precipitation in 2014 was not unprecedented. 1761 What makes the 2012–15+ drought so severe is that it has consisted of four years (so far) of well-below-average precipitation coupled with well-above-average temperatures. So first to the sound-bites:

- As shown in Table 106, average runoff for our four major rivers during the 2012–15+ drought was much less than in the 1999–04 and 2007–09 droughts (34% of average versus 71% and 62%).
- Water year 2012 was drier than average on a statewide basis, in the San Joaquin River Basin, and in the Tulare Lake Basin. As shown in Table 106, total runoff in the Tulare Lake Basin was only 53% of the 1894–2014 average.
- Statewide, water year 2014 was the third-driest water year since record-keeping began in 1895. Only water year 1924 (driest) and 1977 (third-driest) were drier. 1762 The state received less than 60% of average precipitation in water year 2014. 1763 Water year 2015 will almost certainly be drier than any of those years.
- As shown in Table 23, 2015 is projected to be driest water year in the Tulare Lake Basin since record-keeping began in 1894, 2014 was the fifth driest water year, and 2013 was the tenth driest water year.
- Since water year 2013 began in October 2012, it captured the storms that came in October and November of that year. Calendar year 2013 omits that wet period. Statewide, precipitation during 2013 was by far the driest calendar year since national precipitation record-keeping began in 1895. California received only 7.38 inches of precipitation in calendar year 2013. That was 2.42 inches below the previous record dry year of 1898, and 15.13 inches below average. 1764 Possibly it was the driest calendar year since California’s statehood.
- Although the two years 2013 and 2014 were very dry, they were not record-setting; they were not the driest 2-year period we have experienced. The years 1976–77 were the driest 2-year period in the history of California. Based on the combined runoff of the four major rivers in the Tulare Lake Basin (Kings, Kaweah, Tule, and Kern), 1976–77 were also the driest two years in the Tulare Lake Basin since record-keeping began in 1894. The total runoff for those years was 13% less than the total for 2013–14. However, if the current forecast holds, the pair of years 2014–15 will be about 18% less than the record set in 1976–77.
- Water years 2012–14 was the driest 3-year period for statewide precipitation since record-keeping began in 1895. 1765, 1766 (The next five driest 3-year periods were all different combinations of years from during the 1918–34 drought.) Based on dendrochronology, water years 1929–31 were the driest 3-year period on the upper San Joaquin at the inflow to Millerton Lake between 983–2012. The average annual combined runoff of the four major rivers in the Tulare Lake Basin for the 3-year period 2013–15 is projected to be 820,902, 34% less than the previous record set in 1929–31 (1,237,573).
- The weather year is measured from July 1 through June 30; it is very similar to the water year (October 1 through September 30). The weather year is often used for reporting precipitation totals. The 3-year-period July 1, 2011 through June 30, 2014 was the worst 3-year drought in Fresno since record-keeping began in that city in 1878. It broke the 3-year record set in 1931–34, the previous worst dry spell in that city’s records. 1767
- Water year 2014 was critically dry statewide. However, calendar year 2014 was only slightly dry compared to the previous 120 years (1895–2014). 1768 That is because Northern and Southern California (but not the center of the state) received a couple heavy rains during December 2014; that month was part of calendar year 2014, but water year 2015.
- Water year 2015 was unusual in part because relatively narrow atmospheric rivers provided copious precipitation to areas from the Bay Area north to the Shasta drainage in December 2014 and February 2015, and yet the drought continued because of record-setting lack of snowfall in the Sierra.
Droughts aren’t always obvious when they first start. Statewide, the 2012–13 rainy season had an extremely wet start. Parts of the state got a lot of moisture in November 2012. However, no additional significant storms occurred during December 2012 — nor during January–June 2013.

In fact, January–June 2013 was the driest start to the calendar year on record for the state of California since record-keeping began in 1895. Some parts of the state saw virtually no precipitation at all during this period, which made for an especially stark contrast with the extremely wet conditions experienced just a few months earlier.

How did this drastic change occur so quickly? The second half of the water year 2013 saw the development of an extraordinarily persistent region of high pressure in the northeastern Pacific that forced the mid-latitude storm track well to the north of its typical position and prevented winter storms from reaching California. Those storms had to go somewhere; the Midwest had a brutal winter because the storm track was displaced.

Because of its persistence, that region of high pressure became known as the Ridiculously Resilient Ridge (or RRR). The RRR is a persistent region of unusually high atmospheric pressure in the middle levels of the atmosphere centered over the far northeastern Pacific that exists over many consecutive months. While the RRR did become less prominent during the summer months of 2013, it returned with even greater intensity by the fall of 2013. As a result, most winter storms in 2013–14 missed Oregon and Washington as well as California.

During the winter of 2014–15, a high-amplitude atmospheric flow pattern once again developed over the Eastern Pacific and North America, deflecting the Pacific storm track north of its typical cool-season position along the West Coast and allowing repeated intrusions of extremely cold Arctic air to invade the Midwest and Eastern Seaboard.

Parts of New England experienced a recurring nightmare of extreme Arctic outbreaks of cold air and record-breaking snow accumulations. This overall setup — with a big Western ridge and a deep Eastern trough (a region of relatively low atmospheric pressure) — has been a common feature of recent winters in North America. It was a major contributor to the 2012–15+ drought in California.

Large positive ridging anomalies developed along the West Coast during January 2015. By early February, a stronger signal had emerged, and it was apparent that the RRR had returned. January and February turned out to be dry and extremely warm on the West Coast. The start of the 2015 Iditarod Trail Sled Dog Race had to be moved 300 miles north to Fairbanks on March 9 because Anchorage had no snow. The drought and unseasonably warm temperatures along the West Coast resulted from the same RRR that brought extremely cold temperatures and snowy conditions to the Eastern Seaboard.

The severe storms along the Eastern Seaboard during winter 2014–15 were touted by some as proof that global climate change was a myth. Senator James Inhofe, chair of the Senate’s Environment and Public Works Committee, tossed a snowball in the Senate chamber after a February 27, 2015 storm to emphasize his long-held belief that climate change is “the greatest hoax ever perpetrated on the American people.”

This ridging pattern has preceded some of the worst West Coast droughts, including 1934 and 1976.

Part of the reason for the 2012–15+ drought was that the frequency of atmospheric rivers dwindled during this drought. There were three atmospheric river events during the 2013–14 winter, about half of what would normally occur. The lack of these atmospheric rivers was arguably a reflection of the presence of the RRR.

While climatologists attributed the drought to such natural factors, some people had other explanations. For example, California assemblywoman Shannon Grove suggested that the drought may be divine retribution for California providing women with access to abortions. She said that “now God has hold on California.”

The April 1, 2014 snow survey showed that the statewide average snowpack was only 25% of the long-term (1956–2005) average. This tied the record set in 1977 for the lowest April 1 snowpack of record. The April 1, 2014 snowpack in the Tulare Lake Basin was 28% of average. Dave Fox said that the road into the Mineral King Valley was completely melted out, all the way to the trailhead at the east end, by the middle of April 2014.

The April 1, 2015 snow survey found that the snowpack in the Central and Southern Sierra (and the statewide average) was only 6% of the long-term (1956–2005) average. This was the lowest level since record-keeping began in 1950. The previous low April 1 statewide snowpack was 25% of average, set in 1977 and 2014.
As a result of the record-breaking snowpack, water year 2015 is forecast to be the driest year on record for Tulare Lake Basin runoff (see Table 23). As shown in that table, the combined flow of our four major rivers for water year 2015 is forecast to be 22% lower than the previous record year of 1977. Average flow over water years 2012–15 is forecast to be just 34% of the 1894–2014 average.

As California’s long-term precipitation deficits increased during this drought, another dramatic trend became increasingly apparent: an extraordinary string of record-warm days, months, and multi-month periods. Calendar year 2014 was remarkably warm in California, the West, the contiguous U.S., and for the Earth as a whole. This fits within a context of a long-term warming trend that has been going on for several centuries and has been accelerating in recent decades (see the section of this document that describes Long-term Temperature Changes).

This record-shattering warmth had serious implications for the 2012–15+ drought, since warmer temperatures result in greater evapotranspiration. This meant that an even smaller fraction of the already near-record low precipitation was actually available to plants and ecosystems — and as runoff into rivers and streams.

According to the U.S. Drought Monitor, the entire state of California was in some stage of drought from April 2014 through January 27, 2015, then the level dropped to 98% through spring 2015. More than 58% of the state was in exceptional drought, the most severe level, from late July 2014 through late October 2014. Drought conditions eased somewhat during the winter. But by the end of March, 2015, 41% of the state was in exceptional drought and the amount was steadily growing.

Just how severe is the 2012–15+ drought? The combination of exceptional dryness and record warmth acted in combination to produce the most severe drought conditions experienced in California since at least 1934 (and probably longer). PDSI is a widely used indicator of long-term drought severity. See the section of this document that discusses Measurements of Drought on page 43.

Any PDSI value lower than -4 corresponds to extreme drought. As of summer 2014, a large fraction of California was experiencing literally chart-topping PDSI values less than -6. These values — both regionally and on a statewide average basis — were higher than at any other point since at least 1895, according to rankings from the NCDC. From those data, it is reasonable to argue that the 2012–15+ drought was already more intense than any 20th-century drought in California.

Dan Griffin and Kevin Anchukaitis used tree-ring records and other data to compare the 2012–15+ drought with previous droughts that have occurred over the last 1200 years in Central and Southern California. This was not a statewide assessment because it excluded Northern California.

Their study found that while three-year periods of persistent below-average soil moisture are not uncommon, the current event is the most severe drought in the last 1200 years, with single year (2014) and accumulated moisture deficits worse than any previous continuous span of dry years. Tree-ring chronologies reveal that precipitation during the drought were anomalously low but not outside the range of natural variability. The 2012–14 period is exceptionally severe in the context of at least the last millennium and was driven by reduced though not unprecedented precipitation and record-high temperatures.

The authors used blue oak tree-ring records to reconstruct the climate for the past 600 years, back to 1400 A.D. Then they used gridded PDSI reconstructions to infer drought severity for the time period before 1400. Besides the bristlecone pine record, most of those tree-ring collection sites are located outside of the state of California in central Oregon, the Great Basin, and Arizona.

Possibly those are the best data available to reflect the overall drought conditions for Central and Southern California. However, it’s worth keeping in mind that drought conditions vary dramatically for different parts of the state. The Tulare Lake Basin has experienced significantly different drought conditions than the San Joaquin River Basin, which has experience different conditions than the Sacramento River Basin or the Great Basin.

In 2014, Steve Voelker used the redwood data to do an assessment of the 2012–15+ drought for the entire state, including Northern California. He concluded that droughts as serious as the year 2014 have occurred across the state 15 times between 1085 and 1996 (the common interval for the best paleo records). This equates to a 60-year return interval. Droughts as serious as the three-year period from 2012–2014 have

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occurred 10 times, a 90-year return interval. If 2015 turns out to be nearly as dry as water year 2014 (which now seems a near certainty), then we would really be experiencing a millennial anomaly in Steve’s opinion.

There is obviously a significant difference between the blue oak reconstruction and the redwood reconstruction. As Nate Stephenson observed, this points out the need to be cautious when making grand conclusions from one-species chronologies and from areas that do not fully represent the entire state’s climatic envelope.

Another reason that made the 2012–15+ drought worst in terms of measurable harm than previous droughts is that we are demanding more of our water resources than ever before. Our population has never been larger. Prices for California crops have seldom been higher. And our expectations for preserving native fish and wildlife are as high as ever. Never before have we demanded so much of so little water.

The 2012–15+ drought wasn’t limited to North America. It was devastating in much of Central and South America. Guatemala declared a state of emergency in 16 of its 22 provinces in 2014. Crop losses were as high as 70% in some regions and 170,000 families lost almost all their crops. The crop losses were as high as 60% in El Salvador. The price of basic foods like beans quadrupled just between May and August 2014 in Nicaragua; thousands of cattle died that year. In August 2014, the Canal Authority said they might have to restrict ship traffic in the Panama Canal due to the lower levels in Panamanian lakes that feed the locks. That meant ships would have to lighten their loads so that they don’t require as much water to go through the locks. Load limitations in the Panama Canal were eventually imposed in August 2015.

This was the worst drought in the last 30 years in Bolivia; it resulted in 47,000 forest fires in that country just in the first half of 2014. Venezuela had to begin emergency water rationing. The drought triggered unrest in Colombia; some communities there said that they had not seen any rain for two years. It was the worst drought in 84 years in southern coastal Brazil. Sao Paulo, South America's largest city, had to impose severe water rationing. Chile also experienced a huge drought; fires burned about 230,000 acres in the 2014–15 fire season, nearly twice the average amount.

\[\text{Reductions in Water Supplies and Water Deliveries}\]

The drought was felt most severely in the southern part of the Tulare Lake Basin. Total flows for water year 2014 were 32% of the 1894–2014 average for the Kings, 23% for the Kaweah, 11% for the Tule, and 24% for the Kern.

There have been four water years on the Kings drier than 2014: 1977 (386,007 acre-feet, the driest year of record), 1924, 1931, and 1976. Flows for 2015 are forecast to be 341,000 acre-feet.

Flows on the Kaweah (99,481 acre-feet) were the second lowest since record-keeping began in 1894. Only water year 1977 (93,641 acre-feet) was lower. Flows for 2015 are forecast to be 83,700 acre-feet.

Flows on the Tule (14,550 acre-feet) were the lowest since record-keeping began in 1894. This broke the record set in 1977. Flows for 2015 are forecast to be 11,300 acre-feet.

Flows on the Kern (178,159 acre-feet) were also the lowest since record-keeping began in 1894. This broke the record set in 1931. Flows for 2015 are forecast to be 110,000 acre-feet.

The combined runoff of the four rivers in the Tulare Lake Basin during 2014 was 830,549 acre-feet. There have been three water years drier than 2014 since 1894: 1977 (696,572 acre-feet, the driest year of record), 1924, and 1931. Combined runoff of the four rivers in the Tulare Lake Basin during 2014 was 19% greater than in 1977. Looking further back in time, water years 1795 and 1580 were also drier than 2014. Runoff in water year 2015 is forecast to be just 546,000 acre-feet, lower than flows in 1795, 1924, 1931, or 1977 (see Table 23 and Figure 18). Only the year 1580 has been drier than 2015.

The SWRCB did everything it could to accommodate the needs of south-of-Delta water users. In early 2014, the board took emergency action to temporarily change the rules, reducing minimum flows through the Delta, allowing DWR and USBR to hold back more water in their upstream reservoirs. This action was taken under the belief that this action would not unreasonably harm the environment, particularly any threatened or endangered species. This maximized the water available for delivery to south-of-Delta users.

As it turned out, threatened Delta smelt crashed to their lowest level on record in 2014, and 95% of endangered winter-run Chinook salmon eggs and juveniles died upstream on the Sacramento River. Environmentalists have argued that the plight of these fish species was worsened by the 2014 SWRCB water management decisions.
At a hearing in Sacramento on February 19, 2015, Tom Howard, executive officer of the SWRCB, appeared to agree with that position. He said he was “mistaken” and “just wrong” when he concluded on January 31, 2014 that temporarily modifying the conditions of the SWP and CVP water rights permits and licenses to keep more water in reservoirs would not cause unreasonable harm to the environment and harm threatened fish.\textsuperscript{1786}

On January 31, 2014, DWR announced that it was dropping the SWP allocation to zero, seeking to preserve the remaining supplies of water in its reservoirs. The SWP delivers water to 29 local public water suppliers (who supply water to 25 million Californians and roughly 750,000 acres of irrigated farmland). Never before in the 54-year history of the SWP had DWR announced a zero allocation to all 29 public water suppliers that buy from the SWP. The only previous SWP 0% allocation was in 1991 for agriculture, but cities that year received 30% of requested allocations.\textsuperscript{1787} Thanks to rain and snowstorms in February and March 2014, DWR was later able to increase water contract allocations for SWP deliveries from 0 to 5% for both urban users and agriculture.\textsuperscript{1788} (See Table 10 on page 108)

Central Valley Project deliveries were also cut back drastically during the drought. In 2013, south-of-Delta farmers received 20\% of their contracted supply. Most agricultural districts were allocated 0\% of their contracted supply from the CVP in 2014. Municipalities and wildlife refuges also received deep cuts. That was the first time ever that agricultural districts received 0\% of their contracted supply.\textsuperscript{1789} This was also the first Friant Division contractors received a zero allocation of their Class 1 water (see Table 11).\textsuperscript{1790}

As with previous recent droughts, some south-of-Delta agricultural interests put much of the blame for the reduction in water exports on the Endangered Species Act. USBR maintained that there had been very minimal restrictions on pumping due to Endangered Species Act protections. Their position was that, by and large, there just was not enough water in the system to go around.\textsuperscript{1791}

With less available water, irrigation districts had to deliver water for much shorter periods. For example, water delivery from the Fresno Irrigation District (FID) typically lasts six months. But in 2014, they were only able to deliver water for six weeks, beginning June 1. That was the shortest it had been in 37 years. The last time FID was able to run water only six weeks was during the 1977 drought year.\textsuperscript{1792}

In 2015, FID was unable to deliver any surface water to the 6,500 farmers in the 250,000-acre district. It did plan to allocate 90,000 acre-feet to replenishing the groundwater aquifer. This is believed to be the first time in the district’s 95 years of existence — and the first time in the Fresno area’s 145-year history of canal irrigation — that no regular water deliveries will be provided to farms.\textsuperscript{1793}

In April 2014, the Kern Water Bank was proposing a plan to use water that had been stored in the water bank. Under that plan, eight pumps would be installed in the California Aqueduct (which was going to be dry anyway), and the water would be pumped north to water users in the northern part of the district. The state was to decide by the end of June 204 whether the project could proceed.\textsuperscript{1794} There is no indication that the state approved the plan, or that the plan was implemented.

Entities with access to remaining water in 2014 auctioned off their rights for over ten times the long-term average rate.\textsuperscript{1795}

On March 2, 2015, DWR announced that the projected SWP allocation for 2015 would be 20\% of the contracted supply, an increase from the 5\% allocation of 2014.\textsuperscript{1796} The state’s water pumps are in a different location than the federal pumps and were not plagued by the invasive water hyacinth plant, which allowed them to send more water south into San Luis Reservoir during the winter.\textsuperscript{1797} Getting that water into San Luis was fortunate. Otherwise, the state might not have been able to project a 20\% delivery. There wasn’t enough water north of the Delta in Oroville Dam to assure deliveries to the south.

CVP deliveries were projected to be essentially as low in 2015 as they were in 2014. Compounding the problem of limited water, federal export pumps in the south Sacramento–San Joaquin Delta had to be slowed to avoid clogging them with an invasive water hyacinth plant.\textsuperscript{1798} The initial allocation for south-of-Delta farmers was 0\% of their contracted supply. Municipalities and wildlife refuges also received deep cuts.\textsuperscript{1800}

Water Conservation Efforts
Governor Jerry Brown declared a Drought State of Emergency for California on January 17, 2014.\textsuperscript{1801} At that time, he urged the state’s residents to voluntarily reduce water consumption by 20\%. (The 2012–15+ period marked the second time a statewide proclamation of emergency has been issued for drought.\textsuperscript{1802} The 2007–09...
drought was California’s first drought for which a statewide proclamation of drought emergency was issued. Among other things, the governor’s proclamation also ordered local urban water suppliers to immediately implement their water shortage contingency plans.

The Tulare County Board of Supervisors proclaimed a local drought emergency on February 4, 2014.\textsuperscript{1803} Governor Brown issued an executive order on April 25 that called on Californians to take specific actions to avoid wasting water, including limiting outdoor watering of lawns and landscaping to no more than two times a week, and limiting car washing by patronizing local car washes that used recycled water.\textsuperscript{1804} Those calls by the governor for voluntary conservation proved to be inadequate. Data for May 2014 showed that statewide, residential and business water use went up slightly, not down.\textsuperscript{1805} Californians as a whole failed to conserve water during the worst drought in a generation, according to a report reviewed by the State Water Resources Control Board (SWRCB) at its July 15, 2014 meeting in Sacramento.

The report of urban water use described conservation by hydrologic region. Some regions used a good bit more water in May 2014 compared to their three-year average of the same month from 2011–13 while some reduced their use. The Tulare Lake Basin used the same amount as our past average, not managing to conserve any. Overall, urban water use in California rose 1%. That was a long way from the 20% conservation target Governor Brown set in his emergency drought proclamation in January 2014.

As a result of those findings and the ongoing drought, the SWRCB adopted temporary emergency conservation regulations on July 15, 2014. Under those temporary regulations, the SWRCB imposed four individual prohibitions on all Californians. Potable water could not be used for any of the following purposes:

1. Direct application of water to wash sidewalks and driveways.
2. Landscape irrigation that caused runoff to streets and gutters.
3. Washing a motor vehicle using a hose without a shut-off nozzle.
4. Using drinkable water in a decorative fountain unless it recirculated the water.

As part of those temporary regulations, the SWRCB required urban water suppliers to impose mandatory restrictions on outdoor watering if they had not already done so. Some urban water suppliers had the authority to enforce those regulations through fines; others did not. Therefore, the SWRCB gave those agencies an avenue to get that authority. The SWRCB took the unprecedented step of declaring the above four types of water waste a criminal infraction similar to a speeding violation and authorized local urban water suppliers to impose a maximum $500 per day fine for the above four prohibited types of water use.\textsuperscript{1806}

The water conservation effort as reported by the urban water suppliers gradually climbed to 11.6% by August 2014 (compared to usage in August 2013).\textsuperscript{1807} In December 2014, statewide residential and business water consumption was 22% less than use in December 2013.\textsuperscript{1808} This reduction in water use was possible because December 2014 was a relatively rainy month while December 2013 was a very dry month. December was the only month in 2014 that met the governor’s goal; August was the next best effort at 11.6%.

Most urban water suppliers are public organizations. However 113 of the suppliers are investor-owned and are regulated by the California Public Utilities Commission (CPUC). On February 27, 2015, the CPUC took action authorizing the investor-owned water utilities to take the same kind of actions as the public suppliers.\textsuperscript{1809, 1810}

Emergency regulations have a shelf life of only 270 days. The regulations adopted on July 15, 2014 would have expired on April 25, 2015. At its March 17, 2015 meeting, the SWRCB readopted and expanded those regulations. Under the expanded regulations, all Californians were prohibited from the same four actions as in the July 2014 regulations plus one new individual prohibition:\textsuperscript{1811}

5. Irrigating turf or ornamental landscapes during and 48 hours following measurable precipitation.

Commercial businesses were banned from the following uses of potable water:

1. Restaurants and other food service establishments can only serve water to customers on request.
2. Operators of hotels and motels must provide guests with the option of choosing not to have towels and linens laundered daily and prominently display notice of this option.

The temporary emergency regulations that were extended on March 17 also specified that:

- The regulations apply to all public urban water suppliers (regardless of size) and to those urban water suppliers that are investor-owned.
- The reporting requirements do not apply to small water suppliers that serve less than 3,000 connections.
Urban water suppliers are required to implement their water shortage contingency plans to a level that imposes mandatory outdoor irrigation restrictions; voluntary restrictions are not sufficient. Urban water suppliers must limit the number of days per week customers can irrigate outdoors. The limit must either be specified in the water shortage contingency plan; or if the plan contains no specific limit, irrigation is by default limited to no more than 2 days per week. The SWRCB specifically recognizes that a water shortage contingency plan can restrict outdoor irrigation to 3 days per week and still be in compliance.

Under the temporary regulations, all urban water suppliers were authorized to impose a maximum $500 per day fine on customers for the five prohibited types of water use. However, when the Associated Press investigated a sample of urban water suppliers, they found that most have been reluctant to crack down on violators. Warning letters are unusual, small fines are rare, and the $500 a day fine was virtually never wielded.

The temporary emergency regulations (the 5 individual and 2 commercial prohibitions) that were adopted by the SWRCB at its March 17 meeting were approved by the Office of Administrative Law on March 27, 2015. In May 2014, the SWRCB imposed widespread curtailments of diversions in large portions of the Sacramento, San Joaquin, Eel, and Russian River watersheds for holders of junior water rights (those issued by the state after 1914). The order required the affected water suppliers and users to stop diversions from all streams. This was the first time such an action had been taken since the drought of 1977. It appears that no curtailment notices were sent to holders of senior water rights during calendar year 2014. Likewise, it appears that no curtailment notices were issued for the Tulare Lake Basin in calendar year 2014.

On April 1, 2015, Governor Brown issued the fourth in a series of executive orders dealing with the ongoing drought. This one directed the SWRCB to devise a plan that would impose restrictions to achieve a statewide 25% reduction in potable urban water usage through February 28, 2016. Those restrictions would require urban water suppliers to reduce usage as compared to the amount used in 2013. Areas with high per capita use would be required to achieve proportionally greater reductions than those with low use. It was initially estimated that this 25% reduction in potable urban water usage would result in savings of approximately 1.3 million acre-feet of water over the nine-month period June 2015 – February 2016.

Brown’s order was the first-ever mandatory statewide water reduction ordered by a California governor. The order also set up a program to replace 50 million square feet of lawns with drought-tolerant landscaping; created a statewide consumer rebate to replace old appliances with energy- and water-efficient ones; and required big water users like campuses, golf courses and cemeteries to make significant cuts in water use.

The order added two new requirements that would apply to all Californians. In addition to those added on March 27, 2015, it include the prohibition of irrigation with potable water of:

6. ornamental turf on public street medians
7. outside of newly constructed homes and buildings not delivered by drip irrigation or microspray.

The order also imposed new enforcement mechanisms on big agricultural water users, who will be required to report usage to state regulators.

The SWRCB formulated a draft framework to achieve the governor’s 25% conservation goal of April 7, 2015. After soliciting and considering public comment, the SWRCB adopted a final regulation at its May 5 meeting. The specific prohibitions in the emergency regulation took effect immediately upon approval by the Office of Administrative Law on May 18, 2015.

The new regulation assigns the 411 or so urban water suppliers who serve more than 3,000 customers to a tier of water reduction based upon three months of summer residential gallons-per-capita-per-day data. Cities with relatively high per capita water use like Bakersfield and Beverly Hills are required to reduce their water use by 36% while cities with low per capita water use have to reduce their water use by lower percentages.

Smaller water suppliers that serve fewer than 3,000 connections and commercial, industrial, and institutional users that are not served by a water supplier are required to either reduce their water use by 25% or restrict outdoor irrigation to no more than two days per week.

**Impacts of Drought**

By July 2014, every county in the state had been designated by the federal government as a primary natural disaster area due to the drought.
As water tables dropped, many wells (both domestic and agricultural) ran dry. Those that could, re-drilled their wells or drilled new wells.\textsuperscript{1822} The J.G. Boswell Co. drilled five ultra-deep 2,500-foot wells in Kern County in 2013.\textsuperscript{1823} Tulare County led the state in new wells being drilled; in the first nine months of 2014, there were at least 1500 permits to dig new wells in that county.\textsuperscript{1824}

Tulare County had just over 1,000 private domestic wells that were reported to have gone dry as of April 2015. That is 60\% of all the residential wells that had gone dry in the entire state. Most of those were in East Porterville.\textsuperscript{1825, 1826, 1827, 1828} State officials said that there could be hundreds more dry residential wells, with many rural well-owners not knowing whom to contact.

A 2014 study from UC Davis led by Richard Howitt updated estimates of the drought's effects on Central Valley farm production, presented new data on the state's coastal and southern farm areas, and forecast the drought's economic fallout through 2016.\textsuperscript{1829} Among the study's primary findings:

- The drought was responsible for the greatest water loss ever seen in California agriculture, with river water for Central Valley farms reduced by roughly one-third.
- Groundwater pumping was expected to replace most river water losses, with some areas more than doubling their pumping rate over the previous year. More than 80\% of this replacement pumping occurred in the San Joaquin River and Tulare Lake Basins.

Key findings of the UC Davis study on the drought's effects, just in 2014:

- Direct costs to agriculture totaled $1.5 billion (revenue losses of $1 billion and $0.5 billion in additional pumping costs). This net revenue loss was about 3\% of the state's total agricultural value.
- The total statewide economic cost of the 2014 drought was $2.2 billion.
- The loss of 17,100 seasonal and part-time jobs related to agriculture represented 3.8\% of farm unemployment.
- 428,000 acres, or 5\%, of irrigated cropland was going out of production in the Central Valley, Central Coast and Southern California due to the drought.
- The Central Valley was hardest hit, particularly the Tulare Lake Basin, with projected losses of $810 million, or 2.3\%, in crop revenue; $203 million in dairy and livestock value; and $453 million in additional well-pumping costs.
- Overdraft of groundwater was expected to cause additional wells in the Tulare Lake Basin to run dry if the drought continued.
- The drought was likely to continue through 2015, regardless of El Niño conditions.
- Consumer food prices will be largely unaffected. Higher prices at the grocery store of high-value California crops like nuts, wine grapes and dairy foods are driven more by market demand than by the drought.

Due to the prolonged drought, the 2014 wildfire season began much earlier than usual with Northern California experiencing relatively big fires in January 2014.\textsuperscript{1830} That was followed by the destructive San Diego-area fires driven by unusually strong Santa Ana winds in May 2014. The fire season continued with unusually intensely-burning fires in Northern and Central California despite the relative lack of extreme weather conditions usually required to sustain such extreme fire behavior. Several special fuel and fire behavior advisories were issued in the summer of 2014 for much of California due to record-low fuel moistures and potentially explosive wildfire behavior in the coming months.\textsuperscript{1831}

\textbf{Ash Mountain Pasture}

The national parks’ Ash Mountain Pasture has experienced seven multi-year drought since it began being used in 1921:

1. \textbf{1918–34}, a 17-year-long megadrought
2. \textbf{1947–50}
3. \textbf{1959–61}
4. \textbf{1976–77}, the driest two years in the state’s history prior to 2014–15
5. \textbf{1987–92}
6. \textbf{2007–09}
7. \textbf{2012–15+}

The parks’ stock are used to support wilderness operations during the summer. But in the winter, the stock have to be brought back to lower elevation pasture. In the early years, the parks’ stock were kept on the Ash Mountain Pasture during the winter. However, beginning in about 1970, the national parks began sending most or all of their stock outside the parks during the winter whenever they could.
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From about 1975 to the present, most of that winter pasture has been on the Horse Pasture Unit at the Pixley National Wildlife Refuge. Pixley manages their Horse Pasture Unit for a particular conservation objective: maintaining average residual dry matter of 800 pounds per acre at the beginning of summer. This is done for the benefit of two threatened and endangered species that live on this pasture: the blunt-nosed lizard and the Tipton kangaroo rat.

This partnership between the national parks’ stock and the refuge’s conservation objective works reasonably well except when the refuge experiences a dry winter. The Mediterranean grasses on the Horse Pasture Unit don’t grow when the rains don’t come. Therefore, when there is a dry winter on the refuge, the Horse Pasture Unit can meet its residual dry matter objective without any grazing.

When that happens, Pixley’s managers inform the national parks that they have to discontinue putting their stock on that unit. As a result, the parks typically keep most or all of their stock on the Ash Mountain Pasture during severe droughts. Refer to the section of this document that describes the 1987–92 drought for a description of how the Ash Mountain Pasture has come through earlier droughts.

The winter of 2012–13 presented special challenges in stock management. In that winter, the drought was felt much more strongly in the San Joaquin Valley than in the national parks.

Largely unrelated to the drought, the national parks had 10 head of stock die on the Pixley NWR in June 2012 when a water system failed. As a result, Pixley NWR managers told the national parks that they could not put stock on the refuge in the fall of 2012 because the parks’ Memorandum of Understanding (MOU) had expired, and because of concerns about animal welfare. Therefore the parks kept all 89 head of stock on the Ash Mountain Pasture during the fall and mid-winter. The MOU and animal welfare issues were resolved during the winter. However, the lack of mid-winter rains in the San Joaquin Valley in 2012–13 kept the forage low on the Pixley Horse Pasture.

Refuge managers eventually allowed the national parks to put 20 head on the Horse Pasture Unit beginning mid-February 2013; the parks had to keep the remaining 69 head on the Ash Mountain Pasture. The east half of the Lower Pasture and the Tunnel Rock Pasture units were not used for grazing during the winter of 2012–13, at least through mid-March 2013. No monitoring of residual biomass was conducted. The parks had to feed a very large amount of supplemental weed-free hay during the winter of 2012–13.

In 2014, the national parks had 91 head of stock (29 horses and 62 mules). Pixley NWR managers did not allow the parks to put any stock on the Horse Pasture Unit during the winter of 2013–14. (That was the first time that had happened since the winter of 1990–91.) As a result, all of the parks’ stock had to be kept on the Ash Mountain Pasture throughout the winter.

Because of the continuing drought, there was little natural forage on the Ash Mountain Pasture when the stock came back to it in the fall of 2013. The parks had to feed a very large amount of supplemental weed-free hay during the winter of 2013–14 and continuing through 2014. That hay was distributed along the jeep road that leads into the Big Oak Flat unit of the Upper Pasture. The stock had access to the entire Upper Pasture, but they stayed primarily on the Big Oak Flat unit; they stayed near where the feed was. The last of the stock were removed from the Upper Pasture by the end of May 2014.

About 5–8 head of stock were kept at the corrals and on the west half of the Lower Pasture throughout the summer of 2014. The east half of the Lower Pasture and the Tunnel Rock Pasture units were not used for grazing during the winter of 2013–14. No monitoring of residual biomass was conducted on the Ash Mountain Pasture during 2013–14.

Water sources on the pasture

**Springs in the Big Oak Flat unit of the Upper Pasture:**

1. **Bathtub Spring.** This spring is located on the left side of the jeep road where it crosses the seasonal stream just before the old blasting school site. There is a trough at this location that is fed off of the spring. In May 2014, this spring was redeveloped in order keep it flowing: all the vegetation was removed, and the trough was replaced. That is the first time this spring had been maintained since the 1970s. Tyler Johnson said that this spring kept running at a low rate throughout the summer of 2014. No stock were in the Big Oak Flat unit after May of that year, so it is difficult to estimate the actual production of this spring.

2. **A spring on the left side of the jeep road mid-way between Bathtub Spring and Wishbone Creek.** The concrete spring box at this site is pretty deteriorated. Greg Feltis said that this spring continued to run
through at least July of 2013. We don’t know how well this spring did in 2014. Possibly this is an error, and this is really a reference to Bathtub Spring.

3. **Powder Spring.** This is an undeveloped spring on the left side of the jeep road just as you get to Wishbone Creek. It is near where the old blasting school site was located. Tyler Johnson said that this spring continued running until August 2014 and remained wet throughout the summer.

4. **A significant spring downhill from the old blasting school site near the park boundary.** Bill Tweed said that this site gets heavily used by stock and tends to be pretty beat up. We don’t know how well it did in the 2012–15+ drought. National park staff seldom visit this site.

**Wishbone Creek.** Greg Feltis said that one of the branches of Wishbone Creek kept flowing throughout the summer of 2013 near where it crosses the Shepherd’s Saddle Road. This is not the branch of Wishbone Creek that the CCC developed in the 1930s with a spring box and roadside tank.

**Sycamore Creek spring box.** Despite the name, this is really a water catchment on the creek, about ¼ mile upstream from the Shepherd’s Saddle Road. In April 2014, this catchment received some maintenance: sediment was dug out of the box, and four or so of the smaller nearby trees were removed. The larger trees near the spring were left standing, even though the park recognized that these trees are consuming much of the water needed to keep water flowing to the spring box. This is the first time that this catchment has been maintained since 2003. Erik Meyer said that this section of the creek dried up in May 2014, earlier than it normally does. Sycamore Creek was also fenced off at the Shepherd’s Saddle Road to prevent stock from getting into the pool there. This creek began flowing again at Shepherd’s Saddle Road in about early November 2014.

**Indian Head Creek.** This creek drains a large portion of the pasture north and northeast of the Sycamore Corrals. Erik Meyer said that this creek dried up in early summer 2014, about when it normally does.

**Cricket Hollow.** We aren’t sure how well this water source did in the 2012–15+ drought. Erik Meyer’s impression was that this creek barely flowed during 2014. Tyler Johnson thought that stock might be fenced out of this water source.

**Alder Creek.** This is the primary water source for the national parks’ Ash Mountain development. No significant water conservation measures were required during the summer of 2014. A well in Alder Creek was brought on line in 2013–14 to supplement production from surface supplies. Al Damazio said that total production barely kept up with demand from mid-July through mid-September 2014.

**A spring in the brushy draw in Indian Gulch.** There is a tub in the creekbed that was formerly fed by this spring. Roy Lee Davis recalled that the spring was a reliable year-round source of water in most years but ran dry in the severe drought of 1976–77. This spring was still flowing in the late 1980s, but it may not have flowed since. The area around the spring has become overgrown with dense vegetation, including a large cottonwood and a smaller sycamore. Now the spring doesn’t flow even during years of average precipitation. It seems plausible that this was the natural condition in this gulch. If so, this spring may have been manmade rather than natural.

Most developed springs on the Ash Mountain pasture may have been manmade rather than natural. They appear to have been created to support stock use of the pasture.

**Vegetation Response in the Tulare County Foothills**

Phenology is the study of how seasonal changes in biological phenomena are correlated with climatic conditions. During the 2012–15+ drought, there was minimal formal monitoring to document how plant communities responded to the drought. The national parks in collaboration with USGS enlisted volunteer help in the summer of 2014 to take repeat photographs along the Generals Highway in the Kaweah River Basin to document expected increases in oak mortality.

In our area of the Southern Sierra, the foothills are dominated with blue oak woodland (or savannah) with California buckeye, interior live oak, and canyon live oak as typical tree associates in the blue oak woodland. The following represents qualitative observations of what appears to be extended drought impacts on blue oak woodlands. Observations were made primarily in the Kaweah and Tule River Basins. These consolidated observations of vegetation responses were made primarily by Melanie Baer-Keeley, Tony Caprio, Adrian Das, Athena Demetry, Sylvia Haultain, Ann Huber, Jon Keeley, and Nate Stephenson.

- In the foothills and conifer zone of the Kaweah River Basin, whiteleaf manzanitas, and interior live oaks, some over 100 years old, began dying in significant numbers at least as early as mid-December, 2013.
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- There was a substantial dieback of both interior live oak and whiteleaf manzanita in the Kaweah and Tule River Basins in 2014.
- Dave Parsons recalled that he saw nothing like the current die-back happen during the 1976-1977 drought. He and Phil Rundel were paying close attention to the foothills vegetation at that time.
- Noticeably more interior live oak in the blue oak woodlands than usual had full canopies of dead leaves in early spring 2014. Interior live oaks with green leaves often had a branch or more of dead leaves. We were unable to confirm a noticeable change in canyon live oaks. By spring 2015, mortality of interior live oaks seemed to be on the order of 30%-50%.
- Large mature whiteleaf manzanita shrubs began dying in the Middle Fork of the Kaweah foothills in 2014. We do not have active monitoring in place to quantify this (or an established baseline against which to compare mortality rates), but Sylvia Haultain said that the mortality was greater than anything she had seen in the Kaweah foothills in over 30 years. Jon Keeley estimated that about 25% of the manzanitas in his yard and his neighbor’s yard died in 2014. By spring 2015, mortality of whiteleaf manzanitas in the Kaweah River Basin seemed to be well over 50%.
- The Sequoia Riverlands Trust’s (SRT) Kaweah Oaks Preserve on the valley floor near Exeter has one of the last remaining valley oak woodlands in the Central Valley. Several large valley oaks split apart in 2014 without any sign of pest or pathogen attack. SRT staff suspected that the cause was water stress due to drought and lowering of the water table. That association seems plausible. There are only a few references in the literature to trees splitting due to drought; it isn’t clear what the mechanism for this might be. The western portion of the preserve and adjacent lands experienced an abrupt and striking increase in dieback and mortality of many valley oaks and sycamores beginning in about May 2015. Mortality was especially high in very large, older valley oaks and in older, larger sycamores. Ann Huber wondered if this could be due in part to larger trees needing more water to survive. Rob Hansen suggested that the preserve may be witnessing a transition to a sparser valley oak riparian forest that will lose its sycamore component. The sycamores at the preserve are already near the western extent of their range in the Kaweah River Basin.
- There was much earlier leaf drop in blue oaks in 2014 year and noticeably much reduced canopies earlier in the year. This pattern was patchy but widespread in some areas, and was not always explained by aspect.
- Blue oaks started growing new leaves in January 2015. By spring 2015, mortality of blue oaks seemed to be on the order of 5%-20% and varied dramatically across the landscape.
- Susan Mazer at UC Santa Barbara is the lead for the California Phenology Project. She has slides showing 2012–14 changes in blue oaks using the national parks’ Foothill Visitor Center phenological data that show 2014 driest and warmest, and significantly earlier flower onset, open flowers, and pollen release for blue oaks. This type of monitoring may lend itself to broad-scale implementation by landowners in the foothills. With a dedicated container for data storage (Nature’s Notebook), the data would be immediately accessible to researchers and managers.
- The heavy acorn drop of the blue oaks seemed markedly premature in 2014. Initial production in blue oaks seemed like a banner year, but immature, withered acorns fell beginning in August, as opposed to as usual in October and November. Mature acorns continued falling until November, but in lesser quantities.
- Nate Stephenson observed that a noticeable minority of blue oaks held on to green leaves throughout the winter of 2013–2014, on into spring when the new leaves flushed. The green leaves they held onto were a subset; that is, they did shed some leaves, just not all of them. Nate had never seen this happen before; it was very odd. He didn’t know how this might tie into the drought.

According to the California Acorn Survey, 2014 was a medium-to-poor year for acorn production, statewide. Overall, acorn production was the worst it had been in the state since 2003, and less than half the crop production of 2012 and 2013. Valley and blue oaks generally have better acorn crops in dry years than in average or wet years. The valley oak acorn crop was reasonably good in 2014. The blue oak acorn crop was fair to poor. Results were particularly poor for live oaks: 2014 was either the worst or next-to-worst year ever for coast live oaks, canyon live oaks, and interior live oaks.

So what conditions make for an outstanding acorn year? Relatively little is known concerning acorn production patterns of California’s oaks. One of the very few published studies is based on data collected at the University of California’s Hastings Reservation in coastal Central California. That study measured acorn production in five different species of oaks over 15 years.

As expected, acorn crops at Hastings varied widely in size from year to year. For example, blue oaks had an average of less than one acorn per tree in 1986 and over 60 per tree in both 1985 and 1987. More surprisingly, there was no significant correlation between the acorn crops of the different species with the exception of valley and blue oaks, two closely related species. In other words, a good year for coast live oaks is not necessarily a good year for any of the other species.
The environmental correlates of the annual differences in acorn crop size differ from species to species. For valley and blue oaks, the most important single factor is weather in April, the peak month for pollination, with acorn crops being heavier in years when mean April temperatures are warmer. This suggests that conditions favoring more efficient pollination are a principal factor in the annual variation in acorn production by those species.

For coast and canyon live oaks, mean acorn production was positively correlated with rainfall occurring one and (for canyon live oak) two years earlier. None of the weather factors tested was correlated with the mean annual crop size of California black oak. Winter rainfall in the same year as acorn production (the factor most commonly thought to determine acorn crop size) was not correlated positively with mean annual crop size of any of the species studied.

On top of the large annual variability, individuals of the same species varied drastically in their overall acorn productivity. Some individual trees seldom have a good crop of acorns, while others frequently have a good crop regardless of the environmental variables. That suggests that acorn production in most years might be correlated with site-specific factors such as soil nutrient and water availability. Preliminary results indicate that water availability to the root system of individual trees might be a critical factor for valley and blue oaks.

Foothill residents who put out hummingbird feeders went through much more hummingbird feed than normal in spring 2015, an indication of how little natural nectar sources were available. Tony Caprio was going through over one gallon per day at one point. Bill Tweed and Dave Graber each observed five species of hummingbirds (Anna’s, rufous, Costa’s, calliope, and black-chinned) in their yard in a single day.

**Conifer Response in the Southern Sierra**
The following represents qualitative and quantitative observations of extended drought impacts on conifers in the Southern Sierra. Bev Bulaon, a USFS entomologist, provided invaluable observations and context for what has been occurring throughout that wide area. Nick Ampersee, Tony Caprio, Nate Stephenson, and Tom Warner provided observations made primarily in Kings Canyon (aka South Fork of Kings River) and the Middle Fork of the Kaweah within the national parks.

Accelerated conifer mortality began in the Tulare Lake Basin (Hume Lake Ranger District and the southern districts of the Sequoia National Forest) in water year 2007 at the beginning of the 2007–09 drought. High levels of such mortality have been observed in each year since then. The one year of good precipitation during that period (water year 2011) did not significantly reduce already-high beetle populations.

Bark beetles became very active in the High Sierra Ranger District of the Sierra National Forest beginning in 2009. According to Forest Health Monitoring Aerial Detection Surveys (2008–2014), overall bark beetle-associated mortality has been continuous and intensifying in various locations throughout that national forest as the drought has persisted. Some of the highest counts on the Sierra National Forest were detected in 2009 and 2013.

Increased levels of conifer mortality in Kings Canyon probably began in 2012, but were first observed in spring 2013.

Kings Canyon is in the rain shadow of Park Ridge; it is relatively dry compared to many places where ponderosa pine grows in the Southern Sierra. Ponderosa pine generally grows where annual precipitation averages 30 inches. Precipitation in Kings Canyon for the period 1999–2014 averaged only 23 inches. This suggests that pines in Kings Canyon are in borderline drought conditions even during average precipitation conditions, making them particularly susceptible when a drought, especially a warm drought, occurs. This increases their susceptibility to successful bark beetle attack. It is surprising that increasing conifer mortality was not observed in Kings Canyon until at least five years after being observed in the adjacent Hume Lake Ranger District. Presumably either beetles were slow to arrive in Kings Canyon or there was less moisture stress in the canyon than in the Hume Lake Ranger District.

In the conifer zone of the Middle Fork of the Kaweah, mature sugar pines began dying in significant numbers in spring 2014. By February 2015, significant numbers of dead sugar pine, ponderosa pine, white fir, and incense cedar covered the landscape from Big Fern Springs to Crystal Cave and beyond. Since conifer mortality was not observed in the Middle Fork of the Kaweah until the winter of 2014–15, either beetles were slow to arrive in this drainage basin or there was less moisture stress than elsewhere in the Southern Sierra.
Giant sequoias began showing the effect of drought stress (die-back of foliage) in the summer of 2014 in the North Fork Kings Basin and in groves throughout the national parks.

The 2007–09 and 2012–15+ droughts severely stressed conifers throughout the Southern Sierra. Conifer mortality was particularly apparent in the lower montane zone (3000–6000 foot elevation), involving virtually all conifer species except giant sequoia.

Drought-related moisture stress predisposes white firs and pines to successful attack by bark beetles. The current episode of beetle-caused mortality is reminiscent of the 1918–34 drought. Superintendent’s reports from 1924, 1925, 1930, 1932, and 1934 have lots of focus on bark beetle attacks on pines.

Insects are opportunists and are simply responding to highly favorable conditions for expansion and growth. Healthy trees ordinarily produce abundant amounts of resin, which pitch out or eject attacking beetles. But, during a severe drought, stressed trees are less able to produce sufficient resin flow to resist attack. Not all the tree mortality in the Southern Sierra has been associated with insect attack. Some trees died solely from the drought without any insect attack. On the other end of the spectrum, some trees would have survived this drought if it were not for the double whammy of the drought coupled with the insect attack.

The high level of bark beetle-associated mortality was widespread throughout the Southern Sierra. Yosemite National Park south through the Tehachapis experienced mounting levels of beetle-kills during the drought. Bark beetle activity was mostly concentrated on pines, with higher levels of activity (and mortality) in the low elevation ranges of ponderosa pine. Western pine beetle (Dendroctonus brevicomis) was the most aggressive insect pest in ponderosa pine.

Natural stands and plantations continually lost large-diameter pines as beetles moved through and primarily selected bigger trees. (Bark beetles preferentially select larger trees because the phloem and bark thickness provides more protection for developing broods.) There were dramatic changes in many locations, particularly public campgrounds and private homesteads that had few but cherished high-value trees.

Pine plantations in the national forests were particularly hard hit by western pine beetle due to their high proportions of even-aged trees at high densities. Groups of up to 100 trees were attacked annually, 20–50 trees on average. Entire patches of older plantations that would have previously been regarded as lower risk — low basal area, minimal brush competition, and adequate spacing — were completely infested within a single year. Trees along the ridgeline or south-facing slopes were attacked first, but beetles then migrated into all areas where ponderosa pines were grouped.

Patches identified as beetle-cause mortality in the Sierra National Forest have been as large as 500 acres, with 3 attacked trees per acre. Losses of trees per acre have ranged from 2 per acre to as high as 10. Background mortality is considered less than 1 tree per acre annually.

Thousands of ponderosa pine in Kings Canyon have succumbed to the combined effects of drought and bark beetles (primarily western pine beetle) within and adjacent to the national park in Tulare and Fresno Counties.

Over 1,700 trees (of all conifer species) died within developed sites and adjacent roads in Kings Canyon within the national park during 2013–14 as a result of drought and bark beetles. Trees died singly or in group kills of up to over 100 trees. Over-stocking (high stand density) because of many decades of fire exclusion was a factor in some, but not all deaths.

In March 2014, Tony Caprio did a quick survey of mortality on a 35-acre valley floor burn unit prior to a prescribed fire. He observed 302 dead trees of four species; ponderosa pine (42% of the trees that were dead), incense cedar (30%), white fir (23%), and sugar pine (5%). Quite a few of the tree were the small size classes.

The presence of high densities of bark beetles in Kings Canyon represented a huge potential food source for predators. Nick Ampersee observed how woodpeckers responded to this opportunity. Hairy, downy, and white-headed woodpeckers were commonly seen feeding on the beetle larvae under the bark. Nick and Tony Caprio also saw multiple black-backed woodpeckers, and there was evidence of their presence throughout the valley as well as in the Grant Grove area. Everywhere there was a large group of dead ponderosa pines, the bark had been scraped off by what appeared to be black-backed woodpeckers.
Mortality in true firs, incense cedars, and Douglas-fir increased during the drought, but not to the same extent as in pines. Fir engraver bark beetle (*Scolytus spp.*) was the most common insect pest associated with mortality for white fir. Fir engraver bark beetles typically surge 2-3 years into an ongoing drought, and are anticipated to increase in 2015.

Incense cedar does not have a primary bark beetle. Based on the national parks’ fire effects monitoring plots (FMH data), incense cedar experienced about half as much mortality as pines. Sometimes dying incense cedar were associated with insects/pathogens, but sometimes not. This species that was mostly affected by drought.

The mortality of ponderosa pines, true firs, and incense cedars on the west side of the Sierra is expected during droughts or bark beetle outbreaks. From an ecological perspective, this mortality event is roughly equivalent to a hot, patchy fire. Bark beetles, under epidemic population levels, are not very selective thinning agents. They are a crude tool, but the result is more or less what land managers want to accomplish, especially in the face of global climate change. It significantly reduces stand density in a mosaic patchwork.

Drought and beetles also limit the expansion of conifer trees into lower elevation shrub and woodland areas. Trees grow a lot better when stands are not so dense. Fewer trees and greater diversity as the forest grows back could in the long run result in a healthier, more resilient forest better suited to withstand global climate change.

The loss of sugar pines is a different matter. Losses sustained by this species are outside the range of natural variability and are having serious consequences. Sugar pines have been declining for many years throughout California due to multiple factors. White pine blister rust, an exotic pathogen introduced nearly a century ago, has been very slowly decimating most five-needled white pines from western forests. Where populations of white pines are small or scattered, this gradual disappearance is barely noticeable but significant to stand diversity. Compounded with drought, all sizes and age classes are currently being killed by bark beetles. This accelerated loss of legacy-sized sugar pine affected cone production, stand diversity, and composition.

Bev Bulaon said that the impact of the drought on sugar pines is seriously overlooked. It is not just big trees going out; it is also small understory regeneration that is not noticed from aerial surveys. The loss of big seed trees, diversity, and the spread of invasive white pine blister rust is really a tragedy. There are areas in the Southern Sierra where sugar pines will probably not come back after this drought because the stand conditions are not conducive for regeneration. Sugar pines are dying from a combination of stresses: moisture stress, beetle attack, and white pine blister rust. However, the drought is not contributing to the spread of blister rust; drought conditions are probably inhibiting the spread of blister rust.

Mortality of lodgepole and Jeffrey pines in the higher elevations are not as observable, but large polygons with mortality have been detected in the Sierra National Forest in aerial detection surveys. Areas with 200-500 acre patches were noted with 1-3 trees recently dead per acre for both species. Jeffrey pine beetle (*Dendroctonus jeffreyi*) typically outbreak a few years into severe drought events. Lodgepole pine mortality due to mountain pine beetle (*Dendroctonus ponderosae*) in California does not typically reach epidemic status as in Rocky Mountain regions, but can be just as devastating and stand-altering.

We have conflicting mortality data for high-elevation pines (whitebark, limber, foxtail, and western white pine). Bev said that there were some spectacular losses of high elevation pines on the east side of the Sierra, even more so than on the west side. The ecological effect for these species was like a severe wildfire, where certain age/size classes have been wiped away. Presumably Bev located this mortality through aerial surveys, but we don’t have specific information on where she observed this. Jonny Nesmith guessed that Bev saw this mortality in areas like the June Mountains where there have been some relatively large beetle outbreaks in recent years.

The national parks have not observed any increased mortality of high-elevation pines through summer 2014. Nate Stephenson hiked the Cottonwood, Rock Creek, Crabtree, Wallace area in July 2014, and was struck by the lack of any apparent drought-induced tree death; all the high-elevation pines including lodgepole looked healthy. The national parks’ meadow monitors did not report increased mortality in 2014. The Sierra Nevada Network forest monitoring crews did not report any significant mortality in any of the whitebark or foxtail areas in summer 2014.

A 2014 survey of 4,321 mature sequoias in the national parks by Nick Ampere and Kate Cahill found 9.5% had lost 25-50% of their foliage and 1.5% had lost more than 50%. Nate Stephenson and Adrian Das of USGS analyzed this rich dataset. There appeared to be relatively substantial variation in foliage die-back among groves. Even Garfield Grove, a north-facing grove, had high average die-back. There was also substantial
variation in foliage die-back within groves. Even trees next to wet meadows were susceptible to die-back. This level of foliage die-back in giant sequoias appears to be unprecedented in the history of Sequoia and Kings Canyon National Parks. Significant foliage loss was observed in Mariposa Grove in Yosemite National Park in spring 2015; it had surely been present long before this.

Significant mortality of sequoia seedlings was first observed during the winter of 2014–15 in the vicinity of Huckleberry Meadow. Six mature sequoias that were partially girdled by fire a few years earlier died during the drought. All the dead trees were in relatively moist locations (versus drier uplands) which is interesting and not what we would have expected. To the best of the parks’ knowledge, no other mature sequoias have died during the 2012–15+ drought.

The USFS aerial Forest Health Protection Surveys detected a large increase in tree mortality in 2014, especially in the Central Coast and Southern Sierra. An aerial survey of more than 4.1 million acres of the Southern Sierra from Yosemite to the Tehachapis including the national parks was conducted on April 15-17, 2015. The purpose of these surveys is to detect and map recently dead or injured trees. Among the findings of this survey:

- Of the area surveyed, 835,000 acres (20%) had some level of mortality.
- There were an estimated 10,450,000 trees killed in the area surveyed.
- In general, mortality was quite severe in many pine species especially in ponderosa and pinyon at lower elevations and more southern areas.
- Along the foothills, mortality was often widespread and severe especially in ponderosa but also gray pine and likely blue and live oak. The survey was conducted too early in the season to detect oak mortality and dieback.
- On the Stanislaus NF, mortality was scattered in northern areas, but pockets of severe ponderosa and other pine mortality were seen in the southern low areas. Mortality roughly doubled since July 2014 in the areas of the Stanislaus that were resurveyed in April 2015.
- On the Sierra and Sequoia NF, pine mortality, mostly from western pine beetle, was common and severe almost everywhere at lower elevations. Estimated number of trees killed on these two forests together exceeded 5 million. Only about 300,000 trees were estimated killed in 2014 in the same area. This was a 17-fold increase in total mortality.
- Conifer mortality was scattered at higher elevations. The survey was conducted too early to detect the full extent of mortality levels.
- Southeastern portions of the Sequoia NF and wilderness areas further east were also flown, at times intense pinyon mortality was observed and widespread.
- On the Tehachapi Range and on private lands along the Sierra foothills, extensive areas of pine mortality were common. Large areas of oak mortality were also suspect.

Due to the severity and prolonged nature of the 2012–15+ drought, accelerated levels of conifer mortality will probably continue for some time after the drought ends. Beetle activity will persist since populations (and availability of susceptible trees) are very high at the moment. Beetles will continue to disperse to look for new hosts. Tree resiliency will take some time to rebuild — just like after fires. Forest trees will need several consecutive years of average or above-average precipitation to recover vigor and regain sufficient resistance against bark beetles.

The prescription to help forests survive severe droughts is the promotion of vigorous, healthy stands. That involves thinning dense stands so as to reduce resource competition. In some cases, it may also involve increasing stand diversity. Tony Caprio said that some of our forest types had low diversity in the past; that is their natural condition. A resilient forest does not necessarily have to be diverse.

The combination of very low precipitation and high temperatures in the 2012–15+ drought resulted in record low PDSI, creating unprecedented stress on vegetation (see Figure 12). Nate Stephenson thinks that it is possible, perhaps even probable, that the effects of the current drought on native vegetation may be unprecedented in at least the last century. Unfortunately, we have no way to prove that. But an informed opinion from Nate is a pretty good assessment of the situation.

This second edition generally ends with the fall of 2014, the end of water year 2014. Only limited data and projections are available to describe water year 2015. This document covers what we know about the floods and droughts that have occurred within the Tulare Lake Basin over the preceding 2,000 years or so. For a summary and conclusions regarding this material, see the Summary section of this document.
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